

**THERMOBAROMETRY OF ECLOGITE FROM THE
NORDØYANE ULTRA-HIGH PRESSURE DOMAIN,
WESTERN GNEISS REGION, NORWAY**

BENJAMIN MYRER
Benjamin.Myrer@dal.ca

Department of Earth and Environmental Sciences,
Dalhousie University, Halifax, NS

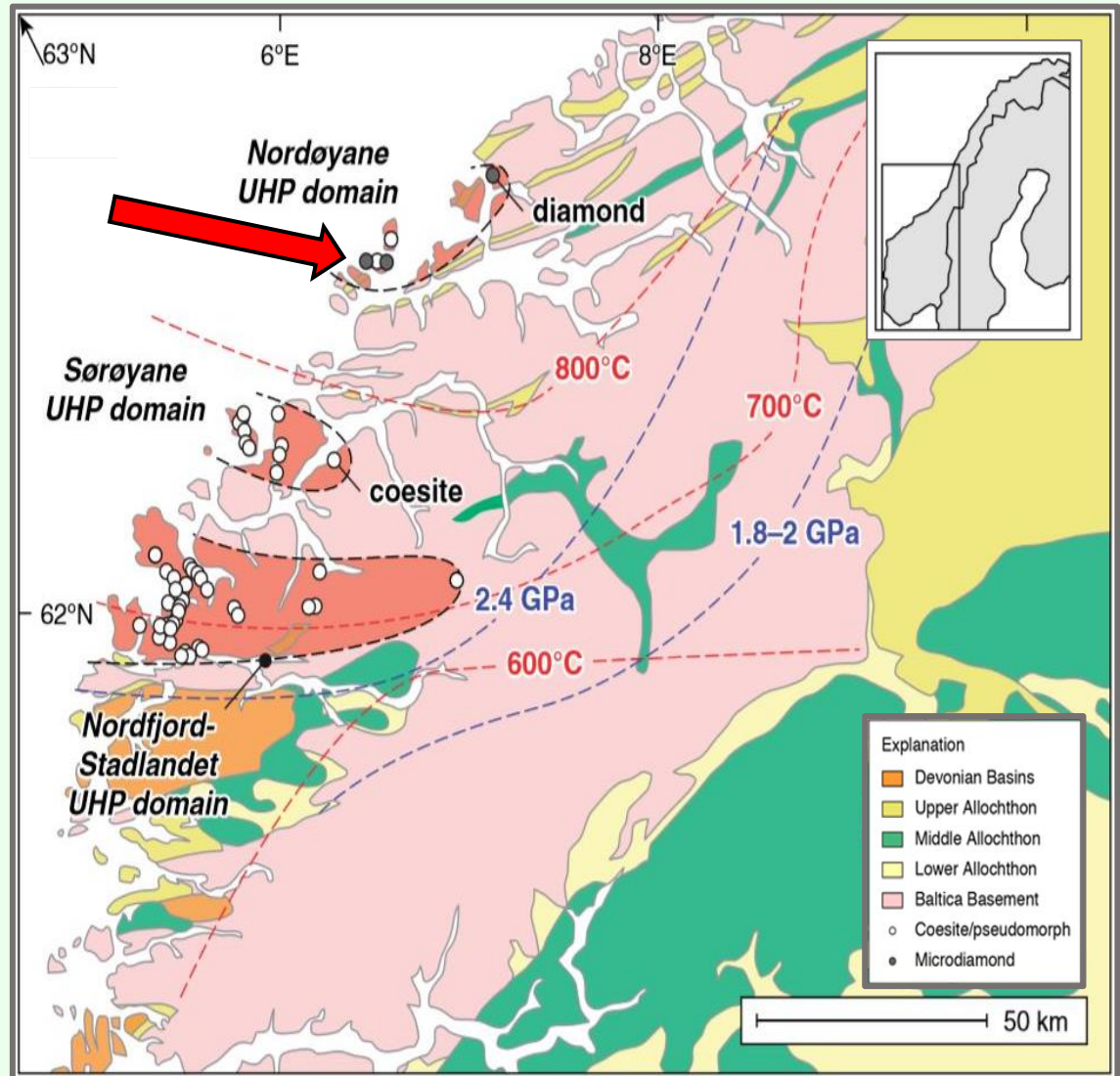
Purpose and scope of study

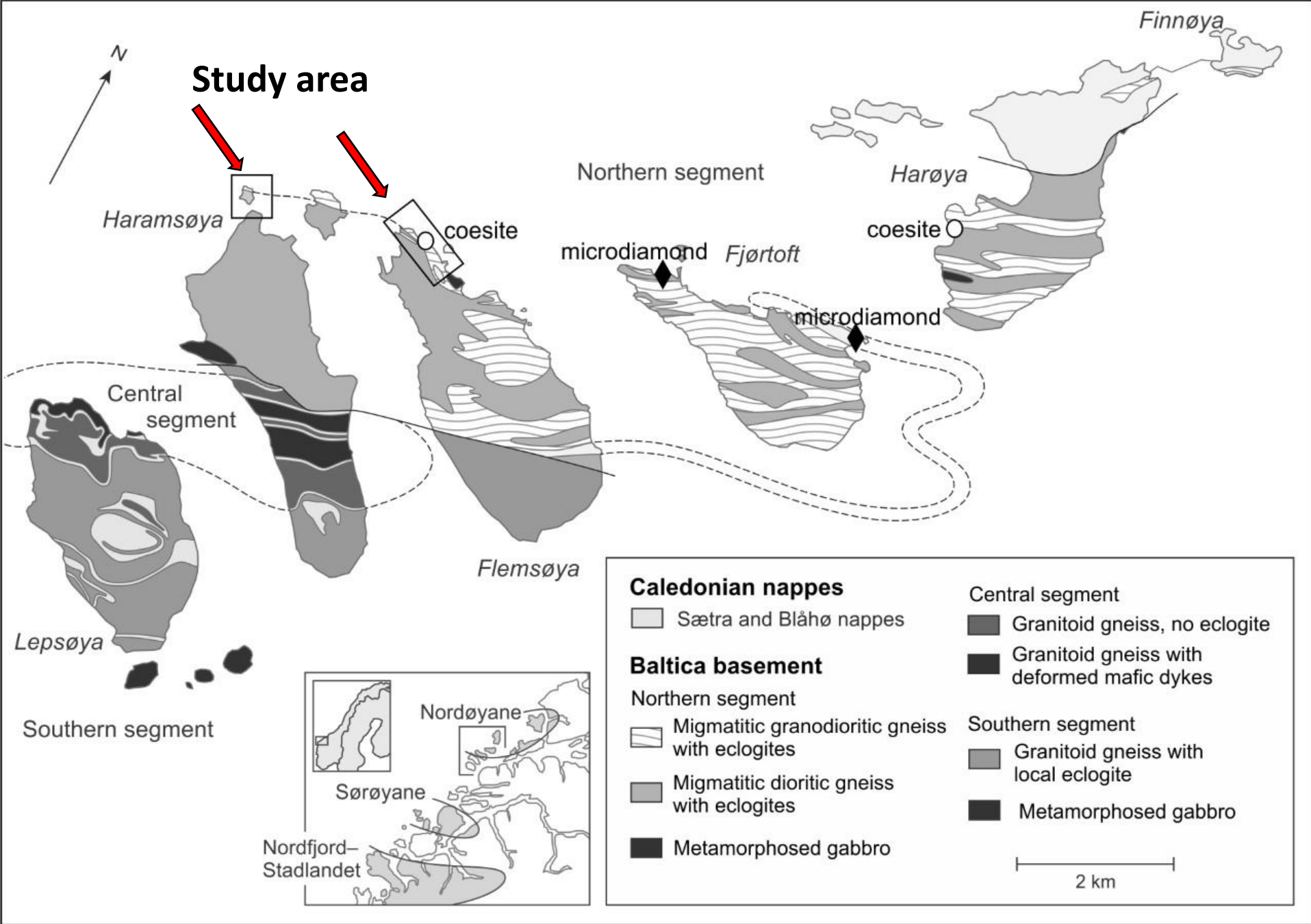
- Analyze petrography of eclogites from Nordøyane
- Obtain EMP analyses from eclogite-facies minerals
- Use conventional and trace-element thermometry
- Construct an equilibrium phase diagram for a representative bulk composition
- Assess which thermobarometric methods are most robust for given mineral assemblage

Do the P-T conditions recorded in the eclogite bodies overlap with the UHP melting range for eclogites and their host rocks?

Regional geology of WGR

- Exposes Baltican Proterozoic basement gneisses
- Result of subduction of Baltican margin beneath Laurentia
 - Scandian phase of Caledonian Orogeny
 - Ca. 410 Ma
- Affected by ultra-high pressure (UHP) metamorphic conditions of ca. 800 °C, 30 kb
 - Corresponds to ca. 100 km depth
- Three UHP domains: Nordfjord /Stadlandet, Sorøyane, Nordøyane
 - All contain coesite-bearing eclogite





Geology of study area

- Eclogite sampled from Haramsøya and Flemsøya
- Dioritic gneiss and augen gneiss comprise majority of study areas
- Minor granodiorites with mafic enclaves
- Eclogite bodies range from massive to layered
 - Some display amphibolite-facies retrogression
- Evidence suggesting UHP conditions supported by discovery of coesite-bearing eclogite in Flemsøya



Photo by G. Chapman

Field relationship

Host migmatitic orthogneiss mineral assemblage records amphibolite facies conditions

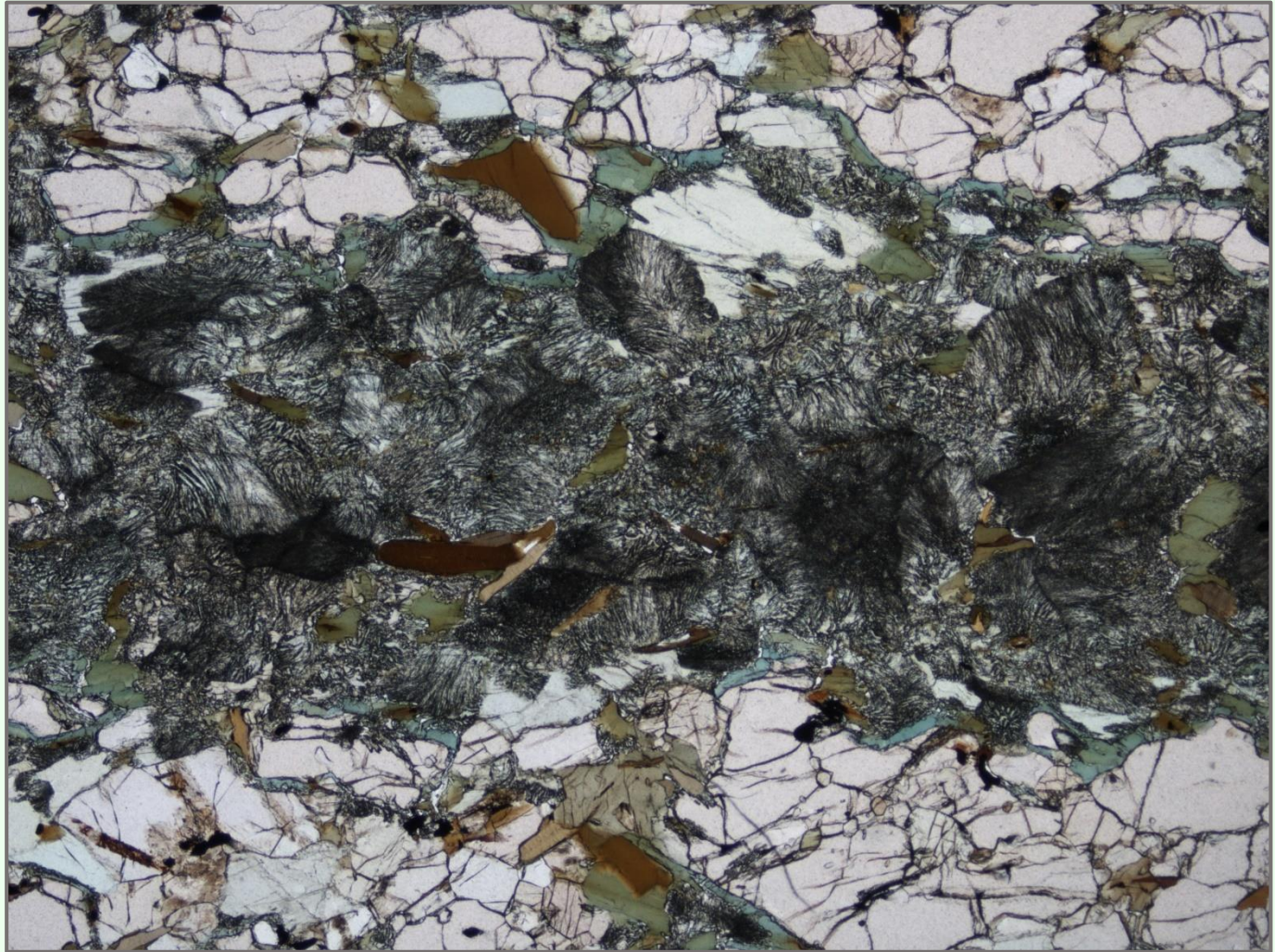
Eclogite mineral assemblage may record peak P-T conditions



Photos by G. Chapman

General methods

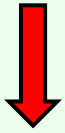
Petrography



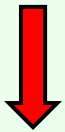
Sample CB15-70 F.O.V 15 mm

General methods

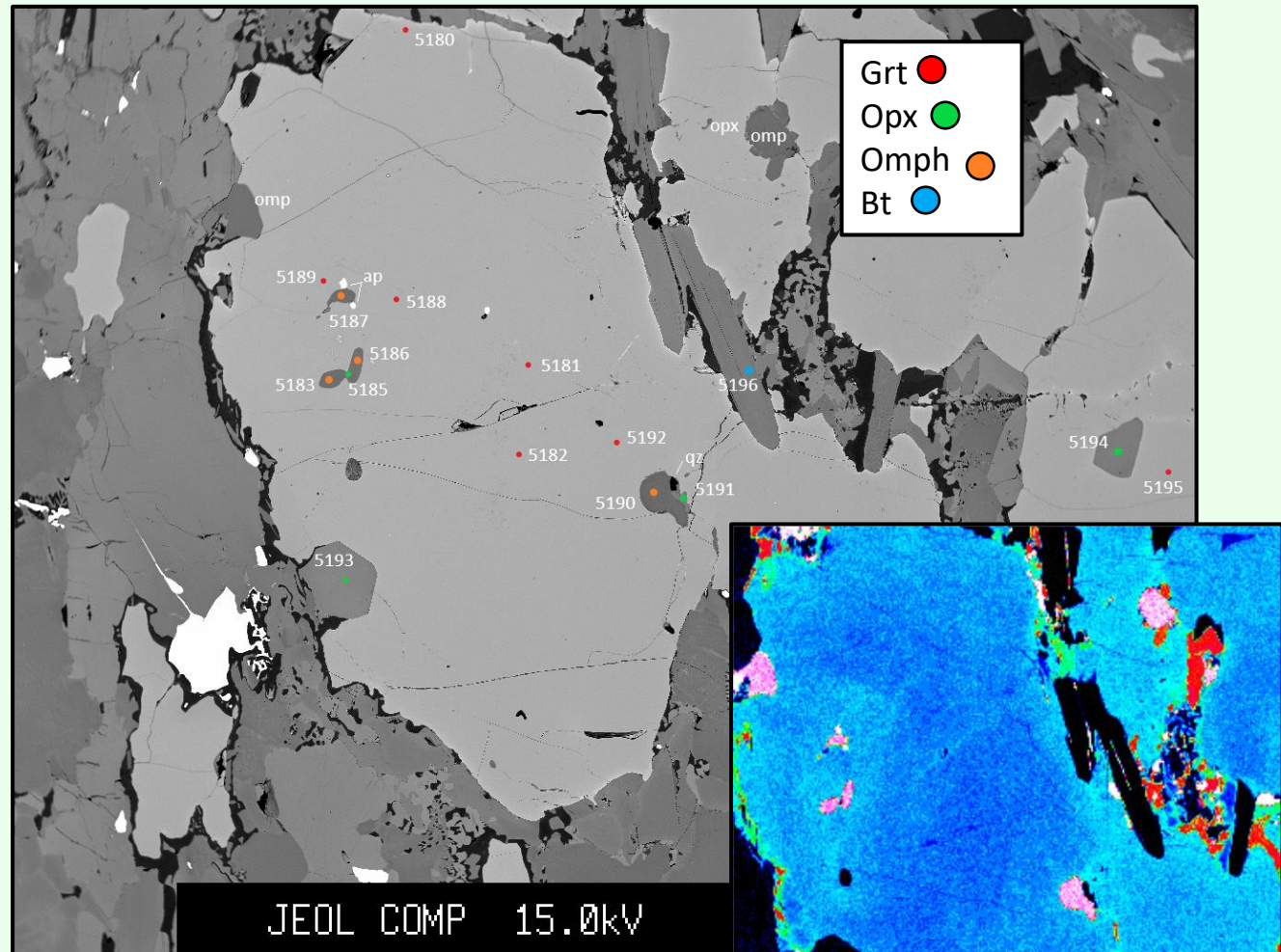
Petrography



EMP and
imaging



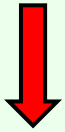
Conventional
thermobarometry



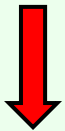
BSE image of CB15-19 showing microprobe analysis spots for various minerals and chemical map of same garnet grain

General methods

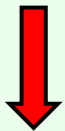
Petrography



EMP and
imaging



Conventional
thermobarometry



Thermodynamic
modelling

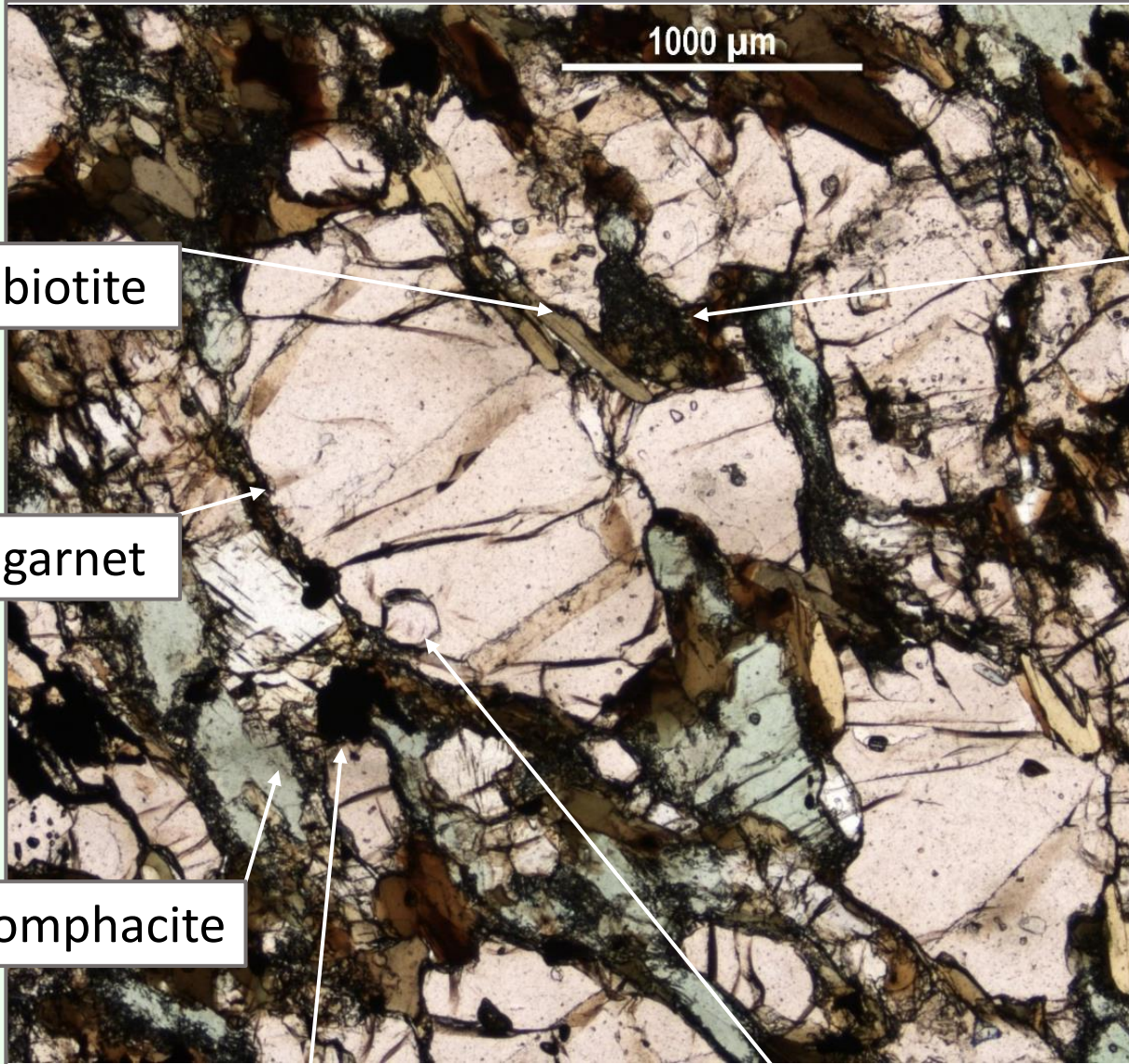
Conventional thermobarometry

- **Reverse** approach
- Assumes **equilibrium compositions**
- **Produces P-T estimates** using equations derived from experimental calibrations

Thermodynamic modelling

- **Forward** approach
- Based on **internally consistent databases**
- **Predicts** equilibrium phase **assemblages for P-T ranges**

Petrography



1000 μm

biotite

garnet

omphacite

rutile

orthopyroxene

Retrograde

clinopyroxene₂ +
amphibole + plagioclase
symplectite

Garnet inclusions:

- omphacite
- rutile
- zircon
- apatite
- idioblastic biotite

Rare quartz, absent
coesite, kyanite,
phengite

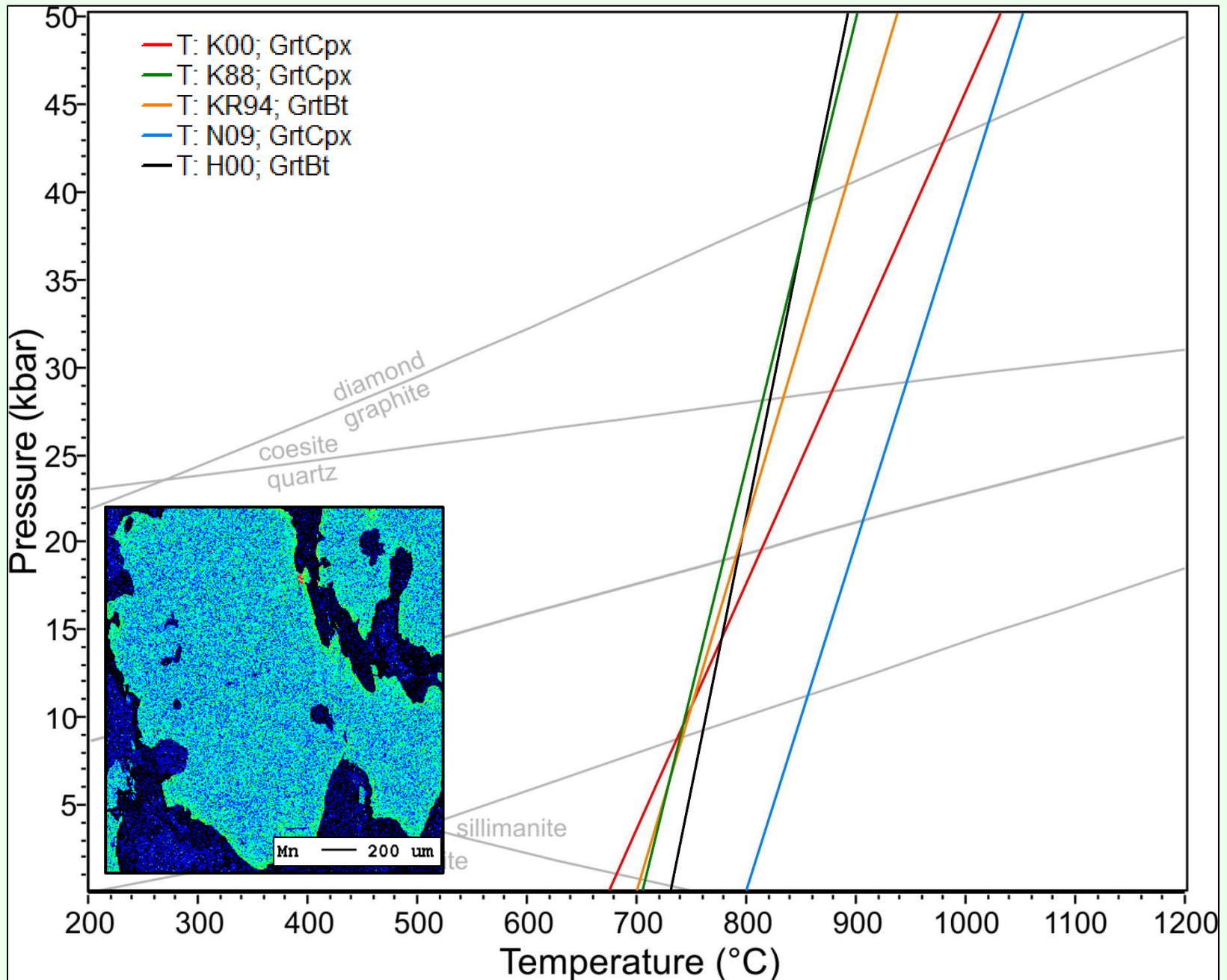
Method 1A: conventional thermobarometry

- P and T dependent equilibria
- Minerals present mean heavy reliance on **Fe-Mg exchange thermometry**
 - Garnet-clinopyroxene (e.g. Krogh Ravna, 2000)
 - Garnet-biotite (e.g. Holdaway, 2000)
- **Issues:**
 - **Unknown Fe³⁺ values**
 - Mineral assemblages **sensitive to retrograde exchange**

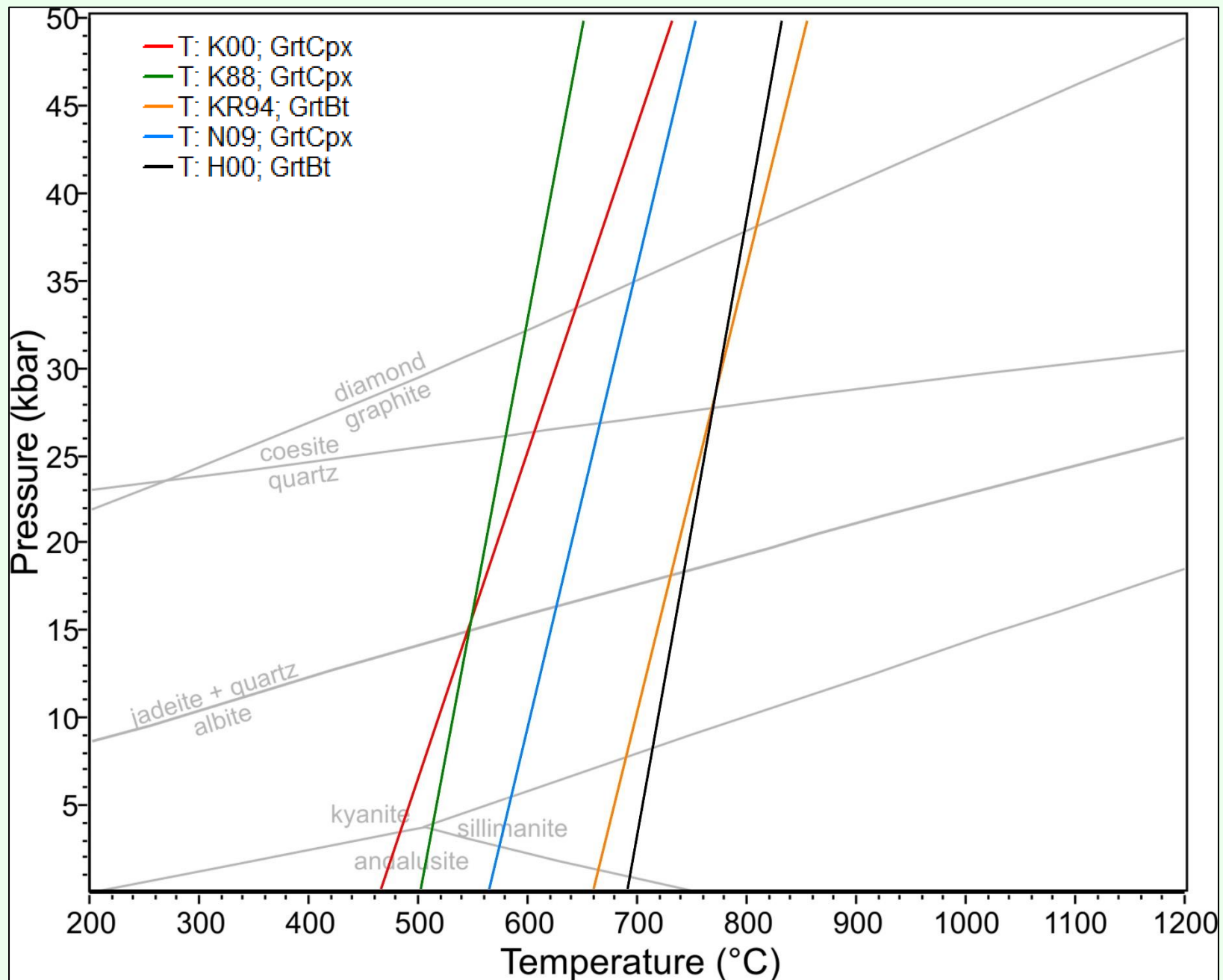
Method 1B: trace-element thermometry

- Type of conventional thermometry based on calibrated reactions
- **Zr-in-rutile thermometer** (Tomkins et al. 2007)
- Exchange between Ti⁴⁺ and Zr⁴⁺ in **rutile + zircon + quartz assemblages**
- Less sensitive to issues such as retrograde exchange

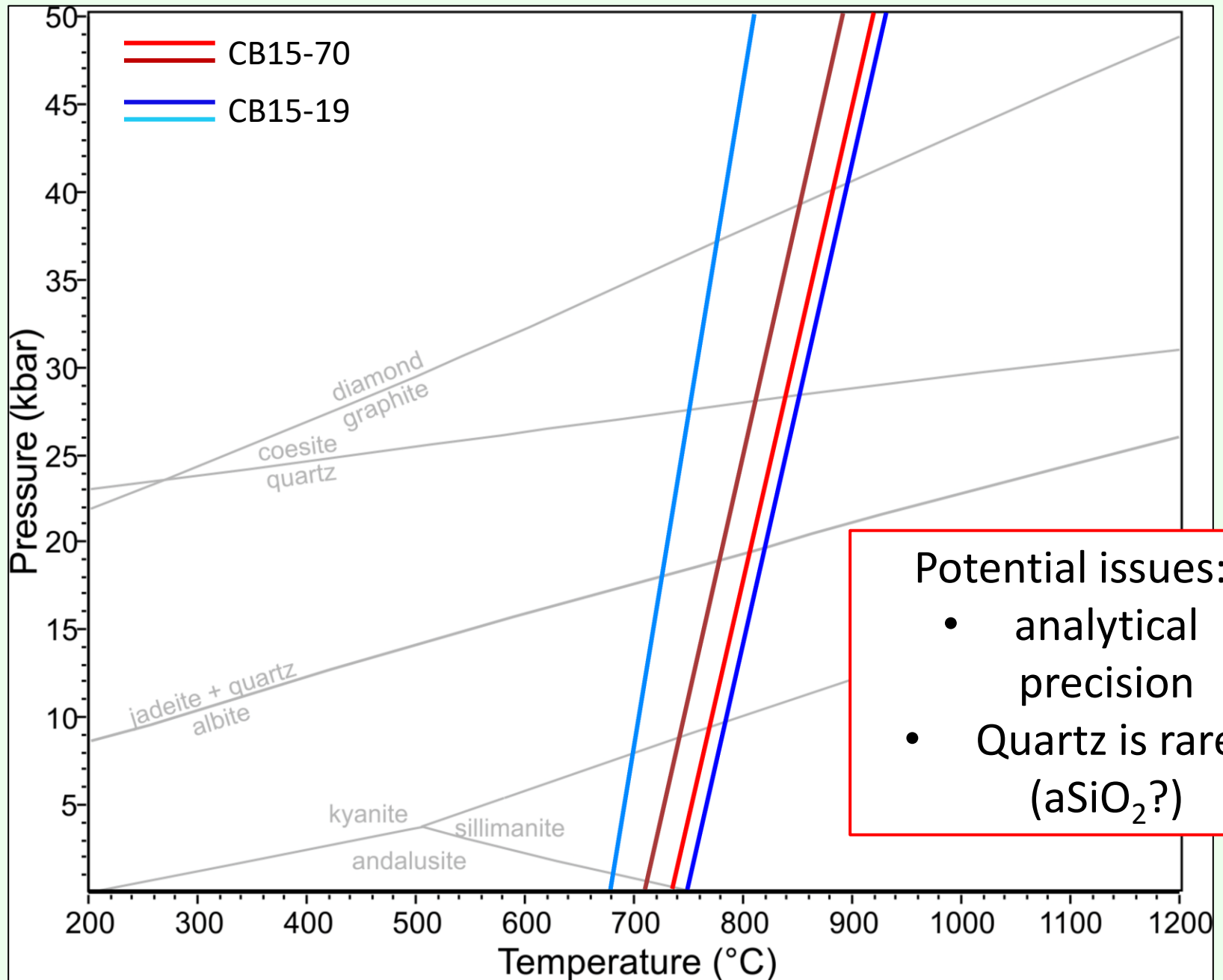
Assuming Fe as Fe²⁺



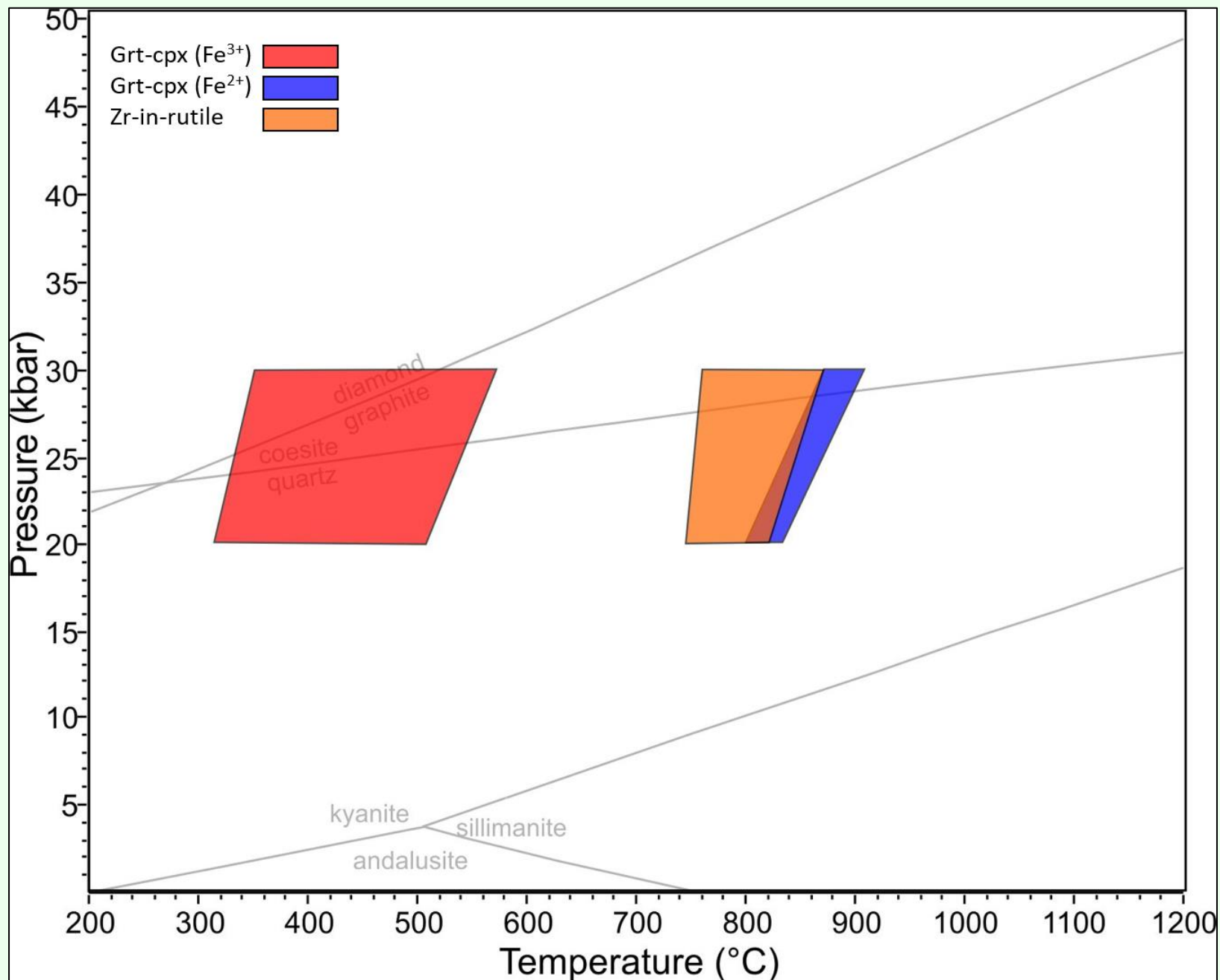
Using calculated $\text{Fe}^{3+}/\text{Fe}^{2+}$



Zr-in-Rutile thermometry



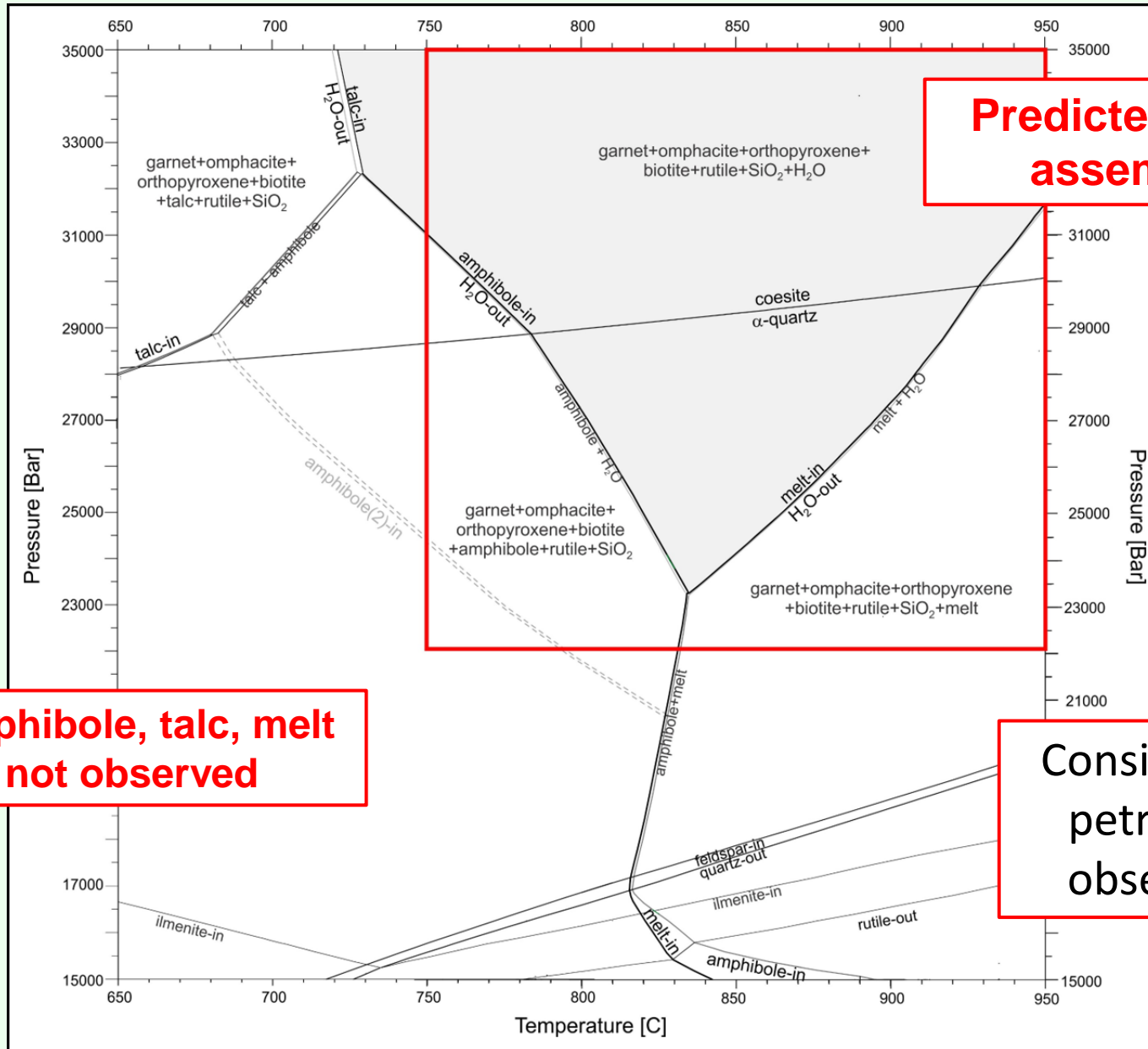
P-T assessment



Method 2: equilibrium phase diagram

- **Given bulk composition**, program predicts mineral assemblage, modal proportions, mineral compositions
- **Independent** of conventional thermobarometry
- Theriak-Domino in family of programs such as THERMOCALC and Perplex
 - Differ in algorithm used, activity models, databases
- Results use TC55 database (Holland and Powell, 1998)
- Input: **measured bulk composition of sample CB15-19**

Method 2 equilibrium phase diagram (Theriak-Domino)

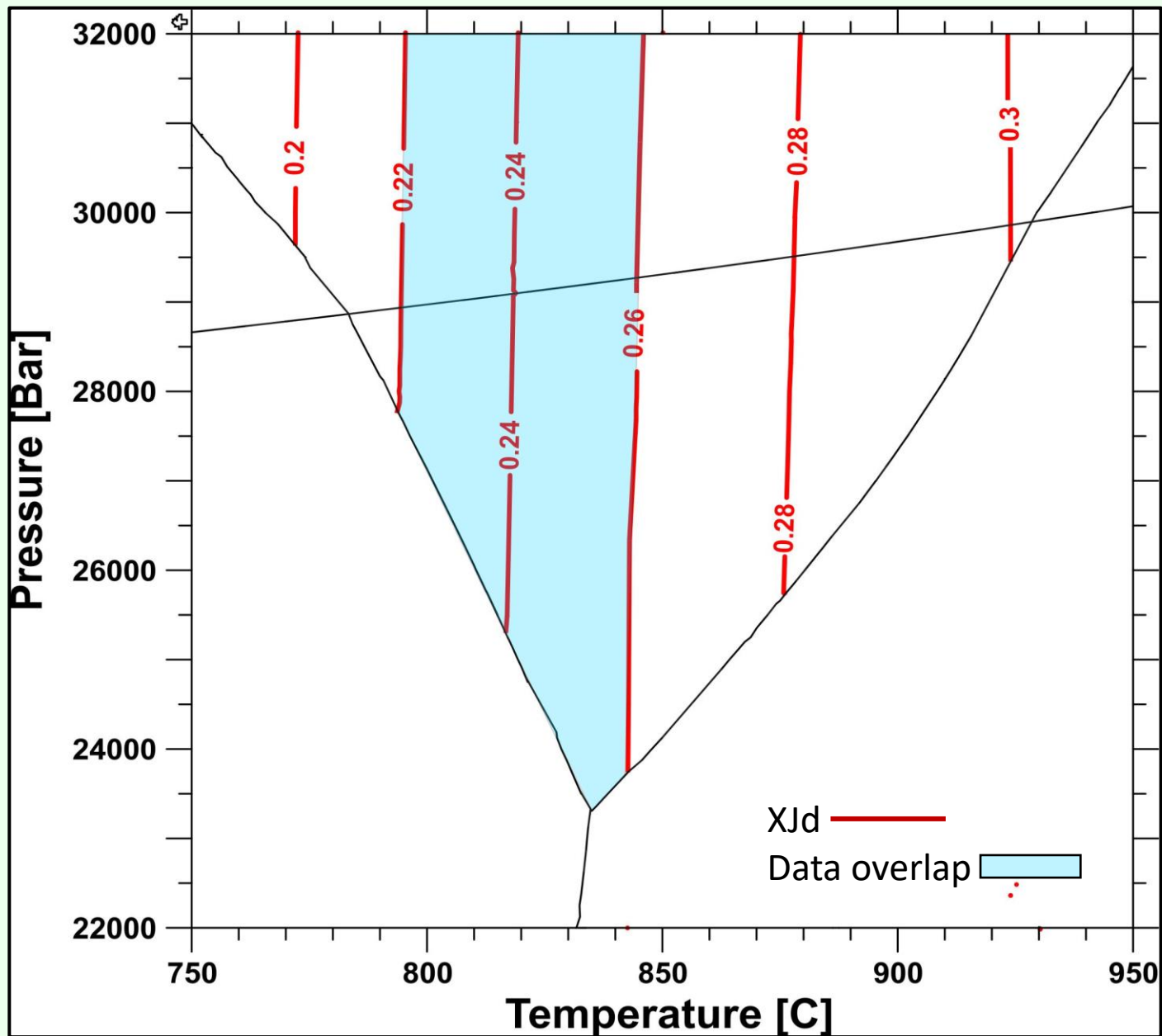


Predicted mineral assemblage

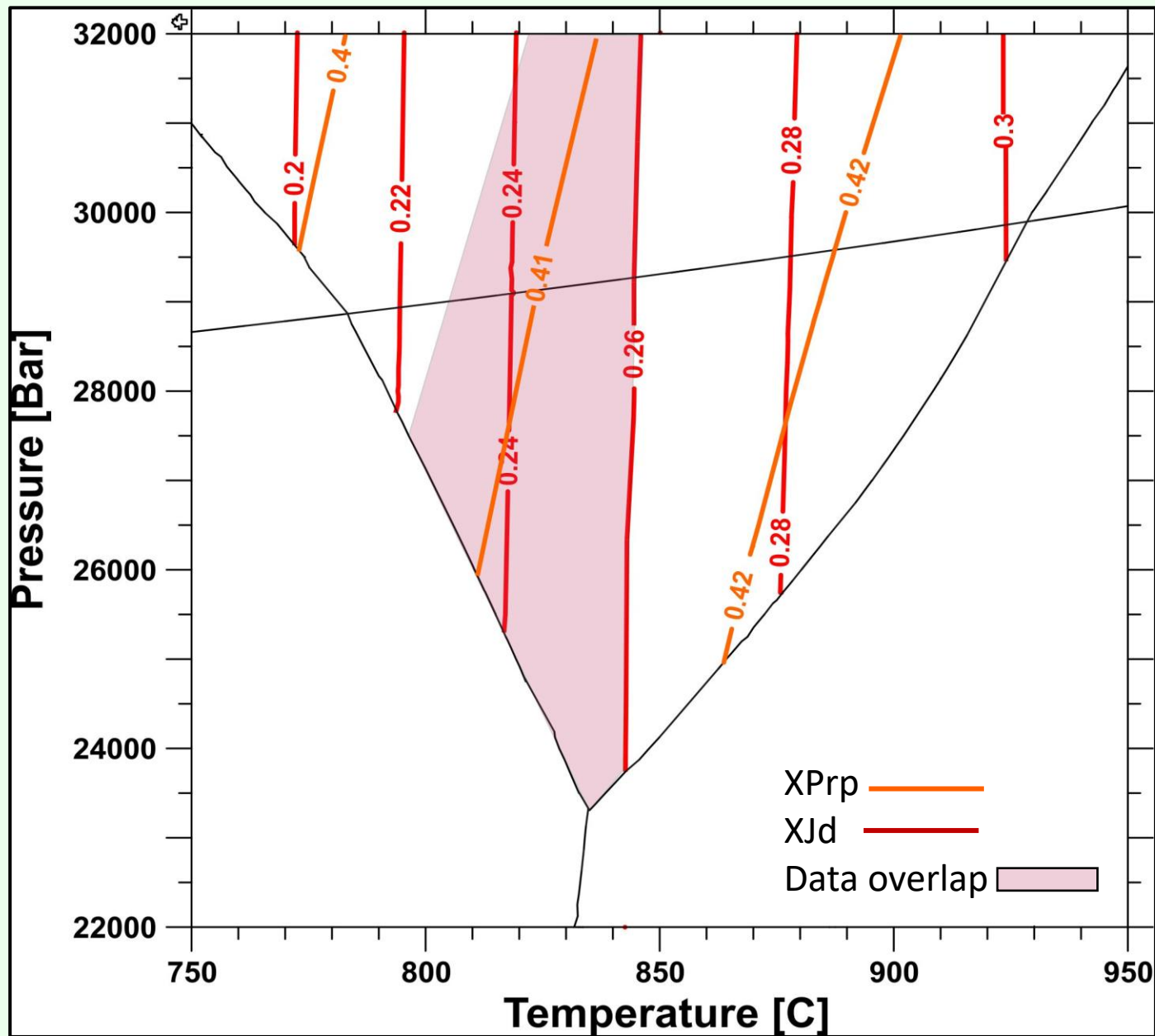
Amphibole, talc, melt not observed

Consistent with petrographic observations

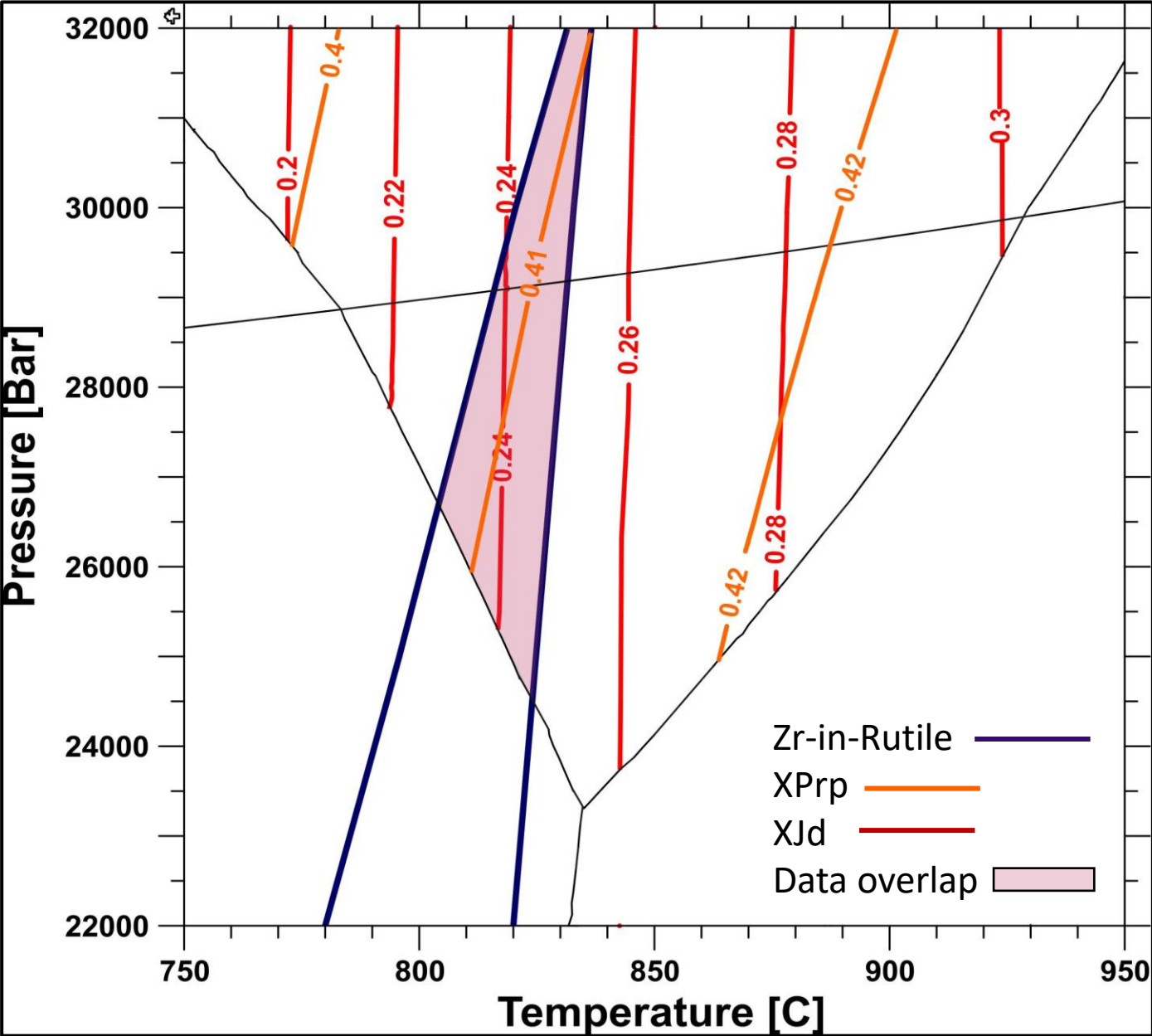
Predicted mole fractions



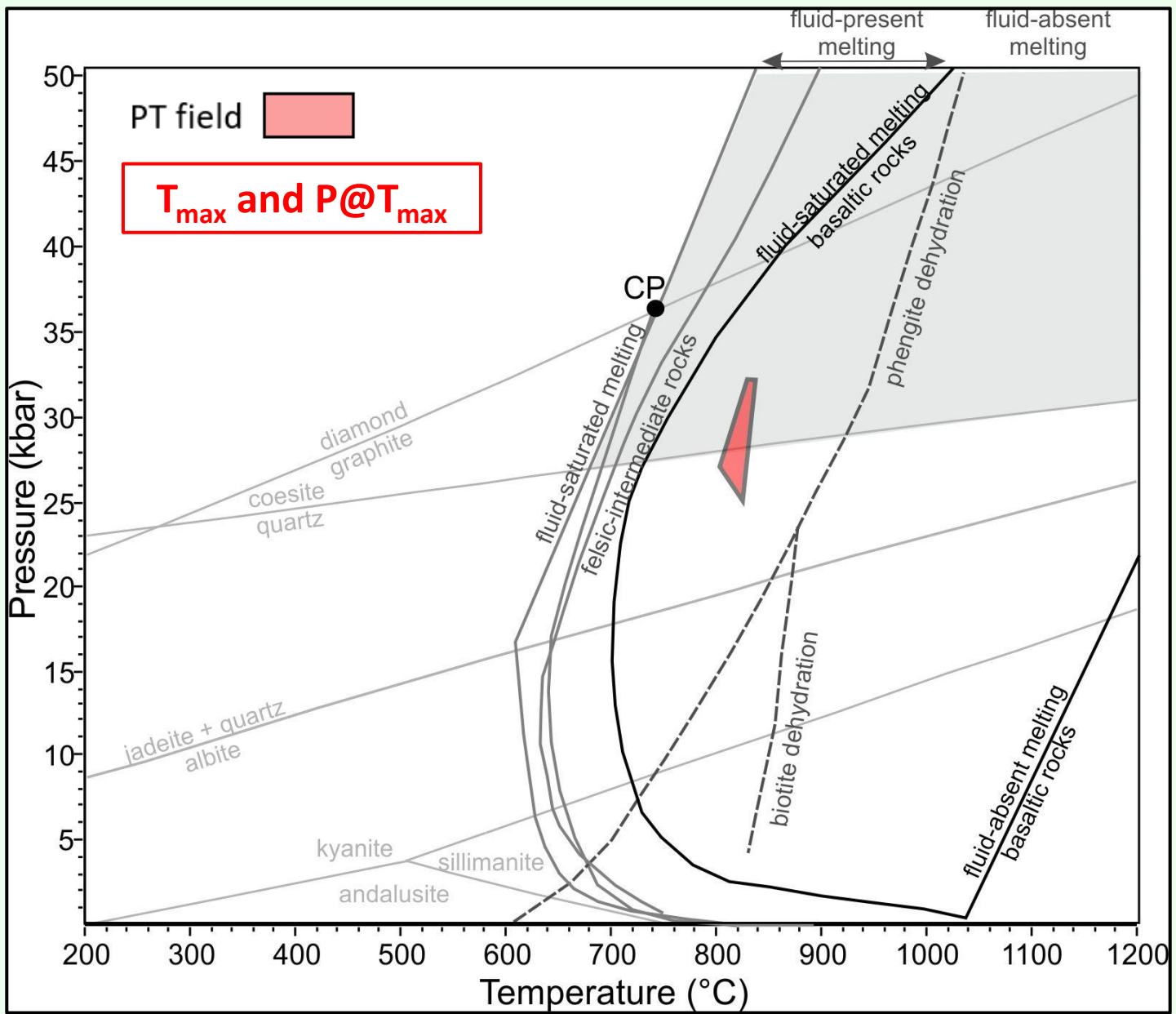
Predicted mole fractions



Predicted mole fractions with Zr-in-Rutile



Fluid-present melting



various melting curves of crustal compositions

(Kessel et al. 2005, Patiño Douce 2005, Hermann & Spandler 2008, Auzanneau et al. 2010)

Conclusions

- Peak mineral assemblage comprises garnet + omphacite + biotite + rutile + orthopyroxene
- Fe-Mg exchange thermometry proved to be unreliable for P-T estimates
 - Two factors: Fe³⁺ and grain scale diffusion during retrogression
- Trace and minor element thermometers yielded T of **780-870 °C** for 20-30 kb
- Thermodynamic modelling coupled with Zr-in-Rutile isopleths suggests P-T field of ca. **800-820 °C** and ca. **23-32 kb**
- **T_{max}** and **P@T_{max}** overlap with field where UHP melting is possible in presence of fluid
- Cannot conclude if UHP melting assisted exhumation
 - No independent P constraint
 - No evidence of *in situ* melting or fluid infiltration

Acknowledgments

Becky Jamieson

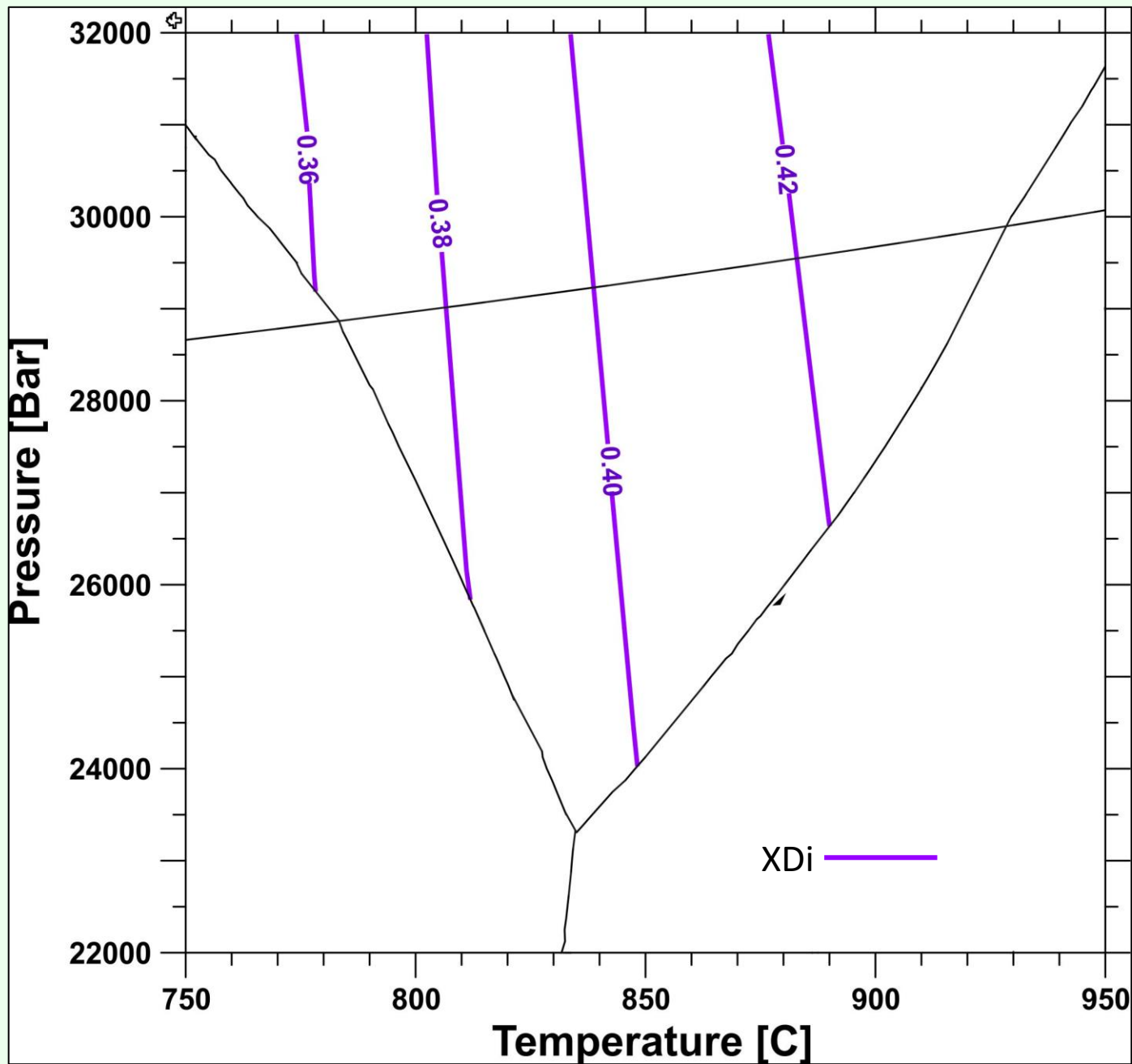
Luke Hilchie

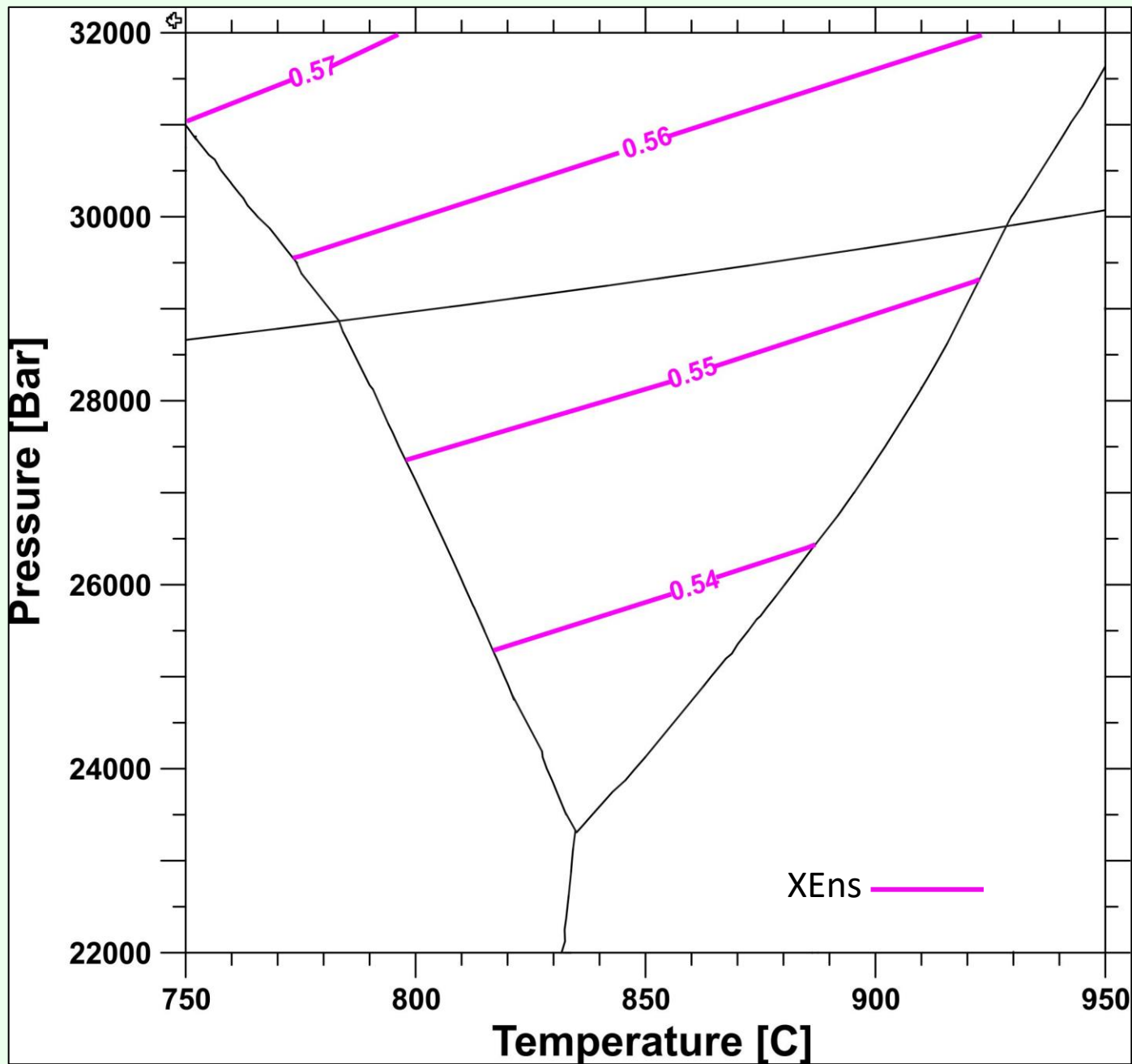
Dan MacDonald

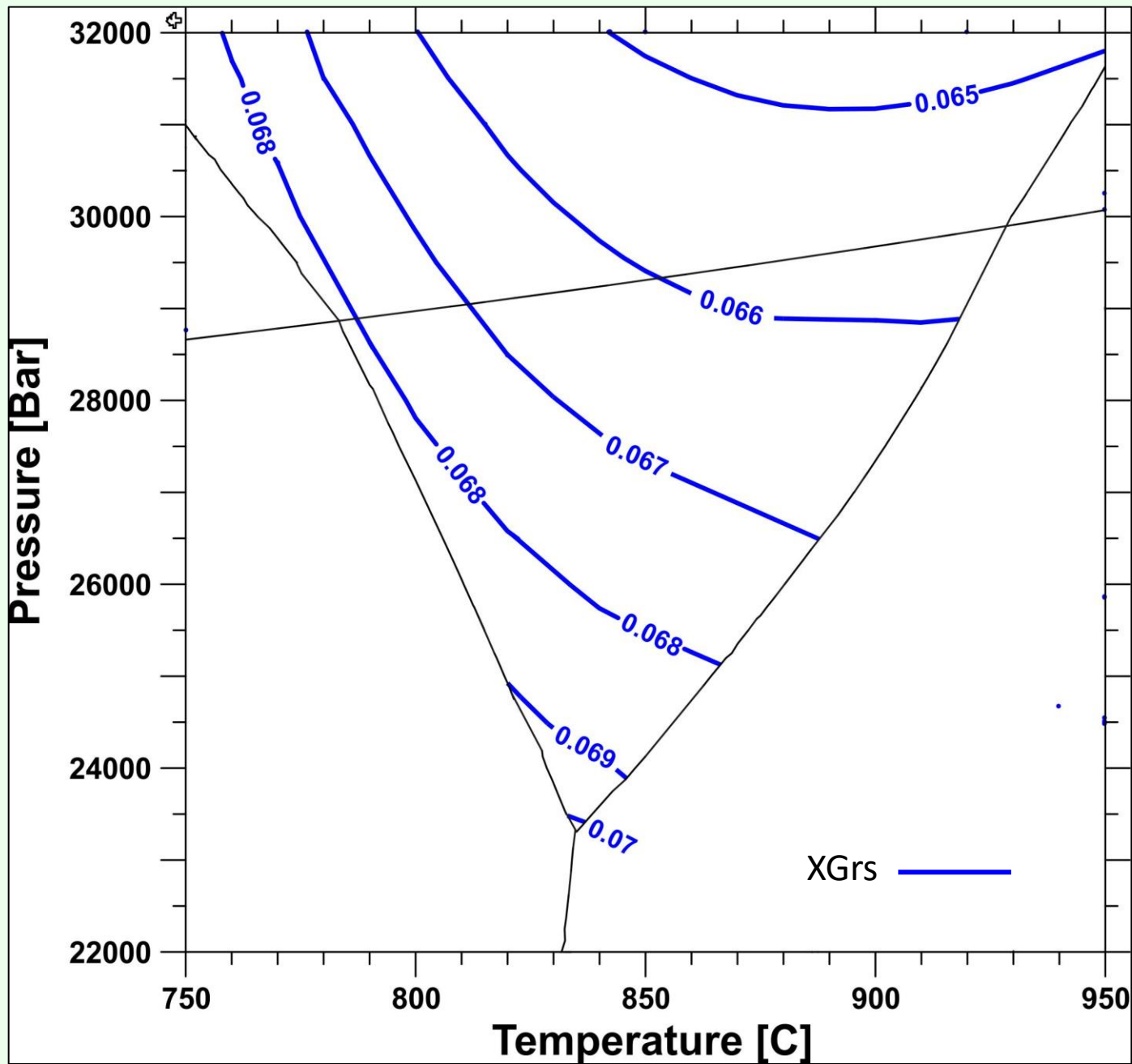
Djordje Grujic and Honours Peers

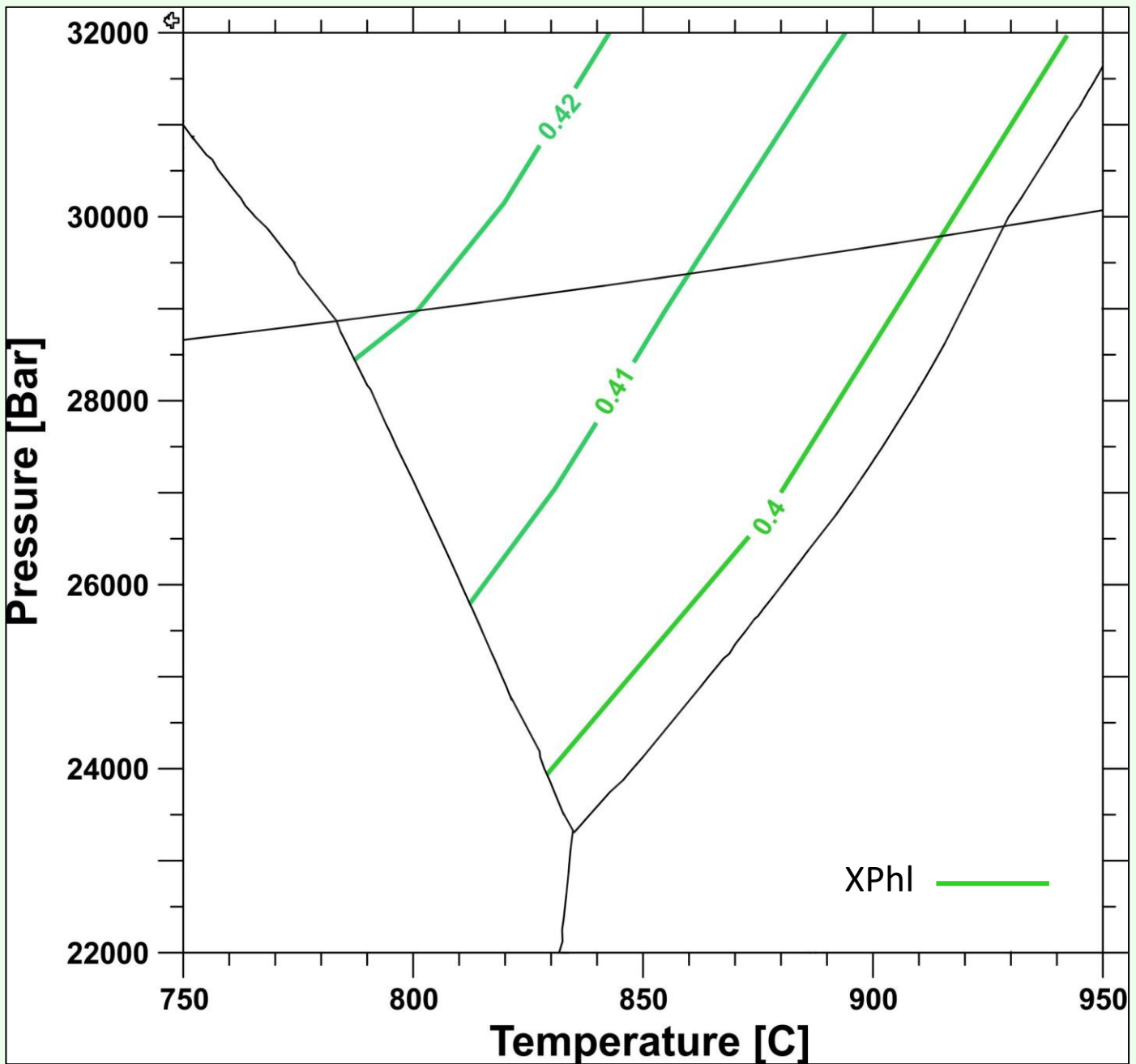
Department of Earth and Environmental Sciences

Thank you!

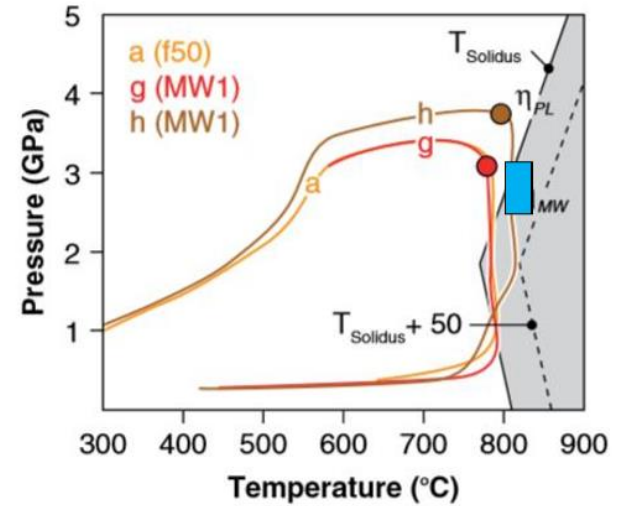
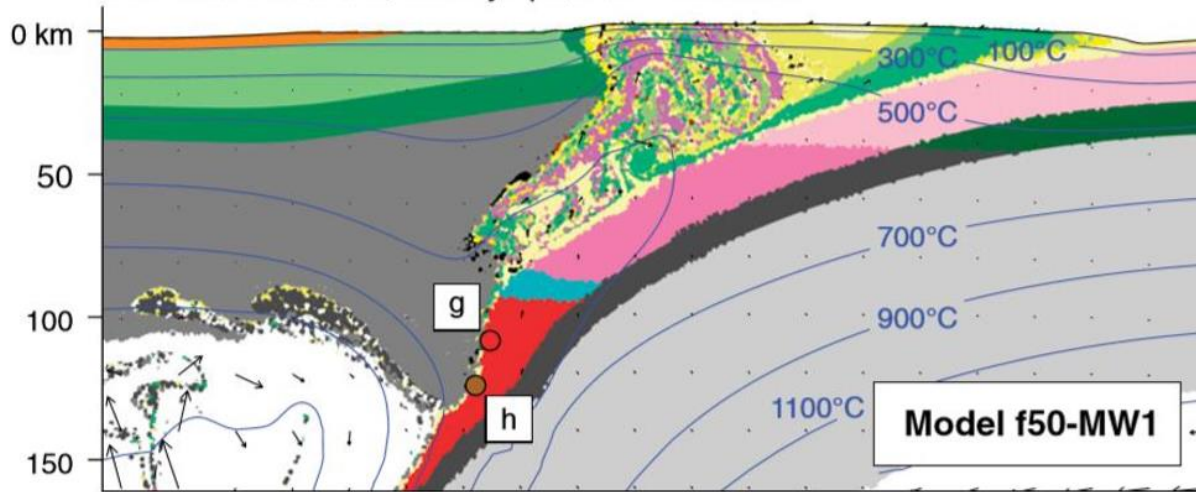




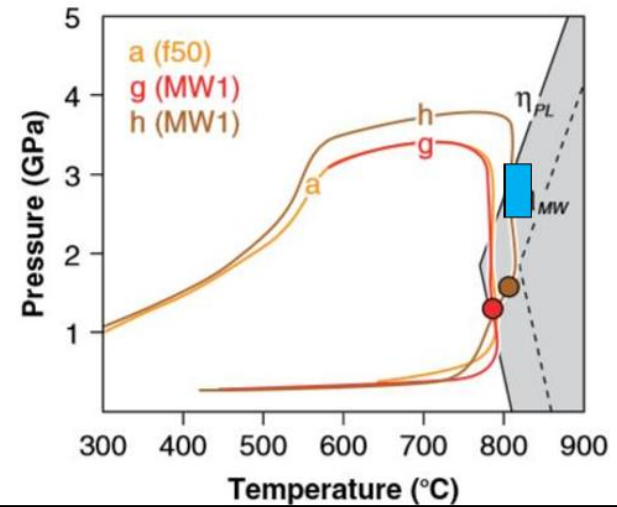
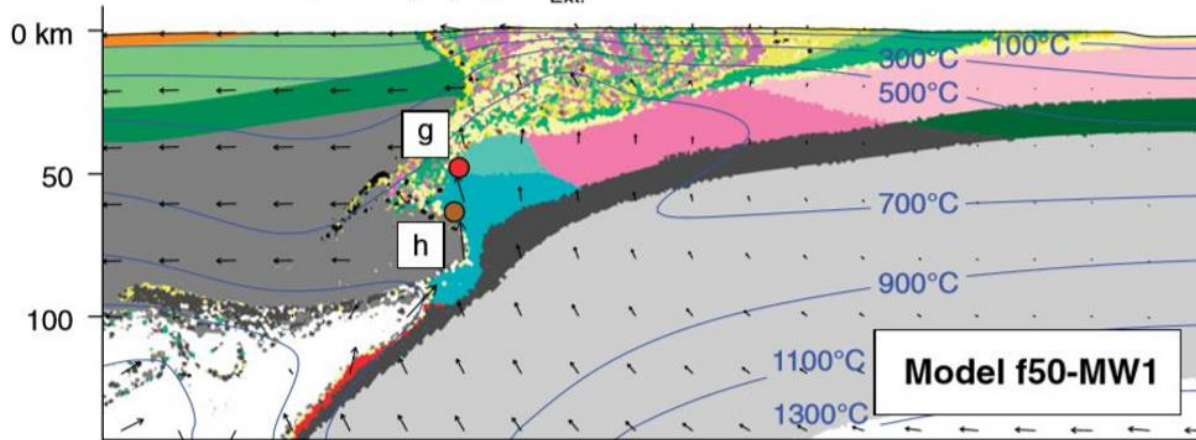




A. End Phase 3; 55 Myr-pc; $\Delta x = 1125$ km



B. Phase 4; 60 Myr-pc; $\Delta x_{Ext.} = 50$ km



Modified from Butler et al (2015)