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BENTHIC FORAMINIFERAL ASSEMBLAGES ON THE  
CONTINENTAL MARGIN OFF NOVA SCOTIA:  
A MULTIVARIATE APPROACH.

By

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Submitted in partial fulfillment of the requirements for the  
degree of Doctor of Philosophy at Dalhousie University  
April, 1983.

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## ABSTRACT

Two hundred and fifty grab samples from the continental margin off Nova Scotia were examined for total (live+dead) foraminiferal content. One hundred and twenty species were recognised, seventy five of which occur with an abundance greater than 3% in any one sample. Q-Mode factor analysis of the raw abundance data (total population) determined 8 shelf and 4 slope factor assemblages, accounting for 87.5% and 77% of the original raw data respectively. Comparison of the live data with the factor assemblages enabled the detection of anomalous assemblages which are un-representative in respect to present day environments.

In the north east of the study area (around Misaine, Canso and Banquereau) an exclusively agglutinated assemblage dominated by Adercotryma glomerata occupies both banks and basins. In the LaHave and Emerald Basins of the central shelf, a predominantly calcareous assemblage occurs with maximum amounts of Globobulimina auriculata and Nonionellina labradorica. Transitional between these two shelf assemblages is an agglutinating assemblage dominated by Saccammina atlantica. Consistently present along the shelf edge is a Trifarina angulosa assemblage. Present in Chedabucto and Gaberous bays and in a few samples near Sable Island is an agglutinating Eggerella advena assemblage. A relict and in some areas a transport affected assemblage is recognised in the south western approaches to Emerald Basin on the Scotian Gulf; this is dominated by Elphidium excavatum. Occupying the hard, rough bedrock, and sandy/gravelly areas of the inner shelf and outer bank regions is a Cibicides lobatulus assemblage.

Four assemblages lie seaward of the shelf break. An upper slope assemblage is dominated by Bulimina exilis, with a local variation dominated by Trifarina occidentalis. A lower slope assemblage is dominated by Uvigerina peregrina. Also present in slope regions is another Elphidium excavatum assemblage which is thought to be the down slope equivalent of the shelf assemblage.

The statistical relationship of these defined assemblages to various aspects of the marine environment (depth, temperature, salinity, percent gravel, sand and mud) were investigated through multiple regression techniques. This indicates that the present foraminiferal distribution patterns off Nova Scotia are mainly a response to the prevailing hydrography. The Adercotryma glomerata assemblage is influenced by the presence of cold, less than normal salinity waters of arctic, Labrador current origin. The central basin assemblage (G. auriculata) is related to warmer more saline waters of slope origin. The transition between these two bottom waters is marked by the Saccammina atlantica assemblage. The primary influence of the hydrography on assemblage distribution patterns on the shelf off Nova Scotia is probably related to calcium carbonate availability which is a function of temperature, salinity, and carbon dioxide content within each water type. Preferred substrate character is found to determine the occurrence of the Cibicides lobatulus assemblage.

Diversity trends within the study area were examined ( $S, H(s), E$ ). Each index used produced similar patterns which indicate relationships to hydrography and environmental stability.

Implications of the present study for further work on Pleistocene to Recent climatic changes are discussed with reference to a piston core from Canso Basin. On the basis of this study the down core changes in foraminiferal populations indicate several distinct variations in

temperature and salinity since the last glaciation; these variations are probably a result of the varying degree of influence that Labrador derived waters have had in this area and is in turn a function of changing ice volumes to the north.

## ACKNOWLEDGEMENTS

Sincere thanks are due to Dr. Franco Medioli for his advice and support as my supervisor. Dr. C.E. Keen of the Bedford Institute of Oceanography is sincerely thanked for her help and advice on matters mathematical and for making available her computing expertise and the modified version of CABFAC. In these respects Dr. Peta Mudie is also thanked together with her reading and discussion of the manuscript which proved invaluable. Drs Gibling, Medioli, Hill and Scott critically read the tome and provided many pertinent comments. Dr. D.B. Scott provided much help and advice. Many thanks to Dr. D. Piper of B.I.O. for helpful suggestions, advice and discussions. Several people both at Dalhousie and the B.I.O. helped with many aspects of the work, providing encouragement, advice and useful discussion; the following are but a few: Dr. A. Aksu; A.A. Miller; A. DeLure; Dr. R. Boyd; R. Courtney; Dr. P. Schenk; Dr. C. Schafer; Dr. G. Vilks; Dr. F. Gradstein; Dr. P. Hill and M.C. Blanchard-Williamson. Dr L. H. King and Mr. G. Fader are thanked for their interest and kind permission to use their sediment data for the Shelf off Nova Scotia.

Thomas Duffett is thanked for invaluable technical support on several cruises and C. Younger for help in the laboratory. Sincere thanks is extended to the Captains, Officers and Crew of the C.S.S. Dawson and M.V. Pandora II for their professional service. The manuscript was typed by M.C. Blanchard-Williamson whose constant encouragement and help is appreciated. Dal Graphics and Stefani Ferenc of photographic services are thanked for their work. Scanning Electron Microscope Pictures were taken by C. Mason.

Previous to this thesis much help, discussion, advice and encouragement was provided by Drs J. Haynes and R. Whatley at the University College of Wales Aberystwyth; ideas, concepts and discussion seen in this work were in no small way formulated in those earlier years. My mother and father have constantly supported me over the years and in many ways are responsible for this work.

Finally, gratitude is due for an Izaak Walton Killam Memorial Scholarship, and to Dalhousie Graduate Studies for financial support that has made my sojourn at Dalhousie very pleasant indeed.

In dedication to Marie-Claude.

...the face of places,  
and their forms decay;  
and that is solid earth,  
that once was sea;  
Seas, in their turn,  
retreating from the shore  
make solid land,  
where ocean was before.

Ovid metamorphoses, XV

## CHAPTER 1: INTRODUCTION

## 1.1 General Introduction

The usefulness of foraminifera to stratigraphic and palaeoenvironmental investigations of earth's history is well documented in the geological literature. Early taxonomic studies of the 19th century gave way in the first half of the 20th century to a recognition of their worth in helping to overcome problems of correlation, palaeoecology and age determinations posed by hydrocarbon exploration. Their usefulness in this respect is a result of their small size, being easily retrieved (if diagenetic conditions allow) from borehole chips in statistically sufficient numbers to provide vital numerical control. This capability of foraminifera to provide representative populations from small samples coupled with their sensitivity to environmental conditions has provided investigators with a valuable tool in the determination of palaeoenvironments and, more recently, with a quantitative control of these investigations.

One of the basic tenets of geological philosophy is that the present is a key to the past. To this end, there is an increasingly vast body of literature that documents the occurrence of foraminiferal species and their influencing environments. The very richness of foraminiferal species and their individual response to environments lead often to confusing and misleading conclusions. One thing remains clear: the limits of accuracy of each palaeoenvironmental study are set by the knowledge, direct or homeomorphic, of present day regimes governing a species distribution. This is providing that the "present is a key to the past" assumption is valid.

## 1.2 Objectives

The aims of this study are:

- a/ to determine, using quantitative methods, the present day assemblages of benthic foraminifera on the continental margin off Nova Scotia, their distribution and faunal trends;
- b/ to determine, utilizing live distribution data, the degree and extent that these foraminiferal assemblages are in equilibrium with the presently occurring environmental regimes (i.e. in situ, transported or relict).
- c/ to establish a quantitative correlation between the defined assemblages and various geological and oceanographic parameters related to presently-occurring environmental regimes.
- d/ to comment on the implications of the above study.

The purpose behind these aims becomes apparent in the light of other research projects underway on the Eastern Canadian coast. These other investigations are concerned with the interpretation of Pleistocene to Recent sediments in terms of the climatic and oceanographic history of Eastern Canadian watermasses since the last glaciation. The overall project is necessarily multidisciplinary in approach, involving the study of pollen ( Mudie, 1982), diatoms, sedimentology ( King 1972; Piper 1975, 1976; Hill 1981), geochemistry (Oxygen isotopes; Alam and Piper, 1980; Aksu 1980), as well as the use

of foraminifera ( Vilks, 1980, Scott *et al* 1979)). The results of this study will expand the present data base and provide a vital quantitative control applicable to palaeoceanographic investigations of the Pleistocene to Recent sediments off Eastern Canada, using benthic foraminifera. A following section will describe the prevailing hydrography of this study area and will show the presence of a major hydrographic boundary extending across the shelf. Waters to the north east of this boundary are essentially of Arctic origin whilst those to the south west are warm, slope, Gulf stream derived. The studies mentioned above endeavour to document the changing nature of these influences since the Pleistocene; the present study provides an ideal opportunity to study the foraminiferal population dynamics across this oceanographic boundary at the present day, and to consider the implications of the results for understanding past and future oceanographic changes.

### 1.3. General Aspects

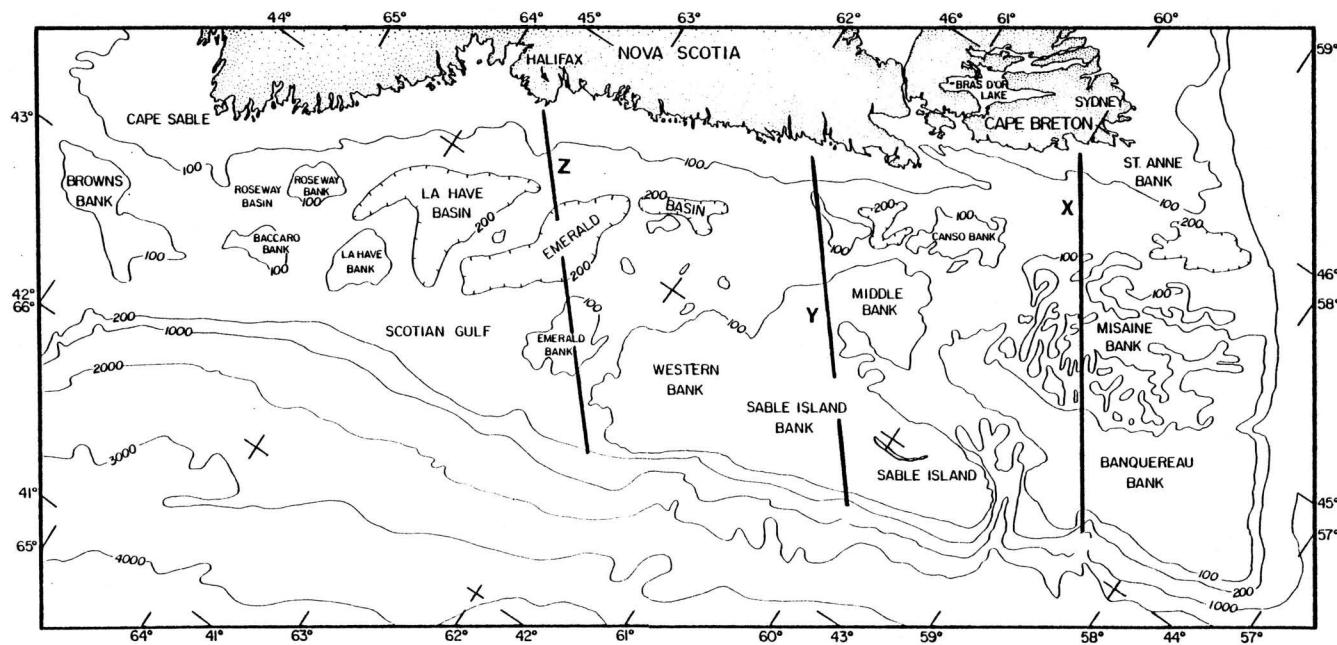
#### 1.3.1. Location.

The continental margin off Nova Scotia is considered here as that submerged area adjacent to its mainland Atlantic coast which lies between 41°N 65° 30'W and 46°N 57'W and extends to depths of approximately 3000 m. It is bounded to the east by the Laurentian channel and to the west by Browns Bank ( Fig. 1.1).

#### 1.3.2. Bathymetry.

Figure 1.1 Bathymetry of the continental margin off  
Nova Scotia.

Sections X,Y,Z, refer to figure 1.4.



The shelf break occurs at about the 200 m contour off Nova Scotia and effectively divides the study area into two physiographic regions: the continental shelf and the continental slope. The shelf is quite variable in width, ranging from 125 km in the west to 230 km in the east. Compared to other shelf areas of the eastern United States, northern Newfoundland and, to some extent, Labrador, the Nova Scotian shelf is topographically complex (Emery, 1972).

King and Maclean (1976) recognized three physiographic zones: an inner zone of rough topography lying between Cape Sable and St-Annes Bank, a central zone of isolated banks and basins and an outer zone of shallow, wide, flat banks. Within this framework, five physiographic and bathymetric subzones are observed:(fig. 1.1)

- 1/ The Banks: These are commonly as shallow as 50 m and include Baccaro, Roseway, LaHave. Emerald, Western, Sable Island, Middle, Canso, St-Annes, Misaine and Banquereau Banks.
- 2/ The Basins: These are as deep as 350 m and commonly 250 m , and include Roseway, LaHave, Emerald, Canso and Misaine Basins.
- 3/ "Between-Bank" depressions: These include the deeper areas between Banks such as the Scotian Gulf and that part of the Gully between the north arms of Banquereau and Sable Island Banks.
- 4/ Transitional areas between Banks and Basins: These include the area between Baccaro Bank and Roseway Basin, the Emerald Bank and Basin, the Nearshore zone and Central Basin complex. Only the Misaine Basin complex shows sharp transitions from banks to basins with deep localized

basins and intervening isolated banks.

5/ Inner Shelf Nearshore area: often of rough bedrock, sands and gravel.

The continental slope lies seaward of the shelf break and in this study, extends to depths of 3000 m. Transitional between the shelf and the abyssal plains, this area appears smooth topographically. However, several canyons dissect the slope in these regions. Hill (1982) has shown that, although smooth and featureless when compared to the shelf, even undissected areas of the slope are often morphologically complex, albeit on a smaller scale.

### 1.3.3. Surficial Geology

Much literature pertains to the surficial distribution of sediments on the Scotian shelf. In a series of papers on the subject, King and his co-workers have described this distribution (Drapeau, 1969; Cok, 1970; King, 1970; Maclean and King, 1971; Drapeau and King, 1972; Maclean, Fader and King, 1977), and recognized, on the basis of their acoustical data and samples, five surficial sedimentary units (Table 1.1). All of these essentially flat lying units unconformably rest upon pre-Pleistocene strata.

The central and outer shelf areas (Fig 1.1) are covered with well sorted sands and gravels and localised boulder fields (Table 1.1: Sable Island Sand and Gravel unit). Peripheral to these banks and predominating in the Scotian Gulf region is a somewhat finer sand unit (Sambro Sand unit) that is thought by King (1970) to be relict. The central basins of LaHave and Emerald, as well as the complex of basins

TABLE 1.1  
SURFICIAL SEDIMENTS OF THE SCOTIAN MARGIN  
(After King 1970)

MAP-UNIT	LITHOLOGY	THICKNESS (m)
Sable Island sand and gravel	medium to coarse grained, very well sorted sand. grades laterally to coarse, sub-rounded, glacial lag gravel. basal transgressive unit.	0-15
LaHave clay	very dark, loosely compacted silty clay, grades locally to clayey silt.	0-20
Sambro sand	medium to fine grained, moderately to well sorted, sub-littoral sand, grades to sandy gravel.	0-80
Emerald silt	poorly sorted clayey and sandy silt. some gravel.	0-200
Scotian Shelf drift	dark, cohesive sandy till. ground and end moraines.	0-90

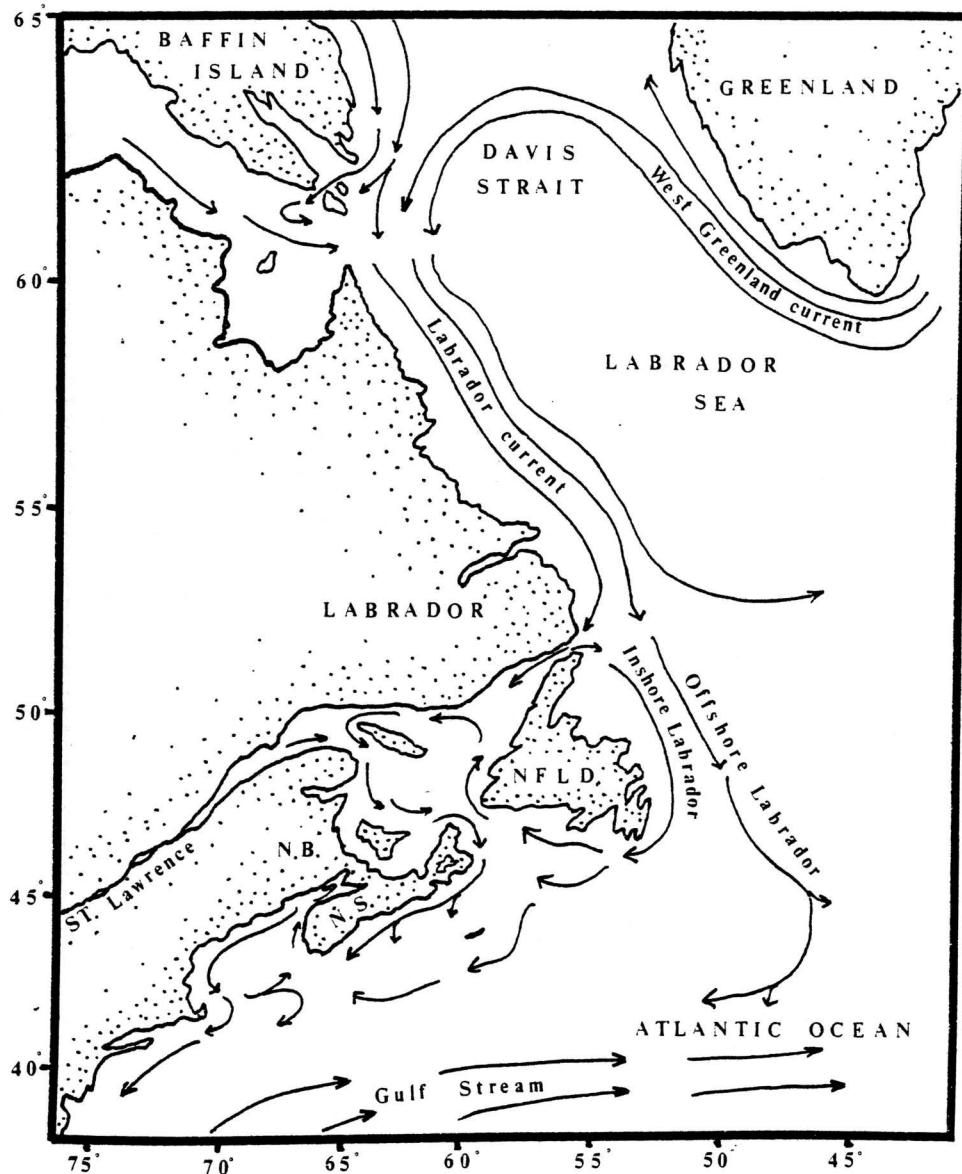
south of Misaine Bank, are carpeted by the LaHave Clay unit (Table 1.1). This clay includes contemporary sediment and fines winnowed from glacial deposits that subsequently filled topographic depressions. The Glacial Till (Scotian Shelf Drift, Table 1), mainly thought to be of Wisconsinan age, occurs in varying amounts on the central and inner zones of the shelf. Deposits of Emerald silt can be found at the periphery of topographic basins and are generally thought to be older than the LaHave Clay units. Similar to the Drift, this unit varies texturally by its absence of gravel. The inner shelf area consists mainly of coarse sands and gravels (Sable Island Sand and Gravel), and locally, of rough exposed bedrock.

#### 1.3.4. Oceanography

The waters off Nova Scotia possess some of the richest fishing grounds in the world. A consequence of this is a large body of oceanographic data and publications. Of direct interest to this study is the nature and distribution of bottom water masses and their influences on benthic foraminiferal distributions. This section briefly reviews the salient oceanographic features of the study area.

Surface circulation off Nova Scotia is dominated by three main elements (Dunbar 1951). The Labrador Current skirts the continental margin off Labrador and flows southward around Newfoundland and into the study area. Some Labrador water flows into the Gulf of St. Lawrence and enters the study area through the Cabot Strait (Fig. 1.2). This cold, relatively fresh water originates in the Arctic and includes some Baffin Bay and Beaufort Sea water (Dunbar, 1951). It continues off Nova Scotia as the Nova Scotian current of cool water flowing southwest. The Gulf

Figure 1.2 Surface currents off eastern Canada.  
( after Robe, 1971 )



Stream flows northeast. Branches of this warm, saline water mix with the Labrador Current to form the third element: the Slope water (McLellen, 1957).

The slope water has been shown to be of two parts. Gatien (1975) distinguished, through detailed temperature and salinity data, between warm slope water and cold Labrador slope water. Mixing of Labrador Current water and Gulf Stream water produces the warm slope waters to depths of 300-400 metres. Mixing of Labrador Current water and North Atlantic Central water produces the Labrador slope water directly overlying the continental slope down to 1200 m. At greater depths lies the North Atlantic Central water.

Bottom waters on the shelf are ultimately a product of mixture and advection of these watermasses. Using methods devised by Miller (1950) and first applied to the Scotian shelf by McLellen (1954), Houghton *et al.* (1978) described the distribution of Nova Scotian shelf waters. This work, together with McLellen's mapping of the bottom waters (1954) and data obtained from the Marine Environmental Data Service tapes, enable the distribution of shelf bottom waters to be shown (Fig. 1.3., Table 1.2.). Cold, relatively fresh bottom water predominates in the northeast of the study area. It overlies the Misaine Bank and Basin complex, St. Annes Bank, Canso Bank and Basin as well as Banquereau Bank. To the southwest, this water is pushed shorewards both by bathymetry (Middle Bank) and by an increased presence of warm, more saline water of slope origin. The cold water is present to the north and northeast margins of Emerald and LaHave Basins and in the Scotian Gulf. Although the presence of this slope derived water is in part due to tropical storm-generated water movements and eddies, Warner (1970) has shown its persistence throughout the year (although its boundary with

TABLE 1.2

WATERMASS CHARACTERS OFF NOVA SCOTIA  
 (SUMMARIZED FROM HOUGHTON et al. (1978))

WATERMASS NO.	SALINITY/TEMPERATURE	CHARACTER AND ORIGIN
1	32.3%	0-4°C Subsurface water from Cabot strait and St. Lawrence. Labrador current origin.
2	31.8%	3-6°C Surface water from Cabot strait and St. Lawrence.
3	33-34%	4-8°C Similar to 1, although warmer, a product of 1 and 4 (slope waters).
4	35%	8-12°C Water dominated by slope characters of maximum temperature and salinity. Some mixing with 1 and 2.
5	34.5-35%	4-6°C Deep Atlantic water.

Figure 1.3 Bottom hydrography on the continental  
margin off Nova Scotia

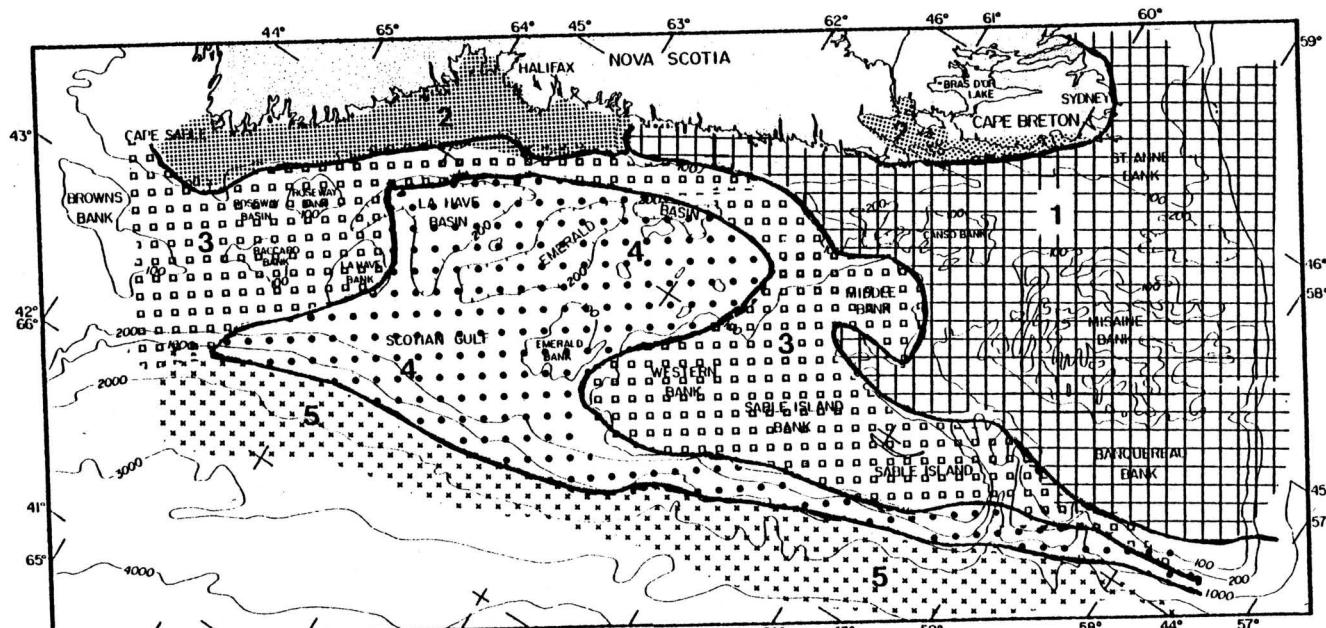


FIG. 1.3

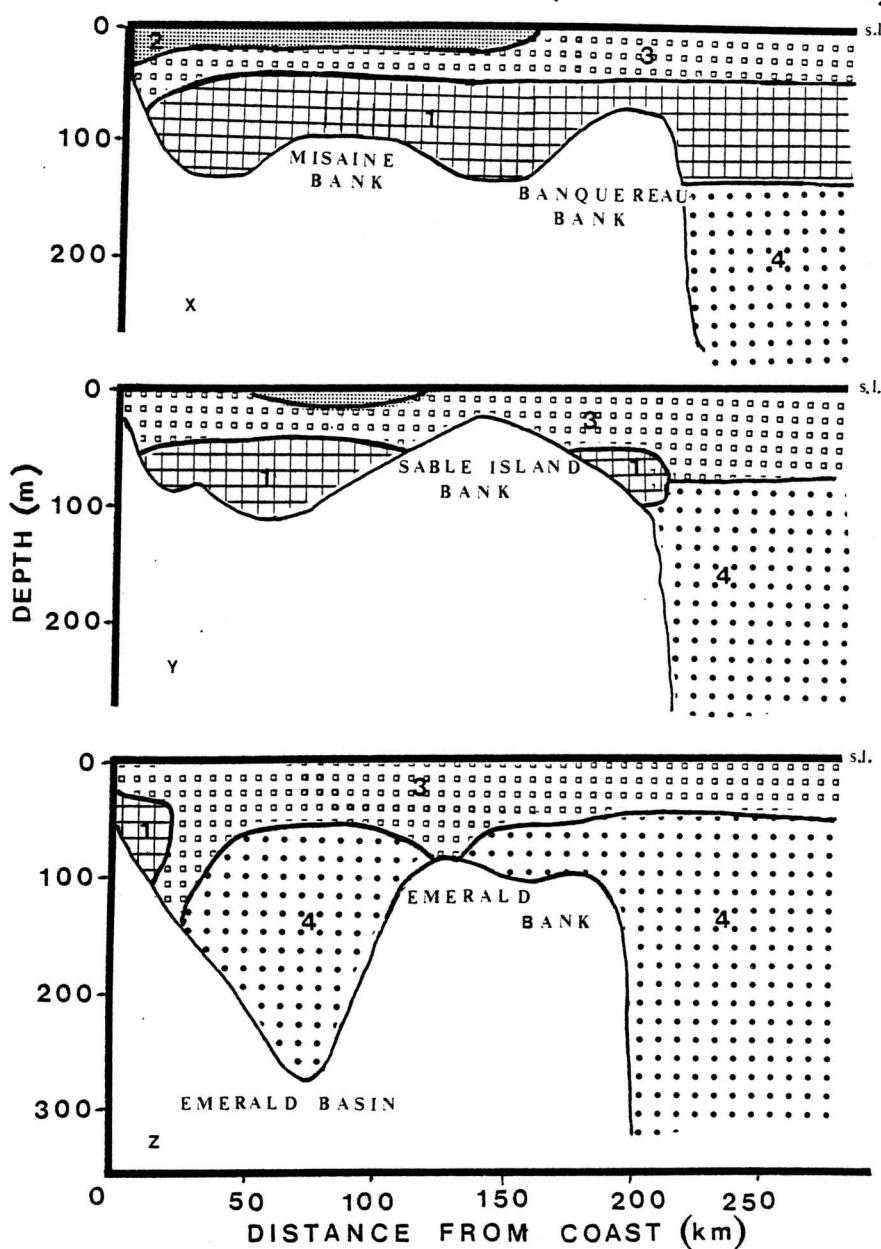
BOTTOM WATER DISTRIBUTION

SEE TABLE 1.2  
FOR EXPLANATION

Figure 1.4 Hydrographic character along sections  
X,Y,Z.

## SHELF WATERS OFF NOVA SCOTIA

(AFTER HOUGHTON 1978)



the cold water is subject to fluctuation). Sections 1,2 and 3, adapted and schematized from Houghton (1978), illustrate these relationships (Fig. 1.4.).

This enrichment of shelf waters with waters of slope origin is paralleled in the distribution of nutrients on the shelf. Fournier *et al.* (1977) have documented the sporadic advection of nutrient-rich slope waters onto the shelf off Nova Scotia.

The significance of this and other oceanographic influences upon the distribution of benthic foraminifera will be discussed later.

### 1.3.5. Sample Locations

As can be seen on the map (Fig. 1.5.), many of the distinct physiographic, sedimentologic and oceanographic regions in the study area have been sampled. Appendix I tabulates the latitude and longitudes for these locations. The means and methods of obtaining these samples will be discussed in a later section.

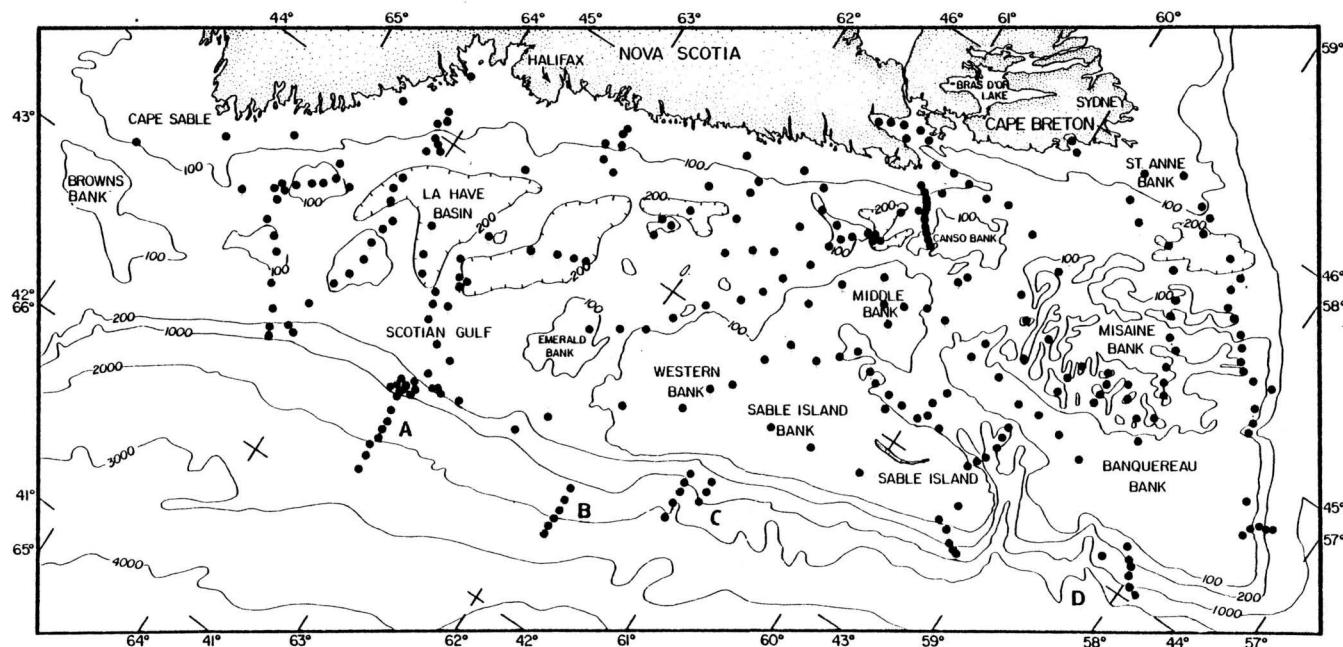
Many studies have documented the changes of species abundance with season ( Walker, 1955; Douglas, 1976; Ellison, 1978). However, the scale and nature of this study render seasonal consideration impractical. The ultimate purpose is to apply the Recent distribution model for foraminifera to time-averaged fossil populations, which tend to aggregate seasonal perturbations ( Scott and Medioli, 1980). Faunal patchiness, as documented by Buzas (1970), must also be considered. Types of sample networks that would account for seasonal and spatial variations in foraminiferal populations would be more restrictive and effectively prohibitive to the stated aims of this study. The possible effects of faunal patchiness and seasonality however, will be discussed

where pertinent.

### 1.3.6. Previous Foraminiferal Distribution Studies on the Scotian Shelf and Slope

The recent summary of Culver and Buzas (1980), documenting and analyzing the data available for recent foraminifera off the North American Atlantic Coast, shows some interesting locality concentrations. A dense coverage south of Georges Bank, with markedly less dense coverage on the margin off Nova Scotia and Newfoundland. The present study essentially fills in the Nova Scotian gap. Previous foraminiferal studies in the Nova Scotian area have hitherto been localized. Early work concerning the documentation of foraminifera off East Canada includes Dawson (1870) and Whiteaves (1901). Cushman, in a series of papers published between 1918 and 1931, recorded East coast U.S. and Canadian occurrences. Little was done until the work of Bartlett (1964), who documented occurrences in St Marys and Mahone Bays and, to some extent, on the inner shelf off Nova Scotia. Hooper (1970) and Shafer and Cole (1978) reported on Eastern Canadian foraminifera to the north and northeast of the present area. Similarly, Sen Gupta (1971) determined foraminiferal patterns on the Grand Banks. Hamdan's (1969) study concerned the ecology and distribution of foraminifera along a section of the shelf, from Halifax into Emerald Basin and the shelf edge, while Barbieri and Medioli's work (1969) concentrated on the western part of the shelf. Other areas studied include Cole and Ferguson's (1975) on Canso Strait foraminifera; Allen and Roda's (1977) on LaHave estuary; Walker's (1976), inshore, near Halifax; Clark's (1971), off Clam Bay; Gregory's (1970) in Halifax harbour and Scott and

Figure 1.5 Grab sample locations. Slope transects  
are marked A,B,C,D.



Medioli's (1979) off Chezzetcook. The most recent work by Rodrigues and Hooper (1982) documented benthic associations from the Gulf of St Lawrence.

All the above studies are similar in that they endeavour to document and show distribution features in relatively localized areas. This study is the first to attempt an overall synthesis of the regional benthic foraminiferal distribution off Nova Scotia. Reference to other studies off the eastern continental margin of North America will be made where pertinent

## CHAPTER 2 : METHODS

## 2.1. Shipboard

All samples were collected from Bedford Institute of Oceanography research vessels on the following cruises: February 1979, CSS Dawson; March 1980, CSS Dawson; October 1980, M. V. Pandora II; March 1981, CSS Dawson; May 1981, CSS Dawson. Cruise numbers and sample localities are shown in appendix I.

Control of sample type and quality is important in a study such as this. The type of sampling instrument used is known to affect the size and quality of the sample ( Douglas, 1978). Piston and gravity core tops have inherent inadequacies in preserving the surficial layer of sediment, especially in silty mud substrates ( Douglas, 1978). This is largely due to the flushing effect of a preceding bow-wave. Ideally, a large box core (100 cm<sup>2</sup>) would have been used. However, due to impracticalities of its use in rough sea conditions, length of time taken to deploy and recover the instrument as well as its unsuitability for sandy substrates, an alternative method was sought.

The device used was a Shipek grab sampler. This consistently produced good quality samples and was employed fast enough to enable the collection of many samples. A grading system similar to that of Douglas (1978) was devised in an effort to provide a quality control of the samples. The grading system is outlined in Table 2.1. All samples examined and analyzed rated 1-3. Other poorer samples were rejected. (

TABLE 2.1

## RATING INDEX FOR SHIPEK GRAB SAMPLES

## Rate

- |   |                                                                                                                          |
|---|--------------------------------------------------------------------------------------------------------------------------|
| 1 | Surface mm present over entire sample; supernatant water clear sub-sample recovery good.                                 |
| 2 | Surface mm present over part of surface only; supernatant water moderately to slightly turbid. Sub-sample recovery fair. |
| 3 | Surface present over part of sample only; supernatant water strongly turbid due to ship recovery and drainage.           |
| 4 | Surface apparent, sufficient to warrant sample; very turbid waters.                                                      |
| 5 | Sample extremely poor in quality and quantity sample taken.                                                              |
| X | Sample not rated, seaweed/laminaria/pebbles.                                                                             |

Comments: On sand and gravelly substrates if the larger clasts appear in situ (ie. as defined by algal growths etc.) rate 1. If macro-benthos are in life position rate 1 Gravity cores rate 4.

It was necessary to use a few gravity core tops to fill gaps in the sample grid. These are noted in the appendix.).

Procedures concerning sample recovery aboard ship were as follows: Immediately that the Shipek came shipboard, the quality of the sample was noted, together with the recording of other location data ( i.e. depth, latitude and longitude, sea state, sediment estimates, etc). Sediment was then scraped off the surface (20 ml) and placed in a plastic vial. This was subsequently topped with seawater and formaldehyde (buffered with borax), final pH of the liquid= 8.0

## 2.2 Laboratory.

In the laboratory, the samples were washed through a 63 micron sieve and stained in Rose Bengal to determine live foraminifera. After 24 hours, the samples were filtered, dried in air and concentrated by flotation in carbon tetrachloride. The samples were floated as many times as necessary to retrieve the maximum amount of foraminifera. The next step involved making a census of foraminifera in each sample or part of a sample. The census involved a determination of the relative abundance of each species in a sample. Other population characters evaluated and recorded at this stage included: the number of species (S), total number of individuals, live/total ratios, total non-live individuals and diversity values of S, H(S), E and  $\alpha$ . The latter will be discussed in chapter 5.

### 2.3. Discussion of Laboratory Methods

Although the methods and procedures outlined have been in use for many years, some discussion is warranted.

Three stages in the laboratory analysis have been described.

1. Initial treatment to clean and remove clay and silt.
2. Staining to determine live foraminifera.
3. Population counts on the samples.

Each stage may vary in approach and method from study to study.

1. All species retained on the 63 micron sieve were examined in this study. Other workers have often used different size fractions. Streeter (1973) and Miller and Lohmann (1981) used 250 microns and it is common practice in industry to restrict an investigation to a particular size range ( Haynes, 1981). There are two main reasons for looking at greater than 63 microns in this study:

a/ to take into account small species such as Fuskenkoina fusiformis, Adercotryma glomerata and Reophax arctica, which may otherwise be lost.

b/ to retain juveniles; important in considering the effect of transportation and other phenomena, i.e. winnowing resulting in juvenile loss.

2. The ambiguous nature of Rose Bengal staining techniques is well documented in the literature ( e.g. Douglas, 1978). The main problems result from staining of bacteria and other organic material in the

chambers of dead tests and difficulty in recognizing stains in thick shelled calcareous species and agglutinates.

Great care was taken in this study in the determination of live species. Individuals were recorded as living only if several chambers were stained. Some of the more heavily agglutinated forms were broken in order to determine the extent of staining.

An alternate method of staining was described by Walker et al (1974) and involved the use of Sudan Black. Difficulties involved with this method, however, precluded its use in this study.

Even with diligent live determinations, some ambiguity remained in the live data. Indeed, the live data proved unsuitable for exhaustive statistical treatment and is mainly used here as a tool to help determine the " *in situ*" nature of each assemblage (Chapter 4).

3. Counts of about 300 individuals have long been accepted as a practical method of accurately estimating species proportions within a sample. The derivation of this figure and other associated counting errors are described in Douglas (1978). Dennison and Hay (1967) and Dryden (1931) statistically estimated that a minimum of 300 specimens must be counted in order to detect a species that constitutes 1% of the total population with a certainty of 95% ( Douglas, 1978).

Some workers have used samples with as few as 100 specimens (and some even less) as a basis for quantitative analysis of ecology and palaeoecology ( Streeter, 1980; Schnitker, 1974; Osterman and Kellogg, 1980). Clearly, any ecological information obtained from low population counts have to be carefully evaluated. In the present study, between 250-350 individuals per sample were counted. Statistical aliquots were obtained using an Otto microsplitter where necessary.

In this study, all samples or parts of a sample were examined whilst dry. Douglas (1976) showed that, in some cases, wet counting procedures tend to introduce errors into the counts and had a low reproducibility. He also documented the difficulties involved in splitting a water/sediment mixture into statistical aliquots. The more efficient procedure and smaller size-shape bias in splits, determined that dry counting methods be employed in this study.

Finally, brief mention should be made of the foraminiferal identification and storage methods used.

The foraminifera were picked from a standard picking tray and placed into a gridded rectangular tray. This tray of around 300 individuals was then subjected to genetic and specific identification. Representative specimens of each species identified were stored in a one-hole slide as a reference collection. Identification methods were standard and are described in Haynes (1981).

#### 2.4. Data analysis methods

Some 382 samples were available for foraminiferal analysis. Low sample quality reduced this number, as did the occurrence of barren samples and others containing less than 200 individuals. Investigation of the remaining 267 samples revealed the presence of 120 species ( of hyaline, agglutinated and porcellaneous types) resulting in a matrix of 267 samples by 120 species (N by n ) represented by their percentage abundance. This initial matrix was reduced to 75 species through elimination from further analyses of species not occurring greater than

3% in any one sample. Clearly, the designation of biofacies or species associations by simple observation was not feasible with this amount of data. As well as unreliable, such an interpretation would not be in keeping with the stated aims. A statistical method was needed to make the huge data set more manageable. Computers provide the means for this data reduction (Buzas, 1970): several procedures and methods have been developed over the past 15-20 years (Joreskog, 1976; Buzas, 1979).

The basic concepts of factor, principal component, cluster and canonical analysis are outlined in Buzas (1979). From the above literature, it is clear that a wide variety of techniques are available to suit particular needs. Cluster analysis is commonly used in palaeontology. First devised for such use by Sokal and Sneath (1963), it has been used successfully by Kafescioglu (1975), Schafer (1978) and Brooks (1973) *inter alia*. Clustering algorithms either compare species with each other and their joint occurrence throughout the stations (R-Mode); or stations are compared based on joint occurrences of foraminiferal species at each station (Q-Mode). The end result is a dendrogram i.e. a visual representation of the objects as grouped by the analysis. Although simple to understand, the method is not statistical and will pair off all the data provided. This, and the fact that it is based on simple presence/absence data or rough qualitative evaluations, rule out its use in this statistical study.

What is required for the present work is a method that will simply and objectively reduce a large body of data to a more manageable sized data set in which the relationships between samples and their species composition can be clearly seen. A quantitative method that can do this is factor analysis, in particular the method of Imbrie and Kipp (1971). Malmgren and Haq (1982) have recently shown that this method (Imbrie and

Kipp, 1971) is in fact a principal component analysis sensu stricto; however in order to be more easily compared with Imbrie and Kipp's results their terminology is retained. The method is capable of simplifying large data matrices by grouping co-varying species into statistically independant assemblages.

The geological application of this multivariate statistical technique has been documented by Imbrie and Kipp (1971) and Joreskog et al (1976), Kipp (1976) Molfino et al.(1982). In particular, it has been shown to be a useful quantitative method of data reduction in works with similar aims to the present ( Imbrie and Kipp, 1971; Osterman and Kellogg, 1979; Streeter and Lavery, 1980; Belanger and Streeter, 1980; Miller and Lohman, 1980; Mudie, et al. 1983; Lohmann, 1979).

The mathematical theory behind Q-Mode factor analysis is described in Imbrie and Van Andel (1964) and Manson and Imbrie (1964). The factor analytical computations in this study were carried out using a CABFAC program devised by Klovan and Imbrie (1971) and modified by C.E.Keen (Mudie et al 1983). This program gives a varimax solution to Q-Mode factor analysis, reducing the large initial data matrix into two smaller matrices. The varimax factor matrix describes the importance of each factor ( or assemblage) at each station, and the factor score matrix reveals the importance of each species in each factor ( or assemblage).

The computational steps required to derive these matrices are illustrated in Imbrie and Kipp (1971). They show that the varimax factor matrix is derived through calculation of the eigenvalues and eigenvectors of the row-normalized N by n matrix. The factor score matrix is a result of rotation to the "best fit" ( according to the varimax criteria) of the eigenvectors. ( The row-normalized matrix is a result of the multiplication of raw % values by a constant so that the

transformed row vectors have unit length). Through the initial normalization is derived the sample communality ( square of the vector lengths) which is a useful measure of the degree to which a sample is adequately described by the factors.

Close inspection of the factor loading values displayed in the varimax factor matrix enables one to determine whether the produced factors ( or assemblages) are statistically and ecologically significant. They are assumed to be so if the factor loadings vary with known bathymetric, sedimentologic or oceanographic environmental gradients ( Imbrie and Kipp, 1971). The corresponding species compositions for each factor are shown by the factor score matrix.

The following chapters (3,4) of this thesis show the results of these analyses on the foraminiferal data obtained from the 267 samples. For reasons to be explained, the shelf and slope data were treated separately. Several maps plot the distribution of the factor assemblages determined ( factor loadings are listed in Appendix II, III) whilst several sections across assemblage boundaries on the shelf illustrate their nature and gradient. The specific compositions of each factor assemblage are summarized by the factor score matrix shown in the appendices.

The proper perspective of this type of analysis should be borne in mind. This is highlighted in Imbrie and Kipp (1971):

" The proper role of the machinery of matrix algebra used in factor analysis is to summarize parsimoniously, objectively and quantitatively general tendencies in a given body of faunal data, to make tendencies clear which are obscured by random error, or other, quantitatively

dominant effects; and to focus attention on ecological or diagenetic questions worthy of further inquiry".

In this way, the techniques described permit the primary objective of the thesis to be attained. The secondary aims ( i.e. assemblage verification using live data) are also reached in chapter4. The tertiary objective, i.e. quantifying the relationship of assemblage and environment, is investigated in chapters 6-9. The factor analysis results can be used to do this in the form of multiple regression analysis techniques. The relevant theory and procedures will be discussed in chapter 6.

## CHAPTER 3 ASSEMBLAGE DISTRIBUTION AND COMPOSITION

## 3.1 Introduction

This chapter describes the results of Q-Mode factor analysis of the total foraminiferal (live and dead) abundance data.

In order to maximise the advantages of the analytical technique and to overcome certain inherent ambiguities within the data, three groups of abundance data were analysed.

- a/. Group One data: 258 shelf and slope stations;
- b/. Group Two data: 223 shelf only stations;
- c/. Group Three data: 43 slope only stations.

The descriptions of each assemblage are taken from the varimax factor matrix and the factor score matrix (appendices II, III). Reference will be made to several maps which plot for each assemblage the appropriate factor loadings ( i.e. assemblage importance) and to several figures which illustrate the variation of these with depth.

## 3.2.1. Group 1 Shelf and Slope Data

Q-Mode factor analysis of 258 stations determined 9 factor assemblages accounting for 85% of the observed variance. Table 3.1 summarizes the factor analysis and interpretive results.

It can be seen that ,with two exceptions, several well defined

TABLE 3.1  
FACTOR SCORES FOR SHELF AND SLOPE ANALYSIS--A SUMMARY

FACTOR ASSEMBLAGE	% VAR.	DOMINANT SPECIES WITH FACTOR SCORES		DISTRIBUTION	SUBSTRATE
1	24.9	<u>Adercotryma glomeratum</u> <u>Recurvirodes turbinatus</u> <u>Spiroplectammina biformis</u> <u>Cribrostomoides crassimargo</u>	8.1 1.7 1.2 1.2	N.E. Shelf between Cape Breton shore and Middle and Banquereau Banks and basins. 80-350 m	sand, gravels, silts, clays.
2	13.1	<u>Cibicides lobatulus</u> <u>Cassidulina reniforme</u>	8.5 1.2	Mainly on banks and innershore. 50-100 m	coarse sands, bedrock.
3	8.6	<u>Elphidium excavatum</u> <u>Uvigerina peregrina</u>	7.1 4.5	Deep waters (>2000 m) and on shelf near Emerald basin.	clays, silts.
4	7.8	<u>Olobulimina auriculata</u> <u>Nonionellina labradorica</u> <u>Buliminella aculeata</u> <u>Brizalina subaenariensis</u> <u>Isonidella teretis</u>	6.4 3.6 2.9 2.1 1.7	Central shelf basin complex. Emerald, LaHave, Roseway. 150-220 m	clays, silts.
5	7.6	<u>Saccammina atlantica</u> <u>Reophax scorpiurus</u> <u>Cribrostomoides Jefferyi</u>	8.0 1.9 1.1	Periphery to central basin complex.	clays, silts, sands.
6	7.4	<u>Trifarina angulosa</u> <u>Brizalina spathuleata</u> <u>Buccella frigida</u>	-7.9 2.0 1.0	Shelf break regions. 150-400 m	sands, silts.
7	4.7	<u>Eggerella advena</u> <u>Spiroplectammina biformis</u> <u>Cribrostomoides crassimargo</u>	8.2 1.6 1.0	Chedabucto and Gabarus Bays, Sable Island Bank.	muds, silts, well sorted sands.
8	5.0	<u>Cassidulina reniforme</u>	8.4	Same as F.2.	see F.2.
9	5.6	<u>Buliminella aculeata</u> <u>Buliminella exilis</u> <u>Furcarenkonia fusiformis</u> <u>Elphidium excavatum</u> <u>Trifarina occidentalis</u> <u>Buccella frigida</u> <u>Nonionellina turgida</u>	3.8 3.4 3.3 3.0 2.3 2.3 1.2	Middle to Upper slope regions. Scotian Gulf near Emerald Basin.	silts.

assemblages are present on the shelf and slope.

Two assemblages, factors 3 and 9, create a problem (Table 3.1) seen in both their distribution and composition. The distribution of factor 3 in deep water of greater than 1000 m. depth; in shallow waters of 150-200 m. around Emerald Basin; and its composition,(ie. dominantly Elphidium excavatum and Uvigerina peregrina) conflicts with what is observed in the raw data. The latter reveals a complete absence of Uvigerina peregrina at any depths less than 900 m. This would be misleading in the light of its occurrence as a major component of factor assemblage 3.

The discrepancies are due to the occurrence of large amounts of Elphidium excavatum present on both the slope (at depths greater than 1000 m.) and the shelf. This distribution enables the analytical technique to 'lump' together two assemblages, one that is clearly deepwater and one that is shallow-water. The problem of combining two assemblages together by virtue of their possessing similar dominant species is an important one to recognise, in view of the potential for relict and transported populations to dilute and flood otherwise dissimilar samples with a common species.

Factor 9 represents a similar anomaly. The result there is more complex due to the greater number of species involved.

In order to more rigorously define these assemblages, factor analysis was applied to "shelf only" (Group 2) and "slope only" (Group 3) data matrices. This removes the misleading common shelf/slope components from both the shelf and slope distributions. Later discussion will show how, in the light of the living data collected, most of the assemblages reflect the distribution patterns of the main species presently living in the area.

The following describes the results and interpretations obtained from this procedure.

### 3.2.2. Group 2 Shelf Data

Factor analysis of 223 shelf stations delimited 8 factor assemblages accounting for 87.5% of the variance. The assemblages are essentially the same (ie. composition and distribution) as those defined from Group One 1 (Table 3.1) The exclusion of slope stations removed the problems highlighted in the previous section. Table 3.2 summarizes Group 2 analysis results; Figures 3.1-3.9 plot assemblage distributions.

The following is a more detailed description of these assemblages, their compositions and their distributions. The assemblages thus defined can be seen to be more in accordance with the observed raw data. This also explains more of the between sample variance and the higher communalities. The assemblages are named after the most dominant species as defined by the factor scores (appendix III).

#### Factor assemblage 1: Adercotryma glomerata

This completely arenaceous assemblage accounts for 29.4% of the observed variance and is the most important assemblage found on the shelf. It is predominant throughout the north east of the study area and is restricted to central and innershelf regions (fig. 3.1) The southern

TABLE 3.2

## FACTOR SCORES FOR SHELF ANALYSIS--A SUMMARY

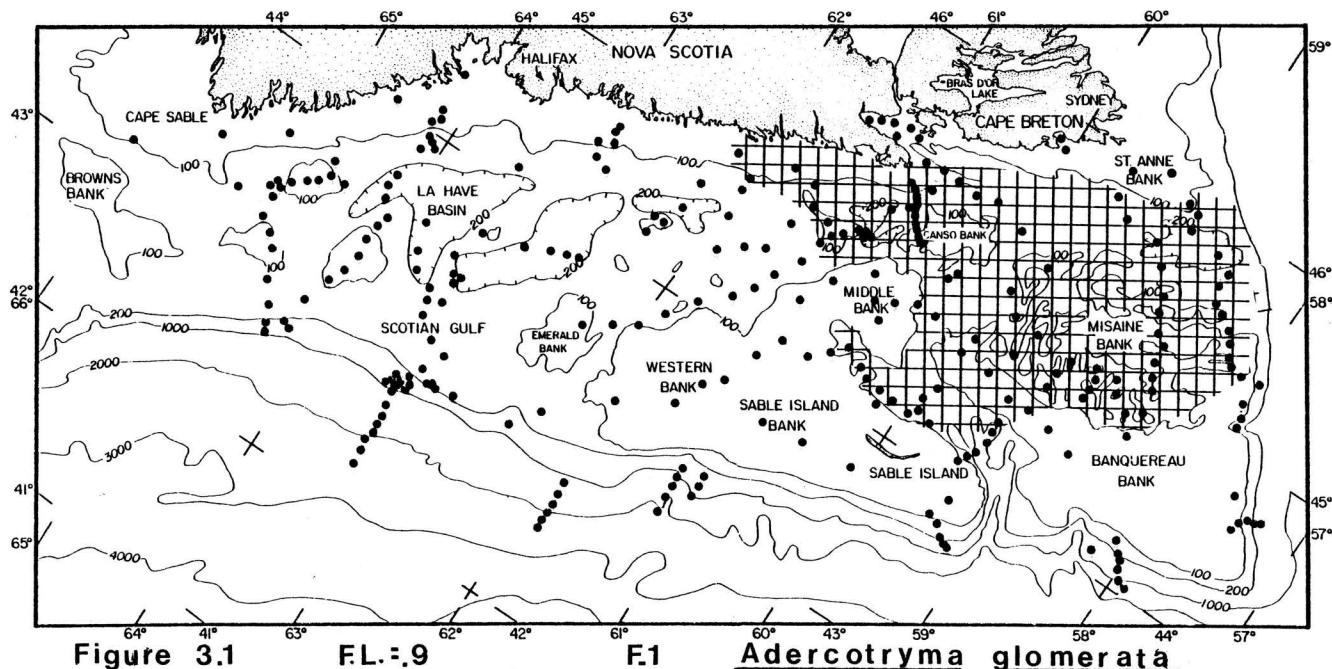
FACTOR ASSEMBLAGE	% VAR.	DOMINANT SPECIES WITH FACTOR SCORES	DISTRIBUTION	SUBSTRATE	
1	29.4	<u>Adercotryma glomeratum</u> <u>Recurvirodes turbinatus</u> <u>Spiroplectammina biformis</u> <u>Cribrostomoides crassimargo</u>	7.7 1.6 1.1 1.2	As for F.1 in Table 3.1.	As for F.1 in Table 3.1.
2	14.9	<u>Cibicides lobatulus</u> <u>Cassidulina reniforme</u>	8.0 1.0	As for F.2 in Table 3.1.	As for F.2 in Table 3.1.
3	9.0	<u>Globobulimina auriculata</u> <u>Nonionellina labradorica</u> <u>Buliminia aculeata</u> <u>Brizalina subaenariensis</u> <u>Islandiella teretis</u>	5.8 3.3 3.1 2.3 1.6	As for F.4 in Table 3.1.	As for F.4 in Table 3.1.
4	8.4	<u>Saccammina atlantica</u> <u>Reophax scorpiurus</u> <u>Cribrostomoides Jefferys</u>	5.8 1.7 1.0	As for F.5 in Table 3.1.	As for F.5 in Table 3.1.
5	7.7	<u>Trifarina angulosa</u> <u>Brizalina spathulaeata</u> <u>Buccella frigida</u> <u>Elphidium excavatum</u>	7.6 1.5 0.6 1.1	As for F.6 in Table 3.1.	As for F.6 in Table 3.1.
6	5.4	<u>Eggerella advena</u> <u>Spiroplectammina biformis</u> <u>Cribrostomoides crassimargo</u>	7.8 1.5 0.8	As for F.7 in Table 3.1.	As for F.7 in Table 3.1.
7	5.9	<u>Cassidulina reniforme</u>	8.0	As for F.8 in Table 3.1.	As for F.8 in Table 3.1.
8	6.1	<u>Elphidium excavatum</u> <u>Buliminia aculeata</u> <u>Furstenkoina fusiformis</u> <u>Buccella frigida</u>	-6.0 -2.7 -2.5 -1.9	Within and approaching the southern extremities of Emerald Basin.	silts, clays, sands.

limit of this assemblage is bounded by the northern edges of Banquereau, Sable Island and Middle Banks. It characterizes both banks and basins within this area (Misaine and Canso). The assemblage is seen in three other smaller areas but with much reduced factor loadings (ie. importance). The relatively wide depth range of this shelf assemblage (fig. 3.14, 50-300 m.) is a reflection of its occurrence both in banks and basins in this region of the shelf. Other significant but less important species include Recurvooides turbinatus, Cribrostomoides crassimargo and Spiroplectammina biformis. Within the limits of its main occurrence (fig. 3.1) this assemblage only becomes sub-dominant in two areas: in shallow bays (Chedabucto, Gaberouse; where factor 8 predominates) and wherever substrate character favours Cibicides lobatulus (assemblage 2).

Factor assemblage 2: Cibicides lobatulus.

Accounting for 14.9% of the variance, this calcareous assemblage is marked by its sporadic but widespread spatial occurrence (fig.3.2) It occurs on parts of the outer banks (Banquereau, Baccaro, LaHave, Emerald), the central banks (Roseway, Canso, Misaine) and the inner shelf areas of the Eastern Shore and St-Annes Bank. The restriction of this assemblage (which is almost monospecific) to coarse sands, rough bedrock and boulders throughout the study area seems to support the many reports suggesting a substrate related occurrence for this species. The absence or slow rate of sedimentation in the areas where this assemblage occurs (mainly between 50-150 m. ; fig. 3.2) is reflected in the very large total foraminiferal counts.

Figures 3.1--3.8 Spatial distribution of each factor  
assemblage (F.1--F.8) on the shelf.  
Areas with factor loadings greater  
than that shown (i.e. F.L.=) are shaded.



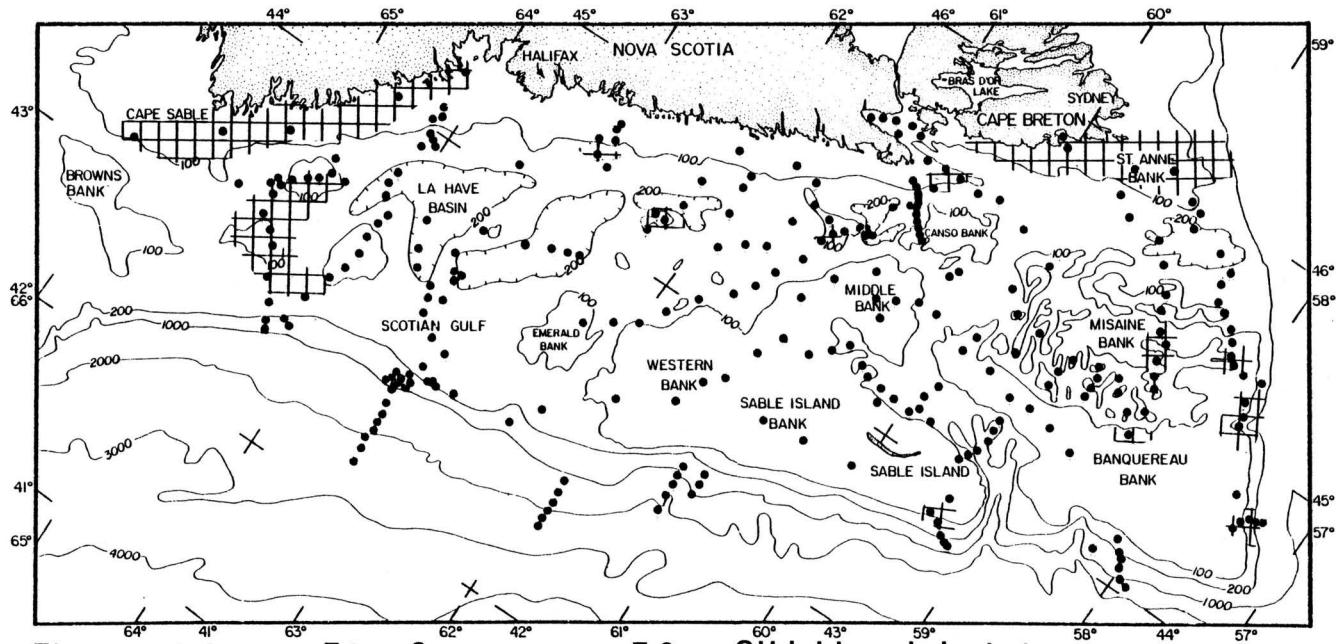
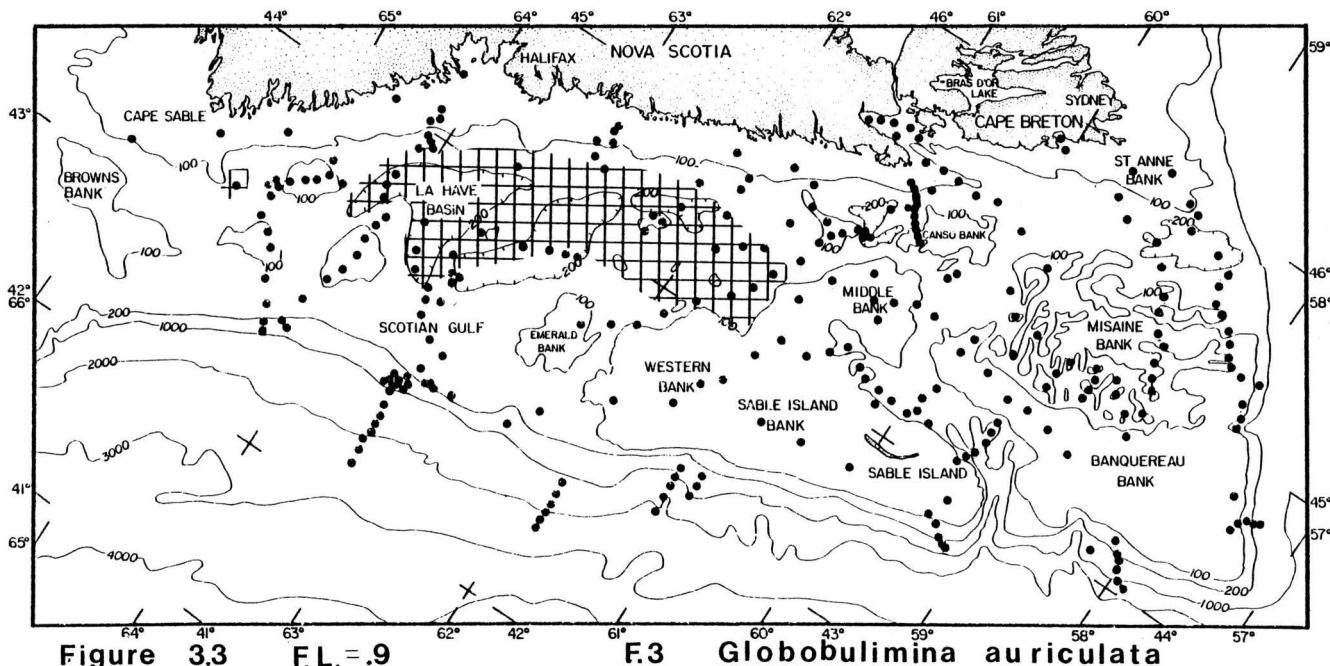


Figure 3.2



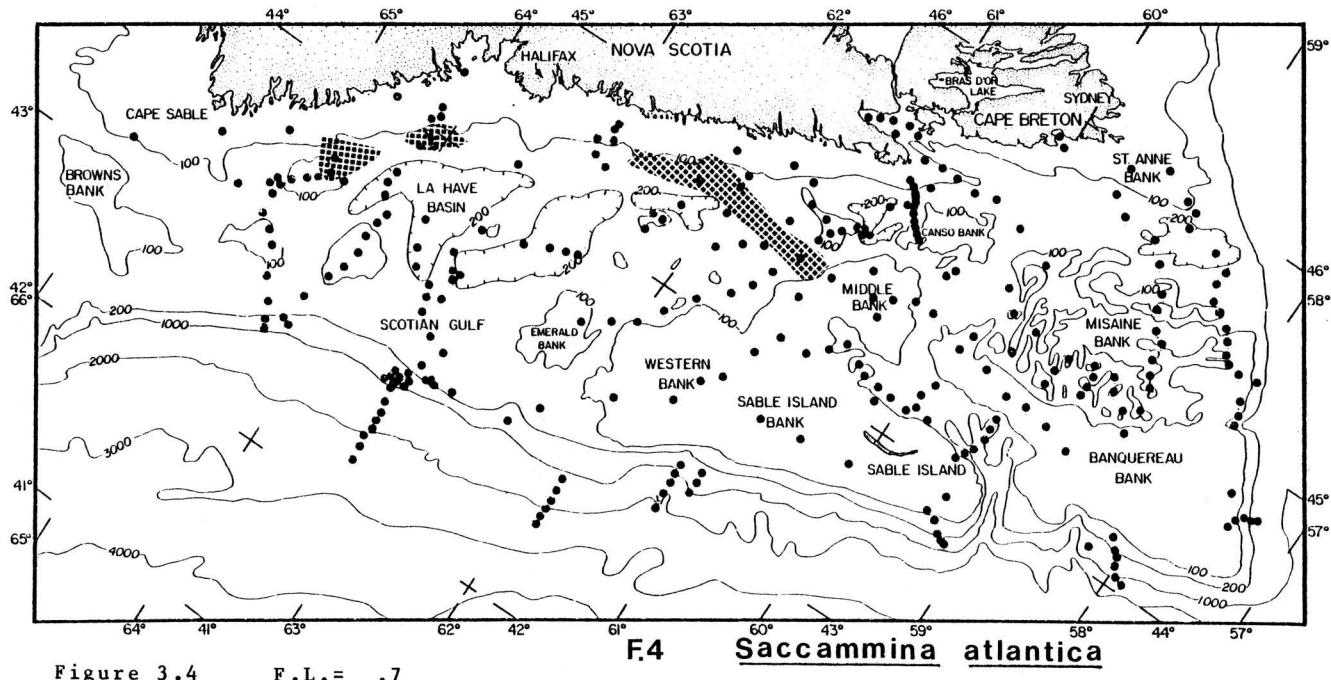


Figure 3.4

F.L. = .7

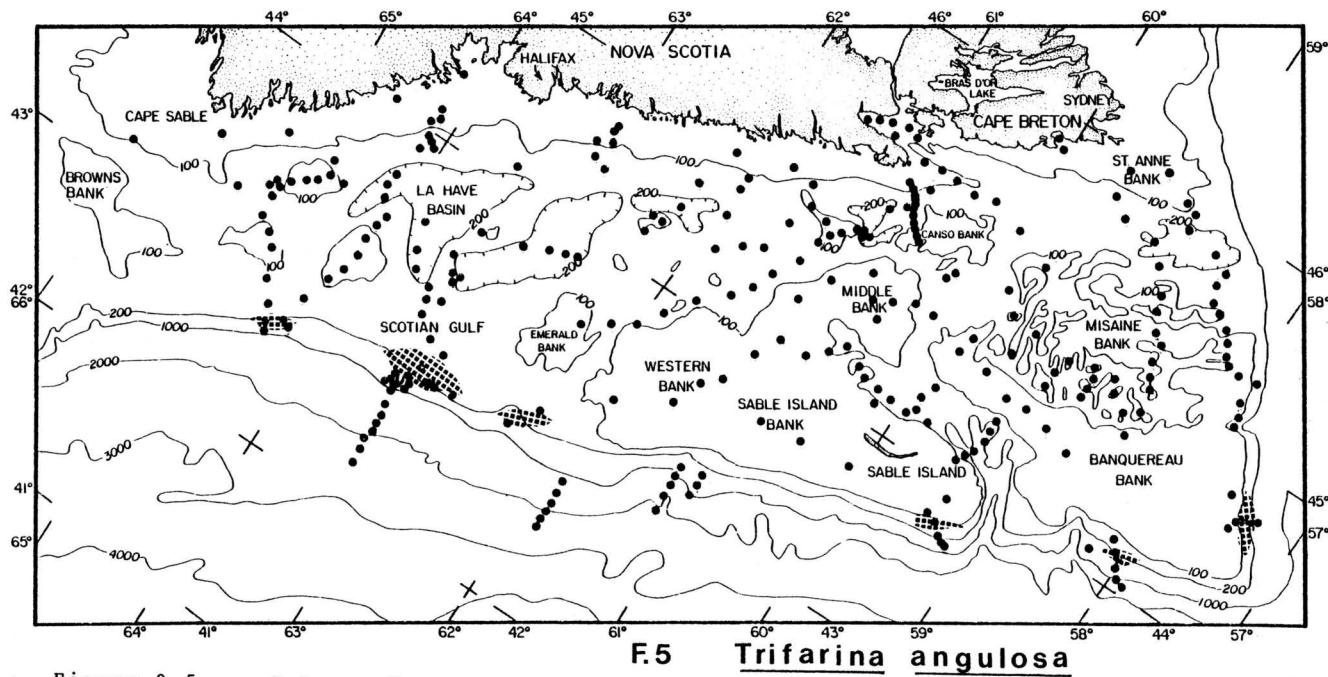


Figure 3.5

F.L. = .7

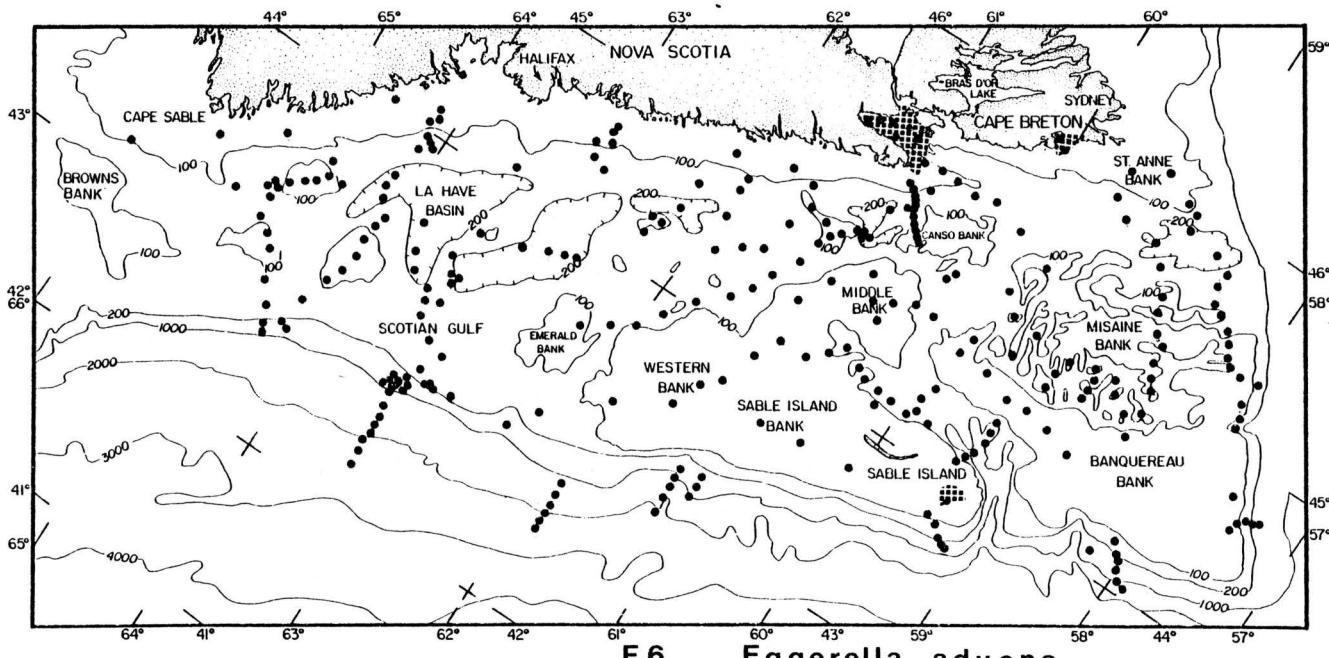


Figure 3.6

F.L. = .9

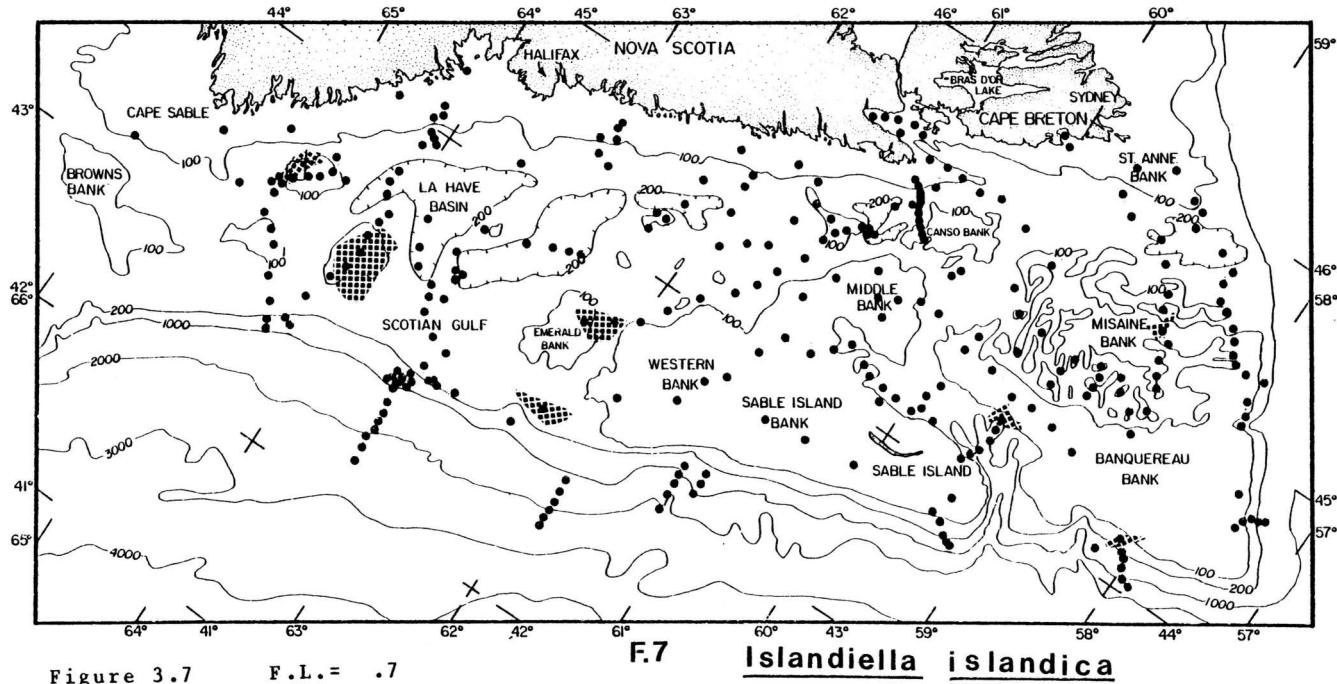


Figure 3.7

F.L. = .7

F.7

*Islandiella islandica*

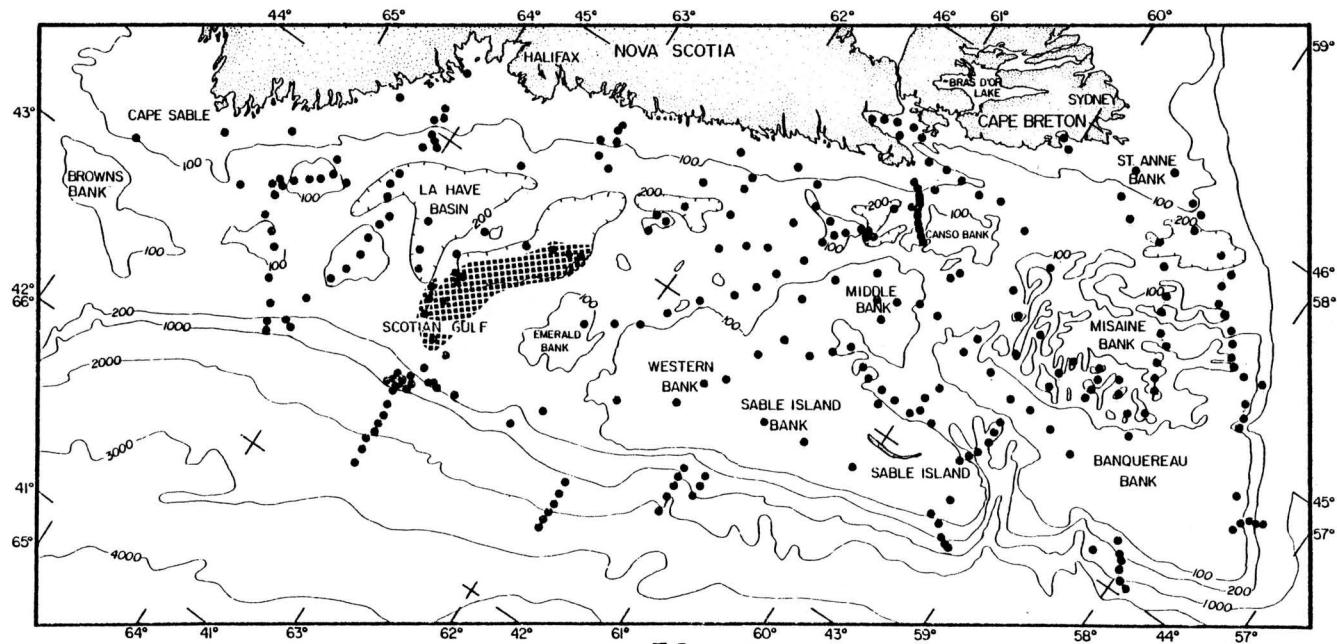
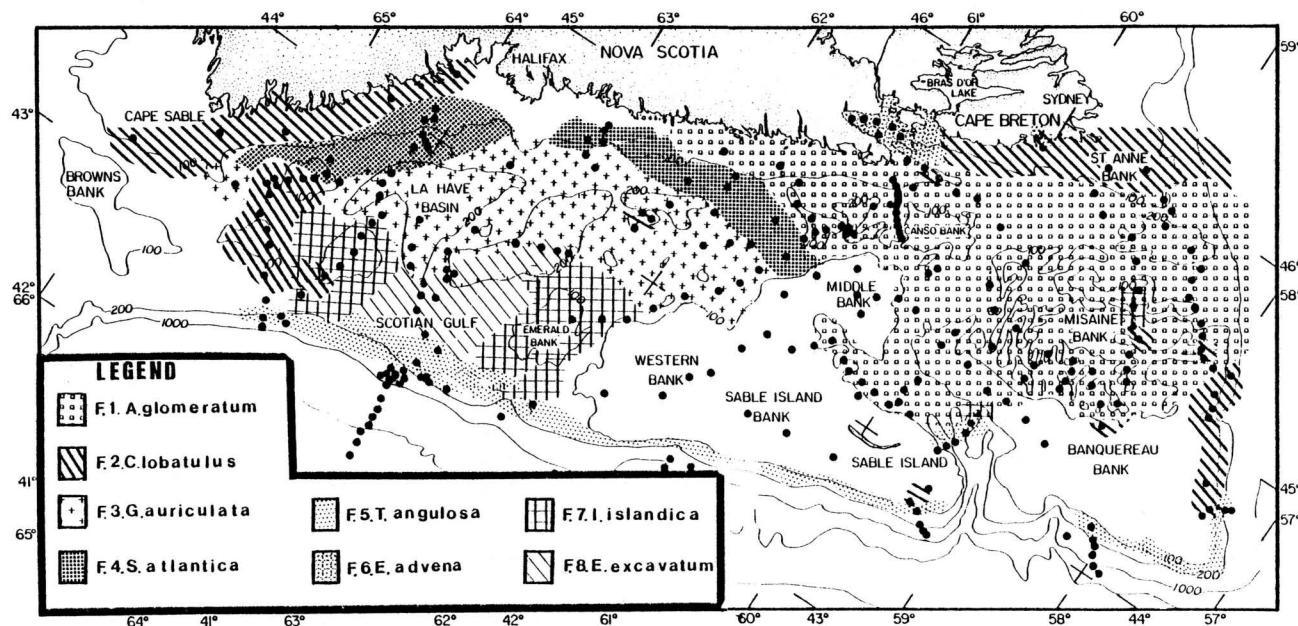


Figure 3.8

F.L. = .8

Figure 3.9 Summary of shelf factor assemblage distributions.



Factor assemblage 3: Globobulimina auriculata.

This calcareous assemblage accounts for 9% of the observed variance. Its species diversity is higher than any of the other assemblages and it occurs only in the central shelf basin areas of Emerald, LaHave and Roseway (ie.fig. 3.2 in LaHave clay and Emerald silts). The northern limit of this assemblage appears gradational with factor assemblage 5, the latter forming a rim to the central basin complex. Other subdominant but important components of this assemblage include Nonionellina labradorica, Bulimina aculeata and Bolivina subaenariensis. The assemblage does not occur at all in those basins to the northeast of the area. Depth distribution is restricted to 150-250 m., ie. basin depths.

Factor assemblage 4: Saccammina atlantica.

Accounting for 8.4% of the variance, this arenaceous assemblage is restricted entirely to the northern periphery of the central basin complex (assemblage 3). It is characterized by low factor loadings (ie. proportional significance) and is not predominant at any station being associated with assemblages 1 and 3 (Appendix II). Although in lesser amounts than the extremely dominant Saccammina atlantica; Reophax scorpiurus and Cribrostomoides jeffervysi also occur. The assemblage is found between 90-200 m. depth.(fig. 3.4)

Factor assemblage 5: Trifarina angulosa.

This calcareous, relatively diverse assemblage, accounts for 7.7% of the variance. It is completely restricted to shelf edge regions, being found slightly downslope in some places (fig. 3.5.) Its depth distribution is mainly 150-250 m. (shelf edge), with deeper occurrences at 300-400 m. Other important species (Table 3.2) include Bolivina spathulata and Buccella frigida.

Factor assemblage 6: Eggerella advena.

This entirely agglutinating assemblage accounts for only 5.4% of the variance. Its occurrence is very restricted, being confined to two areas (fig. 3.6). It dominates in the northeast portion of the study area, in the inner shelf bays of Chedabucto and Gaberouse. Its other occurrence is south of Sable Island, on the well sorted sands of Sable Island Bank. The assemblage is almost monospecific; other species include Spiroplectammina biformis and Cribrostomoides crassimargo.

Factor assemblage 7: Islandiella islandica.

This calcareous assemblage accounts for 5.9% of the variance. This almost monospecific assemblage is restricted entirely to shelf occurrences, mainly on the banks. It generally co-occurs with factor assemblage 2 (Cibicides lobatus), mostly on the outer banks but also on the inner shelf along the Eastern Shore. As with factor 2, this assemblage again occurs upon coarse rocky substrate (fig. 3.7).

Factor assemblage 8: Elphidium excavatum f. clavata.

This calcareous assemblage accounts for 6.1% of the variance. It occurs on the shelf between LaHave and Emerald Banks (Scotian Gulf), in the approaches to Emerald Basin (fig. 3.8) To the north, it gives way to assemblage 3 and southwards, to the shelf edge factor assemblage 5. Other species present include Bulimina aculeata, Fursenkoina fusiformis and Buccella frigida.

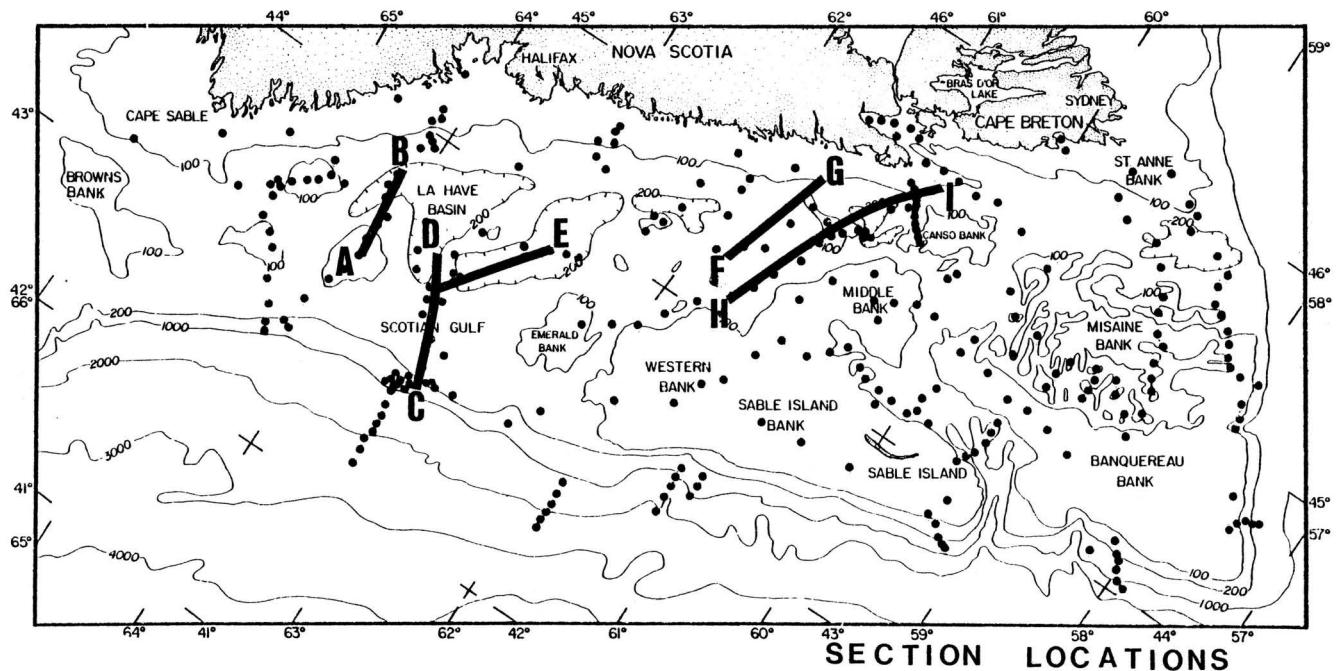
### 3.2.3 Discussion

The assemblage distribution on the continental margin off Nova Scotia is seen on fig. 3.9. However, this map and indeed the previous 8 maps fail to show the nature of the assemblage boundaries. The character of these boundaries (i.e. gradational or of a sharp nature, etc) is an important clue to understanding the response of an assemblage to various limiting environmental gradients. The nature of these assemblage boundaries on the shelf are illustrated by a series of cross-sections.(figs. 3.10-3.15) These were examined with respect to their bathymetry and the corresponding factor loadings. Other parameters, such as the substrate, temperature and salinity gradients along these cross-sections, will be examined in a later chapter.

Section A-B extends from LaHave Bank into LaHave Basin, a decline in dominance of factor 7 as the section enters the Basin is seen. A corresponding increase in importance of factor 3 is observed.

Similarly in section C-D, the shelf edge assemblage 5 gradually

Figure 3.10 Location of faunal boundary sections seen  
in figures 3.11--3.15.



Figures 3.11—3.15 Variation of factor loadings (vertical) of several shelf assemblages with bathymetry and distance (horizontal).

Figure 3.11

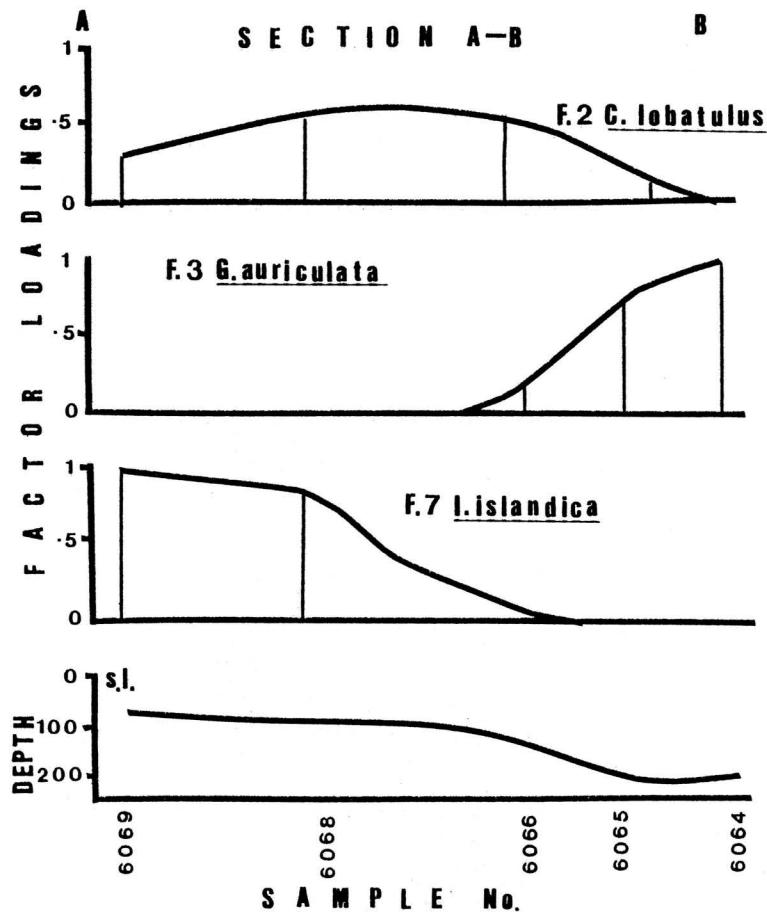


Figure 3.12

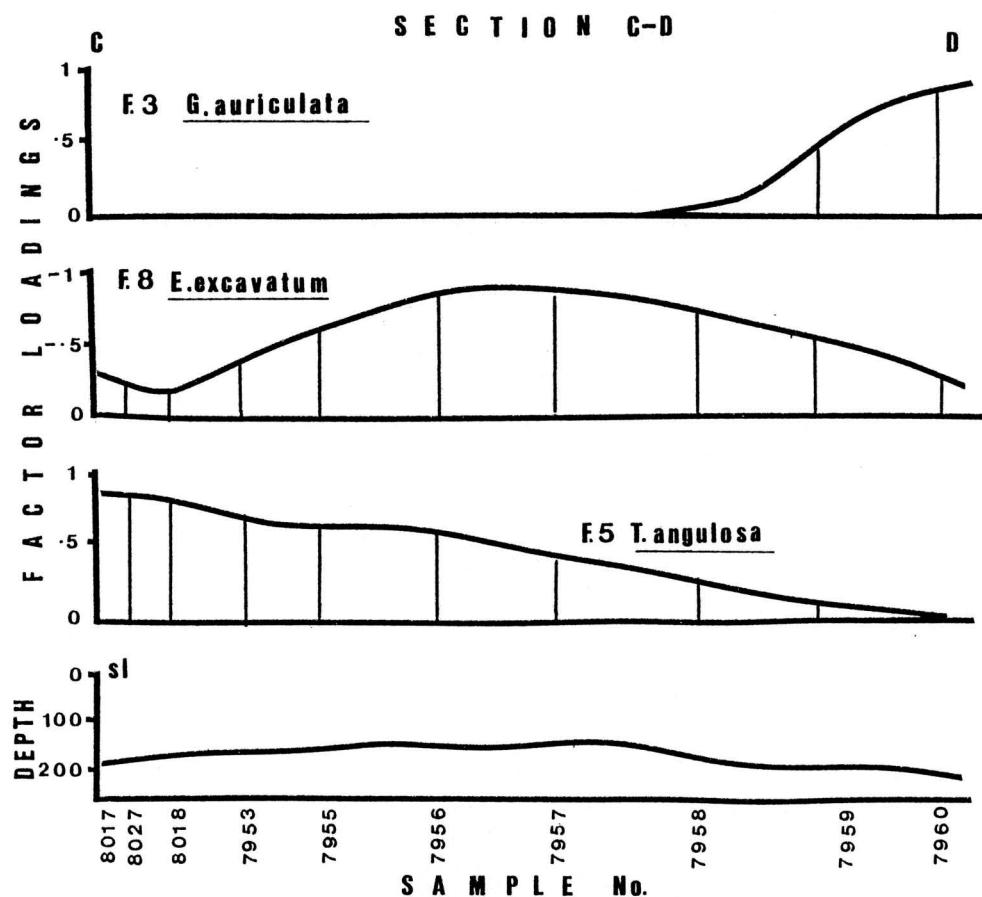


Figure 3.13

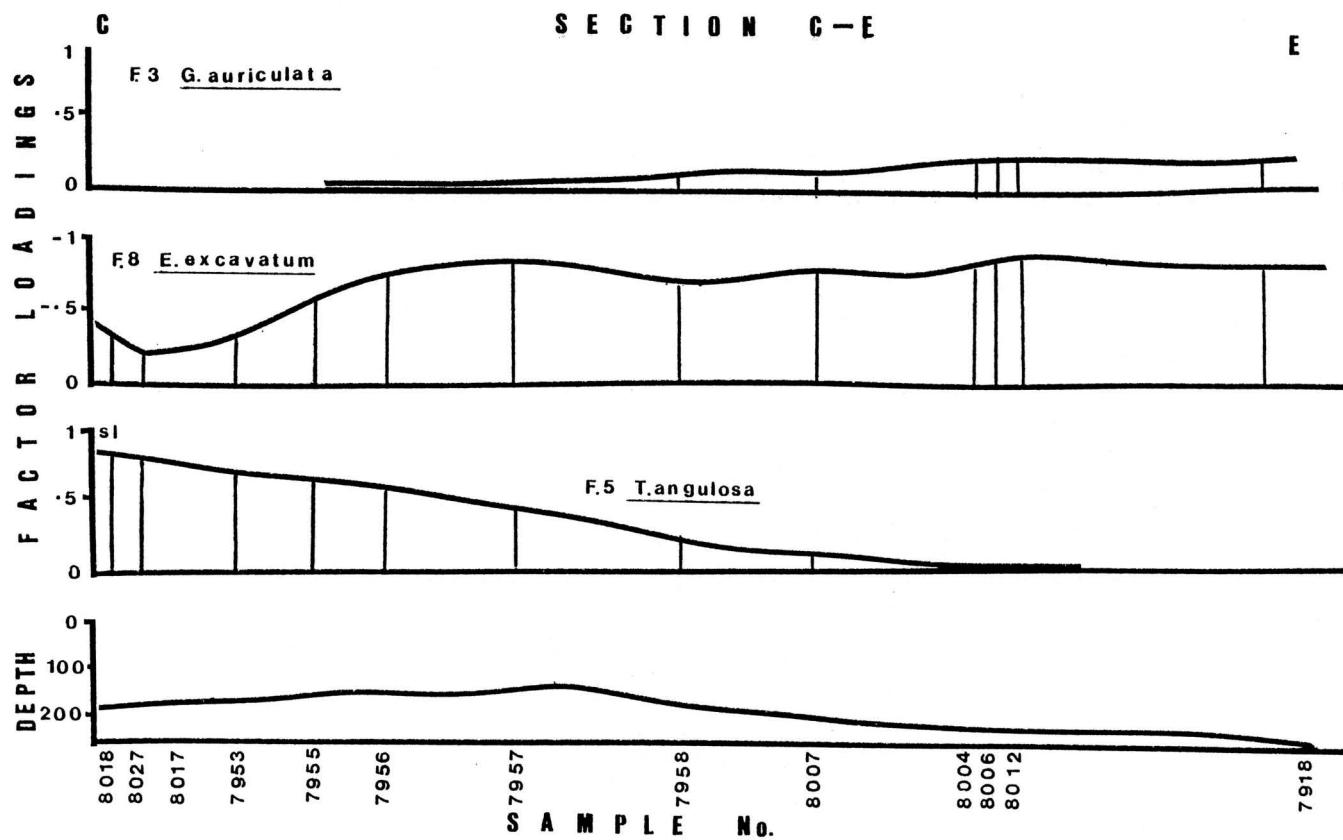


Figure 3.14

SECTION F-G

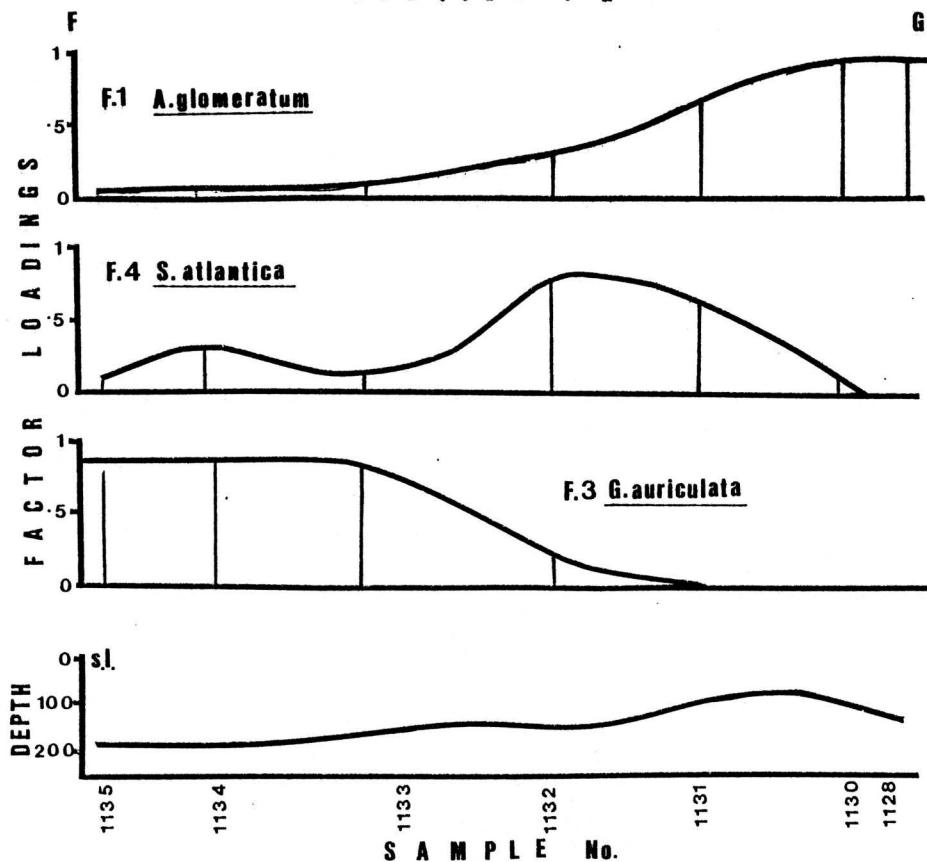
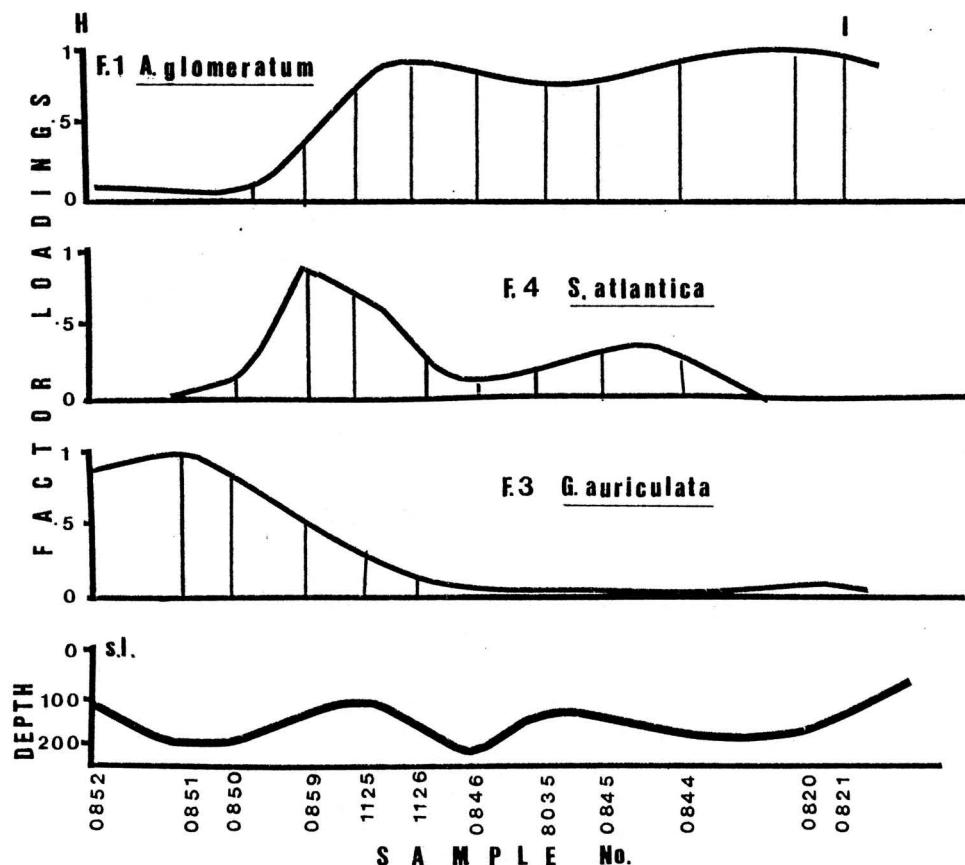


Figure 3.15

S E C T I O N   H - I



decreases in significance shoreward, across the Scotian Gulf and into LaHave Basin. As this factor decreases, there is a corresponding increase in basin assemblage factor 3. Factor 8 peaks in the middle of the Scotian Gulf and for reasons to be explained, is thought to be relict.(fig. 3.12.)

Section C-E is similar, although the section crosses Emerald Basin. Again, factor 5 gradually decreases with a corresponding increase in factor 3. It is believed that factor assemblage 3 is here masked by the occurrence of the relict transported factor 8 in part of Emerald Basin. This will be explained in chapter 5.(fig. 3.13)

Section F-G illustrates rather well the transition of the central basin assemblage 3 into the northeast bank/basin assemblage 1. Assemblage 4 appears to be a transitional assemblage between the two. These relationships are seen also in section H-I (figs. 3.14,3.15).

These sections illustrate the gradational nature of most of the assemblages on the shelf. Further chapters will more thoroughly discuss the ecological implications of these faunal gradients.

### 3.2.4: GROUP 3 SLOPE DATA.

Q-Mode factor analysis of these 43 stations determined 4 factor assemblages accounting for 76% of the observed variance. Table 3.3 shows the nature of the factor assemblages, their composition and distribution. This table is extracted from the varimax factor matrix and the factor score matrix ( Appendix III). The depth distribution of these factor assemblages ( i.e. factor loading versus depth) are shown

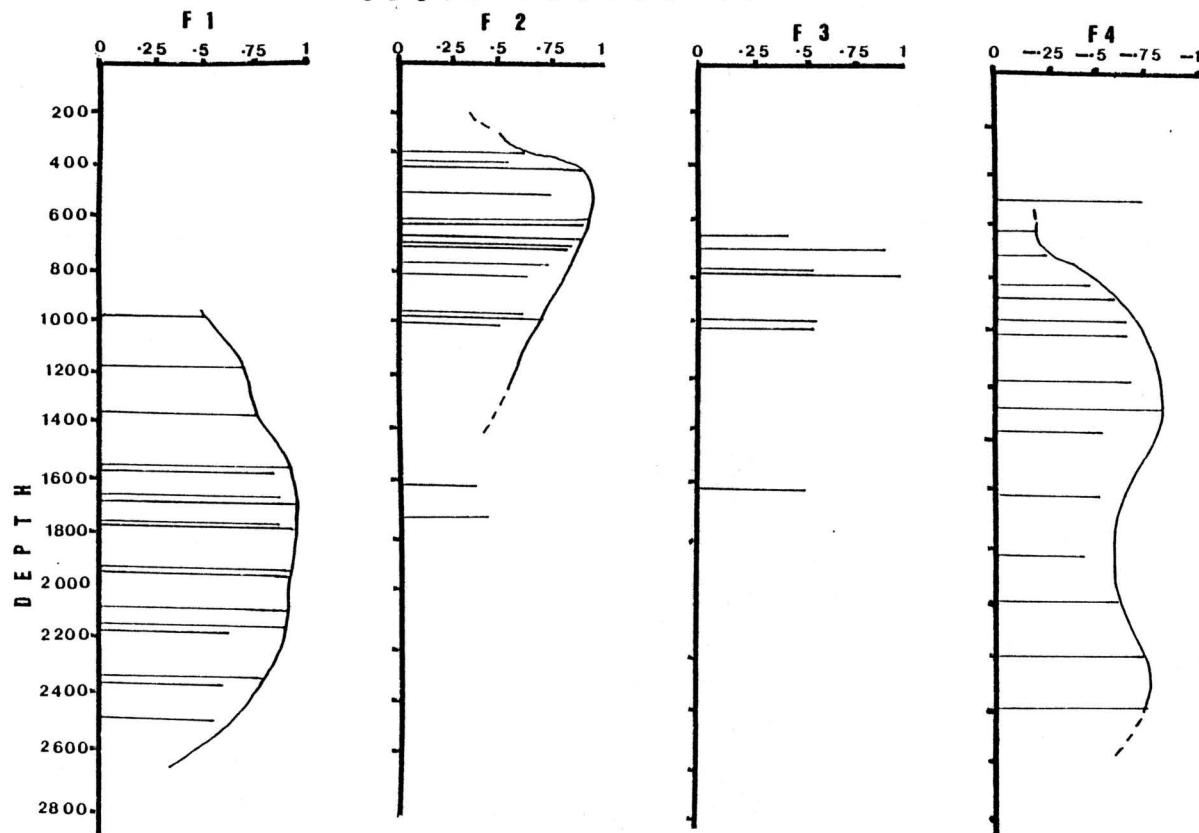
TABLE 3.3

## FACTOR SCORES FOR SLOPE ANALYSIS--A SUMMARY

FACTOR ASSEMBLAGE	% VAR.	DOMINANT SPECIES WITH FACTOR SCORES		DISTRIBUTION	SUBSTRATE
1	27.0	<u>Uvigerina peregrina</u> <u>Elphidium excavatum</u>	7.7 3.1	Deep slope. +1200 m	silts.
2	15.5	<u>Bulimina exilis</u> <u>Bulimina aculeata</u> <u>Cibicides lobatulus</u> <u>Nonionellina turgida</u> <u>Elphidium excavatum</u> <u>Buccella frigida</u> <u>Nonionellina labradorica</u>	7.3 2.2 1.2 1.2 1.1 1.2 1.0	Upper slope. 400-1000 m	silts.
3	8.5	<u>Trifarina occidentalis</u> <u>Nonionellina turgida</u>	8.0 1.2	700-1000 m	silts.
4	18.0	<u>Elphidium excavatum</u>	-7.5	400-2400 m	silts.

Figure 3.16 Depth distribution for slope assemblage  
factor loadings. (vertical=depth in m,  
horizontal=factor loadings)

SLOPE FACTOR LOADINGS



Uvigerina peregrina

Bulimina exilis

Triferina occidentalis Elphidium excavatum

in fig.3.16. Factor assemblages 1 (*U. peregrina*, table 3.3) and 2 (*B. exilis*) are restricted to lower and upper slope regions respectively. Factor 4 (*E. excavatum*) has a wide depth range whilst factor 3 is restricted to a few upper slope occurrences; this latter assemblage is not thought to be important.

The following describes the main features of each assemblage. Maps depicting the occurrence of the highest factor loadings for each assemblage in each slope area (figs. 3.17-3.23) are shown. The insets in these maps are the appropriate factor loading versus depth plot for each assemblage at each transect.

Slope Factor assemblage 1: *Uvigerina peregrina*.

This deep slope calcareous assemblage explains 33.6% of the observed variance. Figs.3.17 and 3.18 show the spatial occurrence of this assemblage ( i.e. occurrence of stations with factor loadings greater than 0.9) and along with their insets reveal the consistency of this assemblage at 1200-2400 m. depth and the gradational nature of its limits.

Slope Factor assemblage 2: *Bulimina exilis*.

This calcareous assemblage accounts for 21.6% of the variance. Figs.3.19, 3.20, and 3.21, reveal its constant occurrence at upper slope depths ( between 400-100 m. ). It is absent in slope area B, i.e. at stations which appear too deep.

Slope Factor assemblage 3: *Trifarina occidentalis*.

Figures 3.17--3.23. Spatial distribution of slope factor assemblages (F.L.--+.8). Inserts show depth distribution of assemblage factor loading at each transect.

Figure 3.17. Distribution of slope factor 1 *Uvigerina peregrina*

3.18.	"	"	"	"	1	"	"	"
3.19.	"	"	"	"	2	<u><i>Bulimina exilis</i></u>		
3.20.	"	"	"	"	2	"	"	
3.21.	"	"	"	"	2	"	"	
3.22.	"	"	"	"	4	<u><i>Elphidium excavatum</i></u>		
3.23.	"	"	"	"	4	"	"	"

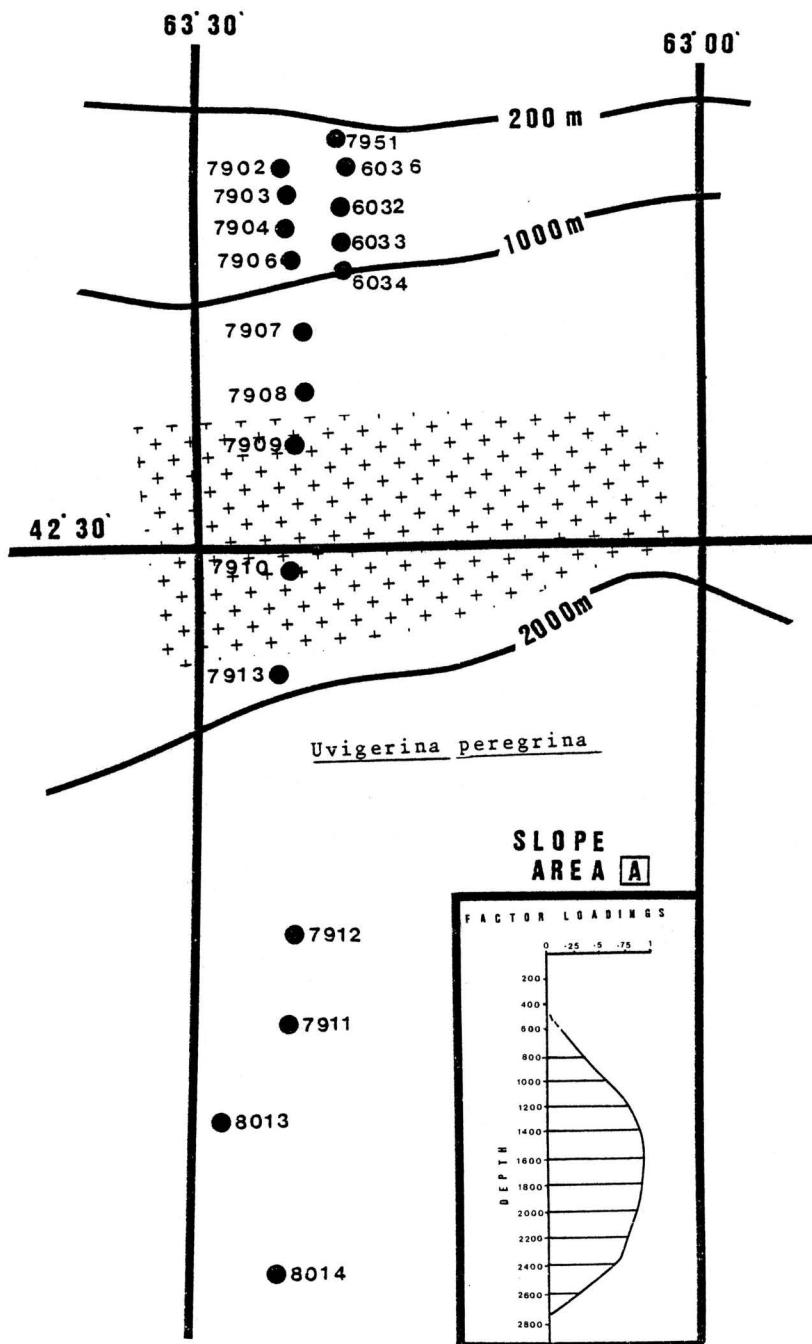


Figure 3.17.

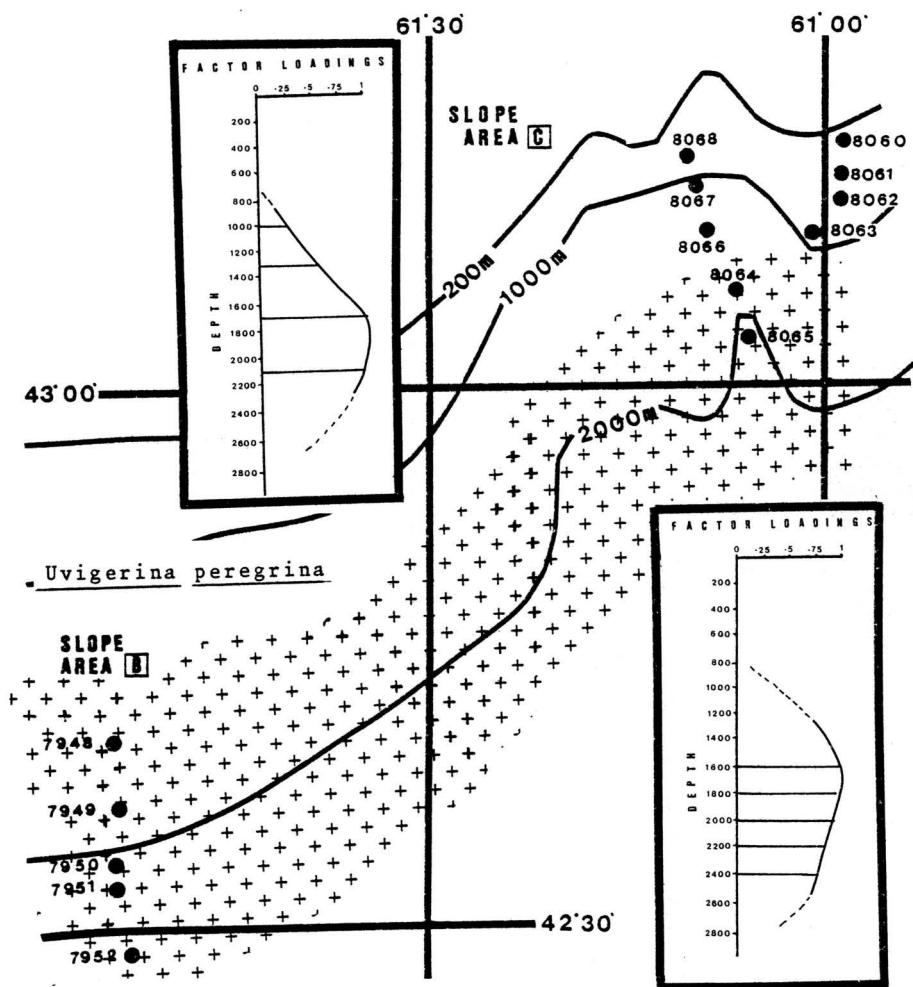


Figure 3.18

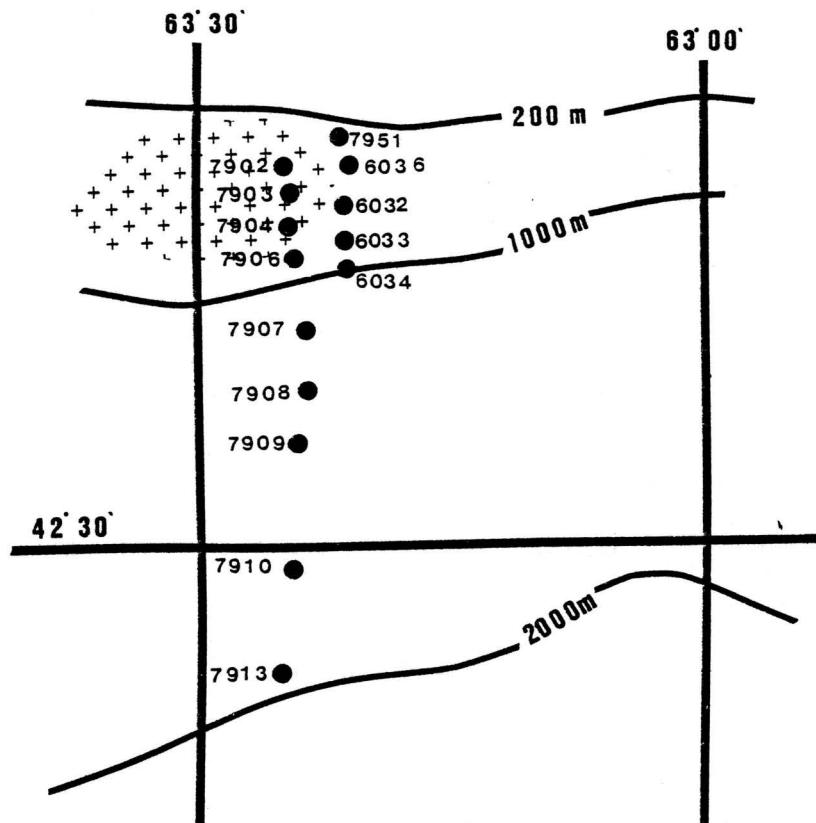
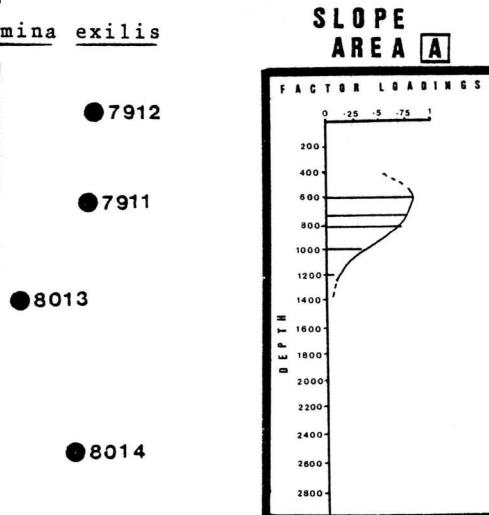


Figure 3.19 Bulimina exilis



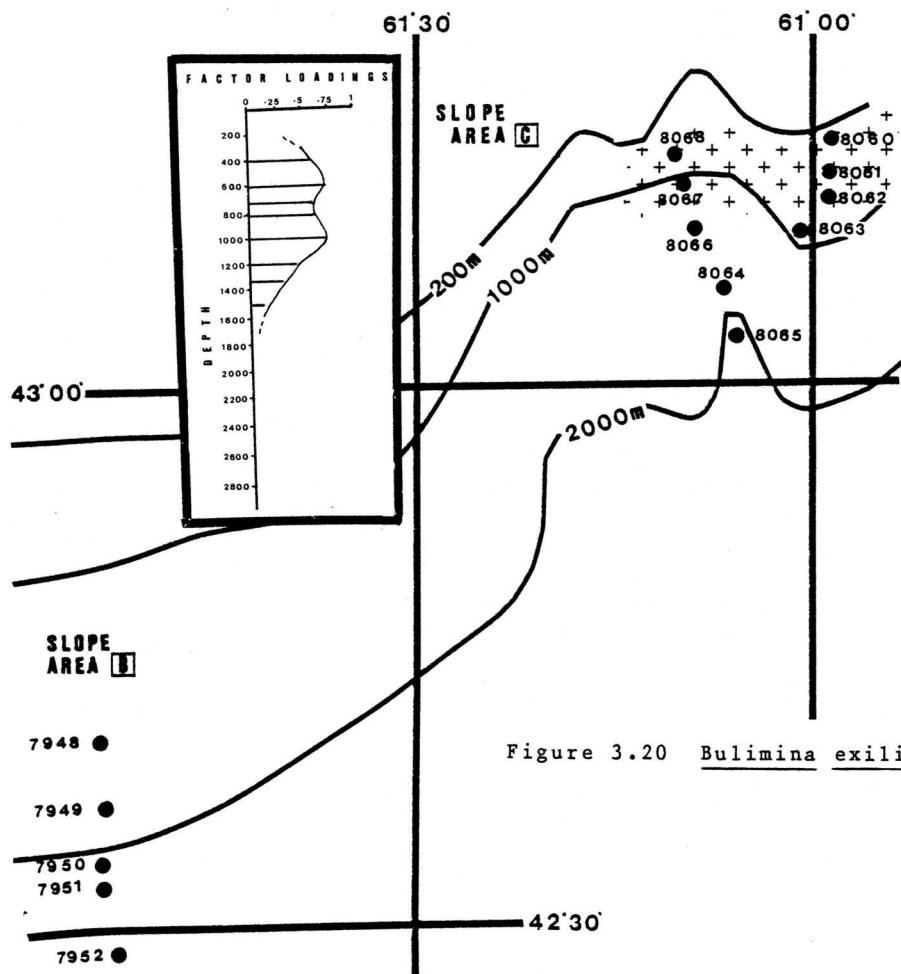


Figure 3.20 Bulimina exilis

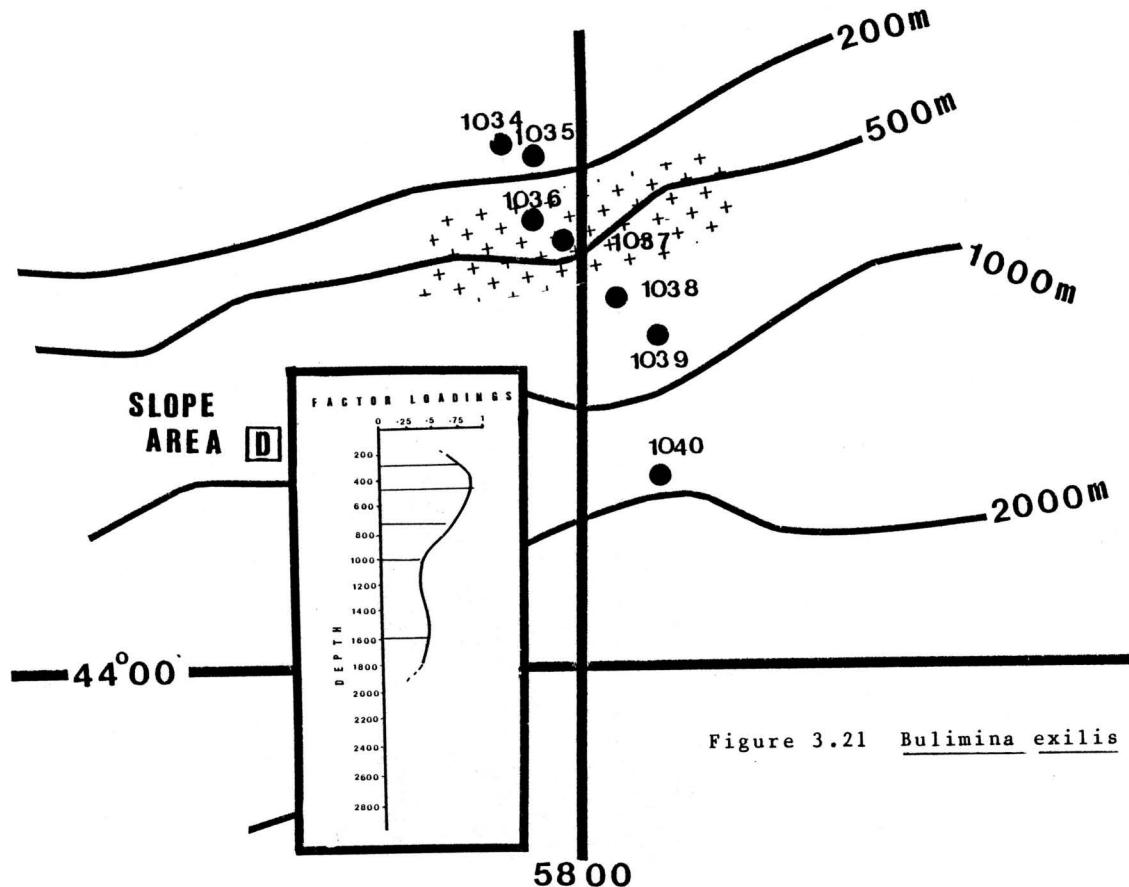
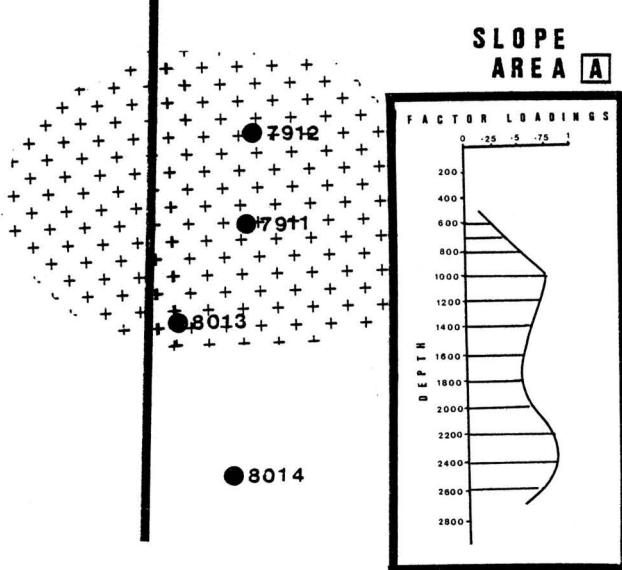
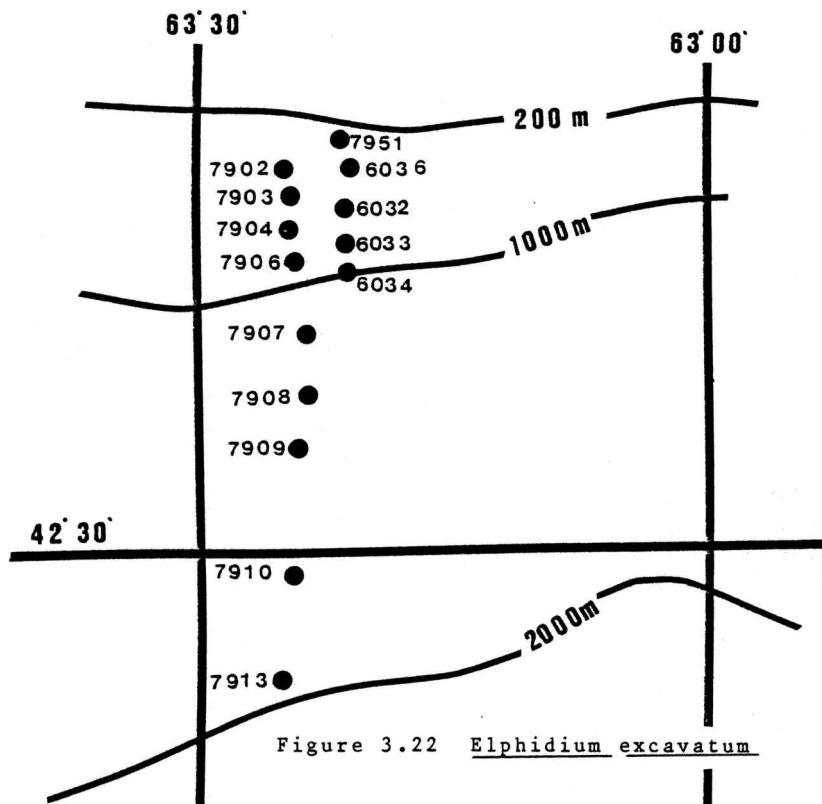
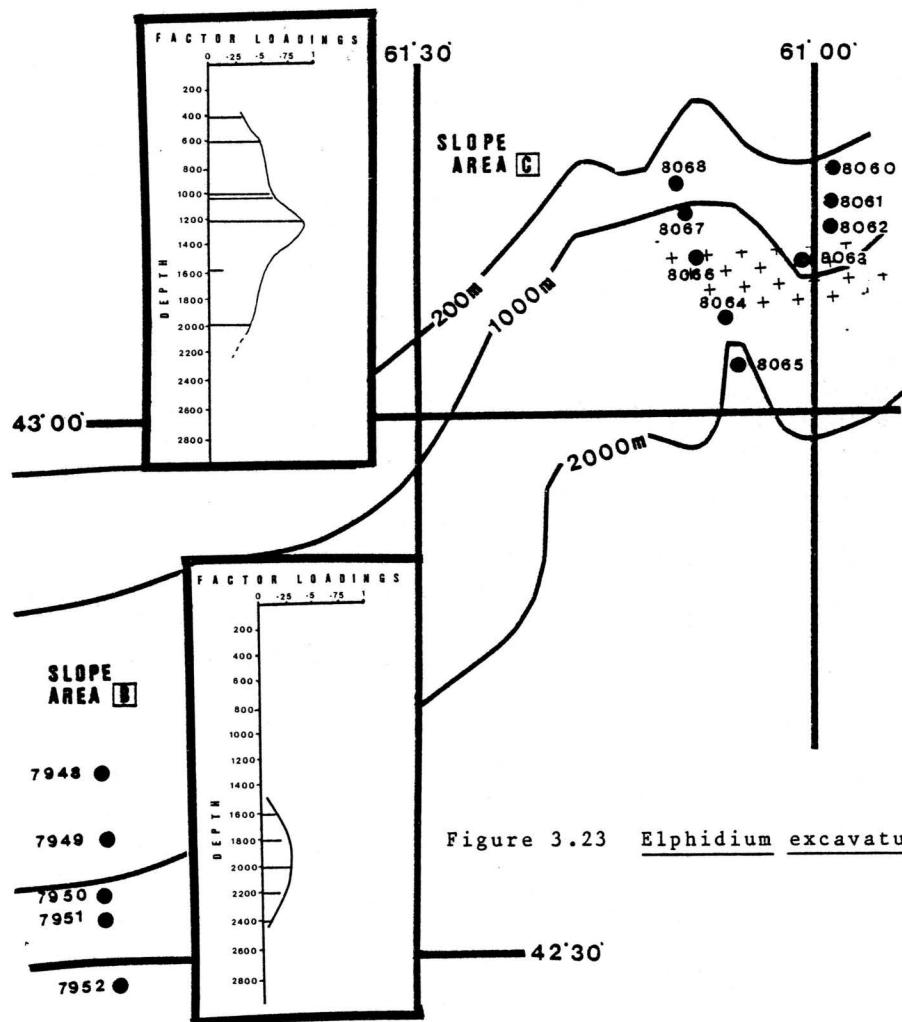


Figure 3.21 Bulimina exilis





Accounting for only 9.0% of the variance, this assemblage is restricted to a few stations in the upper slope areas, and could possibly be a local variation of assemblage 2. It is very likely that this assemblage is dominated by a megalospheric form of T. angulosa

Slope Factor assemblage 4: Elphidium excavatum f. clavata.

This calcareous assemblage accounts for 12% of the observed variance and is entirely dominated by E. excavatum. Its main occurrence is in slope area A and less importantly in slope area C (Figs. 3.22, 3.23). Its non-occurrence in other transects will be discussed later.

### 3.2.4: Shelf and Slope Distribution Summary

The factor analysis techniques have described and summarized the shelf and slope raw data. Their validity is shown by close scrutiny and comparison of the raw data and the factor assemblages produced. That the factor assemblages explain well the initial data input is shown through the high communalities of most stations (i.e. the degree to which a sample is adequately described by each factor). The total variance, which is the % of the original data explained by the technique, is also high.

In summary, eight shelf and four slope factor assemblages are produced by the analysis and have been described. They represent an accurate appraisal of the total (i.e. live + dead) foraminiferal distribution characteristics. Separate analysis of the slope and shelf

data proved more accurate, removing certain ambiguities inherent to the data.

In the next chapter, the live data is used to investigate whether the described assemblages are in equilibrium with presently occurring oceanographic regimes. This is a necessary precursor to the sections on environmental relationships, etc

## CHAPTER 4: ASSEMBLAGE MODIFYING PROCESSES

## 4.1 Introduction

The scientific objective of the distribution analysis, as undertaken in this study, is to attempt accurate, quantitative correlations with the causative and constraining oceanographic parameters that influence assemblage occurrences. If conclusions drawn from this study are to be applicable to the present day and for palaeoecological settings, account must be taken of processes that may affect an assemblage composition, and render such conclusions invalid.

The present section investigates the extent to which the defined assemblages are in equilibrium with the acting hydrographic regimes.

## 4.2 Assemblage modifying processes.

Most of the processes which modify assemblages have been well documented (Murray, 1976; Douglas and Heitman, 1979; Douglas et al, 1980; Murray et al, 1982; Schafer and Prakash, 1982). These works explain observed disparities between living and total foraminiferal populations as the result of relictness, shell dissolution, transportation, fossil reworking, predation or variations in rates of shell production, sample type and inadequate staining. Each of the defined assemblages was examined for evidence of these modifying influences through an appraisal of the raw data for living organisms. Unfortunately, the live data presented in this study was insufficient for exhaustive quantitative methods. They do, however, allow subjective

appraisal of the validity of each assemblage.

Before the live data is examined, further discussion on these modifying influences is warranted.

a/ Relictness: Fossil Reworking

Relict assemblages or relict species within an assemblage are the result of a low sedimentation rate over a period of time. Some or all of the fossils lived under hydrographic conditions not now operating.

Relictness is indicated by the occurrence of species or assemblages that have no apparent living representatives, either within the immediate limits of the assemblage or throughout the study area. The physical appearance of the test may also indicate whether an assemblage has been long exposed at the sediment-water interface and therefore perhaps, relict. Certainly, the sediments on parts of the shelf off Nova Scotia have previously been described as not being in equilibrium with the present hydrographic conditions (Drapeau 1971, Swift et al., 1972). Modern, post-glacial sediment dispersal processes have produced a "spillover" effect of reworked relict sands from shelf to slope environments (Swift, 1971). The potential then, for some or parts of the defined assemblages to be relict and not indicative of the present regimes is quite great. The extent to which they are is discussed in a later section.

Reworked Tertiary or older fossil material being present in an assemblage should also be recognized. Being in disequilibrium with the hydrography, they have the same effect as relictness and imply derivation through the stripping of the Quaternary cover, leaving fossil assemblages exposed.

b/ Dissolution

Active shell destruction resulting from solution is not thought to be a problem within this study area. Etching of the test surface does occur on some species and possibly indicates relict sediments. Discussion on the effect of various levels of calcium carbonate availability is reserved for a later section.

c/ Transportation

Several effects of the sediment transportation of foraminifera must be discussed. Transportation may deplete one or more of the components (or rather a particular size range of its components), on the other hand, an assemblage may become enriched with particular species. A whole assemblage may become mobilized and then selectively transported; whole scale slumping may result in assemblages occurring at deeper levels than usual or depth restricted bathyal microfaunas may become uplifted to shallower depths (Silver, 1971; Murray, 1976; Douglas and Heitman, 1979; Douglas *et al.*, 1980, Coulbourn, 1980). Murray (1982) has recently described the occurrence of benthic foraminifera throughout the water column being actively transported landward in near surface waters or seaward as bottom load. He observed that the tide and storm-swept seas of the Western Approaches off Britain may serve as a model for other continental shelves subject to vigorous tidal currents and wave attack (e.g. Nova Scotias exposure to western Atlantic seas). The extent and magnitude of this type of transportation process off Nova Scotia needs further examination.

In this study, the recognition of transported assemblage components involves the use of live population counts as well as a close examination of the test sizes themselves i.e. size sorting of juveniles or adults, etc.

d/ Sample type and staining inadequacies.

Both these features may introduce ambiguities into the live counts. Their effects have been discussed in an earlier section (Chapter I).

#### 4.3. Comparison of Shelf Factor Assemblages with Live Data.

This section examines each factor assemblage produced through the analysis of total population data and compares its composition with observed live data. Stations within a factor assemblage area that exhibit similar live species rankings to factor score species rankings suggest that the factor assemblage is representative of the actual bioceanosis. In this way it is possible to determine which assemblages were in equilibrium with and thus good indicators of the present oceanographic regimes. Discrepancies between the live data and the factor assemblages are explained with reference to the previous section.

Appendix IV Tables 1-8 show stations from within each factor assemblage area, their respective factor loadings and dominant species components. For each dominant species within a given factor, the corresponding live per cent. value is given per station. Living species that do not figure prominently in the factor score matrix are noted.

a/ Factor assemblage 1: Adercotryma glomerata

The main species components defined for this factor reflect closely the observed live percentage data (i.e. rankings in factor score matrix are essentially the same as the live rankings, Appendix IV Table 1). Two of the subdominant species within this factor assemblage, Spiroplectammina biformis and Cribrostomoides crassimargo, lack significant live representation. The difficulties concerning the observation of stains in agglutinating foraminifera, as explained in chapter I, may explain this feature (in particular for small forms such as S. biformis). The second most dominant species within this factor, R. turbinatus, occurs in greater proportions living, indicating that it may play a greater role within the assemblage than is otherwise implied. It is clear then, that factor assemblage one reflects the observed live data to an extent that it is representative of the presently occurring oceanographic regime in the area. In other words, relictness and transportation are not a problem here.

b/ Factor assemblage 2: Cibicides lobatulus

This C. lobatulus dominated factor compares well with the observed living percentages (Appendix IV Table 2). Several other live species occur but do not figure prominently in the factor scores. Many of these, however, do not occur consistently in the assemblage throughout the stations (e.g. B. frigida and E. advena, Table 2). The live data suggests there are rather more species indigenous to this assemblage than the factor scores imply. It is thought here that the Bank and Nearshore regions, to which this factor is commonly associated, provide

large amounts of attachment surfaces for its main species, C. lobatulus. The death and subsequent deposition of this form tend to flood the assemblage, thus masking other live components. All these species must be considered in an ecological study.

c/ Factor assemblage 3: Globobulimina auriculata

This assemblage (Appendix IV Table 3) reflects the observed live data reasonably well. Several discrepancies, however, are seen. Both G. auriculata and N. labradorica are prominent in the factor scores and the live data. Their values in the latter vary quite appreciably. There are also a large number of species that occur living but do not figure prominently in the factor scores. Several are consistently present living (e.g. F. fusiformis and C. crassimargo) whereas the remainder are not so persistent.

The species G. auriculata and N. labradorica are both characterized by globular, lightweight tests. Kontrovitz (1978) has shown that high operational sphericity (i.e. globularity) and thin light weights are correlated to low traction velocities and thus low velocity currents would be required to quickly put these two species into suspension. This may partially explain the variation observed in the live occurrences of these two species. The transportation may have occurred on a small local scale, concentrating the species and perhaps making them appear more important within the assemblage than they really are.

B. aculeata does not occur live at all, even though it is an important component in the assemblage. Absence of spines and general worn nature may imply transportation. Hamden (1969) has also noticed predominance of dead tests of this form. In general terms, this factor

does reflect the live distribution. However, the assemblage is more diverse in terms of the live data. Ecological study will have to consider this.

d/ Factor assemblage 4: Saccammina atlantica

This rather subdominant assemblage is restricted to peripheral regions of assemblage 3. Despite the difficulties in recognizing staining in heavily agglutinated forms, the live data compares well with the factor scores. Confusion between Saccammina atlantica and broken chambers of R. scorpiurus may have produced error. Several species occur live but do not figure in the factor scores. A. glomerata is consistently present and G. auriculata is present in some stations. This is to be expected as factor 4 occurs between factor 3 and 1.

Despite these problems, this factor is thought a reliable indicator of the oceanographic conditions prevailing.

e/ Factor assemblage 5: Trifarina angulosa

This assemblage occurs consistently along the shelf edge margin, and is dominated by Trifarina angulosa. However it is not often seen in the live assemblage. Table 5 shows that the live assemblage is dominated by Bolivina spathulata. Other species, although not figuring prominently in the factor score matrix, do occur in the live data. Buccella frigida, Cibicides lobatulus, Cassidulina laevigata and B. subaenariensis all occur consistently in the living population. The assemblage then, when the living data is considered, is rather more diverse than the factor score matrix suggests and is in reality less

dominated by T. angulosa.

Several explanations can account for large amounts of dead T. angulosa. The shelf/slope waters off Nova Scotia have recently been described by Fournier (et al, 1977) as being highly productive, nutrient-rich and of enhanced biological activity. On the shelf edge, this could be reflected by a rapid turnover of T. angulosa- perhaps seasonal - with a resulting large accumulation of dead tests. Alternatively the oceanographic conditions that lead to the dominance of T. angulosa may not be operating, implying relictness. The majority of the tests, however, do not appear etched or decayed; neither is there much evidence of transportation downslope or shorewards into the central basin areas (as will be described later in further detail).

Transportation would be expected if the T. angulosa component was relict on the shelf edge, which appears to be an axis for sediment dispersal (Stanley, 1972). It is thought here that the assemblage does reflect present conditions well, although consideration of the live data is essential to an ecological analysis of this assemblage. No shortage of live T. angulosa has been observed in several other works (Hamden, 1969), leading to the conclusion that, in the case of this study, seasonal effects might explain why T. angulosa is rarely found living.

#### f/ Factor assemblage 6: Eggerella advena

The factor score results in Table 6 show that this assemblage, which is restricted to nearshore bay areas, compares very well with the observed live data, i.e. E. advena dominates with subdominant S. biformis. A wide variation in live percentages of both these species is seen. This is most likely due to local variation, sample difficulties

and the previously discussed errors in recognizing stains in agglutinates. It is clear that this assemblage, based on the total data, is representative of the live occurrences and has no relict or transported components.

g/ Factor assemblage 7: *Islandiella islandica*

This assemblage is similar in occurrence to assemblage 2 and is almost monospecific. The living data however (Table 7), reveals that it is less dominant. Other species are consistently present living but do not figure in the factor score matrix. These include *C. lobatulus*, *F. fusiformis* and *E. advena*. Low sedimentation coupled with a high production of *I. islandica* results in a masking of the total and live occurrences mentioned. The factor assemblage is partly a reflection of the actual live distribution but consideration of the other live species is necessary in any ecological study.

h/ Factor assemblage 8: *Elphidium excavatum*

The factor score matrix indicates that this assemblage is dominated by *Elphidium excavatum* with several subdominants of *F. fusiformis* and *B. aculeata*. The live data reveals only these subdominants together with *C. laevigata*, *C. lobatulus*, *B. subaenariensis* and *R. guttifera*. The dominant species, *E. excavatum*, does not appear live anywhere within this assemblage or indeed, within the study area. Several explanations can account for this.

The etched and ragged nature of the Elphidium tests indicate that they have been lying on the sea bottom for some time and are possibly

relict. However, a large part of this assemblage lies within Emerald Basin (Chapter III, fig. 3.8) where, at the present time, active sedimentation is in progress: clays or silts would have long since covered relict sediments.

The southern part of this assemblage lies on the outer bank area around the Scotian Gulf where little or no sedimentation is presently occurring. This area is also periodically swept by strong currents capable of putting sediment including E. excavatum into suspension and transporting it (Hill, 1981).

From the above, it appears that assemblage 8 is dominated by a relict species which has been transported from its source area to the south of the assemblage and deposited in and on the approaches to South Emerald Basin. Further evidence to this is provided through an observation of the size distribution of E. excavatum in the relevant samples. Samples to the south (i.e. in the "source area") have a higher average size (maximum diameter) and greater standard deviation than those to the north, approaching and within Emerald Basin. The decrease in mean and standard deviation is gradational. This essentially means that the source area has a mixed population of E. excavatum that includes juveniles and adults. Subsequent sorting of this relict population has occurred and smaller fractions (i.e. juveniles) were deposited further north.

B. aculeata does not appear living and is here thought to be a transported assemblage also. This is reflected in the relative predominance of adult forms, all etched; broken spines.

In summary, this assemblage does not reflect the observed live data very well. If one takes into account that the dominant species of this

assemblage is relict transported, then ecological conclusions may be reached.

#### 4.4 Comparison of Slope Assemblages with Living Data

Unlike the shelf assemblages which exhibit a close, albeit subjective similarity to the living, the slope assemblages are not directly comparable to the live data. Significant liva data are completely lacking. The following considers each slope assemblage and, through a process of elimination of the previously discussed modifying influences on an assemblage, will determine its reliability with respect to the present environment. This is not an eminently desirable method, but considering the paucity of live data in these areas, it is the only one possible. Relictness, transportation, sample methods, staining methods, etc, are but some of the possible explanations for the depauparate live data.

##### a/ Slope Factor assemblage 1 Uvigerina peregrina

This deep slope assemblage becomes dominant at 1200 m., peaks at 1800 m. and generally declines in importance at depths greater than 2400 m. It is almost monospecific. The presence of well preserved tests (i.e. not etched) with both juveniles and adults, seems to preclude relictness and transportation. The nature of the test, thick with coarse costae often infilled with mud, may have prevented observation of staining. The surface sediment may have been lost in sampling; most of the samples from these stations are core tops and the problems encountered with good core top recovery have been discussed ( Chapter

I).

This species is rarely reported living. ( Streeter and Lavery, 1981; K. Miller, 1982), perhaps for the reasons outlined. Available evidence suggests that this assemblage is in place and in equilibrium with its present environment.

b/ Slope Factor assemblage 2 Bulimina exilis

This upper slope diverse assemblage also has no live data. Like assemblage 1, there are good proportions of adults and juveniles, with well preserved tests ( i.e. no etching). Lack of apparent sources on the shelf edge and shelf or apparent displacement into deeper slope waters seem to rule out transportation. Some tests appeared slightly stained but not enough for positive identification. All this implies an in situ, assemblage. Lack of live data may be explained by the highly productive surface waters and the warm, nutrient-rich slope waters. This would lead to increased and rapid turnover of foraminiferal populations which would tend to dilute the living representatives.

c/ Slope Factor assemblage 3 Trifarina occidentalis

This small, relatively minor assemblage, often co-occurs with assemblage 2. It is thought reliable for the same reasons.

d/ Slope Factor assemblage 4 Elphidium excavatum

This monospecific assemblage reveals a large depth range, 400-1200 m., although it is only dominant from 800-1200 m. It is marked by its

restriction to the A transect ( fig. 1.5).

Unlike other slope assemblages, several features of this factor ( or of the E. excavatum monospecific assemblage) cast doubt on its validity, both with respect to being "in situ" and the extent of its present interaction with the environment.

The etched and worn nature of the tests would imply both transportation and relictness. Most of its previously documented occurrences restrict this species to shelf depths. Several workers, however, have found Elphidium species living in deeper slope waters ( Matoba, 1976; Schnitker, 1980; Schafer, 1981). Based on this, Streeter and Lavery (1981) and Miller and Lohman (1981) have recently assigned their slope assemblages as being in situ. Data presented here, however, implies transported relict assemblages of E. excavatum from shelf depths. A source area is the previously defined relict assemblage on the Scotian Gulf ( this chapter). The mechanism and process of sediment transportation over the shelf edge in this area has recently been determined by Hill (1981). The non-importance of this assemblage in any other of the slope transects, all of which do not have a corresponding prospective source area similar to the Scotian Gulf, would support this notion of transportation. The size distribution of measured diemeters of E. excavatum is similarly supportive, i.e. low s.d. would imply some size sorting of juveniles and adults.

It would appear that the use of E. excavatum fauna to characterize presently operating hydrographic regimes would lead to erroneous results.

#### 4.5 Summary of assemblage distribution and composition

Eight shelf and four slope assemblages have been recognised with the aid of Q-Mode factor analysis. The shelf assemblages appear restricted to several physiographic and hydrographic regimes. Seven of the shelf assemblages compare well with the observed live data and are thus assumed to be representative of the presently occurring environmental conditions. An eighth assemblage is thought to be relict and transported. The degree to which these assemblages are a result of the environment will be discussed in later chapters.

Slope assemblages appear to show better depth limitations and with one exception are consistently present along contours. No slope assemblage shows significant live foraminiferal data, three of them however, for reasons discussed earlier do appear representative of existing conditions. The fourth is thought to be the down slope equivalent of shelf assemblage 8.

A consideration of the more general faunal trends, such as diversity, is an important precursor to any investigation of the relationships between these statistically independant assemblages and the limiting oceanographic influences. This will be discussed in the next chapter.

## CHAPTER : 5 FAUNAL DIVERSITY, TRENDS AND CHARACTERS.

## 5.1 Introduction

The aims presented in the introduction are directed towards establishing a quantitative distribution model of foraminiferal occurrences on the continental margin off Nova Scotia. Hitherto, the approach to this study has involved analysis at the recurrent species or assemblage level, enabling distinct groupings of foraminiferal associations to be established. Much useful information can also be obtained from the general faunal aspects which illustrate the regional trends and characteristics of the population data. Furthermore, it is necessary that any ecological conclusions made from the distribution of foraminiferal assemblages be compatible with these general population data.

The methods used to determine faunal trends in this study involve an analysis of diversity, standing crop determinations and Live Total ratios. Prior to showing the results of these investigations, short discussion is warranted concerning the usefulness and nature of these various measures.

## 5.2 Diversity.

The diversity of an assemblage has two distinct components: the number of different species, and the relative frequency of each species (Deevey, 1969). Several indices of diversity have been used in foraminiferal studies, the calculation, use and relative advantages of which are well documented in Buzas and Gibson (1969), Murray (1973) and

Buzas (1979). Indices frequently employed include: the species number (S), Yule-Simpson Index, Fisher Index, Information Theory Index and an index using cumulative curves.

The Yule-Simpson Index has been shown by Murray (1973) to have an abundance bias whereas measured angles on cumulative curve slopes are time consuming according to an earlier study (Murray, 1968).

The Fisher Index ( $\alpha$ ) is favoured by Murray (1973) although this index is known to increase with sample size.

$$\alpha = \frac{n_1}{X}$$

where X = constant value less than 1;  
 n = a value calculated from  $N=(1-X)$  where  
 N = size of population

The calculation of  $\alpha$  is simplified through reading off a graph (Murray, 1973)

The advantage of the Fisher index is that it takes the rare species into account (Williams, 1964). When the underlying distribution is log series in form, which it often is, Buzas (1982), it is a consistent index.

Several workers have used the information theory to obtain measures of diversity. One such index, known as the Shannon-Weiner information function, H(s), (Patten, 1962) measures the average diversity.

$$H(s) = - \sum P_i \ln P_i$$

$P_i$  = proportion of the  $i^{th}$  species;

s = no. of species in the sample.

Maximum values are reached when all the species are equally distributed. It has been shown to be informative for palaeoecological studies by Williams (1964), Murray (1968), Beerbower and Jordan (1969), Buzas and Gibson (1969), Douglas (1979) and Sen Gupta (1981). It also accounts for both components of diversity, as noted by Deevey (1969).

A derivative of the  $H(s)$  value is the species equitability  $E$  as defined in Buzas and Gibson (1969); this gives a better view of the relationship of the dominant and rare species to total populations:

$$E = \frac{eH(s)}{S}$$

where  $e$  = base of natural logarithms

$S$  = number of species in sample

As the value of  $E$  approaches 1.0, the distribution of the species within a sample approaches equality.

Many workers have simply used  $S$ , the number of species per sample, as a measure of diversity. As this does not take account of relative proportions, great care must be exercised with its use and interpretations.

There are thus several well used measures of diversity, all measuring different relationships within a sample. Buzas (1979) notes that no single index is sufficient to describe and measure differences (between samples) in diversity and recommends that:

"... $S$ ,  $H(s)$  and  $E$  or their counterparts should be measured to ascertain the nature of these differences".

This study investigates spatial variations of  $S$ ,  $H(s)$ ,  $\alpha$ , and  $E$ . These values for each station are recorded in Appendix V. Figures 5.1-5.4 show the main variations in these values across the shelf and slope.

### 5.3. Diversity patterns.

The figures (5.1-5.4) intend to show the major groupings (as discerned by eye) of the diversity values recorded at each station. The data has been simplified into three classes: low, medium and high diversity. Low values include those barren stations that are confined

mainly to Western and Sable Island banks. Within each region of the map characterized by a particular diversity value, very few anomalous values occur

#### 5.3.1 Species number S (fig. 5.1)

Lowest values of species numbers are seen on the outer Banks, Central Banks and nearshore coastal areas. The whole of the northeast portion of the shelf (including Misaine Bank-Basin complex and Canso Bank-Basin) is characterized by low values for S. Medium to high values are found within LaHave and Emerald Basins, Roseway Basin and slope areas greater than 1800 m. depth. Highest values are seen along the shelf edge and upper slope and lower slope depths off Banquereau Bank (fig. 5.1). High values are also seen pushing into the shelf region around the Scotian Gulf.

#### 5.3.2 Shannon-Weiner index H(s) (fig. 5.2)

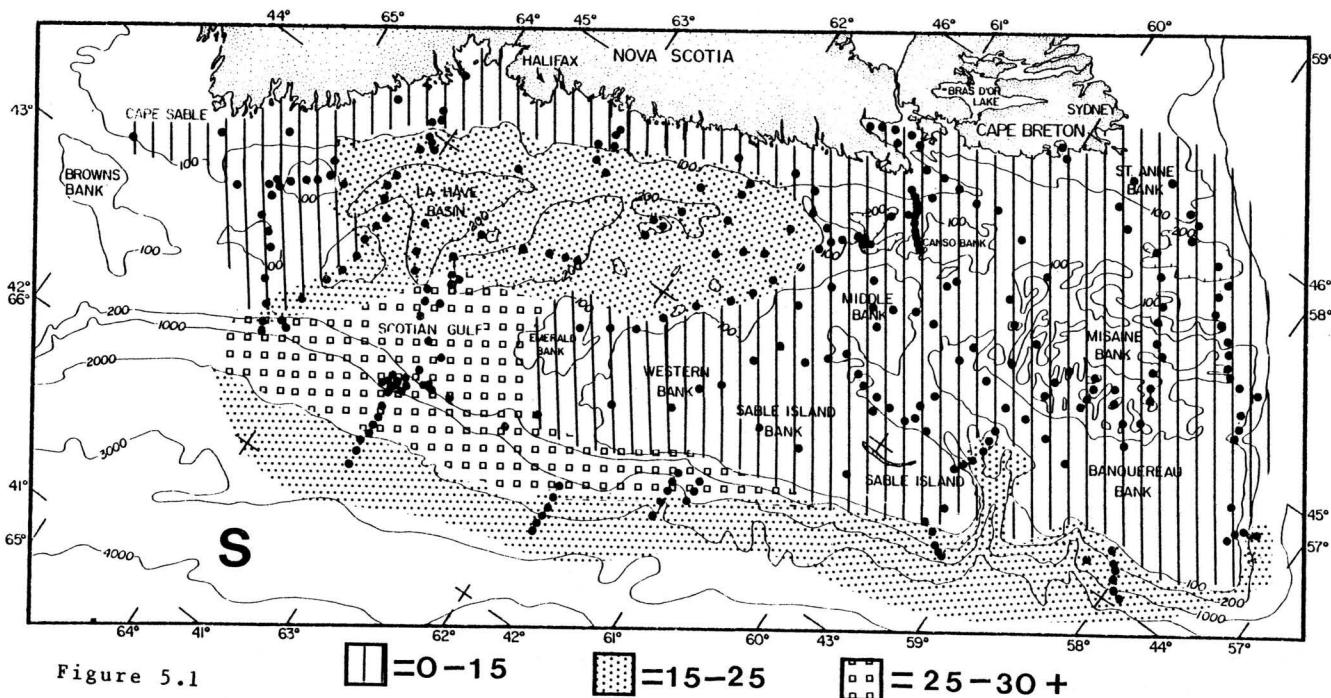
The lowest values for this index also occur on the outer, central banks and inner coastal areas (reflecting S distribution). Several discrete areas of medium values are seen within this low H(s) region. Medium values occur in LaHave, Emerald and Roseway Basins, the Scotian Gulf and along the shelf edge. Highest values are observed on the upper slope regions. Again, lower values are seen in lower slope areas.

#### 5.3.3 Fisher index $\alpha$ (fig. 5.3)

Low, medium and high values of this index again seem to reflect the

Figures 5.1--5.4. Diversity trends on the continental margin off Nova Scotia.

Figure 5.1    S    Species number  
5.2    H(s)    Shannon-Weiner index  
5.3    E    Equitability  
5.4     $\alpha$     Fisher index



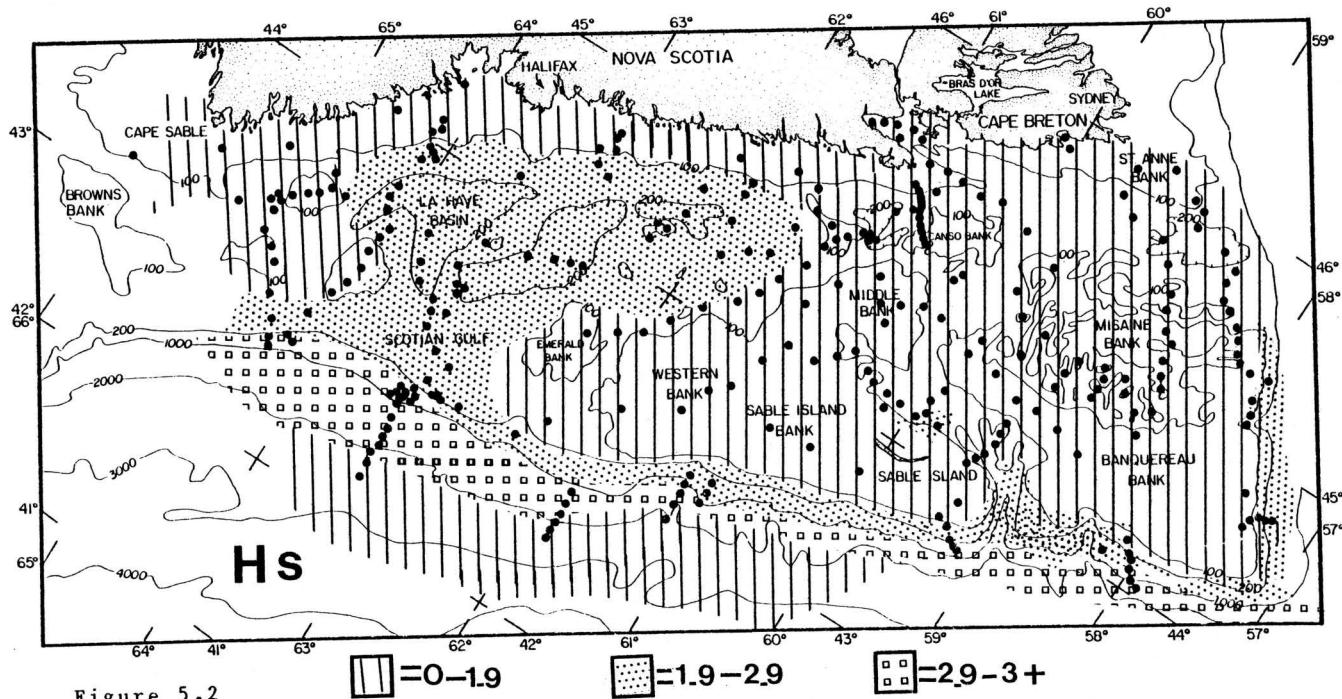


Figure 5.2

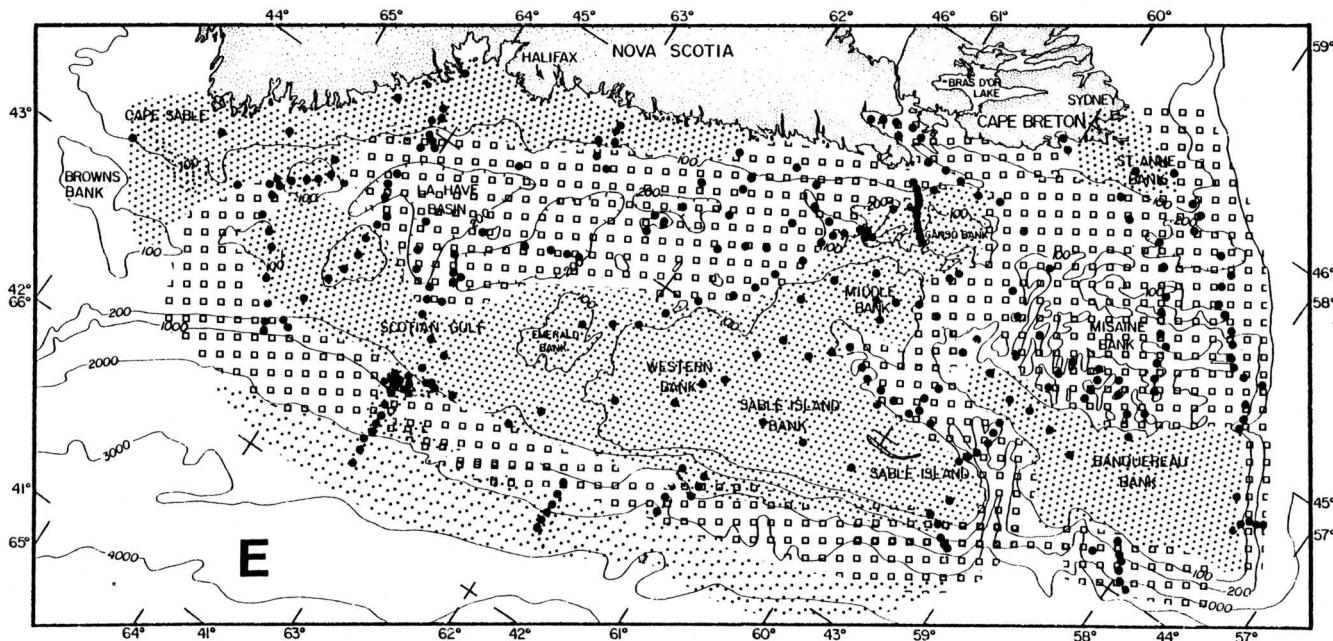
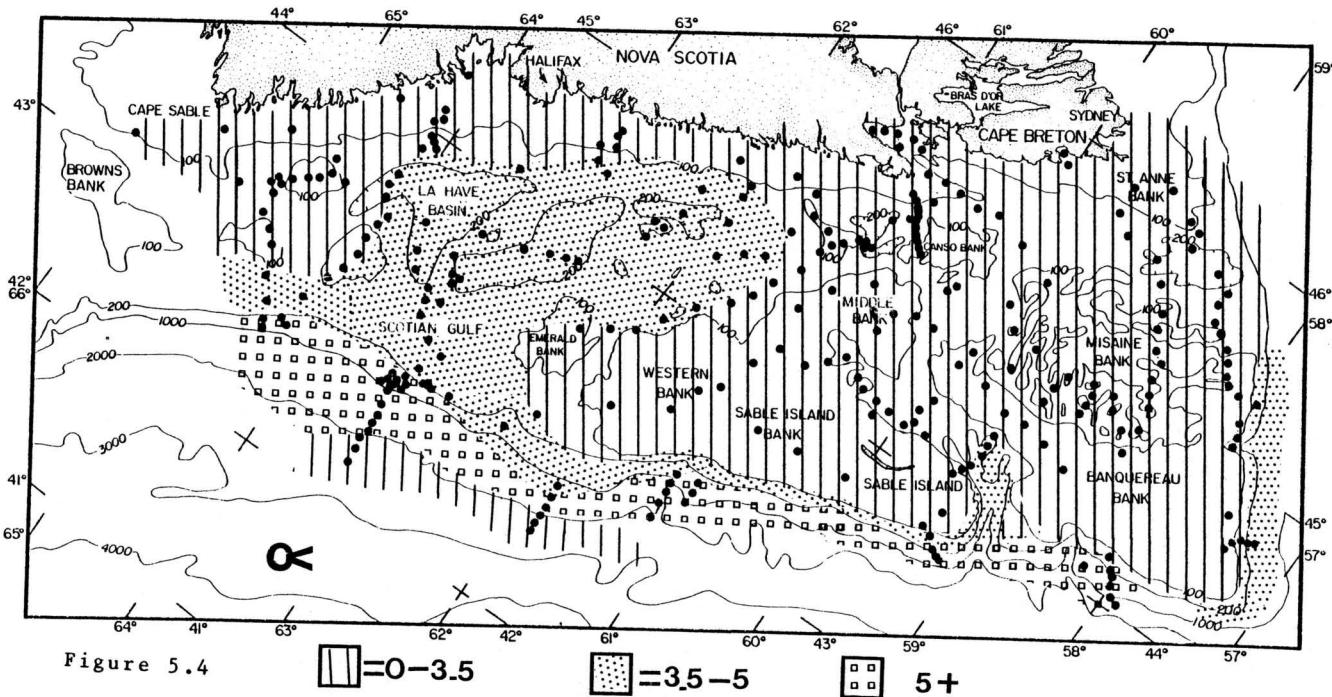


Figure 5.3



regional patterns of H(s) and S, with the highest values in the upper slope regions.

#### 5.3.4 Equitability E (fig. 5.4)

Values of equitability show more variation between stations than the other indices; it was consequently harder to define important trends. (Values from 0-0.4 were taken as low trends and 0.4-0.7, as high). Lowest values (i.e. lowest equitability, and therefore more likely to be mono- or duo specific stations) occur on the outer banks, most of the inner shore areas and lower slope regions. Higher values are found on the shelf edge-upper slope, LaHave and Emerald Basins and Misaine Bank-Basin regions.

### 5.4 Summary of Diversity Patterns.

Figures 5.1-5.3 enable several generalizations to be made concerning species diversity in this area.

Whichever index is referred to, the lowest values are consistently found in the nearshore, outer and central Banks and the northeast Bank and Basin complex. These areas are for the most part shallow (less than 80 m.) with a boulder, gravel or sandy substrate (exceptions to this lie in the northeast section around Misaine; some of the basins there exceed 300 m. and have silty clay sediment substrates). Values increase in the central basin area as does depth (greater than 180 m.) with silty clay sediments becoming predominant. Maximum values are seen along the shelf edge and upper slope regions with fine sand-silty substrates. The Scotian Gulf area also maintains a medium to highly diverse fauna.

Values for E are not as well defined. Table 5.1 shows four broad geographic areas that may be defined in terms of H(s), E, S and  $\alpha$ .

### 5.5 Standing crop and live determinations.

Standing crop is the total number of benthic foraminifera found living at any instant in a specified sample volume. Various sample sizes for measuring this value have been used as described by Gregory (1971). This study measured the live totals in a 20 ml sample.

Live Total ratios are the relationship between numbers of living and total populations in each sample. Both these values have, in the past, been found to exhibit great variability within areas of interest (Bartlett, 1964; Murray, 1968; Nichols and Norton, 1969; Gregory, 1971; Douglas, 1978). The present study area is no exception (fig. 5.6)

Several factors are thought responsible for this variability (Douglas, 1978). The effects of sampling techniques on live recovery have been described in an earlier section (chapter 1). Low rates of sedimentation would tend to lower both values, as would unstable and reworked sediments (Gregory, 1971; Douglas, 1978). Variations in rates of reproduction such as seasonal effects result in different values obtained at different times of the year. Relict sediments would also have an effect on the values obtained.

These factors essentially reduce the importance of Live Total ratios and standing crop in defining general faunal patterns. Since it is important to know the extent of these error producing factors, consideration will be given to their spatial variability.

### 5.5.1 Standing crop.

The largest living populations are present (Fig. 5.5) along the Eastern boundaries of the study areas, St-Ann's Bank, Misaine Bank and Basin as well as inner shelf areas East of Halifax. Other high values are seen near Emerald Basin and Cape Sable. As can be observed on the map, large areas are devoid of any live populations. These include most of the outer Bank regions (very low or no sedimentation) and the Central shelf bank areas (as well as some basins). Variability in the values is apparently not correlated with depth. It remains to discuss the variability shown within this area and the low values of standing crop when compared with those of other areas.

Phleger (1956, 1960, 1964) has previously noted large living populations near river mouths. It is generally thought that these are a response to high organic productivity associated with mixing of nutrient-rich fresh water. This might explain the high values for standard crop in and around the approaches of Chedabucto Bay. The high values off St-Ann's Bank may equally be associated to an upwelling along Cape Breton Island's southern shore (Cok, 1970) or the adjacent St-Lawrence outflow. Upwelling of nutrient-rich waters has been cited as a possible factor in producing high standing crop values (Phleger, 1964; Douglas, 1978). Certainly, relatively high values for standing crop are seen in some places along the shelf edge, where upwelling is in operation. The cause for increased biological activity in the shelf/slope region is not well understood. Fournier *et al* (1977) have shown that it may be the result of nutrient concentrations caused by increased vertical turbulence. If so, this would explain these "highs".

The low values associated with much of the outer bank areas is a

Figure 5.5 Standing crop distribution on the continental margin off Nova Scotia.

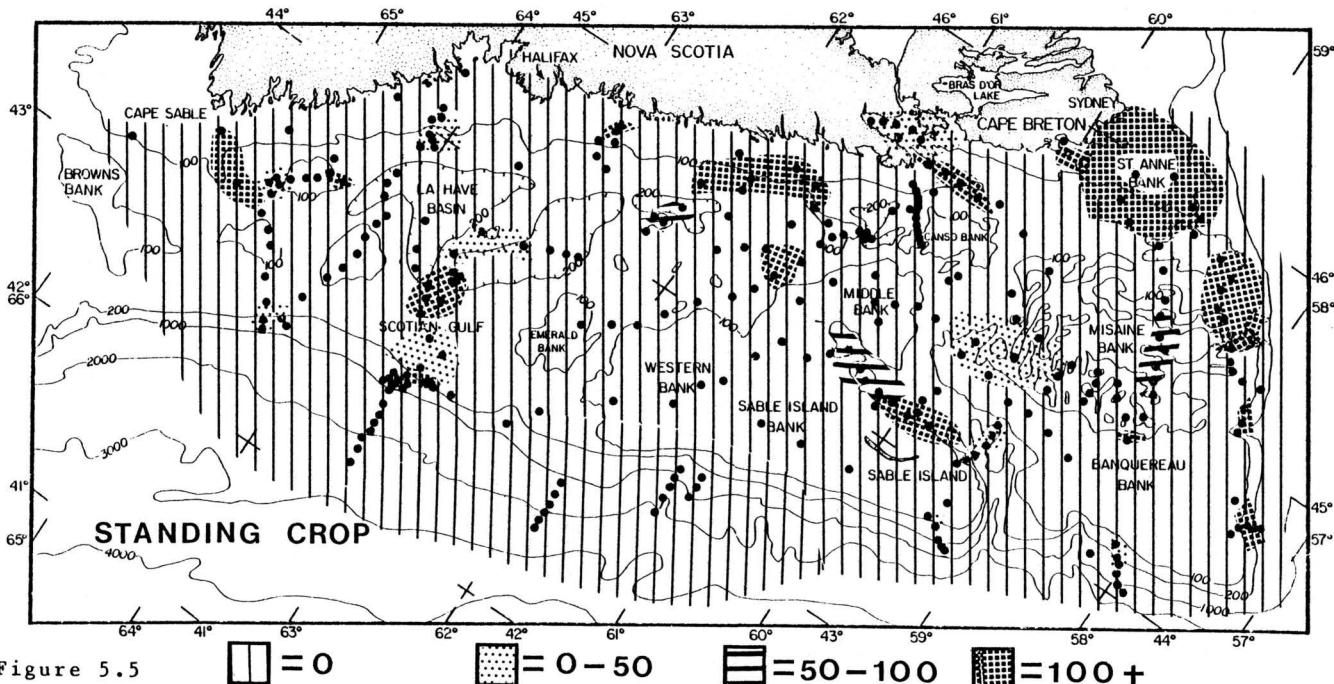


Figure 5.5

Figure 5.6 Live/Total ratios off Nova Scotia.

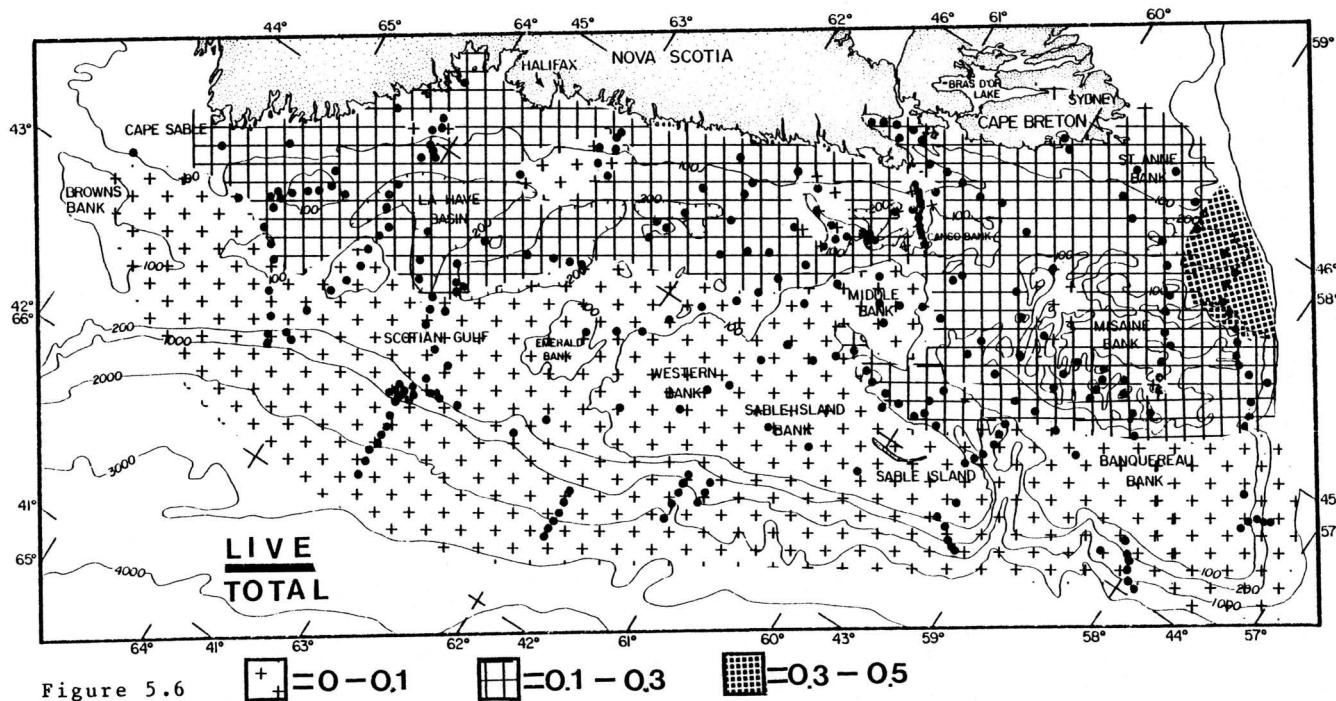


Figure 5.6

result of the overall lack of foraminiferal populations on these banks due mainly to their reworked nature. Some of the coarser inner banks are dominated by Cibicides lobatulus and show higher standing crop. Values are very low or absent on the slope and is a feature commonly recorded for slope assemblages (Miller and Lohmann, 1982). This may be explained by dilution through transportation, relictness or increased turnover.

Compared to other areas on both the eastern and western seaboard of continental North America values observed for standing crop in this area are low. Ellison (1979) recorded higher densities of foraminifera on the Mid-Atlantic Bight with large variability from the swales (depressions)— high values, to the ridges— low values. He also recorded a general increase from shallow to deeper waters. Akpati (1975) described high values from Long Island Sound although again with a high variability. This variability in standing crop is a consistent feature within areas and has been documented in Ellison, 1977; Akpati, 1972; Murray, 1969; Douglas, 1979; and Buzas, 1965, 1969. Gregory's data (1971) from Halifax Harbour Nova Scotia also exhibits great variability and overall low values when compared with other areas. Low values of staanding crop on the slope off Nova Scotia are most likely a result of low sedimentation rates which increase the relative abundance of dead tests. On the shelf itself low values could be the result of several factors including (as seen in chapter 4) relictness, transportation etc. and perhaps the type of reproduction; whether sexual or asexual (the former capable of producing millions of offspring which subsequently would be added to the dead population).

Not much importance should be attached to the distribution of standing crop seen in this study as these values are known to vary with

season. For reasons stated in chapter 1 and 2 seasonality of the foraminiferal occurrences could not be directly examined in this study.

#### 5.5.2 Live/Total Ratios.

Lowest values of this ratio occur consistently on the outer banks and some central shelf banks (fig. 5.6). This is likely to be a reflection of the low level or absence of sedimentation in these areas. Medium to high values occur landward of the outer banks, within the central basin complex, inner shelf areas and Misaine Bank/Basin complex. Live/Total ratios have been used in the past to estimate rates of sedimentation. In this study, such estimates would be of doubtful validity. Not only has there been transportation and winnowing of some of the tests but there is also a lack of data on species' life cycles. This, together with the observed erratic variations in values on the shelf, would preclude its use in estimating rates of sedimentation. Bartlett (1964) and Gregory (1971) came to similar conclusions. Gregory (1971) also documented, as did Bolstovskoy and Lena (1969), some infaunal habits of benthic foraminifera which seem to further affect these estimates.

#### 5.6. Discussion of faunal trends.

Previous observations (Deevey, 1969; Hessler and Sanders, 1967; Schnitker, 1971.) show that under extreme adverse environmental conditions, individuals of only a few species in a faunal assemblage will comprise a large number of the population and that, under normal conditions, individuals of the dominant species are less abundant,

allowing a greater number of ecological niches and resulting in an increase in diversity. Within these assumptions lies the usefulness of faunal diversity measure.

It is the purpose of this section to account in some part for the geographical variations in diversity values. Consideration will be made further, in a later chapter following the discussion, of the ecological constraints controlling assemblage distributions.

The full nature and extent of these variations should be considered. Buzas (1979) has shown that there are a limited number of possibilities when comparing such indices as S, H(s), E and  $\alpha$ . Areas with apparently similar diversities may not be directly comparable in terms of the general ecological constraints influencing these values. For example, low values of H(s) in the deep slope areas and on the northeast Misaine Bank/Basin areas. These values are low for differing reasons, indicated with reference to their S and E values. It appears that the slope region has a greater number of species (S) and lower equitability (E) than the Misaine area i.e. the greater number of species is less evenly distributed. This implies that slope areas are more dominated by one species with large numbers of less frequent species. This further implies a less stable environment than the relatively higher equitable Misaine area. Observation of the lower slope assemblages (Chapter 4) has revealed a dominance of Elphidium excavatum and a belief that this form has diluted an original Uvigerina peregrina assemblage (via transportation of relict fauna from the shelf). The low E values seem to support this idea. On the other hand, the high E values in the Misaine area tend to reflect a more stable assemblage, unaffected by transportation processes, etc.

A similar effect is seen in the Scotian Gulf area. Figures 5.1-5.4

show very high S values, with medium to high values of  $\alpha$  and H(s). Values of E, however, are low and indeed, these stations are dominated by E. excavatum with large amounts of less frequent species. It is thought here that the E. excavatum component of the assemblage is relict, imprinting and masking an otherwise diverse fauna.

Similar diversity values are seen for the northeast Misaine Bank area and the inner shelf nearshore regions west of Halifax. Both regions exhibit low H(s),  $\alpha$  and S. Values for E, however, are much lower in the inner shelf area. It has previously been shown that the shelf assemblage 2 Cibicides lobatulus (Chapter 3) is predominant in this area. The monospecificity of this assemblage is a direct cause of the low E value and possibly a reflection of the relatively extreme environmental conditions prevalent here (e.g. shallow, turbulent, bedrock-coarse sand-gravel environment). The higher values of E in the Misaine area imply a stability associated with more "open", marine conditions. Low values also occur within this Misaine, northeast area but are restricted to St-Ann's Bank which is the only pocket in this region that maintains a C. lobatulus fauna. Finally, the central basin complex, upper slope and shelf edge, which consistently exhibit medium to high values of S, H(s),  $\alpha$  and E imply stable diverse faunas.

According to time stability theories (Hessler and Sanders, 1967), this could mean that these areas have been submerged longer. The same case can apply to the upper slope but is unlikely for the central western Basins. It is probably more correct to invoke the characters of the slope waters in the latter area, with its relatively high temperatures and salinities, nutrient-carrying capabilities and ability to frequently incuse into the central Basin areas. This will produce a less stressed environment with a resultant increase in diversity relative to the

remainder of the shelf. This explains why the diversities in the eastern central shelf Basin complex (around Misaine) are lower when compared to central basin areas: the colder, more restricted circulation in the former (St-Lawrence bottom waters) causing more adverse conditions.

Previous studies have well documented diversity trends off the North American eastern continental margin. A general increase in diversity values with distance from shore has been observed by Schnitker (1971), Ellison (1979) and Arnold and Sen Gupta (1981) in transects off Delaware, North Carolina and Georgia's continental margins, respectively. These studies also show peaks of diversity at the shelf edge.

Gibson (1966) has related population diversity to environmental variability. This is possible for large scale gradients. On the other hand, Gregory (1971) was unable to correlate local environmental changes within Halifax inlet with diversity indices. Murray (1968), working through the literature, observed low  $\alpha$  values (averaging 1-4) in Tracadie, St-Margarets, Miramichi and Buzzards' Bays as well as in Long Island Sound. All inner shelf areas are similar in diversity to this study.

In these respects, the present study reflects previous observations. However, as has been shown in the maps and discussion, the simple idea of an increase in diversity with distance from shore and increasing depth is not strictly tenable for this investigation.

By way of a summary, several aspects of the varying physiography, bathymetry, sedimentology and oceanography appear to explain the variation in diversity values and the occurrence of four geographic regions on the shelf, each with distinctive values.

It has been seen that where cold St-Lawrence bottom waters lie on the shelf and the inner shelf area west of Halifax (areas 4 and 2 on Table 5.1), the populations have a low diversity (cold, poorly circulating, harsh environment, lower niche availability). However, local effects of substrate and wave energy (i.e. bedrock, wave-swept area of the inner shelf west of Halifax) lead to an even lower diversity (as seen, for example, in values of E).

Those areas that are influenced by warmer slope waters; (the upper slope, shelf break and Central Basin complex) by incursions through the Scotian Gulf are more diverse due to increased availability and optimum conditions of temperature, salinity, nutrients, etc. Low values of diversity observed on the lower slope are believed to be a result of faunal mixing through transportation of Elphidium excavatum downslope, creating a virtually monospecific population.

The basic control then, of diversity values on the continental margin off Nova Scotia, appears to be hydrographic. Local conditions, however, such as substrate and wave energy, may act to negate this influence, locally depressing diversity values.

In a general way, these observations fit in with the idea of Slobodkin and Sanders (1969) whose "unpredictable" environments (Sen Gupta and Kilbourne, 1974) display a greater variability of physical energy concurrent with depressed values of diversity (as observed in the inner shelf and outer Bank areas of this study).

A later section more thoroughly discusses the relationship between diversity and the environmental parameters.

### 5.7. Implications of diversity for Pleistocene to recent studies.

It has been seen, that values of diversity (including E) are associated in this study area with certain properties of the environmental regimes, in particular, high diversity correlate with warm nutrient-rich waters. The implications of this for palaeoceanographic study of the Pleistocene to Recent distribution of these watermasses are clear. At the very least, they will provide qualitative support to more quantitative studies based on assemblage distributions. At the most, they could provide independent quantitative evidence concerning palaeoceanographic distributions of watermasses off Eastern Canada. These aspects will be further investigated after the ecological section.

## CHAPTER 6: ASSEMBLAGE / ENVIRONMENT RELATIONSHIPS

## 6.1. Introduction

The philosophy and evolution of thought concerning foraminiferal ecological studies, from Natlands' (1933) impetus to present day quantitative methods, is well reviewed in Douglas (1979). This work shows the development from purely depth related assemblage considerations to the complex analyses of many interacting physical and biological variables. All these historic and more recent works aim to apply the present to the past. They attempt to observe present day distributions and their causative factors and apply them to geological occurrences. Similarly, the stated aims of this present work are based on this concept. This decidedly geological approach has often led biologists to argue that it neglects rigorous scientific treatment and observation.

The art of geological science, in particular for those of us who are interested in palaeoenvironmental interpretation of past climatic and geologic events, in both the Recent and the more distant geological past, requires us to make some assumptions which are untenable to most biologists and all theoretical ecologists. For instance, purely biological or ecological study of foraminiferal habitat requires intense seasonal and controlled study. Such an approach, in a study whose aim is to make and apply estimates of past conditions on the basis of time averaged populations (Walker and Bambach, 1971), would be time consuming and prohibitive for regional studies of an area as large as the shelf

off Nova Scotia.

Basic biological and ecological theory states that each species has its own tolerance range for a particular environmental factor under consideration. Biologists then, would rightly argue that many computer-identified biofacies are not true natural entities (Pielou, 1979) and that to assume that they are (i.e. in terms of the environment) is erroneous. The fact that many palaeoenvironmentalists seek a "safety in numbers approach" (i.e. assemblages) is due to the fossil material to be dealt with and its inherent problems of selective preservation, transportation, vertical mixing etc.

The previously stated aims of this study are to identify recurrent foraminiferal assemblages and to relate them quantitatively to the present day watermasses off Nova Scotia. No assumptions are made concerning the relationships of individual species to strict environmental ranges. The only assumption is that the species together within an assemblage find their optimum conditions there, and together, reflect these conditions.

Several tenets of ecological theory may help to put the present study in better perspective.

The hitherto defined factor assemblages are assumed to be communities in that they comprise a system of species "living and linked together by their effects on one another and their responses to the environment they share" (Whitaker, 1970). Exceptions to this are seen in some of the assemblages, the main components of which are relict, i.e. not in equilibrium with their environment or with the other species in the assemblage.

Community theory shows two aspects of community composition:

1. Taxonomic composition, i.e. a list of all species found in the community;
2. The rank and relative abundance of these species.

The nature of these aspects are controlled by environmental factors and chance (the latter being bias and error introduced through sampling and spatial inhomogeneities within communities). More specifically, in division 1, the taxonomic composition is controlled by one or more of the following: space, light, salinity, trace elements, temperature, pressure, substrate, dissolved gases, nutrients, biological interactions (predation, parasitism, symbiosis, commensalism) and factor combinations of turbidity, abrasion, etc (Walker and Bambach, 1974). Biological activity is capable of depleting some of these (space, light, O<sub>2</sub>, nutrients); these are termed "limiting resources". It is the competition for these limiting resources (plus chance) that controls the nature of division 2, i.e. the relative and rank abundance of the species. Walker and Bambach (1971) show that elements of chance are reduced due to the time averaged nature of fossil communities.

Another, often neglected, aspect of community study is community function, or its autecology. This is the relationship of an individual species to its physical and biological environment (Walker, 1974). Autecological studies involve the investigation of feeding, metabolic, reproductive and protective mechanisms within the organisms.

With this hierarchical approach mentioned above in mind, it remains to investigate the degree and extent to which the assemblages on the continental margin off Nova Scotia (as outlined in chapter 3, and 4 of

this study) are related to their environment. Particular aspects of the environment to be investigated include: depth, temperature, salinity, gravel, sand and mud contents. Their intercorrelation will be examined quantitatively. In addition, several aspects of nutrient, organic carbon, O<sub>2</sub> and their potential for affecting distributions will be examined in a qualitative manner. The source of these oceanographic data, methods and procedures of correlation, together with their results, will be discussed in the following section.

## 6.2 Environmental and Oceanographic data.

The nature and spatial distribution of various environmental variables have been described in a general way in Chapter 1 (i.e. Bathymetry, Surficial Geology and Oceanography). These particular environmental regimes, presently occurring on the shelf and slope off Nova Scotia, are the main variables under investigation in this study, in terms of their influence on the benthic foraminiferal distributions.

Data pertaining to percentage values of gravel, sand and mud contained at each station were obtained from the unpublished files of Dr. L.H. King at the Bedford Institute of Oceanography. This data formed the basis of the surficial sediment maps of King, (1969). Temperature and salinity values were extracted from the Oceans IV files of the Canadian Marine Environment Data Service. These were made available, along with help and advice, by Dr. C. E. Keen also at the Bedford Institute of Oceanography, Dartmouth. These data, together with information provided in various publications (McLellen, 1957; Houghton

et al, 1977; Fournier, 1978) provided a good coverage of the study area in terms of temperature, salinity and substrate characters. Values of temperature and salinity are annual averages and represent time averaged measurements of bottom waters up to 20m above the sea bed. Reasons for this were described earlier and include the averaged nature of the assemblages, time averaged fossil populations to be considered and the limited and patchy nature of the data available for seasonal analysis.

The remainder of this chapter endeavours to illustrate the degree and significance of assemblage environment associations. Section 6.3 describes the statistical methods involved and is followed (in 6.4) by a section that describes the more obvious assemblage environment associations, such as are seen in correlation matrices and through simple plots of assemblage importance (factor loadings) versus observed values of each environmental variable. The statistical significance of these relationships are presented and discussed towards the end of the chapter (6.5, 6.6).

### 6.3 Statistical method.

The problem at hand then, is to investigate the nature of the relationships between the previously defined, statistically independent factor assemblages and a set of observed environmental variables.

In order to be directly comparable with similar studies underway off the east coast of Canada (Mudie et al, 1983), these relationships are best investigated through regression analysis. Input to the

regression analysis consists of a set of dependent variables and a set of independant variables. The former are taken from the varimax factor matrix and comprise the factor loadings at each station for 8 shelf and 4 slope factors. The independent variables consist of a set of observations or extrapolations on 6 environmental variables at each station (depth, temperature, salinity, substrate nature). The regression analysis program used in this study (from SPSS) employs the method of least squares to obtain equations which relate the independent variables (environmental variables) to dependent variables (factor assemblages). Imbrie and Kipp (1971) obtained both linear and curvilinear solutions for their planktonic foraminifera data with the curvilinear equations containing the squares and cross-products of the factor assemblages. They found the curvilinear solution gave a more accurate fit between observed and predicted environmental variables. The ecological meaning of the curvilinear solutions however remain rather vague. In a recent paper, Mudie et al (1983) compared the results of the two solutions in a study of benthic foraminifera and found an increase in the standard errors of the estimates when the curvilinear solution was used. For this reason, the more robust and conceptually simpler linear solution was used in this study. Lohmann (1979) also found the linear solution more useful for study of slope foraminifera.

The statistics and methodology of regression analysis are well treated in many texts such as Mendenhall (1971) and Lapin (1975). The regression procedures employed in this study uses standard stepwise multiple regression techniques (Imbrie and Kipp, 1971) involving the addition, at successive stages, of single "factor" terms to the equation for a selected environmental variable. The program selects the variable

(factor assemblages) that contributes most to the efficiency of the prediction as the first step, then proceeds to add the variable accounting for the next closest fit to a regression line etc.

Statistical evaluations are made of the success of each equation at each step in the analysis. In this way, one is able to determine which assemblage is the best criterion for each environmental variable, the second and third best, and so on, and subsequently to characterize each assemblage in terms of its environment.

The application and use of regression analysis in similar studies to the present is described in Imbrie and Kipp (1971), Molfino et al (1982) and Mudie et al (1983). In simple terms, the regression procedure fits a "best line" to the data input. The line is such that the sum of squares of the vertical deviations from all points to the line is a minimum. From the regression equations are derived estimates of the particular environmental variable under examination, the standard error of these estimates being equal to the square root of the mean squared deviations. The coefficients produced are derived through minimizing the sum of squares of errors of the estimates. An F test indicates the significance level of each multiple regression equation; a t test quantifies the significance of the individual regression coefficients. The multiple correlation coefficient indicates the correlation between the observed measures and their estimates. The proportion of variance of the observations explained by the equation is equal to the square of the multiple correlation coefficient. Unexplained portions are, according to Imbrie and Kipp (1971) a result of random error, this aspect will be discussed later.

Prior to a description and discussion of the regression analysis results, the more obvious features of assemblage - environment

relationships will be examined in the next section.

#### 6.4. Main Assemblage - Environment Associations.

Some idea as to the main relationships may be obtained from the correlation matrices. These correlation matrices only show the strength of associations. Large positive or negative values do not indicate whether high values of one caused the others to be large, etc, i.e. do not measure causality. They do, however, permit some general trends to be observed and as such are useful.

For shelf stations (Table 6.1) increasing depth is accompanied by an increase in mud % (i.e. as observed from bank to basins) and a decrease in sand %. Temperature values appear associated with salinity but neither of these vary much with depth. The correlation matrix (Table 6.2.) for the slope samples shows strong associations of temperature with depth and salinity but not much between salinity and depth.

From the shelf correlation matrix (Table 6.1.), Factor assemblage 1 shows some association with temperature and salinity but little else. Factor 3 appears positively related to depth, temperature, percent mud and salinity but not to gravel. Factor assemblage 4 is marked by its apparent lack of association with any other variable. Factor assemblage 5 shows relatively high correlation with temperature, salinity and percent sand whereas factor 6 is only associated with depth and sand. Factor 7 is associated in the correlation matrix with percent gravel, unlike factor 8, which exhibits some association with temperature,

**Table 6.1 Correlation matrix for shelf assemblages:environmental parameters.**

	DEPTH	TEMPERAV	SALINIAV	PERCGRAV	PERCSAND	PERCMUD	F1	F2	F3	F4	F5	F6	F7	F8
DEPTH	1.000													
TEMPERAV	.264	1.000												
SALINIAV	.359	.829	1.000											
PERCGRAV	-.362	-.095	-.180	1.000										
PERCSAND	-.333	-.089	.037	-.128	1.000									
PERCMUD	.502	.135	.065	-.421	-.844	1.000								
F1	.084	-.579	.448	-.174	-.080	.172	1.000							
F2	-.386	-.088	-.071	.463	.263	-.496	-.475	1.000						
F3	.350	.359	.281	-.209	-.387	.464	-.356	-.268	1.000					
F4	.076	.018	-.127	-.199	-.156	.247	.151	-.335	.022	1.000				
F5	.120	.333	.421	.049	.348	-.342	-.427	.285	-.157	-.222	1.000			
F6	-.306	-.007	-.111	-.045	.240	-.193	-.067	-.190	-.168	.037	-.131	1.000		
F7	-.256	.027	.070	.352	.158	-.338	-.279	.368	-.135	-.182	.111	-.075	1.000	
F8	-.285	-.636	-.637	.058	.120	-.142	.424	.333	-.248	.233	-.237	.167	.115	1.000

Table 6.2 Correlation matrix for slope assemblages:environmental parameters.

	DEPTH	TEMPERAV	SALINIAV	F1	F2	F3	F4
DEPTH	1.000						
TEMPERAV	-.715	1.000					
SALINIAV	-.300	.747	1.000				
F1	.696	-.638	-.133	1.000			
F2	-.649	.483	.109	-.597	1.000		
F3	-.222	.074	-.169	-.368	-.092	1.000	
F4	.280	-.363	-.412	.044	-.066	.128	1.000

salinity and little else.

Similarly, slope factor assemblage 1 shows high correlation values with temperature and depth but not with salinity (Table 6.2.). Slope factor 2, however, is more associated with depth than temperature. Factor 3 appears weakly related to depth. Finally, slope factor 4 shows association with salinity and, more slightly, with temperature and depth.

The correlation matrices for the shelf and slope therefore suggest some assemblage - environment associations. Further detail on these associations is provided in Figs. 6.1-10. These diagrams plot factor loadings (assemblage importance) for each assemblage against observed values for some of their associated environmental variables (Raw data is given in appendix VI). The diagrams permit comparison of the differing response of an assemblage to various aspects of environment, and show the degree of sensitivity of a particular assemblage to certain conditions of that environment. As described in Molfino *et al* (1982) and seen in an earlier section of the present work (Chapter 3), these diagrams allow more detailed comparison of whole assemblages (i.e. including low and high factor loadings from each assemblage) and their individual response to environment than would a consideration of high factor loadings only (i.e. as are used to produce factor assemblage maps - Chapter 3). The lines shown on Figs. 6.1-6.10 are fitted by eye and are meant only to point out any obvious associations or non associations. It will be the purpose of a later section to statistically assess the strength of these associations. The following describes the main features illustrated by these Figures.

#### 6.4.1 Shelf Assemblages.

##### Factor assemblage 1: Adercotryma glomerata.

Factor loadings of this assemblage appear to be related to temperature and salinity (Fig. 6.1; a,b). From these diagrams, it appears that factor 1 may be more responsive to temperature change than to salinity, the latter showing greater amounts of scatter. The highest factor loadings lie within a range of 1.5-2.3oC and 31.5-32.5‰/oo salinity. Similar plots for factor 1 against depth and substrate characters (not shown) show large amounts of scatter with no discernable trends. Considering the distribution of this assemblage, this is not entirely unexpected. It is found both in deep basins and on shallow banks (and therefore on muds, gravels and sands) as well as inner shelf nearshore areas.

##### Factor assemblage 2: Cibicides lobatulus.

An inverse trend with depth can be distinguished for assemblage 2 (Fig. 6.2a). Highest factor loadings occur at 50 to 100 m depth, as would be expected in view of this assemblage's distribution in nearshore areas or on shallow banks. The importance of this assemblage can be seen to decline steadily with depth. A less clear trend is seen with temperature (Fig. 6.2b) highest factor loadings seem related to a wide temperature range of 1.5-6oC, with lesser values (factor loadings) at higher temperatures (7-10oC). The increased scatter implies lesser dependence on temperature. Figures 6.2, c and d, may partly explain this. These diagrams show a direct relationship between the assemblage and substrate characters of percent sand and gravel, high percent values

Figures 6.1--6.10 To graphically illustrate the main assemblage environment associations, or non-associations. Plots of assemblage factor loadings (vertical) against values of; temperature, salinity, depth, and substrate character (horizontal).

Figure 6.1 Assemblage 1 A. glomerata with temperature, salinity.

Figure 6.2 Assemblage 2 C. lobatulus with temperature, depth gravel and sand.

Figure 6.3 Assemblage 3 G. auriculata with temperature salinity depth and mud.

Figure 6.4 Assemblage 5 T. angulosa with depth, temperature salinity.

Figure 6.5 Assemblage 6 E. advena with depth and sand.

Figure 6.6 Assemblage 7 I. islandica with salinity and gravel.

Figure 6.7 Assemblage 8 E. excavatum with temperature and salinity.

Figure 6.8 Slope assemblage 1 U. peregrina with depth temperature and salinity.

Figure 6.9 Slope assemblage 2 B. exilis with depth.

Figure 6.10 Slope assemblage 4 E. excavatum with temperature and salinity.

Lines fitted by eye

Figure 6.1

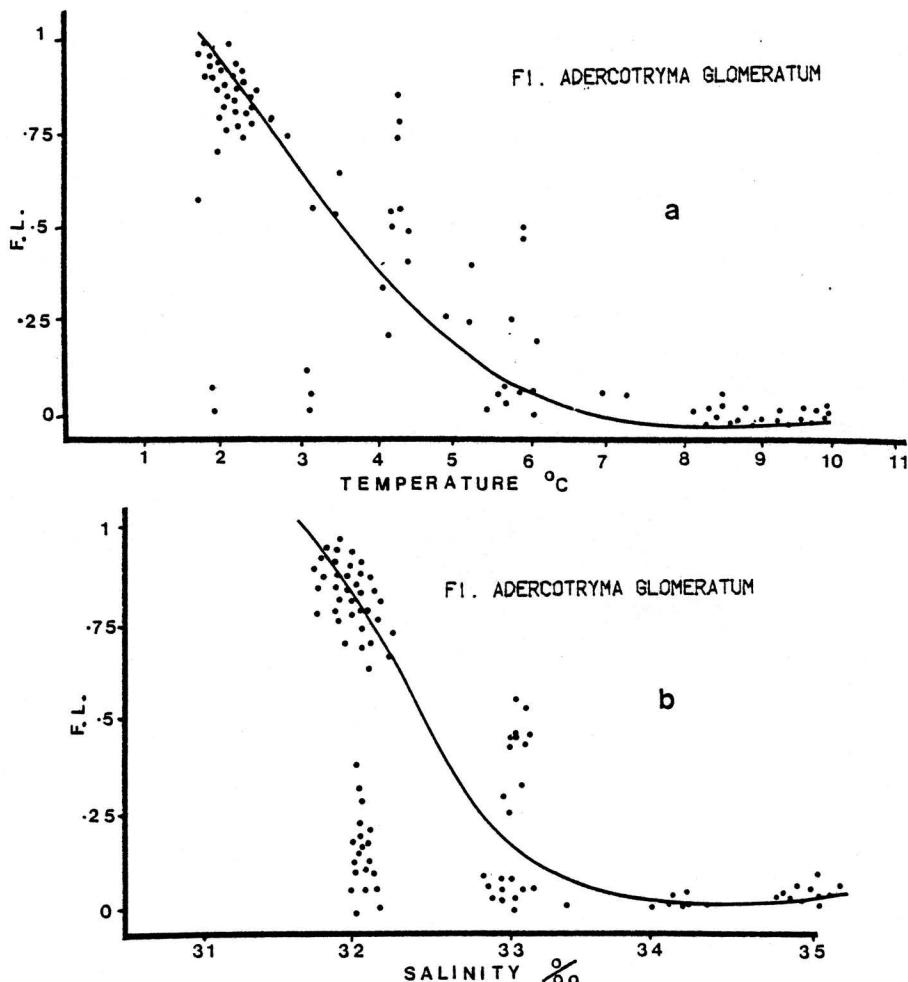


Figure 6.2

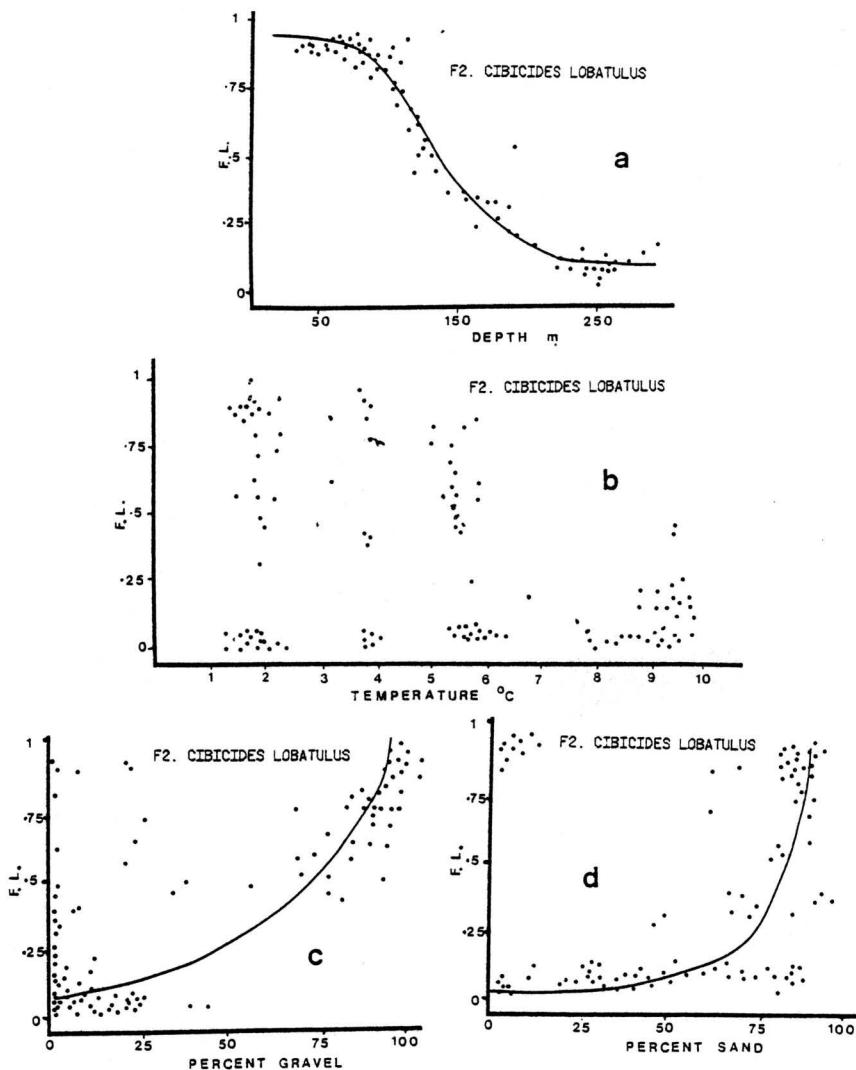


Figure 6.3

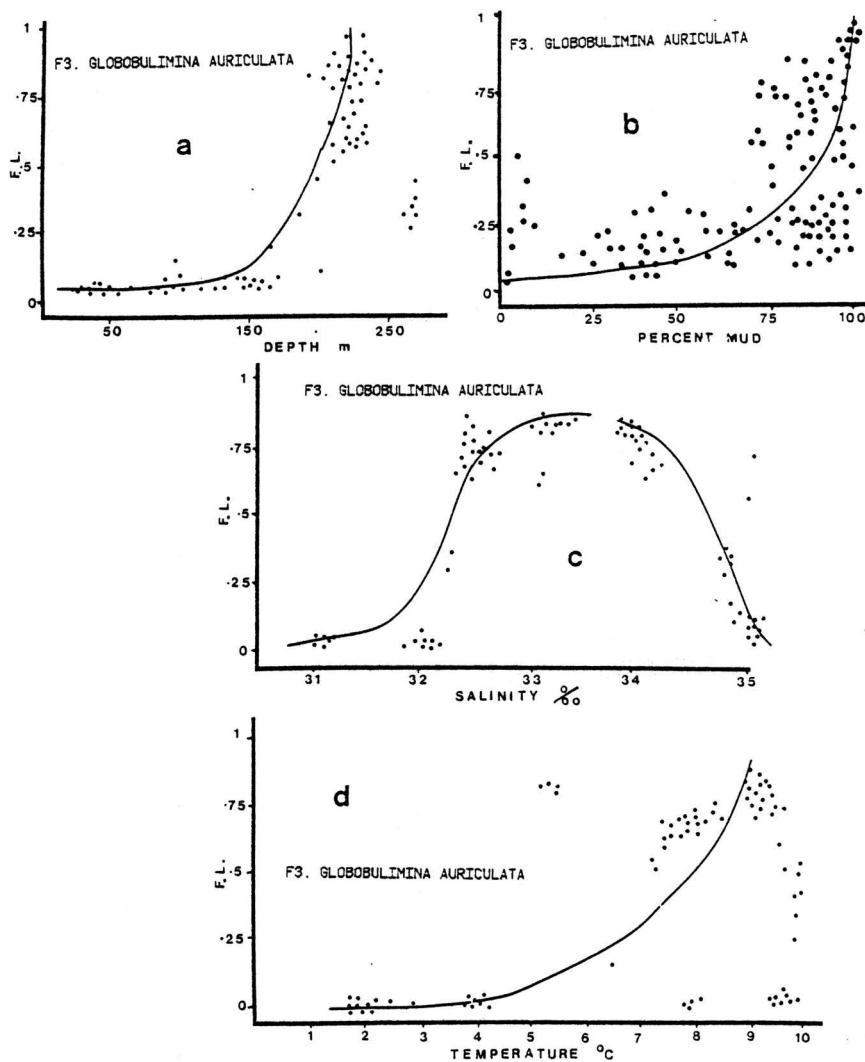


Figure 6.4

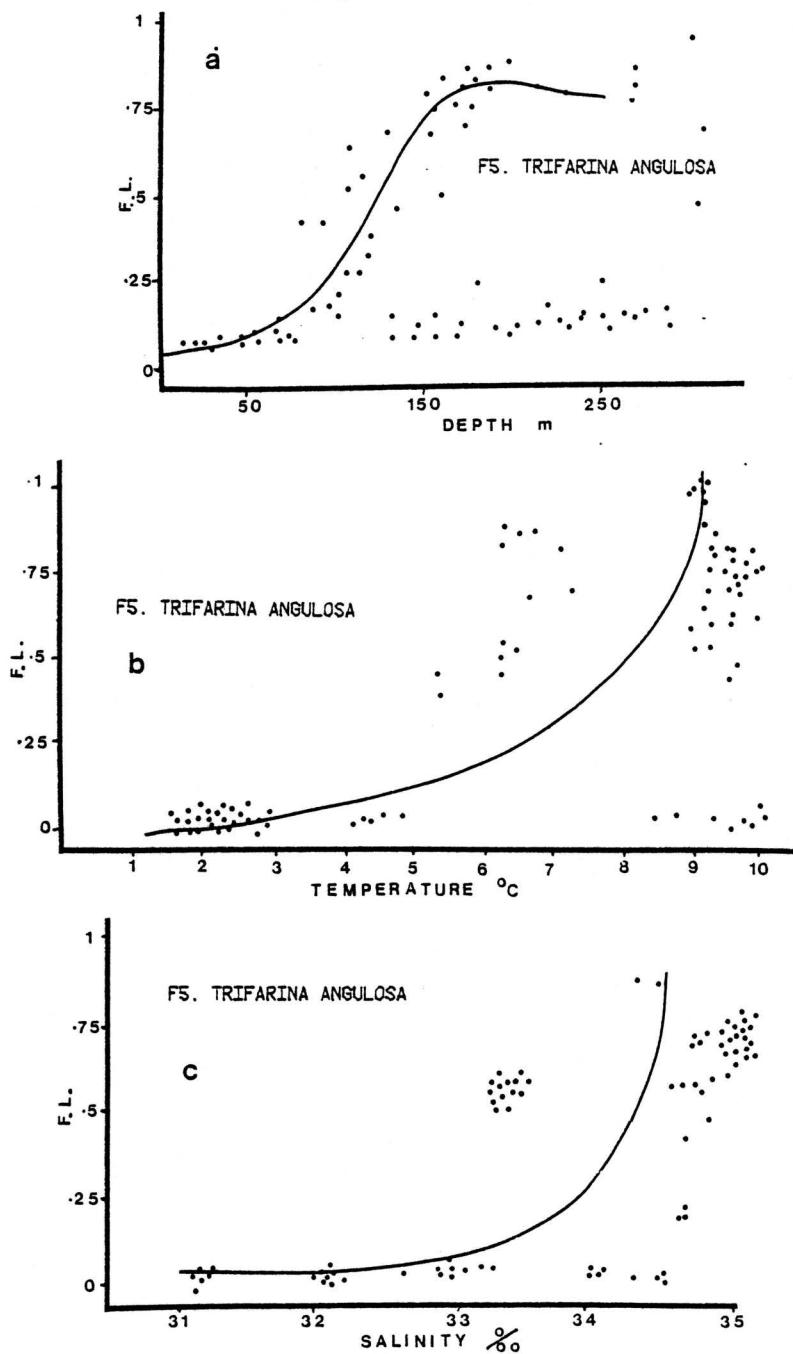


Figure 6.5

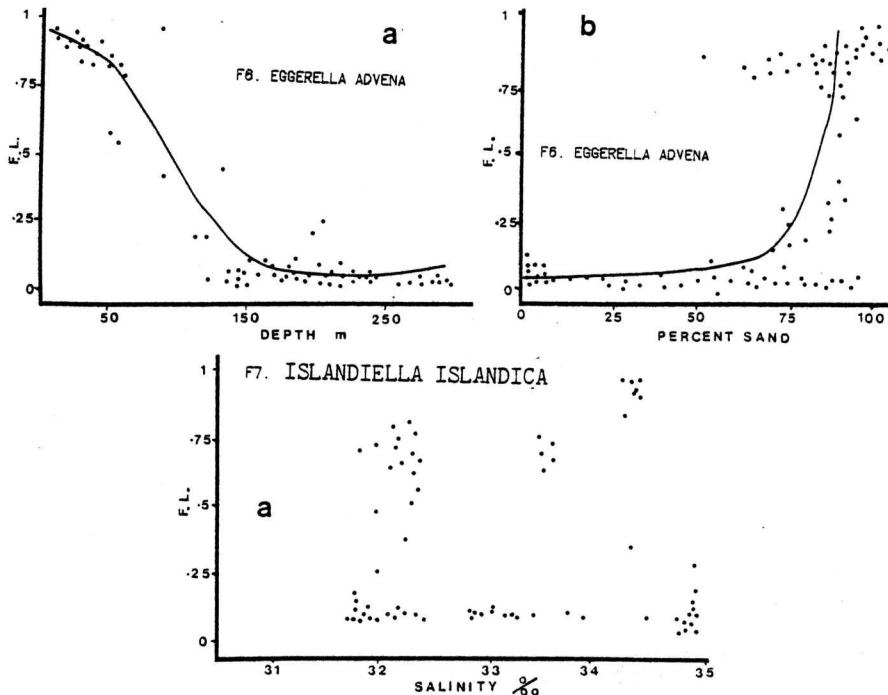


Figure 6.6

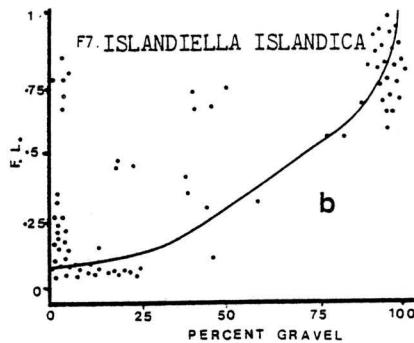


Figure 6.7

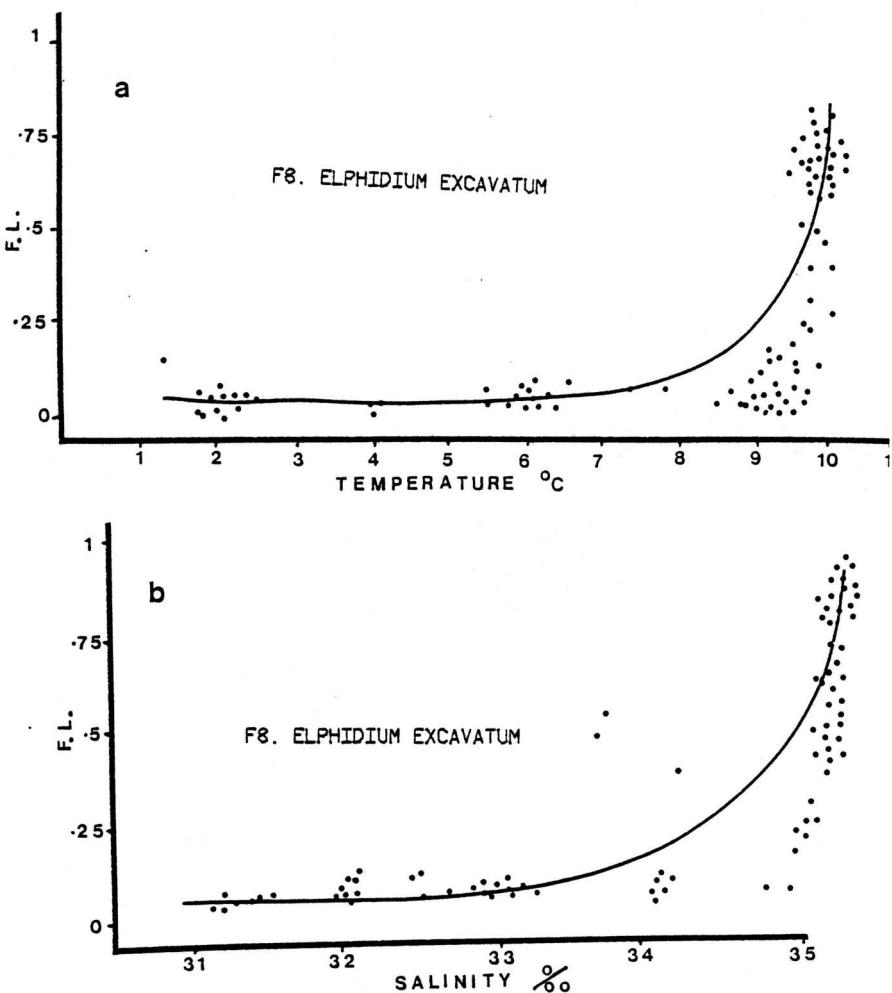


Figure 6.8

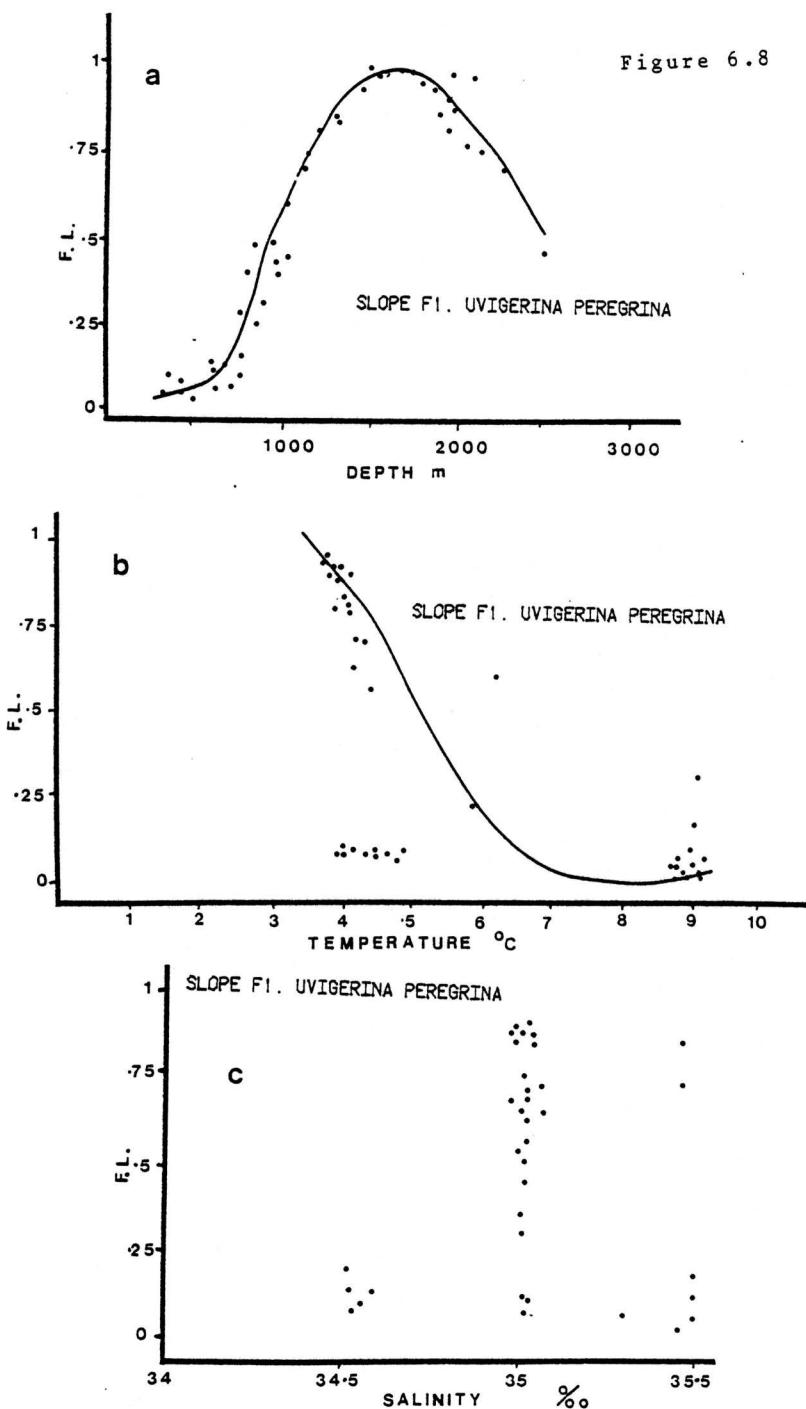
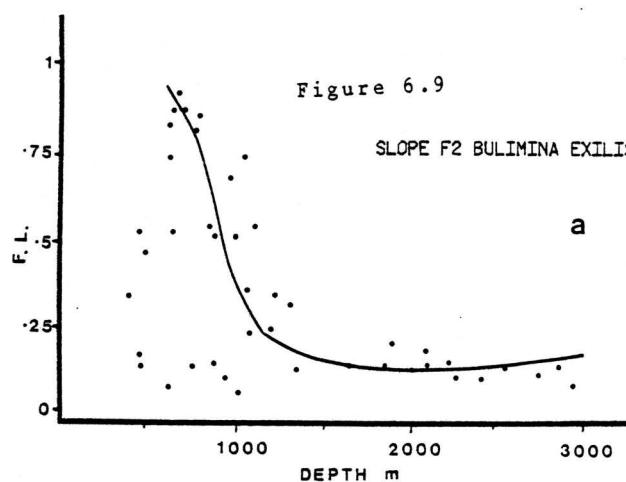


Figure 6.9

SLOPE F2 BULIMINA EXILIS



SLOPE F4 ELPHIDIUM EXCAVATUM

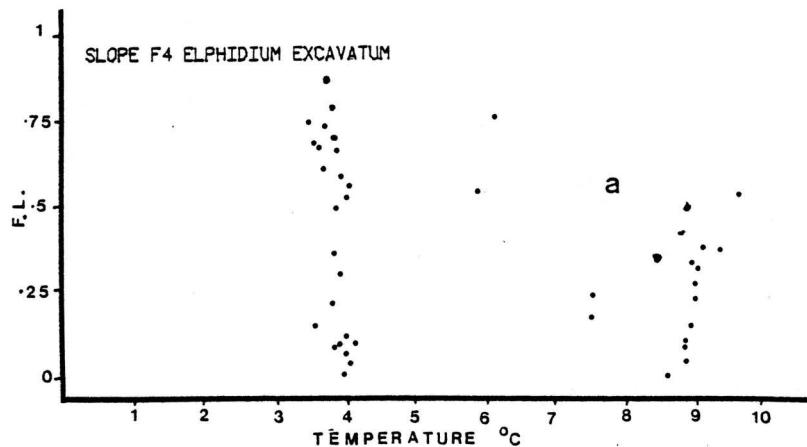
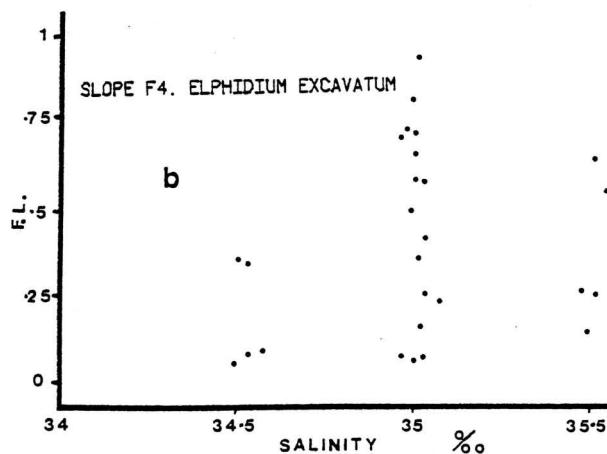


Figure 6.10

SLOPE F4. ELPHIDIUM EXCAVATUM



of both being associated with high factor loadings. It appears that this assemblage is associated with +70% sand and +70% gravel. This is consistent with previous reports (Gregory, 1971; Phleger, 1952; Poag *et al* 1980) linking the main species of this assemblage with attachment surfaces afforded by coarse substrate. No discernable trends were seen with salinity.

Factor assemblage 3: Globobulimina auriculata.

This central shelf basinal assemblage shows a direct relationship (Fig. 6.3a) with depth, maximum factor loadings occurring between 160 and 230 m. depth. A similar relationship is observed (Fig. 6.3b) with % mud: high loadings are associated with +65% mud. The only other discernable trend shows an association with salinity values ranging from 32.5-34°/oo, and an association of high factor loadings with temperatures of 7-10 °C.

Factor assemblage 4: Saccammina atlantica

Wide scatter and no discernable trends are observed in the plots concerning this assemblage. The significance of this will be discussed in a later section.

Factor assemblage 5: Trifarina angulosa.

This shelf edge assemblage, as would be expected from its distribution features, is associated with 150-250 m. depths (Fig 6.4a). It shows a strong association with temperature and, to a lesser degree, with salinity. No noticeable trends are immediately apparent between this assemblage and substrate characters. For high factor loadings, temperature and salinity ranges are from 6-10°C and 33-35°/oo. The

presence of tight clusters of points outside the main trends on Figs. 6.4a and 6.4c suggest reworked populations.

Factor assemblage 6: Eggerella advena.

The only discernable relationships apparent for this assemblage are those between depth and % sand. Figure 6.5 a, shows an inverse association between depth and high factor loadings equating to depths of 15-60 m. A strong association is observed between % sand (60-100%) and high factor loadings. Temperature and salinity show large scatters of points implying lack of association.

Factor assemblage 7: Islandiella islandica.

The only discernable trends of this assemblage appear to be with % gravel and to a lesser extent, with salinity (Fig. 6.6; a,b).

Factor assemblage 8: Elphidium excavatum.

From Figure 6.7 a,b, relationships are observed between temperature and salinity. Temperature and salinity ranges between 8-10°C and 34.5‰ are associated with high values for the factor loadings, etc. The assemblage shows no other trends when plotted against the remaining environmental variables.

#### 6.4.2 Slope Assemblages.

Factor assemblage 1: Uvigerina peregrina.

This lower slope assemblage appears positively associated to depths of 800-2000 m. and negatively related to depths of 2000-3000 m. Maximum values for the factor loadings occur between 1300 and 2000 m. (Fig.

6.8a). A temperature range of 3.5-4.5°C is associated with high factor loadings and there is a strong relationship between factor loadings and higher temperatures. (Fig. 6.8b). Relations to salinity are less clear although it appears that 35‰ is most favourable to this assemblage (Fig. 6.8c).

Factor assemblage 2: Bulimina exilis.

This assemblage shows an inverse trend with depth, maximum values of loadings occurring between 600 and 900 m (Fig. 6.9a). No discernable trends are determined with respect to temperature or salinity.

Factor assemblage 3: Trifarina occidentalis.

No apparent associations are determined for this assemblage; further discussion will be presented in a later section.

Factor assemblage 4: Elphidium excavatum.

The assemblage shows no apparent depth association over its range between 400 and 2800 m. Similarly, temperature and salinity relationships are not particularly clear (Fig. 6.10; a,b) although there is an association with temperature range of around 4°C, and salinity at 35‰.

#### 6.4.3 Discussion.

From the above observations, and with reference to Figures 6.1-6.10, it is clear that there are some obvious relationships between the factor assemblages and certain measured oceanographic variables. It is equally clear that some assemblages show no apparent associations

with these variables. Temperature for example seems not to exert a great influence on many of the assemblages, in fact only for shelf assemblages 1, 8, and 3, together with slope assemblages 1,<sup>4</sup>.

Depth seems to be the main influence on shelf factor assemblages 2,3,5,6 and slope assemblages 1 and 2. This, however, is more likely the indirect result of a response to other depth related environmental variables, the significance of which will be discussed below.

Concerning the substrate characters, percent sand appears to be the main influence on shelf factor 6 and, moderately so, on shelf factor 5. Percent mud appears to be the main control of shelf factor 3, although again, as for depth, this may be a response to other mud related variables such as the amount of organic carbon. Percent gravel seems to influence factor assemblages 2 and 7 on the shelf.

Some assemblage/environment associations and their implied relationships appear to be typically parabolic (or partly so) in response to the observed environmental variables. Imbrie and Kipp (1971) pointed to the large amount of field and experimental data from Bradshaw (1957) and Tolerlund (1969) which supports the notion that a species (or assemblage) has an optimum for a particular environmental variable, with a decrease in abundance as the measured variable increases or decreases.

It is the purpose of the regression analysis to investigate the statistical nature and significance of these assemblage/environment associations and to determine, through stepwise procedures, which assemblage provides the best criterion for each measured environmental variable. In this way, the main controlling aspects of the environment that influence an assemblage distribution can be discerned and discussed.

The next section describes the results of the regression analysis and is followed by an ecological discussion.

### 6.5. Regression Results.

The following summarizes the results of the regression analysis.

Each environmental variable is described in terms of its relationship to each assemblage (i.e. which assemblage is the best criterion for the particular variable). Regression statistics are presented in several tables (Tables 6.3-6.8) which show the cumulative results of the stepwise multiple regression procedure and significance of each factor in the final linear equation. These tables illustrate the regression results produced after applying a cut off value of 0.01; ie. factor assemblages explaining less than 0.01 of the variance of a particular variable were excluded from the stepwise procedure. This procedure gave better results in terms of lower standard error of estimates and higher significance of equations when compared to the statistics produced after applying a cut off of 0.001. The regression statistics produced in the latter case (cut off of 0.001) are included in Appendix VIII for comparison.

Several diagrams (Figures 6.11-6.19) visually show the predictive ability of each equation through plots of observations against estimates for each variable. The first diagram in each figure (ie. 11a, 12a....) plots the observations and estimates for all the factor assemblages. The quality of the fit for the data points reflects the efficiency of the multiple regression equations. Similarly, Figures

6.11 to 6.19 plot observations against estimates for regional groups of stations (i.e. assemblages as seen in Chapter 3, Fig. 3.9). Clearly, those assemblages which the regression analysis statistically deemed the best criterion for a particular variable will show tighter clusters and less scatter about the line than those deemed of lesser significance.

#### 6.5.1. Regression results of shelf and slope assemblages with:

a/ Depth.

The coefficients, standard errors and statistics produced through the regression of the 8 shelf assemblages with respect to depth are shown in Table 6.3. This Table and the others concerning the remaining environmental variables are so arranged as to show the order of stepwise selection (as in Imbrie and Kipp, 1971). At the top is the factor assemblage which is the best single criterion of depth, followed by the next best; critical values for  $t$  in the multiple regression equation are shown for the 0.001 and 0.01 levels. Of the 8 factor assemblages, only those shown to be statistically significant at a certain level are included (mainly at 0.01 level). For depth, an  $F$  value of 24.2 shows a statistically significant regression at the 99.9 confidence level. Similarly,  $t$  values show that three of the factor assemblages have significant regression coefficients at the 0.001 level and at the 0.01 level.

Factor assemblages 2 and 6 (Cibicides lobatulus and Eggerella advena) are therefore the most successful (significantly) at predicting depth estimates, explaining the largest proportion of depth variance. The standard error of estimate for the multiple regression equation is

Table 6.3

## DEPTH REGRESSION STATISTICS

SHELF FACTOR ASSEMBLAGE	REGRESSION COEFFICIENTS (RC)	CUMULATIVE VARIANCE (CV)	MULTIPLE CORR. COEFFICIENT (MCC)	STANDARD ERROR REGRESSION COEFF. (SERC)	STANDARD ERROR ESTIMATE (SEE)	t-value
F2	-88.565	.149	.386	13.14	63.2	-6.7
F6	-99.592	.298	.546	17.51	57.7	-5.6
F5	59.089	.341	.579	16.26	56.1	3.6
F3	46.384	.380	.609	15.38	54.7	3.0
F7	-35.142	.393	.618	17.55	54.3	-2.0
F8	-34.336	.404	.636	17.69	54.6	-1.9

Intercept for multiple regression equation = 160.56918

Critical values for t = 2.576 at 0.01  
= 3.291 at 0.001 levels

SLOPE FACTOR ASSEMBLAGE	RC	CV	MCC	SERC	SEE	t-value
F1	902.93	.484	.696	244.1	517.0	3.6
F2	-934.16	.569	.747	333.5	478.5	-2.8

Intercept for multiple regression equation = 1584.79816

Critical values for t = 2.704 at 0.01  
= 3.551 at 0.001 levels

quite large (54.15) and is 1/6th of the total depth range. Errors seem to increase with depth (Fig. 6.11). The small cumulative variance and large standard errors imply other environmental variables as being of greater influence upon foraminiferal distributions.

Values expressed in Table 6.3 are shown visually in Figs. 6.11 a-i which plot observations for depth against estimates produced by the regression. Fig. 6.11a shows large scatter within and outside the marked confidance bands and reflects the rather poor predictive ability of the depth equation for the combined shelf assemblages listed in Table 6.3. Figure 6.11 b-i show the distribution of this scatter within each factor assemblage, i.e. for the most statistically significant indicators of depth. In this way, the good predictive ability of factor 2 is seen (Fig. 6.11b) and the roughly successively decreasing capabilities of the remaining factors are shown (Fig. 6.11f-i) as tendancies for the data points to fall on a horizontal line (no correlation with depth). Factors 7 and 4 however show some correlation with depth perhaps reflecting their correlation with factor 2 (see table 6.1).

Ecological implications will be discussed in a later section. Suffice it to say the above observations are consistent with those of an earlier section that directly compared the importance of an assemblage and its environmental characters.

Regression of the 4 slope factors with respect to depth explained greater amounts of the depth variance than did shelf factors (Table 6.3). The  $F$  and  $t$  values show high significance of the regression and regression coefficients at the 0.001 level for factor 1 (Uvigerina peregrina). Factor 2 (Bulimina exilis) is significant only at the 0.01 level. The overall relative predictive abilities of the slope stations

Figures 6.11--6.19. Environmental parameter observations versus regression estimates for each assemblage.

Figure 6.11 Regression estimates Vs. observations of depth from shelf stations.

Figure 6.12 Regression estimates Vs. observations of depth from slope stations.

Figure 6.14 Regression estimates Vs. observations of temperature from slope stations.

Figure 6.13 Regression estimates Vs. observations of temperature from shelf stations.

Figure 6.15 Regression estimates Vs. observations of salinity from shelf stations.

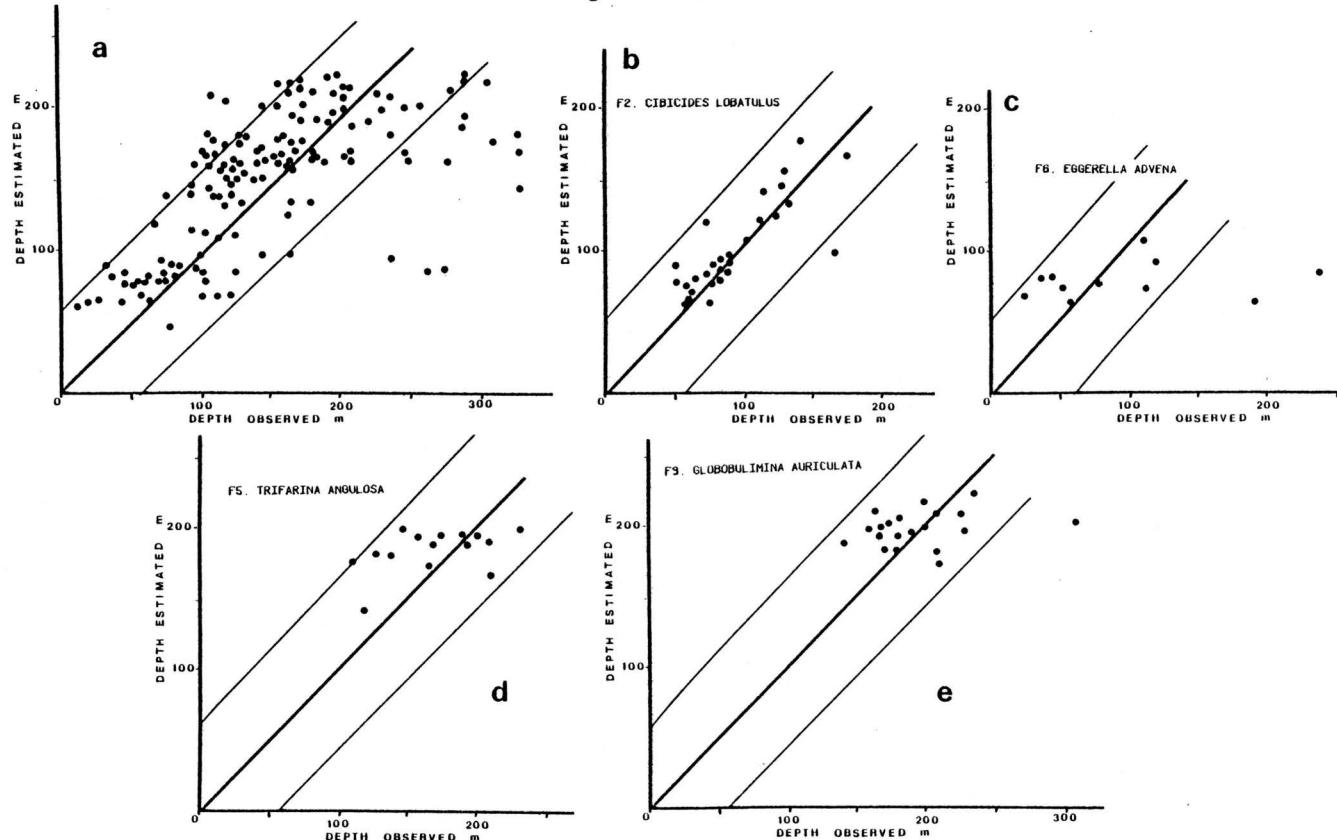
Figure 6.16 Regression estimates Vs. observations of salinity from slope stations.

Figure 6.17 Regression estimates Vs. observations of % gravel from shelf stations.

Figure 6.18 Regression estimates Vs. observations of % sand from shelf stations.

Figure 6.19 Regression estimates Vs. observations of % mud from shelf stations.

Figure 6.11



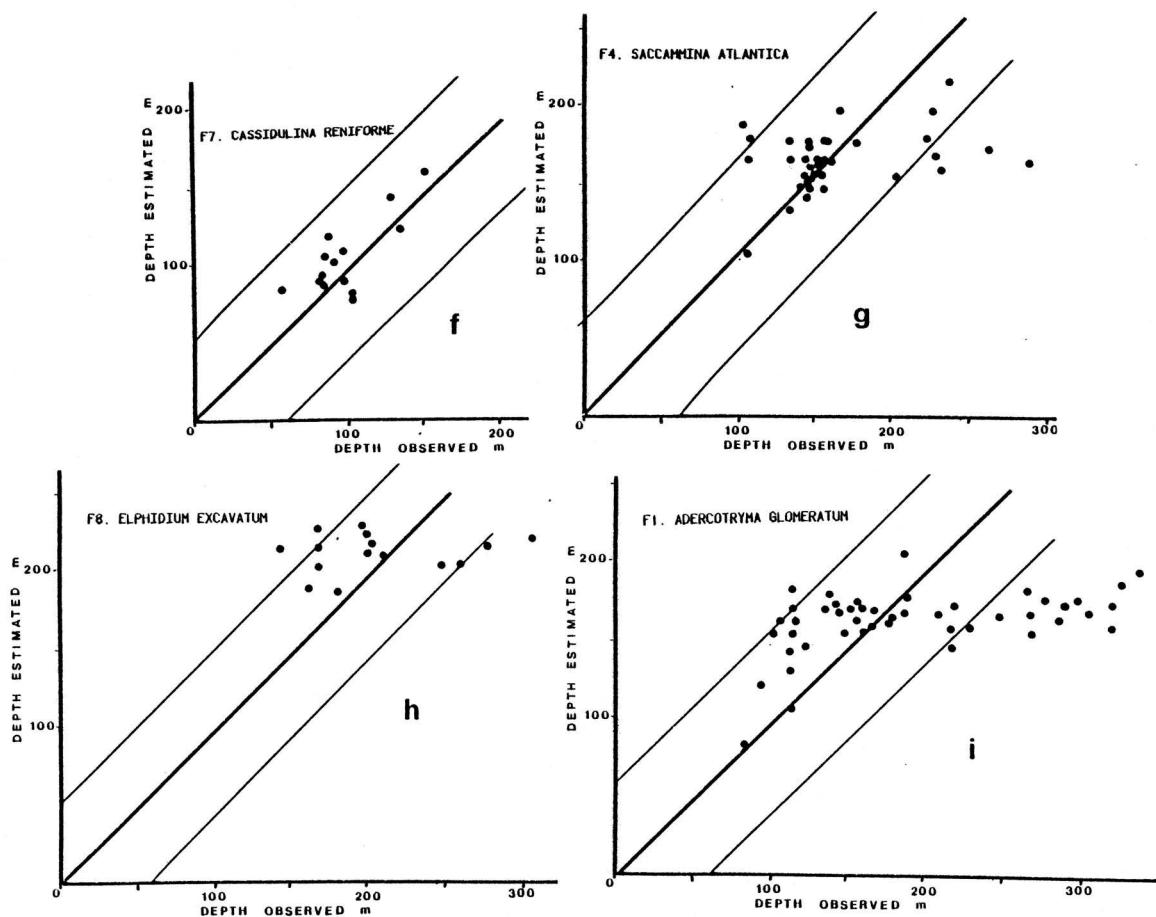
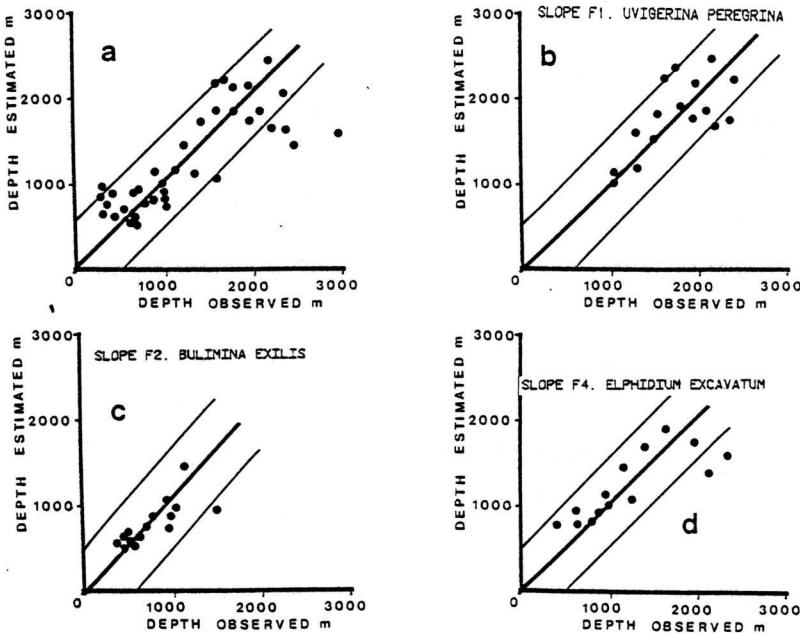


Figure 6.11

Figure 6.12



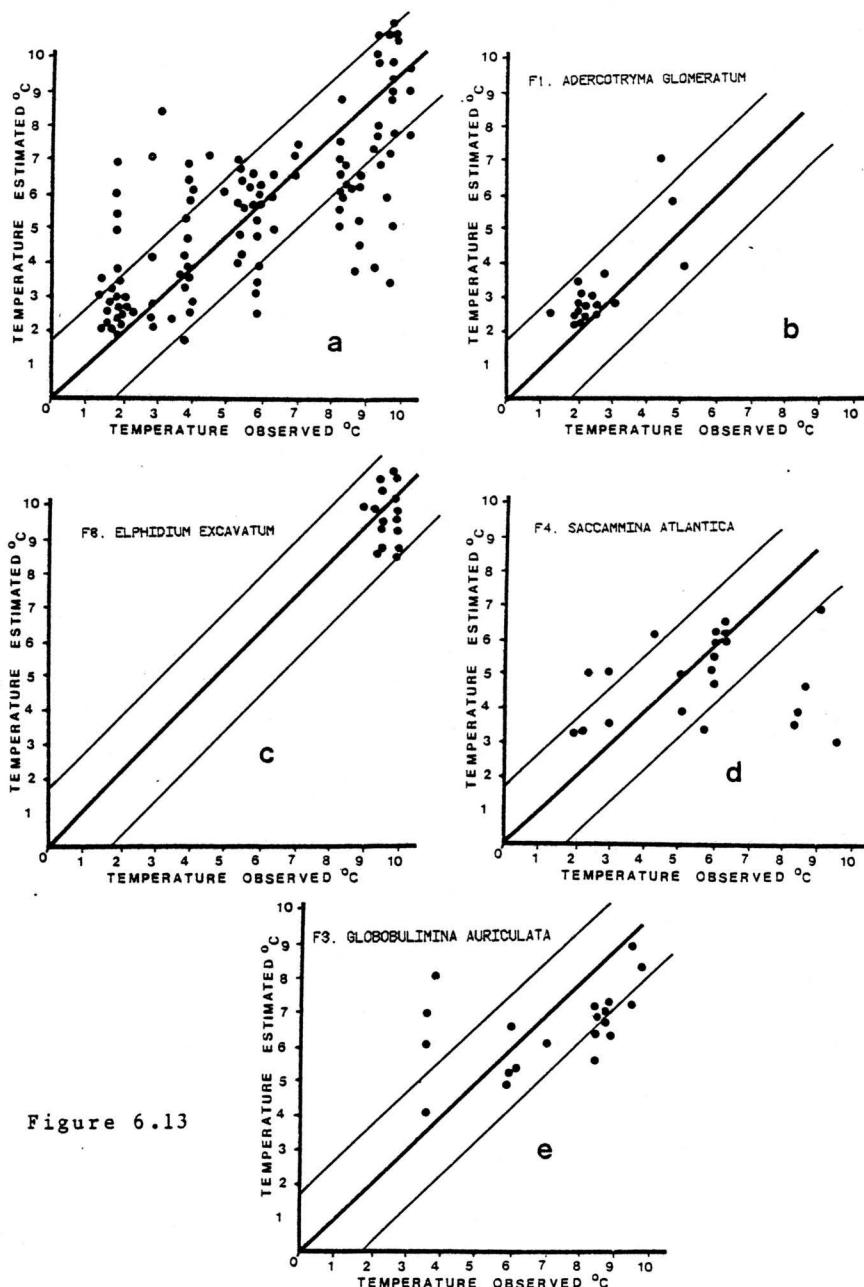


Figure 6.13

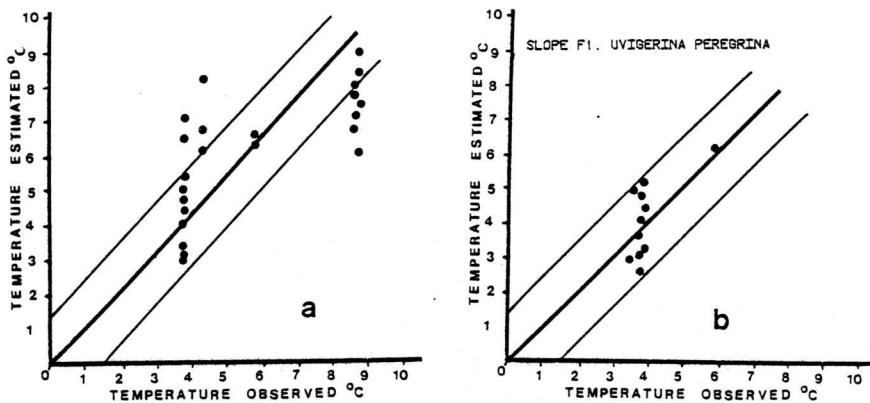


Figure 6.14

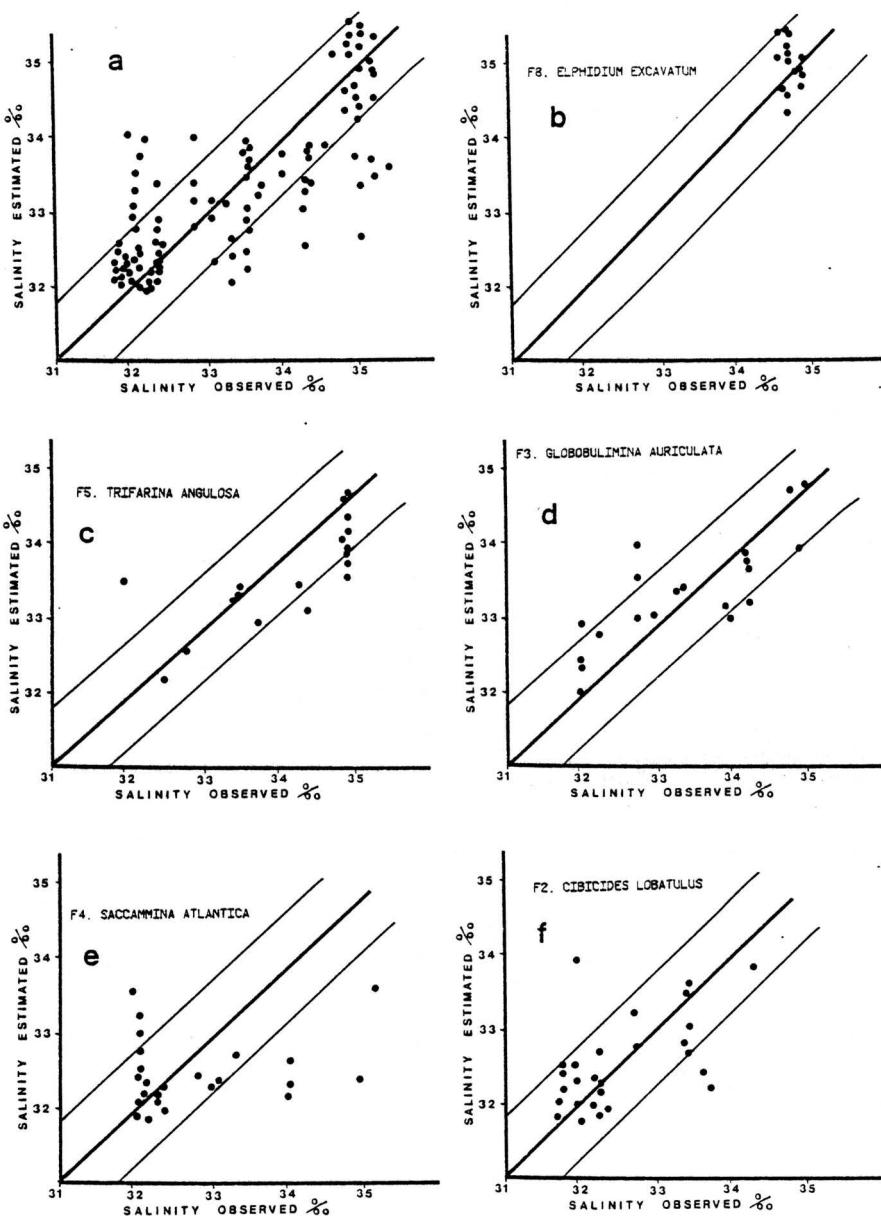


Figure 6.15

Figure 6.16

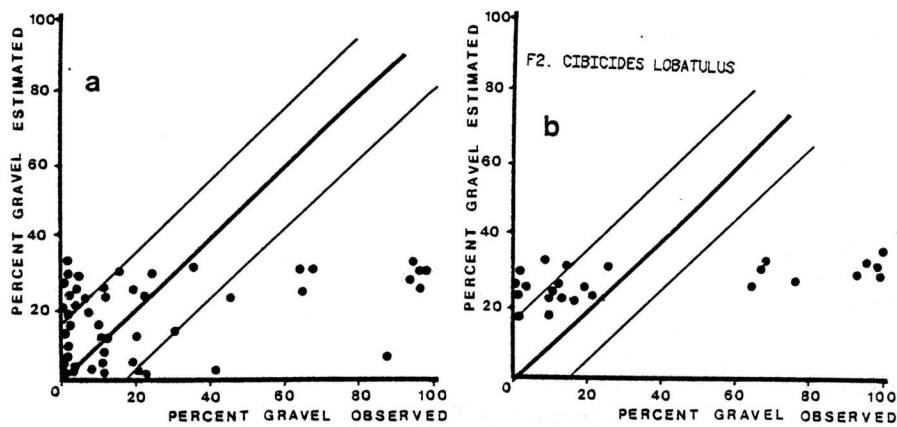
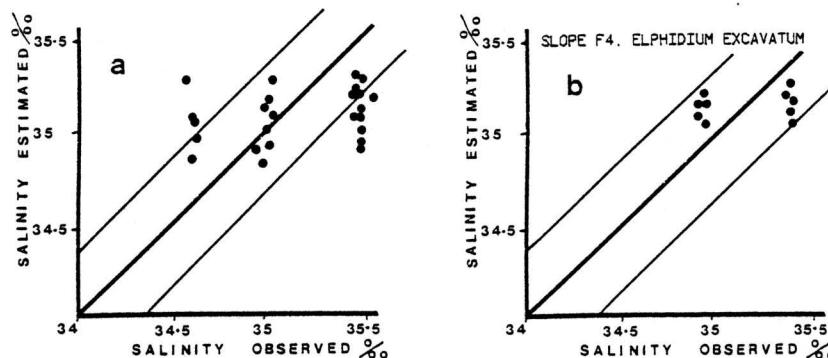
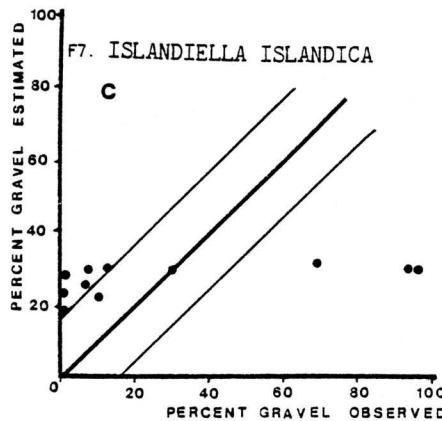


Figure 6.17



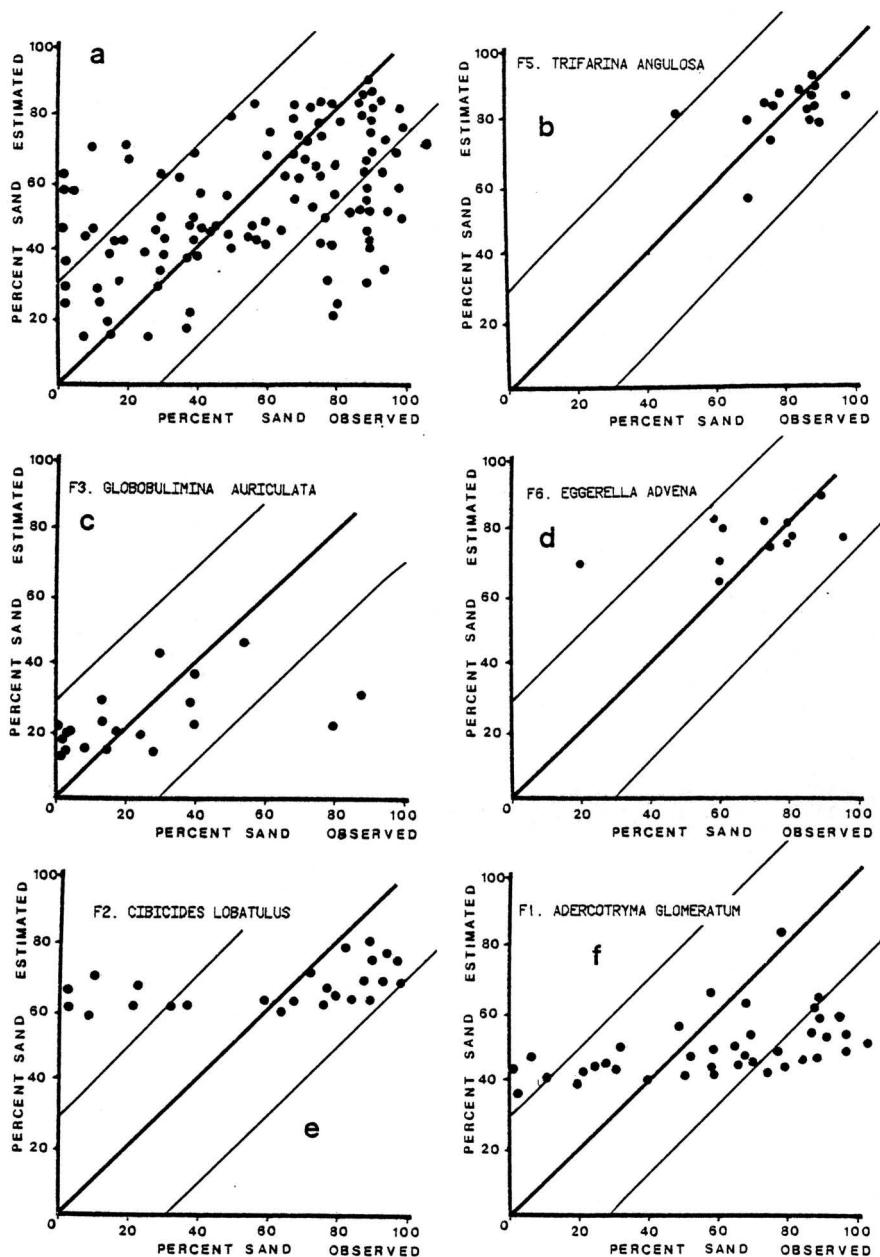


Figure 6.18

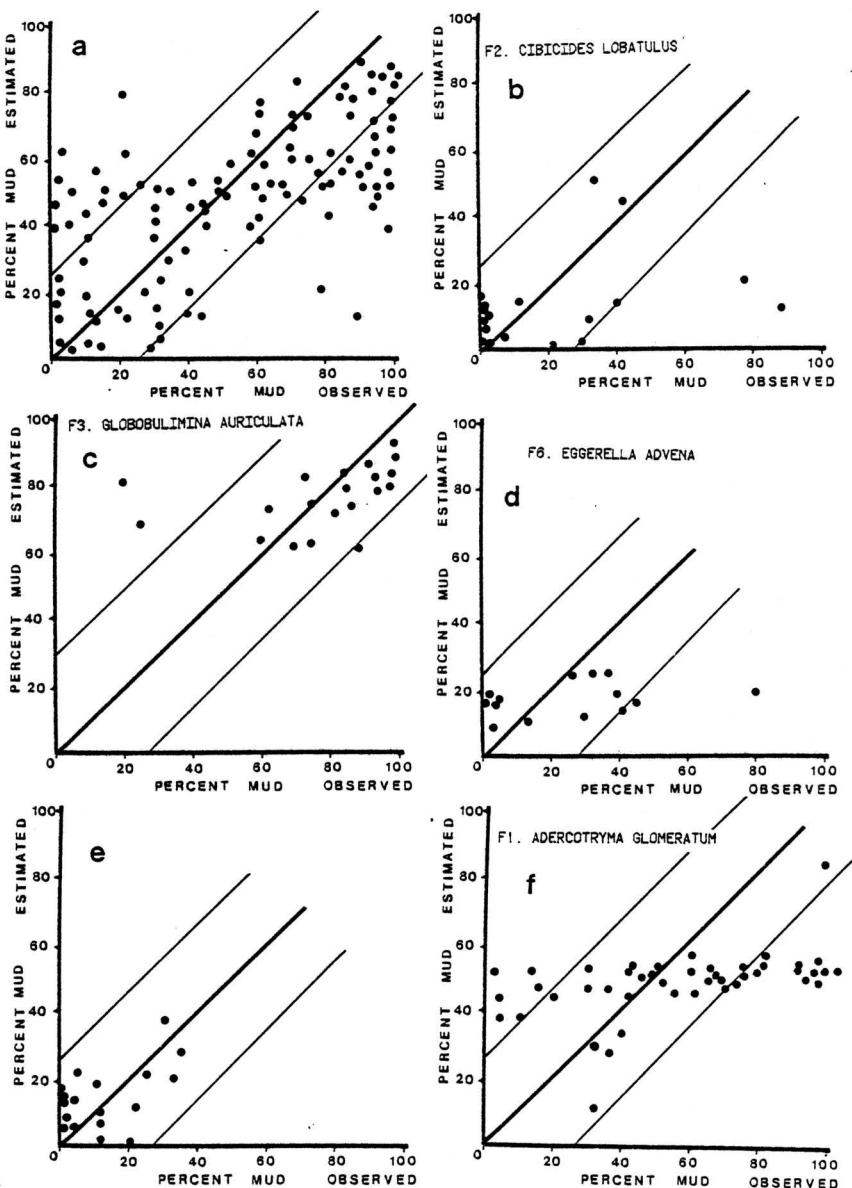


Figure 6.19

are shown in Figs. 6.12a,b,c & d.

b/ Temperature.

Regression of shelf factors against bottom temperatures explains a greater proportion of the variance and has a larger multiple correlation coefficient than the depth regressions (Table 6.4). The SEE for the multiple regression is 1.80C. Several of the regression coefficients are significant at the 0.001 level. Shelf assemblages 1 (Adercotryma glomerata) and 8 (Elphidium excavatum f. clavata) show the most significant relationships to temperature. These relative predictive abilities are illustrated in Figs. 6.13a-e. Again, large amounts of scatter are seen within the confidence bands (6.13a) for the multiple regression equation. Figs. 6.13b,c for factors 1 and 8 display less scatter and better predictive capability whereas Figs. 6.13d,e show the most scatter and a greater tendency to lie in a horizontal line. (ie. poor or non-existent predictive ability with respect to temperature).

Only slope factor 1 (Uvigerina peregrina) shows significant relationship to temperature (Table 6.4). This is seen in Fig. 6.14b where it exhibits less scatter when compared to 6.14a. Standard errors for shelf and slope estimates are similar at around 1.70C.

The observations made on regression of shelf and slope assemblages with respect to temperatures are consistent with earlier more subjective considerations.

c/ Salinity.

Regression results for shelf and slope assemblages with respect to

Table 6.4

## TEMPERATURE REGRESSION STATISTICS

SHELF FACTOR ASSEMBLAGE	RC	CV	MCC	SERC	SEE	t-value
F8	-4.9	.405	.636	.656		-7.4
F1	-4.1	.521	.720	.423	2.0	-9.7
F2	-3.4	.610	.779	.471	1.8	-7.4
F5	1.4	.621	.785	.577	1.8	2.5

Intercept for multiple regression equation = 6.40369

Critical values for t = 2.576 at 0.01  
= 3.291 at 0.001 levels

SLOPE FACTOR ASSEMBLAGE	RC	CV	MCC	SERC	SEE	t-value
F1	-3.8	.416	.638	0.684	1.83	-5.677
F4	-3.2	.519	.712	1.665	1.69	-3.065

Intercept for multiple regression equation = 7.085

Critical values for t = 2.704 at 0.01  
= 3.551 at 0.001 levels

Table 6.5

## SALINITY REGRESSION STATISTICS

SHELF FACTOR ASSEMBLAGE	RC	CV	MCC	SERC	SEE	t-value
F8	-2.6	.406	.637	.249	.873	-10.6
F5	1.6	.483	.693	.230	.818	7.1
F3	.8	.521	.719	.213	.719	3.7
F7	.9	.536	.728	.249	.782	3.6
F2	-.6	.559	.742	.182	.767	-3.3

Intercept for multiple regression equation = 32.39513

Critical values for t = 2.576 at 0.01  
= 3.291 at 0.001 levels

SLOPE FACTOR ASSEMBLAGE	RC	CV	MCC	SERC	SEE	t-value
F4	-.61	.169	.412	.211	.332	2.891

Intercept for multiple regression equation = 35.02067

Critical values for t = 2.704 at 0.01 level

salinity are shown in Table 6.5. Five shelf factors have significant regression coefficients (at 0.001 level). Two of these, factors 8 and 5 (E. excavatum and T. angulosa) show more significant relationships than the remainder. The multiple correlation coefficients are high. Similarly, approximately 60% of the salinity variance is explained in this regression. The relative predictive abilities of these relationships are seen in Figs. 6.15a-e. Fig. 6.15a shows the shelf assemblage predictive ability as a whole. Large amounts of scatter can be seen, especially when compared with Figs. 6.15b,c, & d, which show the relative increased capability of assemblages 8,3 and 5 (Elphidium excavatum, Globobulimina auriculata and Trifarina angulosa) to predict accurate salinity values ( ie. tighter clusters about the central line). Figs. 6.15e,f show the source of scatter i.e. predictive inability as seen in the tendency to form in a horizontal line.

Slope assemblage-salinity regression results are shown in Table 6.5. Only slope factor assemblage 4 shows significant relationships (at 0.001 level) but with small correlation coefficients and cumulative variance. General predictive ability and scatters for slope assemblages are shown in Fig. 6.16a & b.

d/ Substrate characters: % sand, gravel and mud.

Table 6.6 summarizes the regression analysis with respect to gravel %. Only 2 assemblages show significant relationships (at 0.001 level). The cumulative variance explained is small and the associated correlation coefficients, lower. Factor assemblage 2 appears most related to % gravel (Cibicides lobatus) with assemblage 7 (Islandiella

Table 6.6

GRAVEL REGRESSION STATISTICS

SHELF FACTOR ASSEMBLAGE	RC	CV	MCC	SERC	SEE	t-value
F2	23.64	.214	.463	3.85	17.385	6.1
F7	18.47	.252	.499	5.53	17.036	3.3

Intercept for multiple regression equation = 1.04

Critical values for t = 2.576 at 0.01  
= 3.291 at 0.001 levels

Table 6.7

SAND REGRESSION STATISTICS

SHELF FACTOR ASSEMBLAGE	RC	CV	MCC	SERC	SEE	t-value
F3	-32.1	.189	.387	7.74	30.67	-4.1
F5	40.3	.234	.480	8.12	29.2	-4.9
F6	40.7	.286	.528	9.06	28.3	4.4
F2	16.8	.307	.546	6.46	28.0	2.6

Intercept for multiple regression equation = 45.08144

Critical values for t = 2.576 at 0.01  
= 3.291 at 0.001 levels

Table 6.8

MUD REGRESSION STATISTICS

SHELF FACTOR ASSEMBLAGE	RC	CV	MCC	SERC	SEE	t-value
F2	-39.5	.246	.496	6.4	31.5	-6.0
F3	37.8	.364	.601	7.3	29.0	5.1
F6	-42.3	.412	.638	8.6	28.0	-4.9
F5	-32.0	.454	.668	7.7	27.2	-4.1
F7	-27.5	.479	.685	8.6	26.7	-3.1

Intercept for multiple regression equation = 52.43668

Critical values for t = 2.576 at 0.01  
= 3.291 at 0.001 levels

Table 6.9  
A summary of shelf cumulative variance and multiple correlation coefficients for each equation.

	Cumulative variance	Multiple correlation coefficient
DEPTH	.404	.636
TEMPERATURE	.621	.785
SALINITY	.559	.742
% GRAVEL	.252	.499
% SAND	.307	.546
% MUD	.479	.685

Table 6.10  
A summary of slope cumulative variance and multiple correlation coefficients for each equation.

	Cumulative variance	Multiple correlation coefficient
DEPTH	.569	.747
TEMPERATURE	.519	.712
SALINITY	.169	.413

islandica) of lesser significance. Fig. 6.17a shows a large amount of horizontal scatter and reflects the inability of the majority of the shelf assemblages to predict accurate values of gravel %. Figs. 6.17b & c show lesser scatter and increased predictive ability.

Relationships to % sand are summarized in Table 6.7. Factors 3,5 and 6 (Globobulimina auricalata, Trifarina angulosa and Eggerella advena assemblages, respectively) are significantly related at the 0.001 level and factor 2 (C. lobatulus) at the 0.01 level. As for the gravel analyses, only low values are seen for the cumulative variance. Fig. 6.18a shows generally large amounts of horizontal scatter and, overall, a poor predictive ability of shelf stations with respect to % sand. Figs. 6.18b,c,d show less scatter and increased predictive capability; Fig. 6.18e,f, on the other hand, display relatively large amounts of scatter and reflect lesser significant relationships.

Four shelf assemblages display significant relationships (0.001 level) with % mud (Table 6.8) with reasonably high cumulative variance and multiple correlation coefficients. Figs. 6.19b & c show relatively good predictive ability compared to 6.19g & h and 6.19a. These reflect the overall regression statistics in Table 6.9.

Factor assemblages 2 and 3 appear to be most significantly related to % mud. Inverse relationship of factor 2 is to be expected considering its occurrence and previously described relationship to more sandy substrates. Factor assemblage 3 (G. auriculata), a central basin assemblage, is positively correlated to % mud and this may be a response to the higher organic carbon content of the muds. This particular aspect will be discussed in a later section.

## 6.6. Discussion of the Regression Results

The previous section determined the relative usefulness of each assemblage in predicting values of a particular environmental variable. The success of each equation in accounting for the observed original information is summarized in Tables 6.9 and 6.10. They show the final cumulative variance and multiple correlation coefficients for each equation. It is clear that around 75% of the original temperature and salinity information was explained by the model: over 60% of the depth and mud as well as approximately 50% of the gravel and sand information. Perfect linear relationships (one to one correspondence) between observed and predicted values would be manifest in 100% explanation. Random errors, both in the environmental observations and the assemblages, account for part of the "non-explanation" i.e. non-linearity, but not for all. Imbrie and Kipp (1971) showed that their estimates were slightly more precise with a curvilinear solution. The fact that this solution did not substantially improve their estimates, together with the nature of the environmental data available and the somewhat obscure ecological meaning of the curvilinear model, as mentioned earlier warranted the use of the more robust linear solution in this study.

The deviations from linearity are best explained with reference to other environmental variables, both measured and unmeasured. Many works (Douglas, 1978; Lohmann, 1978) have shown the subtle nature of intercorrelations between environmental variables and assemblages, i.e. an assemblage generally occurs within a certain environmental variable's limits but may sometimes be found outside of those limits if

some other aspect of the environment locally overrides the importance of this particular variable (therefore reducing the linearity of the relationship).

Substrate characters consistently have greater amounts of unexplained information (or non-linearity), substantiating the well documented lack of dominant substrate-assemblage relationships. In fact, in this study, only one assemblage (Cibicides lobatulus) is controlled in its occurrence by substrate. Temperature, salinity, depth and to a certain extent percent mud, appear to be more important influences on assemblage distribution (possibly because of increased amounts of organic carbon in mud, implying abundant food and nutrients).

With this in mind, it is possible to discuss each assemblage in terms of the main environmental controls influencing its distribution; and any other feature of the environment that may affect this distribution to a greater or lesser degree. If these relationships are to be valid ecologically and have palaeoecological implications, then some concordancy with previous observations and studies is to be expected. The following chapter will investigate these aspects.

## CHAPTER 7: ASSEMBLAGE DISTRIBUTION CONTROLS : A SYNTHESIS

## 7.1 Introduction

The spatial occurrence and composition of shelf and slope foraminiferal assemblages on the margin off Nova Scotia were described in Chapter 3. It is the purpose of this chapter to discuss these distribution patterns in terms of their limiting controls, the significance of which were investigated in the previous chapter. Reference to figure 3.9 (summary map of shelf distribution) and figs. 3.16-3.23 (slope distribution) will aid in the following discussion.

## 7.2 Shelf Factor assemblages.

Assemblage 1: Adercotryma glomerata.

The distribution of this widespread assemblage, dominant in the northeast shelf areas of this study, is probably controlled by temperature. It shows no significant relationships with any other of the measured parameters. The temperature dependance of this assemblage is clear also from its distribution on banks, basins and nearshore areas that vary in depth and substrate. Temperature limits for this assemblage are from, 1.8 °C to 3 °C. Salinity varies from 31.3 ‰ to 33 ‰. The main species from this assemblage have previously been found on the Labrador shelf and in the Canadian Arctic (Vilks, 1969, 1980, 1982), in waters with temperatures of 2-4 °C and salinities of 31-34 ‰. Leslie (1965) reported their occurrence in the Hudson Bay, in waters of similar temperature but with lower salinities. Similarly, Schafer and Cole (1978) described their occurrence in Chaleur Bay in estuarine conditions of low salinity (28-31‰) and temperature (2-4°C). The wide salinity

tolerance of species in this assemblage is further shown by their occurrence under normal marine waters (34-35‰) of the Newfoundland slope (Schafer and Cole, 1982) and on the Grand Banks; (34‰, 30°C; Sen Gupta, 1971). Both Osterman and Kellogg (1979) and Milam and Anderson (1981) have documented Antarctic occurrences in very cold waters (0-1 °C) and variable salinity. Rodrigues and Hooper (1982) recently recorded the dominance of this type of assemblage under waters of similar temperatures and salinities (in fact under the same watermass type) just to the north east of the present study area in the Bay of St. Lawrence.

The occurrence of this assemblage as documented in the literature seems to comply with the data presented in this study and indicates that temperature is the primary control, salinity being relatively unimportant. Palaeoecological implications of its occurrence at depth in piston cores from this area (Scott, 1982; pers. comm.) and the Newfoundland shelf (Mudie and Guilbault, 1982) will be discussed later.

#### Assemblage 2: Cibicides lobatulus.

This inner shelf (nearshore) and bank assemblage is one of 2 assemblages in this study whose controlling influence appears to be predominantly substrate character. Its most significant, direct relationship is with percent gravel and sand (therefore inversely related to percent mud). Thus it also shows a good relationship to water depth, which controls sediment texture; it occurs on sandy gravel banks where the water is shallow as opposed to deep mud filled basins which offer no attachment surfaces. Some relation to temperature is observed in this study (ranges from 4-8.5°C) and is presumably due to a high

tolerance of temperature range. The salinity relationships seen may be a response to the fact that substrate (preferred) availability is essentially restricted to nearshore waters and corresponding low salinities.

Cibicides lobatulus has a world wide distribution and its occurrence has previously been related to substrate (Phleger, 1952; Parker, 1952; Sen Gupta, 1971). It has been described on the eastern U.S. shelf and off Maine (Cushman, 1931); the Georgia continental shelf (Sen Gupta and Kilbourne, 1976); and from the North Carolina continental shelf (Schnitker, 1971). On the eastern Atlantic continental shelf, Murray (1973) has recorded it in the Celtic Sea; Haman (1966) from Tremadoc Bay, North Wales, and Serjup (1981) from the Norwegian continental shelf. Other occurrences include Halifax Harbour (Gregory, 1971); Long Island (Akpati, 1975); St. Margarets and Mahone Bays, (Bartlett, 1968) and in the Ross Sea, Antarctica (Osterman and Kellogg, 1979). In the Bay of St. Lawrence to the north of this area this assemblage is similarly associated with substrate type and large annual range in temperature and salinity (Rodrigues and Hooper, 1982).

Its occurrence in the North Atlantic from the Georgia shelf to Arctic waters implies, as observed in this study, that there is no significant relationship between its distribution and temperature. In the references mentioned above it is reported for salinities ranging from normal marine to estuarine which also indicates that salinity is not a major controlling influence.

This assemblage then, exhibits strong relationships to the nature of the substrate, occurring upon gravel or coarse sandy bottoms. This appears to pertain to its distribution on a global scale.

Assemblage 3: Globobulimina auriculata.

This central shelf basinal assemblage occurs throughout most of Emerald, LaHave and Roseway basins. It is more diverse when compared to the other assemblages described above. Its most significant relationships are with salinity, percent mud and depth. No direct association with temperature is seen in the study area. The assemblage is usually found in sediments of 70-100% mud composition, under waters of 5-12 °C and 34.5 ‰.

The main species in this assemblage (Globobulimina auriculata, Bolivina subaenariensis and Nonionellina labradorica) have been reported widely in the literature. They occur in the Bay of St-Lawrence (Hooper, 1975), on the Grand Banks of Newfoundland (Sen Gupta, 1971), in Long Island Sound (Parker, 1952; Murray, 1969), on the shelf off Massachusetts (Kaefeleisciou, 1975), on the North Carolina continental shelf (Schnitker 1971), on the Florida-Hatteras slope (Sen Gupta and Strickart, 1982), and on the Norwegian shelf (Sejrup, 1981). These occurrences, together with those documented by Culver and Buzas (1981) cover a wide latitudinal range and therefore imply a general independence of temperature control. Similarly, all these works document occurrences in essentially normal marine waters (34-35‰). Both these tendencies are apparent on the Scotian shelf. It appears that this central basin assemblage requires normal marine salinities (34-35‰) which may partly explain why it does not occur in the basins to the northeast of the study area which have salinities of about 31-33‰. The higher salinities in the central basin area of the shelf are directly due to the influence of the slope water (Warner, 1970) and in this way, shelf assemblage 3 reflects the areal extent of this slope water incursion on to the continental shelf. The significant relationship between this assemblage and mud

content implies other controlling influences acting in conjunction with salinity. This is probably a function of increased amounts of organic carbon in these sediments (Vilks, 1976) and other nutrients (Fournier *et al.*, 1978) which are advected, mixed and moved onto the shelf in this area. Sen Gupta and Strickart (1982) and Sen Gupta *et al.* (1981) have previously documented large amounts of Bolivina subaenariensis co-occurring with high mud and organic carbon content samples in nutrient enriched waters which result from Gulf Stream upwelling. Similarly, Miller and Lohmann (1982) and Mullineaux and Lohmann (1981) find Globobulimina spp. associated with sapropelic sediments. The high species diversity in the present study area would also be explained by the higher productivity in the surface waters. Previous work then, support the present contention of a salinity/nutrient control of shelf assemblage 3.

#### Assemblage 4: Saccammina atlantica.

This assemblage, which is composed entirely of agglutinating forms is found peripheral to the northern boundary of assemblage 3 in the central basins. It is remarkable for its apparent lack of significant controlling influences in the study area (or at least of those investigated in this study). Consistent with these observations is the fact that the main species: Saccammina atlantica, Reophax scorpiurus and Cribrostomoides jeffreysi have been reported from a wide range of salinities, temperatures, depths and substrates including the Georgia shelf (Sen Gupta and Kilbourne, 1976); Rhode Island Sound (Belanger, 1976); Buzzards Bay, Vineyards Sound (Murray, 1968,1969); Chaleur Bay (Schafer and Cole, 1978); Florida Bay (Bock, 1971); the mid- atlantic bight (Ellison, 1977); Labrador slope (Schafer and Cole, 1982); Lake

Melville and Labrador shelf (Vilks *et al.*, 1982). Clearly, these occurrences ranging from sub-tropical to arctic or brackish waters to fully marine, require other environmental explanations than the ones investigated hitherto.

Assemblage 4 lies peripheral to assemblage 3 and from the factor score matrix (appendix II), it appears that if the dominant components of assemblage 3 are removed (i.e. G. auriculata, N. labradorica, B. aculeata, B. marginata and I. teretis), the dominant species as observed in assemblage 4 would now be dominant. Thus the selective removal of these components (of assemblage 3) either by transportation or dissolution could produce the observed assemblage 4. Dissolution on the shelf will be discussed in a later section. Transportation is a stronger option with complete removal of the calcareous species from these peripheral areas and redeposition in deeper parts of the basin being possible. This kind of effect is well documented by Murray (1969) and Haynes and Whatley (pers. comm. 1981) as occurring on a large scale in the Celtic and Irish Seas. The idea is supported by the fact that the remaining assemblage has an underlying presence (ie. a background presence in the surface samples) in the deep basinal assemblage 3.

If transportation did produce the observed patterns however, one would expect at least some of the calcareous species to remain in the depleted assemblage. In particular, those species which, to be removed, would require different traction velocities (i.e. transportability in assemblage 3 differs from species to species; compare, for example, the species G. auriculata or N. labradorica to B. subaenariensis and I. teretis; these two groups are clearly different in terms of their operational sphericity which, according to Kontrovitz 1977, defines their susceptibility to transportation). Similarly, one would expect

some presence of shallower nearshore species (e.g. C. lobatulus) in both assemblage 3 an 4, if transportation was responsible. This is not observed.

The most probable explanation for the distribution of this assemblage is that it occurs in an area of intense temperature and salinity gradients. Shown by MCLellen (1957), these gradients are present to a greater or lesser degree year round (Warner, 1970). They are a result of the slope derived central basin waters being in juxtaposition with the Labrador surface and St. Lawrence waters. This results in temperature gradients ranging from 10°C in the basins to 1.8°C in the northern periphery, with salinity varying from 35‰ to 32‰. These gradients presumably preclude the basin assemblage 3 on the basis of adverse salinity conditions (see previous pages) and assemblage 1, because of adverse elevated temperatures.

The ecological niches made available by the non-occurrence of these assemblage 1 and 3 are filled by assemblage 4, the components of which are, judging by the cited literature, rather opportunistic in nature. Further work is needed to determine whether this assemblage can be used in palaeoecological studies to trace the occurrence and spatial fluctuation of this watermass boundary in Pleistocene to Recent deposits.

#### Assemblage 5: Trifarina angulosa.

This assemblage is persistently present in the shelf break region. Temperature ranges from 8.5-10°C, with salinity constant at 35‰. Significant relationships are evident between occurrence, salinity and depth; there is a lesser correlation with % sand and inverse relationships with % mud. No apparent dependence on temperature is seen

in this study. Again, these observations are consistent with previous documentation. The main components of this assemblage have been described from the warmer waters off the southern United States (Schnitker, 1971; Ellison, 1977) to cold arctic waters (Vilks *et al.*, 1982), antarctic waters (Osterman and Kellogg, 1979; Milam and Anderson, 1982) and intermediate areas (Murray, 1969; Schafer and Cole, 1978). Consequently, no temperature dependence is implied or has been observed in the present study. All these previous occurrences, however, are associated with normal open marine conditions of salinity. Similarly, many are associated with shelf edge locations or areas of upwelling. The main controls influencing the distribution of this assemblage appear to be salinity in concert with high nutrient content and productivity (produced through upwelling along the shelf edge).

**Assemblage 6: Eggerella advena.**

Restricted to Chedabucto, Gaberous, Mahone and St. Margarets bays and to shallow banks off Sable Island, this assemblage shows no significant correlations with temperature or salinity. Its main correlations are with depth, percent sand and, inversely, with percent mud. The observations made in this study are consistent with those reported in the literature. Schafer and Cole (1978) have described similar assemblages from estuarine deposits noting optimum substrate conditions as being well sorted fine sands (i.e. as in this study). Similarly, Bartlett (1964) noticed a depth dependency in St. Margarets and Mahone Bay, and a similar association is seen in the present study. Several works have described the presence of the main species of this assemblage in warmer, more southern waters (Schnitker 1971, North Carolina shelf; Ellison 1977, Mid-Atlantic Bight; Parker 1952, Long

Island Sound; Murray 1969, Cape Hatteras; Sen Gupta and Kilbourne 1976, Georgia shelf). The main species have been recorded from fully marine to brackish waters (Vilks 1982). The temperature and salinity independence recorded in the literature is thus consistent with present observations.

Assemblage 7: *Isandiella islandica*.

This assemblage is completely dominated by *Isandiella islandica* and is prevalent on isolated outer banks and shallow depressions on the shelf. It appears to replace factor 2 *Cibicides lobatulus* in these areas and is not an important faunal component on the shelf. Its most significant relationship (0.001) is to percent gravel and in this respect, is similar to *C. lobatulus*. The live data in this study imply an "in situ" assemblage on a coarse sand/gravel substrate, and is further implied by the lack of occurrence of elements of this assemblage in the central shelf basins of this area. This assemblage has been described on the sandy banks to the north of this area and under similar temperature and salinity conditions (Sen Gupta, 1969).

One question concerning its distribution remains. Why is it that this assemblage is not associated in some parts with *C. lobatulus*? (Raw data or factor score matrix). It is predominantly monospecific in its occurrence. One would expect some amount of *C. lobatulus* in light of their similar substrate requirements. Competition edge seems unlikely as this species does not have a widespread occurrence on the shelf. On the other hand, *Isandiella islandica* does form a small portion of the factor 2 *Cibicides lobatulus* assemblage. The most likely explanation for this is local transportation of the tests within the bank environment, concentrating large amounts of *I. islandica* in small areas. This local transportation effect would explain the observation of live

I. islandica in parts of the assemblage. That the transportation is local is indicated by the non-occurrence of the species in the basins. The available data thus implies that factor assemblage 7 is a local variation of the Cibicides lobatulus assemblage produced through local transportation.

Assemblage 8: Elphidium excavatum f. clavata.

In view of the fact that this assemblage was earlier described as relict and transported, the regression results, showing a relationship to temperature and salinity, appear anomalous. The ragged, worn and size sorted nature of the tests is a testament to their relictness as is the complete lack of live representatives. Previous reports of its occurrence show that it is predominantly a cold water, restricted salinity species (Vilks 1979, 1982). More recent work, however, documents a broader range of salinity (Miller *et al* 1982, Mudie *et al* in press). Further discussion on the distribution of the relict components of this assemblage on the shelf and slope will follow later. It remains to explain here the apparent correlation to temperature and salinity values.

The observed response may be to the sub-dominant but "in situ" live occurrence of Furstenkoina fusiformis, a form previously described from waters of similar temperatures and salinities, 6-8 °C, 34 ‰ (Schnitker 1971; Ellison 1977). The restricted occurrence of the whole assemblage to a small area on the shelf overlain by waters of consistent temperature and salinity values may also explain in part the observed relationships. It is most likely a combination of these two effects that is responsible for the observed correlation.

### 7.3 Slope Factor assemblages.

#### Assemblage 1: Uvigerina peregrina.

This slope assemblage consistently peaks at 1600 m. within the study area. Its total depth range is 800 m. to 2500 m. Regression analysis reveal significant relationships to temperature, depth but not with salinity. The temperature optimum for this assemblage is 40C. This is consistent with optimum values of 3-40C for the occurrence of this species throughout the North Atlantic (Streeter 1973). Uvigerina peregrina has been reported from many depths and is marked by its apparent inconsistency in this respect. Miller and Lohmann (1982) for example, observe its occurrence off the U.S. continental margin at 1000-2000 m.; Cushman (1913-31) observes its presence at 200-4000 m.; Phleger (1942) and Phleger and Parker (1951, off the Gulf of Mexico) at 300-4000 m.; Pfium and Frerichs (1976) at 300-1500 m. and Lohmann (1978) in the South Atlantic at 2000-4000 m.

This depth inconsistency and the fact that the species has been associated with colder (less than 30C) deoxygenated waters (Streeter and Shackleton, 1979) led Miller and Lohmann (1982) to conclude that temperature was not the most important limiting control on the species distribution. This assumption is, at the outset, unwarranted. The fact is that Uvigerina peregrina has rarely been associated with waters warmer than 40C. Thus, it appears that the species has an upper temperature limit below which it may grow and reproduce. That the species is related to, and to some extent, controlled by organic carbon content (Miller and Lohmann, 1982) and low oxygen values (Streeter and Lavery, 1982) is not in dispute here; this association being described many times. Although the present study has no organic carbon/oxygen

content data available, previous work has shown that below 1300 m. there is an oxygen minima (Gatien, 1970) and below 600 m., the proportion of silt and clay in the sediments increases (Hill, 1981). This implies, in a qualitative way, some correlation with low oxygen, and high carbon content. However, if these were the primary limiting controls of the distribution of this species, one would expect to find its occurrence in large numbers in the organic carbon-rich (Vilks and Rashid, 1976), low oxygen sediments and waters prevalent in Emerald and LaHave Basins. That this is not observed may be due to a primary temperature control i.e. basin waters at 7-10°C are too warm for *U. peregrina*.

Further work documenting the occurrence of this species and its covarying environmental parameters is required before significant conclusions can be reached concerning its palaeoenvironment.

Implications to palaeoceanographic studies will be discussed in the next chapter.

#### Assemblage 2: *Bulimina exilis*.

This diverse calcareous assemblage is restricted to upper slope depths (400-800 m.). Regression analysis revealed significant relationships only to depth. It occurs under normal marine salinities in warm slope waters (35‰; 6-9°C). The main components of this assemblage have been previously documented from areas to the South: N-E continental shelf of the U.S. (Phleger, 1942); the Massachusetts shelf (Parker, 1948); Gulf of Maine (Faas, 1971); and from the areas North of the present study area: the St. Lawrence (Vilks, 1968); the Tail of the Grand Banks (Sen Gupta, 1971); and in Chaleur Bay (Schafer and Cole, 1978); i.e. in waters colder and warmer than the present waters. All of its previously reported occurrences are from normal marine waters.

Bartlett (1964) described occurrences of the main species (Bulimina exilis) from the nearshore area within the present study limits.

Therefore, neither temperature or depth appear to be limiting parameters. In the present study, B. exilis is found also at shallow shelf depths as a subdominant component of the Central basinal assemblages.

Further data is required to evaluate the significance of this assemblage on the slope. However, its occurrence on the shelf in the warm nutrient-rich basins and along the Eastern shore (Bartlett, 1964) in an area of upwelling, together with its main occurrence on the upper slope under some highly productive surface waters, implies that nutrient availability in some part controls this assemblages occurrence. This would also explain the high values for diversity in this area.

#### Assemblage 3: Trifarina occidentalis.

A very restricted occurrence (about 3 samples) in several upper slope regions renders this assemblage difficult to interpret. It shows no relationship to any of the measured parameters. It seems to be similar to the B. exilis assemblage in composition but is less diverse. Until further slope data is available, no valid conclusions can be made concerning its role in the slope environment.

#### Assemblage 4: Elphidium excavatum f. clavata.

This assemblage has a large depth range on the slope and occurs in transects A and C. It appears to be the slope equivalent of shelf assemblage 8 and shows a significant relationship to salinity only. The shelf occurrence of the main species of this assemblage has been shown to be relict and in some areas, transported. This was based upon the

ragged and worn nature of the test and the lack of live representatives.

Diameter size data revealed the possibility of transportation from a source area in the Scotian Gulf to the southwest approaches of Emerald Basin. It is necessary now to discuss its overall distribution in terms of the causal factors.

It is unlikely that the slope assemblage E. excavatum is in place, in equilibrium with its present environment. Streeter and Lavery (1982) however, decided their E. excavatum fauna was in place on the basis of a recorded instance (Matoba, 1976) of slope living Elphidium batialis. Similarly, Schafer *et al* (1981) reported the presence of live E. excavatum off the Newfoundland slope but concluded that they were displaced into the slope environment where they continued to live as adults.

The main problem is to adequately explain its presently observed distribution. On the shelf, reworking and transportation have been invoked to explain its occurrence, the mechanism for this probably being the effect of several storm events. It is possible that these same storms put E. excavatum into suspension in the surface layers, with subsequent transportation seaward and slopeward by the net southward surface current. Settling through the water column would allow deposition on the slope adjacent to the source area in the Scotian Gulf. The fact that this assemblage is dominantly present in transect A, i.e. adjacent to the Scotian Gulf, would tend to support this, but does not explain its presence, although subdued, in transect C which is not adjacent to a shelf source. One would also expect significant amounts of the shelf edge fauna to be displaced, and this is not observed to such a degree.

A more likely possibility is the reworking and subsequent

transportation and deposition of relict Emerald silt-type deposits on the upper parts of the slope. The occurrence of such deposits has been documented by Stanley *et al* (1972) and has previously been described as a source for relict E. excavatum (Vilks, 1976). This process, together with cross-shelf transportation in the Scotian Gulf area, could explain the observed distribution. More data is clearly needed to evaluate the degree and extent of this reworking. This will be discussed in a later section.

Each assemblage, then, has been summarized in terms of its environmental controls as indicated through the regression analysis. The next chapter will discuss the relative importance of other aspects of the environment not considered quantitatively in the work so far, and then present a unified assemblage/environment model.

## CHAPTER 8: ASSEMBLAGE DISTRIBUTION CONTROLS: A DISCUSSION

## 8.1. Introduction

The work so far has determined the spatial occurrence of foraminiferal assemblages and related them statistically to various measured aspects of their environment. There are however, several other potentially important aspects of the marine environment which have not been measured in this study and have only been qualitatively described in previous sections. It is the purpose of this section to define and discuss the relative importance of these environmental parameters and arrive at a unifying model concerning assemblage-environment relationships within this study area. Subsequent to this, chapter 9 will consider implications of these results for Pleistocene to Recent studies on the Eastern Canadian continental margin.

## 8.2. Qualitative factors

An earlier section made reference to the community approach in palaeoecology and listed several features of an environment that control the taxonomic composition of an assemblage. Features not directly considered in this study include: space, light, trace elements, pressure, dissolved gases, nutrients, calcium carbonate availability and predation. There is thus a possibility that some or all of these may be more directly responsible for foraminiferal distribution within the study area than the measured depth, temperature, salinity and substrate characters. At the least, it is probable that the nutrient and other factors have a minor effect. Some of them have been considered

indirectly.

Light and pressure are a function of depth, so the effect of light is probably not that great in this area because of the occurrence of several assemblages in this study area in waters deeper than the photic zone. However some assemblages may in an indirect way be related partly to light, the Cibicides lobatulus assemblage for instance is significantly related to depth and substrate. It was observed earlier that the depth relationship was probably a response to the occurrence of suitable substrate at shallow depth only. However, part of this relationship may be due to the availability of light. Similarly, occurrences of certain types of algae and seaweeds are light-related and these also provide attachment surfaces for Cibicides lobatulus.

Hydrostatic pressure is not often considered as a directly limiting influence on foraminifera. It is known that at depth, pressure restricts the availability of CaCO<sub>3</sub> with resulting increase of agglutinating fauna. Belanger and Streeter (1979) considered the difficulties in evaluating the distribution effects of pressure on foraminiferal assemblages. They believed that temperature and pressure together regulate the occurrence of Planulina wuellersdorfi in the Atlantic. That these pressure-related effects should be studied more is indicated in Wirsén and Jannasch (1975) who record the effects of increased pressure on the metabolism of most micro-organisms. No pressure relationships are envisaged in the present study due to the observed independence of most assemblages to depth, etc.

Another aspect of the environment omitted in the regression analysis was dissolved gases, in particular the O<sub>2</sub> content of the overlying waters. Lohmann (1979) and Streeter and Shackleton (1978) pointed to the importance of O<sub>2</sub> (and temperature) as a control in deep sea

distribution; Belanger and Streeter (1980) however, believe that the relationship may be only indirectly linked. They refer to Hessler and Sanders (1969) observation that  $O_2$  has not been previously recorded as a limiting parameter in other deep sea biological studies. Douglas (1977) recorded a restriction of macro-benthos and removal of predatory pressure with low  $O_2$  values (anoxic conditions) and consequent benefits for foraminiferal abundances.

Much speculation exists on the position of foraminifera within the trophic structure of the marine benthos. Unfortunately however the speculation is for the most part not corroborated by scientific documentation of the extent and selectivity that foraminifera are preyed upon. Buzas (1979) alluded to the scant nature of predation studies concerning foraminifera and documented some of the more recent works (Lipps and Valentine, 1970; Mageau and Walker, 1976; Buzas, 1978). Controlled experimentation on the effects of predation in the Indian River Florida, (Buzas, 1979) determined that deposit feeders at least, do not selectively feed on foraminifera and that foraminifera are ingested because of their occurrence in the meiofaunal size range. The present study has no data pertinent to predation effects. Macro-benthic activity on the margin off Nova Scotia does not seem abnormal even in the two areas of oxygen minima (lower slope and central basins) Mills and Fournier, 1979 (ie. restriction of macro benthos and removal of predatory pressure as described in Douglas, 1978). The lack of foraminifera however on large parts of Emerald and Sable Island Banks, previously explained with reference to the nature and mobility of the substrate may in part at least be an effect of predation by the large stocks of fish associated with these spawning and feeding grounds. Daniels and Lipps (1978) however concluded that foraminifera are not a

significant food source for most fish in the Antarctic; that they might be important for smaller or juvenile fish is suggested in Buzas, 1979.

Nutrient availability has been subjectively discussed as probably influencing the composition and diversity of the shelf edge, upper slope and central basin faunae. That it is not a dominant factor is reflected by the high nutrient occurrences along the Eastern shore near Yarmouth, due to centrifugal upwelling (Garrett and Loucks, 1976) and circulation-induced upwelling along the Cape Breton shore, which are not characterised by major changes in the assemblages.

It is clear that most of the environmental parameters not considered in the regression analysis are unlikely to be the main causes for the observed foraminiferal distributions ( $\text{CaCO}_3$  availability will be discussed later). It is also known however that complex interaction exists between environmental parameters and biological populations therefore no one to one relationships are likely to exist or indeed should be sought. This statistical study has shown, however, the environmental parameters that show a strong correlation with species distributions and are therefore probably important controlling aspects of these distributions. It is the sensitivity of the assemblages to these that provide insight to palaeoenvironmental studies.

### 8.3. Shelf and Slope Assemblages: Their controlling environment

In the context of the correlations mentioned above, the regression analysis allows a summary of the margin assemblages in terms of their limiting environments:

The temperature and salinity values of the impinging watermasses appear to be the main influence on the shelf. In particular, a distinct

faunal and hydrographic boundary can be seen between the northeast of study area including Misaine, Middle and Canso Banks and Basins, and the Central shelf basinal areas of LaHave, Roseway and Emerald. The hydrographic boundary is between cold, less saline bottom waters from the St. Lawrence and Cabot Strait; and the warmer, more saline waters of slope origin. The associated faunal changes are from the Adercotryma glomerata, Recurvoides turbinatus and Cribrostomoides crassimargo assemblage (factor 1) to the Globobulimina auriculata, Nonionellina labradorica and Bolivina subaenariensis(factor 3). The former assemblage is restricted by a low temperature optimum and the latter, by a high salinity optimum. Both these statistically defined associations are consistent with the literature describing the species in other areas.

A shelf transition zone is recognized both in terms of the hydrography and fauna. To the northern periphery of the central basin area, observed temperature and salinity gradients for the bottom waters are high (i.e. 35-32 ‰ salinity and 9-1.9 °C). These gradients are present to a greater or lesser degree throughout the year (McLellen, 1957). The variable temperature and salinities in this area appear to preclude the occurrence of the A. glomeratum assemblage on the basis of adverse temperature and the G. auriculata assemblage on the basis of adverse salinity. The niches available appear to be taken up by the opportunistic S. atlantica and R. scorpiurus assemblage. The relationship to percent mud, observed for the central basins assemblage, is believed to be a function of increased amounts of organic carbon. This contention is further implied by the absence of this assemblage on the shelf edge where temperatures and salinities are similar but sediments are coarser with less organic carbon. Imprinted on this

hydrographic (temperature and salinity) control of distribution is the occurrence of a temperature and salinity independent assemblage, Cibicides lobatulus, whose presence is dictated by substrate suitability.

Assemblages of the slope and shelf edge also appear hydrographically controlled but seemingly to a lesser degree than on the shelf, being influenced also by nutrients, productivity of surface waters, organic carbon and oxygen content.

#### 8.4. Foraminiferal response to oceanographic variability

In most studies of benthic foraminiferal distributions it has been thought sufficient to determine, the relationship of assemblage and environment with little attention being paid to those aspects of foraminiferal physiology and metabolism that actually respond to these environments. This in part reflects the aims of the studies involved and the greater emphasis put on palaeoenvironmental rather than ecological or biological studies. In fact, knowledge concerning the life cycles, physiological needs etc. is minimal for all but a few well studied foraminiferal species.

Similarly, this study reveals dominant temperature, salinity and substrate controls of assemblage occurrences but did not investigate which aspects of foraminiferal biology is actually controlled. The most common explanation is that the assemblage occurs in that environment which provides optimum conditions for growth and reproduction. These optimum conditions give an assemblage an edge in competing for the limiting resources. This appears to be the case in the central basin complex where the presence of elements of the A. glomerata assemblage

are extremely suppressed by the more successful G. auriculata fauna. There is, however, an overall lack of elements of this central basin fauna in the northeast of the area. It has been suggested in this study that low salinity precludes the occurrence of the central basin fauna. Even so, its absence in total from all the Misaine Basins and its nearly complete absence elsewhere in the northeast is rather difficult to explain. Less nutrient flux compared to the central basins area and a more restricted bottom circulation in the Misaine basin areas may partly explain this aspect. Another reason that should be considered, especially in the light of the virtual absence of calcareous foraminifera in this area is that of calcium carbonate availability.

The literature suggests that active dissolution of calcareous tests in this area is not a possibility, the calcium carbonate compensation depth being present only at around 4000 m in this part of the Atlantic (Li et al, 1969; Broecker, 1974).

Many physico-chemical factors affect the solubility of  $\text{CaCO}_3$  in sea water. Revelle (1934) used the mass law equation

$$[\text{Ca}^{++}] \times [\text{CO}_3^{--}] = k \text{ CaCO}_3$$

in his discussion on the controlling influences of  $\text{CaCO}_3$  solubility. These influences include: the Ca and  $\text{CO}_3$  ion concentration; the value of a temperature dependant constant k. Both of these are in turn a function of salinity, temperature, hydrostatic pressure, amount of  $\text{CO}_2$  and the concentration of hydrogen and hydroxyl ions (i.e. pH). From this it was suggested (Revelle, 1934) that  $\text{CO}_2$  content in the water affected most the  $\text{CaCO}_3$  solubility when dealing with near normal marine salinity. Sverdrup et al(1942) and Revelle and Fairbridge (1957) added to the discussion and concluded that a temperature increase would lower the solubility product of  $\text{CaCO}_3$  and increases the  $\text{CO}_3$  ions; similarly

evaporation would increase the concentration of Ca ions and thus the carbonate alkalinity: Both these processes facilitate  $\text{CaCO}_3$  precipitation. The converse, low temperature, lower salinity, no evaporation, increased amounts of dissolved  $\text{CO}_2$  will suppress precipitation, which results in a lower availability of  $\text{CaCO}_3$  for foraminifera.

Working in the Antarctic, Anderson (1975), Osterman and Kellogg (1979) and Milam and Anderson (1981) have suggested that ice cover restricts the normal uptake of  $\text{CO}_2$  by photosynthesis, leaving greater amounts in the water and thus (from the above) decreasing  $\text{CaCO}_3$  availability. In the Antarctic, high concentrations of  $\text{CO}_2$  because of pack ice apparently lead to dissolution of hyaline tests.

Although it is clear from the above that the inter-relationships between  $\text{CaCO}_3$  chemistry and temperature, salinity, pressure, dissolved  $\text{CO}_2$  are complex (and thus difficult to evaluate in a study such as this), Greiner (1971) derived a simple model relating benthic foraminiferal distribution patterns to availability of  $\text{CaCO}_3$ . It is useful to discuss the distribution patterns observed in this study in terms of the above model and the above discussion on carbonate chemistry.

Chapters 6 and 7 have shown the association of the completely agglutinated assemblage A. glomerata with low temperatures and salinities ie. two factors that lower  $\text{CaCO}_3$  availability. A look at the winter ice cover in the present study area (Forward 1954; Cok 1970) shows a spatial occurrence similar to the limits of this agglutinating A. glomerata assemblage. Little is known concerning the chemical oceanographic results of this ice cover on waters off Nova Scotia. What is apparent is that the circulation within the Misaine Basin complex is

restricted and would tend to accumulate any dissolved CO<sub>2</sub> that had not been taken up through photosynthesis. There is thus strong evidence that the absence of calcareous tests and the dominance of an agglutinating assemblage is a direct result of corrosive arctic derived waters, which at the least would give competitive advantage to these agglutinating forms. Further evidence is seen in the presence of planktonic foraminifera in the surface waters of this area (Aksu, pers. comm. 1983) and not in the bottom sediments. To the south west in the central basins calcareous forms dominate under warmer more saline ice free waters. A convenient test for this tentative hypothesis is seen in areas to the north of the present study, in particular in the Arctic. One would expect similar patterns in the more northern, heavily iced arctic waters. Carsola (1952) indicated a combination of low salinity, high CO<sub>2</sub> and low temperature as increasing the solubility of CaCO<sub>3</sub>, and thus the prevention of calcareous tests. Arctic Bay foraminifera (Marlowe and Vilks 1963) show a dominance of agglutinate types in areas of low CaCO<sub>3</sub>. Calcareous foraminifera however are present in the Arctic and form dominant populations under heavily iced areas (Lagoe 1979; Green 1960; Vilks 1969). All of these areas however are of normal salinity and, more importantly, lie in deeper well-circulating waters which would tend to remove excess CO<sub>2</sub> from an area; thus two factors that reduce CaCO<sub>3</sub> availability are lessened in their influence.

If this calcium carbonate dependence can be quantitatively evaluated, i.e. threshold levels of CaCO<sub>3</sub> availability above which hyaline tests can form without undue stress, etc., then it would add evidence to Greiners (1970) model concerning foraminiferal distribution. The latter work postulates that the progressive restriction of hyaline types is primarily governed by CaCO<sub>3</sub> availability, explaining their

correlation to temperature and salinity. Bartlett (1966) studied the test composition of arctic, cool temperate and warm saline environments and noted a predominance of agglutinates in the Arctic, hyaline in the cool temperate and miliolids in the warmer saline. The cool temperate (hyaline) bay environments were from Tracadia Bay in the Northumberland Straits within the Bay of St. Lawrence, and therefore represented the same climate as in the present study area. The presence of hyaline types, and of the Cibicides lobatulus assemblage that occurs in the coastal zones to the north east can perhaps be explained as the result of its distribution in the seasonal, turbulent layer where greater amounts of O<sub>2</sub>, less CO<sub>2</sub> and warmer summer temperatures predominate making CaCO<sub>3</sub> more available.

The lack of scientific study concerning the effects of increased CO<sub>2</sub> on CaCO<sub>3</sub> availability in this area renders conclusions regarding its control on foraminiferal distribution unwarranted. One can say, however, that this control is related to temperature and salinity characters and that the evidence suggests that certain oceanographic conditions to the north east (e.g. sluggish circulation, heavy ice cover, less than normal salinity, etc) to prevail together, within this area, to determine the presence or absence (or suppressed presence) of calcareous foraminiferal tests.

#### 8.5 Diversity response to oceanographic variability.

In an earlier chapter, the observed trends in species diversity were described, illustrated and briefly discussed in terms of environmental influences. The results of this and the previous two chapters now allow further insight into these trends. Briefly, what has

to be explained are the apparent associations of:

- a/ low diversity ( $H(S)$ ) and high equitability ( $E$ ) with cold Labrador derived waters of less than normal salinity;
- b/ low  $H(S)$  and low  $E$  with inner shelf areas;
- c/ high  $H(S)$  and high  $E$  with warm slope derived waters in the central basin and shelf edge areas;
- d/ low  $H(S)$  and low  $E$  with deeper slope, cool water environments.

In particular, it is hoped to discuss these in terms of their concordance with various ecological hypotheses of diversity, stability.

In order to explain lowered diversity levels on the shelf and slope of Georgia, Arnold and Sen Gupta (1981) examined three processes:

- a/ reworking
- b/ biotope overlap
- c/ environmental control

Reworking has been shown to be a problem only in certain areas of shelf and slope off Nova Scotia and has been invoked previously (Chapter 5) to explain locally depress  $H(S)$  and  $E$  values. It does not, however, explain the general trends.

Similarly, biotope overlap could not explain these observed general trends. Each defined assemblage is statistically independent in terms of its distribution with the main assemblages involved having very few co-occurring species.

The observed diversity trends then, appear in some way related to environmental factors. With the community approach in mind, this is not surprising. It has already been assumed that certain environmental factors determine the number of species potentially able to live in a particular region; for the purpose of this study these are mainly temperature and salinity characters. Whether or not species become dominant within an assemblage depends upon their competitive ability with respect to the limiting resources. Clearly, the higher number of species (diversity) present in the central shelf basin and shelf edge areas is a function of the warmer, more optimum hydrographic conditions allowing the reproduction and growth of calcareous as well as agglutinating foraminifera. This also results in high E values. Similarly, the low H(S) values to the northeast, which are associated with colder arctic derived waters, are low because less species find optimum conditions. This latter aspect has been shown to be related to lower  $\text{CaCO}_3$  availability and therefore greater physiological stress and the exclusion of  $\text{CaCO}_3$ -secreting foraminifera. High E values in this area would also imply a certain level of environmental stability for the well adapted species.

To a large extent, these observed patterns reflect Sanders (1968) and Slobodkin and Sanders (1969) hypothesis on environmental predictability and species diversity. They contend that predictable environments with "non-extreme physical properties" will be inhabited by a diverse, somewhat specialized population. Any change in the "predictability" of an environment could result in local extinction of specialized forms. To all intents and purposes, if both the central basin area (warm, saline) and the area to the northeast (cold, relatively fresh) are predictable environments, why then is one a high

diversity area whereas the other is a low diversity area? Quite simply, and this is noted in Slobodkin and Sanders (1969), the predictability of an environment is not purely an environmental phenomenon, it is also dependent in part on the organism itself. In effect then, the physiological needs of calcareous foraminifera, i.e. the need to secrete  $\text{CaCO}_3$ , means that the northern area (cold, relatively fresh) is for this group an "unpredictable" environment. This exclusion of calcareous forms has lowered diversity in this area, whilst the agglutinates present have proceeded to become "biologically accommodated" (Sanders 1969) as seen in the high E values.

Further evidence of environment stability and its influence on species diversity is seen in the central basin area and its periphery. The peripheral assemblage (Factor assemblage 4, Saccammina atlantica) has been shown to occupy an environmentally unpredictable area of relatively intense temperature and salinity gradients (of locally fluctuating areal extent). This region is a buffer between the two dominant shelf waters and, as such, is prone to fluctuations in areal extent and intensity seasonally and otherwise (i.e. after storms, McLellen 1957; Warner 1970). This unpredictability then precludes to a large extent calcareous forms present in the basin which are presumably rather specialized and unable to compete with the opportunistic species (S. atlantica). This situation is also supportive of Connell and Slaters (1977) view which allows for an optimum amount of environmental variability, which enables opportunistic species to compete with specialized species (i.e. as in central basin assemblage Factor 3, Globobulimina auriculata where S. atlantica also occurs). At higher levels of disturbance (i.e. in the peripheral, high salinity, temperature gradients zone) opportunistic species (S. atlantica

assemblage) dominate at the expense of more specialized species.

The low H(s) and E values of the inner shelf coastal stations are coincident with coarse substrates, turbulent high oxygen waters and are thought here to be a result of the inability of many species of foraminifera to cope with this kind of stress; the subsequent niche gap is occupied almost monospecifically by the Cibicides lobatulus assemblage which, through its attached mode of life is able to survive and reproduce very well in this environment. On the other hand the low H(s) and E seen in lower slope stations is probably a result of flooding of the assemblage through transportation of Elphidium excavatum again producing an almost monospecific assemblage. It is an important distinction to make then between areas which have a low diversity in response to physiological constraints and those areas which are low because of physical action of the environment.

In summary, the general diversity trends observed in this study are a function of imposed physiological stress related to overall hydrographic characters. These relationships however, as seen in the above discussion are far from simple, and any patterns observed are a result of the inter action of the factors discussed. The patterns seen in this study do seem to follow several hypothesis relating species diversity to environmental stability.

## CHAPTER 9: IMPLICATIONS AND APPLICATIONS OF THIS STUDY

This chapter considers the main ramifications of the present work for:

- 9.1. Ongoing investigation of Pleistocene to Recent deposits off Canada's eastern continental margin.
- 9.2. Future foraminiferal study in this area and elsewhere.
- 9.3. Interpretation of geologically older foraminiferal assemblages.

9.1. Implications for Pleistocene to recent studies off Eastern Canada.

One of the stated aims at the beginning of this study was to objectively synthesize present day assemblage-environment relationships and thus provide data for palaeoceanographic studies off the Eastern Canadian continental margin. The multivariate methods that have been used to attain this are the same as those employed by CLIMAP (1976) studies (Imbrie and Kipp 1971) which are concerned with the wider aspects of reconstructing the palaeoceanography of the Atlantic. Most of these studies (e.g. Imbrie and Kipp 1971; Molfino *et al* 1982) concentrate on information obtained from planktonic foraminifera. More recently, however, studies with similar aims and methods have been undertaken on deep water benthic foraminifera (Miller and Lohmann 1982; Lohmann 1978; Streeter and Lavery 1982; Belanger and Streeter 1981). These are for the most part restricted to deep sea investigations, with little attempt at examining (with similar methods) the continental shelf or neritic foraminiferal distributions. This is all the more surprising in light of the fact that the watermass and circulation changes that

have taken place since the last glaciation will have had more effect on the nearshore sediments and fauna than the deep sea. Only Osterman and Kellogg (1979) in the Antarctic and more recently Mudie *et al.* (1983) off eastern Canada have attempted palaeoceanographic studies of the neritic zone using CLIMAP methods.

The implications of this present study for palaeoceanographic investigation of the Pleistocene to Recent off Canada's eastern continental margin can only be considered in the light of current knowledge concerning this period of time.

Several workers have investigated the palaeoclimatic and palaeoceanographic history as represented in the sediments of the Eastern Canadian margin (Fillon 1976; Alam Piper and Cooke 1983, Mudie 1980). Using palynomorph data from Emerald Basin cores, Mudie (1980) has tentatively recognized several climatic phases from the last glacial event up until the present day. A glacial phase at 15-13 Ka. B.P. was followed by a later glacial phase of warmer waters (13-10 Ka B.P.). At 10-7 Ka B.P. occurred an early Holocene phase of rapid climatic warming with the closer proximity of tropical airmasses resulting in a stronger influx of warm saline waters over the Scotian shelf. The Holocene Hypsithermal phase followed (7-3 Ka. B.P.) and was a period of maximum climatic warmth. During this time period, other workers have shown an oceanographic cooling trend in the Labrador current (Fillon 1976; Alam 1979; Hill 1979). This cooling event is thought by Fillon (1976) and Alam (1979) to represent the increased influence of the shallow shelf component of the Labrador current. Finally, from 3 Ka to the present day, is a climatic deterioration phase due to cooler surface waters and possibly a further increase of the Labrador current influence (Mudie 1980).

At the present time, the margin off Nova Scotia lies at the limits of direct influence of the shallow shelf component of the Labrador current. An earlier chapter (chapter 1) described how its effect is presently squeezed shoreward on the shelf both by bathymetry and more importantly, by the slope waters intrusion. The results of this present work have shown how the broad hydrography is statistically reflected in the spatial distribution of benthic foraminifera on the shelf and thus provides the potential for further quantitative appraisal of piston cores and subsequent interpretation in terms of the most recent geological history off Nova Scotia.

The most convenient method to do this, considering the nature of the present data, would be to apply paleoecological transfer functions (Imbrie and Kipp 1971; Mudie *et al* 1983) to down core assemblages and from these to make quantitative estimates concerning depth, temperature and salinity. It is beyond the scope of this thesis to do this; however, some discussion is warranted on the nature of these core successions and what they may imply.

Core 80-004-33 provides foraminiferal data (Scott, D.B., unpub. data) from sediments within Canso Basin (fig. 9.1). A tentative interpretation of the foraminiferal successions observed is offered. The basis for these are the results of the present work, in particular the observed present day assemblage-environment relationships.

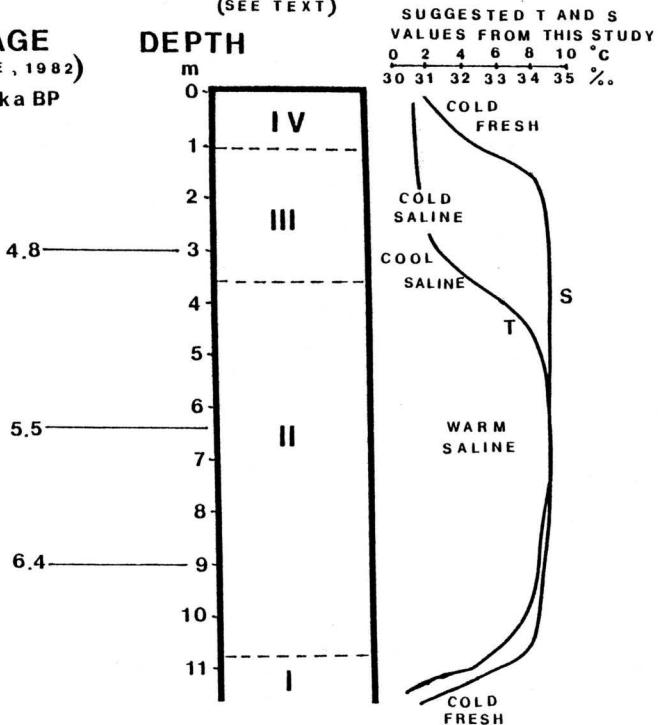
At the base of the core is an Elphidium excavatum f. clavata assemblage (assemblage I in fig. 9.1) almost monospecific and previously associated with adjacent ice margins (Vilks 1976). It is an assemblage such as this that has been recognized in an earlier chapter as being relict and reworked. Dates from nearby Emerald Basin cores put the time interval dominated by this assemblage at 15000 + 280 years B.P. (Vilks,

## FORAMINIFERAL ASSEMBLAGES

(SEE TEXT)

AGE  
(MUDIE, 1982)  
ka BP

DEPTH  
m



## CLIMATIC EVENTS

(MUDIE, 1980 ; FILION, 1976)

CLIMATIC DETERIORATION

HOLOCENE  
HYPSITHERMAL

Period  
of  
maximum  
climatic  
warmth

rapid warming  
EARLY HOLOCENE  
LATE GLACIAL ?

Figure 9.1 Summary diagram for core 80-004-33

1976) which, according to Milliman and Emery (1968) is when sea level was at its lowest stand. This fauna is replaced above by a Globobulimina auriculata, Nonionellina labradorica assemblage (II in fig. 9.1) with some amounts of Bolivina subaenariensis. This assemblage presently occurs at the surface over a large part of the central Basin (Emerald and LaHave) and has in this study been statistically related to the normal marine (34-35‰) salinities and warmer waters (8-12°C) of slope origin that gain access through the Scotian Gulf. This assemblage is replaced up the core by an assemblage dominated by Islandella teretis, Cassidulina reniforme and N. labradorica (III in fig. 9.1). B. subaenariensis disappears and G. auriculata is severely restricted. A similar assemblage to this has been related to normal marine salinity (34-35‰) but rather cooler waters (3-4°C) by Mudie *et al.* (1983) on the Labrador shelf to the North. At the top of the core (30-0 cm) this assemblage (III in fig. 9.1) is replaced by a modern day assemblage, (IV in Fig. 9.1) Adercotryma glomerata which is associated statistically to cold (1-2°C) and relatively fresh (31-33‰) waters of Labrador current origin.

The hydrographic changes then, as represented by the foraminiferal successions in this core, appear to be from cold (1-2°C) relatively fresh (29-31‰) waters; to warmer (6-9°C) more saline (34-35‰) waters; to colder (3-4°C) saline (34‰) waters and finally to colder (1-2°C) less saline waters (31-33‰).

The extent to which these oceanographic events fit into the framework described by Mudie (1980) from Emerald Basin cores can be illustrated with reference to the Canso core pollen data (Mudie 1982, pers. comm.). This data, calibrated with the dates of Livingstone (1968), provide a working chronology (Fig. 9.1). It is clear, both from the age

and pollen zonation that glacial and late glacial deposits have not been penetrated. The early Holocene, rapid warming phase (10-7 Ka.) is perhaps represented in this core by the warm saline assemblage (G. auriculata, N. labradorica). The Holocene hypsithermal phase (7-3 Ka), or period of maximum warming, is represented in the oceanographic record initially by a warmer saline assemblage (including B. subaenariensis) and progressively by a cold saline assemblage (N. labradorica, C. reniforme, I. teretis). This period of maximum climatic warmth at 5-6 Ka is coincident with Fillon's (1976) and Alam's (1979) oceanographic cooling events and is due to the increased influence of the shelf component of the Labrador current. The trigger for this increased cooling influence is presumably the dissipation of the northern ice volumes by the maximum atmospheric warming associated with the hypsithermal. Following this period is the climatic deterioration phase and is (3 Ka. - present) represented by the cold, relatively fresh assemblage A. glomeratum. The significance of this last phase still has to be thoroughly investigated. From this study, it is apparent that the assemblage IV in fig. 9.1 replaces the other (III in 9.1) cooler more saline assemblage (with I. teretis, N. labradorica and C. reniforme) because of decreased salinity. The impetus for this may be the overall climatic cooling trend that could have increased the temperature and salinity gradients between the Labrador influence and the slope water influence. Mixing of these two waters may, in the past, have been responsible for the elevated salinities in the northeast region and therefore for the N. labradorica, I. teretis fauna at depth in the core.

The synthesis above is rather simplistic and is meant only to describe in qualitative terms what is represented oceanographically by the observed piston core foraminiferal assemblages. What is required

further is corroborative evidence from other cores at key sites on the shelf and slope off Nova Scotia. Application of biological transfer functions in the manner of Imbrie and Kipp (1971) and Mudie *et al.* (1983) using the quantitative results of this study would provide quantitative estimates of palaeodepth, temperature and salinity values down these cores. Together with an accurate chronological framework, such an approach will provide valuable insight into the nature of the paleoceanographic changes marking the Pleistocene to recent interval off Eastern Canada.

#### 9.2. Implications for future foraminiferal study.

Several interesting lines of investigation are available for further foraminiferal study within this area. Research could be well employed in a more detailed and systematic analysis of the three main shelf assemblages. The central basin assemblage (Globobulimina auriculata), its peripheral assemblage (Saccammina atlantica) and the northeast shelf assemblage (Adercotryma glomerata) are presently in equilibrium with the oceanographic conditions; a more rigorous definition of these limits and oceanographic controls using the same methods employed in this study would be invaluable for a more accurate palaeoceanographic study. Further study should involve several sample transects across these faunal boundaries together with the collection of alkalinity and CO<sub>2</sub> data which would determine if CaCO<sub>3</sub> availability really is a primary limiting control of these aforementioned shelf assemblages.

The number of statistically independent assemblages defined on the shelf (8) reveals the importance of an adequate sample control. Any lesser amount of samples would probably have failed to distinguish some

of the smaller, but important assemblages such as the peripheral assemblage (S. atlantica), etc. Similarly, it was important to use multivariate methods in this study as consideration only of indicator species from this amount of samples would have proved rather inaccurate.

The recognition of transported and relict components of assemblages is important in any study. Their recognition in this study was made easier through a consideration of live foraminifera and a sieve size that retained juvenile components. The present work located and defined to some extent the degree of relictness and transportation on the shelf. Further work could consider this problem closer, paying particular attention to source areas of relict deposits. Effects of transportation of foraminiferal tests could be more systematically studied. An approach similar to that of Murray (1982) who found benthic foraminifera in the whole water column (following storms) in the southwestern approaches to the British Isles, is envisaged. Of particular interest would be the area known as the Scotian Gulf, southward down the slope and northward into Emerald Basin. Current meter measurements in this area (Hill 1982) have recorded several peaks of water movement throughout the year. These are storm related and are of sufficient velocity to put sand grains (and foraminifera) into suspension. An investigation designed to measure the degree and extent of any cross-shelf movement of foraminiferal tests in this area would be important for a more thorough understanding of the recent sedimentological and foraminiferal history of this area.

Finally, more slope transects would add to the comprehensive understanding of slope foraminiferal assemblages, their variability and limiting parameters.

### 9.3. Interpretation of geologically older foraminiferal assemblages.

Any palaeoecological study is essentially concerned with the reconstruction of the watermass, climatic and bathymetric history of the deposits. The direct application of the present day quantitative models to fossil populations is not theoretically valid for rocks older than Miocene (Douglas 1979); the stratigraphic base of many present day species. At the generic and familial levels, one would be able to go back to the late Cretaceous, and early Cretaceous respectively (Sliter 1972; Douglas 1979) these being the limits of direct taxonomic approaches.

Palaeoecological study based upon analogous or homeomorphic species assumes that neither the physiological needs nor the bathymetric habits have changed substantially with time. In view of the changing nature of the character of the worlds oceans (i.e. development of the psychrosphere in early Tertiary) and theories of speciation and evolution, these assumptions may be misleading. Blake and Douglas (1979) have shown that the depth limits of Melonis pompilioides have shifted during the Pleistocene to Recent, as a response to changes in the watermass depth limits. Similarly, Douglas (1979) summarized the results of Streeter (1973) and Schnitker (1974) who show the changing bathymetric distributions of Atlantic assemblages over several thousand years. These emphasize the caution required for palaeodepth interpretation.

The absence, presence and relative abundance of agglutinated foraminifera is perhaps a more useful tool for interpretation. Their usefulness in this respect lies in the inability of calcareous foraminifera to exist or compete successfully in chemical opposition to

the environment (Greiner 1970). The recent distribution of agglutinated foraminifera dominated assemblages has been shown to have distinct depths, diversity and morphological characters (Scott *et al* 1981) that have remained consistent through a long period of geological time. Gradstein and Berggren (1981) and Miller *et al* (1982) have developed some useful palaeoecological models with this in mind. The use of calcareous assemblages is more difficult and prone to misinterpretation. Bandy (1960) realized the potential of convergent adaptation in palaeoecological interpretation, contending that morphological similarity is related to environmental affinity. Such a correlation of form and environment is seen in Pflum and Frerichs (1966) and Haq and Boersma (1977). The usefulness of buliminids and uvigerinids in this respect has been pointed out by Douglas (1979). Certainly, in this study, the size, shape and ornamentation of buliminids varies with bathymetry. In shallow shelf depths (0-200 m.) the distinctive Bulimina marginata and Bulimina aculeata are the main forms. They are replaced at upper slope depths (250-800 m.) by Bulimina exilis and finally, by Bulimina striata at lower slope depths (1000 m. +).

In conclusion, consideration of all aspects of fossil foraminiferal populations, their diversity and equitability characters, the presence, absence or dominance of agglutinated types, including their morphology, etc, should be rigorously pursued in order to gain the most meaningful palaeoecological results.

## CHAPTER 10: CONCLUDING STATEMENTS

1. Examination of bottom grab samples taken from the continental margin off Nova Scotia revealed some 120 foraminiferal species. Seventy five of these occurred in abundances greater than 3% at any one sample.
2. Q-Mode factor analysis methods of Imbrie and Kipp (1971) delimited eight shelf and four slope statistically independent foraminiferal assemblages on the continental margin off Nova Scotia. The distributions of these assemblages are seen in figure 3.9. The raw data input comprised relative abundance values for the 75 main species (i.e. species making up more than 3% of the total in any sample) at 250 sample sites.
3. Comparison of the factor assemblages with the living foraminiferal abundance data shows that most of the assemblages are presently in equilibrium with the prevailing marine environment. Others were found to be reworked and or relict.
3. Multiple regression analysis shows that the principal shelf assemblages are significantly related to some aspects of the prevailing hydrography. The central shelf basin assemblage ( factor 3; Globobulimina auriculata and other predominantly calcareous species) is controlled in its distribution by its apparent requirement of normal marine salinities, and to a lesser extent, to the higher temperatures associated with the impinging slope derived waters. The north east shelf assemblage (factor 1, Adercotryma glomerata and completely

agglutinated) is significantly related to the cold, less than normal marine salinities of the inner margin Labrador current. This arctic derived water reaches its limit of continental shelf influence over the Misaine and Canso Bank-Basin area.

4. A transition assemblage (factor 4, Saccammina atlantica) is delimited by the factor analysis and occurs between the central shelf basin and the north east shelf assemblages. This assemblage occupies an area of steep temperature and salinity gradients associated with the boundary between the cold arctic derived waters and the warmer slope derived waters.

5. The observed response of the main shelf assemblages to the hydrography may in part be a function of decreased  $\text{CaCO}_3$  availability in the arctic waters and the subsequent inability of hyaline foraminifera to secrete calcareous tests.

6. Imprinted on these main shelf distribution patterns are two substrate controlled assemblages (factor 2, Cibicides lobatulus; factor 7, Islandiella islandica). The nature of the substrate (coarse and gravelly sands) appears to locally override any importance of the hydrographic constraints.

7. Four measures of species diversity ( $H(s)$ , S, E,  $\alpha$ ) show essentially similar trends. Highest values are found along the shelf edge, upper slope and in the central basins. All these areas are overlain by warm, saline waters with high nutrient capacity. Low values of diversity are seen in the north east shelf and along the inner shelf

zone. These values appear related to low temperature, low salinity conditions. The low values along the inner shelf may be a result of the coarse substrate upon which only a few species can survive (eg. C. lobatulus). The lower slope area also has depressed diversity values which is probably due to dilution with a monospecific reworked assemblage (E. excavatum).

8. Several slope assemblages are recognised. An upper slope assemblage shows an association with depth and high water productivity (slope factor 2, Bulimina exilis). A lower slope assemblage (U. peregrina) appears to be associated with a narrow range of temperature and salinity conditions (2-4°C, 34-35‰/00).

9. The foraminiferal response to hydrographic conditions on the shelf off Nova Scotia has been shown to be statistically significant and has potential for further quantitative studies of the palaeoceanographic evolution of this area.

10. Using the results of this objective study, qualitative estimates of past temperature and salinity conditions have been made through an interpretation of the foraminiferal assemblages present in a core from Canso Basin. From the base of the core upwards oceanographic conditions change from cold fresh, to warm saline water, followed by cool saline, to cold saline and finally to cold fresh, water over a time period of 15 Ka to present (fresh being less than normal marine salinity, about 30-33.5 ‰).

11. The indicated oceanographic sequences recognised above fit in

with previous palaeoceanographic interpretations and palaeoclimatic data. It appears that the oceanographic history is a result of the interchanging influences of Labrador current and the Gulf Stream waters.

12. The data, observations and discussions reported in this thesis provides a suitable base for quantitative palaeoceanographic studies of the continental margin off Nova Scotia, and for future, more precise investigation of the present foraminiferal distributions.

## CHAPTER 11 SYSTEMATIC TAXONOMY

Many of the species present off Nova Scotia have been previously well described and documented. For this reason, it was thought unnecessary to include species descriptions in this study. Taxonomic problems are presently the subject of separate studies (Miller 1983). Classifications adopted in this study follow that of Loeblich and Tappan (1964).

The following chapter includes essential systematic data for the seventy-five species used in the statistical analyses. Reference is made to initial descriptions, subsequent nomenclatural changes, more recent documentation and illustration and to several plates which figure the main species found in this study. The photographs were taken with a scanning electron microscope. Reference is also made to forty-two other unfigured rarer species not used in the analysis. Reference slides are housed in the Department of Geology, Dalhousie University, Halifax, Nova Scotia.

Superfamily	AMMODISCACEA	Reuss 1862
Family	ASTRORHIZIDAE	Brady 1881
Subfamily	HIPPROCREPININAE	Rhumbler 1895
Genus	HYPERAMMINA	Brady 1878

Hyperammina elongata Brady

Plate 1 Fig. 1

Hyperammina elongata Brady, 1878, p. 433, pl. 20, figs. 2a, b; Parker, 1952, p. 395, pl. 1, fig. 10; Vilks, 1969, p. 42, pl. 1, fig. 4;

Gregory, 1971, p. 161, pl. 1, figs. 2, 3.

Distribution: Rarely living and often found broken this species is restricted to central shelf basin areas and middle slope regions.

SACCAMMINIDAE    Brady 1884

SACCAMMININAE    Brady 1884

SACCAMMINA    Sars in Carpenter 1869

Saccammina atlantica (Cushman)

Plate 1 Fig. 2

Proteoina atlantica Cushman, 1944, p5, pl. 1, fig. 4; Phleger, 1952, p. 85, pl. 13, figs. 1, 2.

Saccammina atlantica (Cushman) Vilks, 1962, p. 43, pl. 1, fig. 13; Barbieri and Medioli, 1969, p. 853, pl. 1, fig. 4; Cole, 1981, p. 13, pl. 1, fig. 14.

Distribution: Dominant on the inner shelf, in assemblages 5 and 6. It occurs also in association with other agglutinated assemblages.

Saccammina difflugiformis (Brady)

Plate 1 Fig. 3

Reophax difflugiformis Brady, 1879, p. 51, pl. 4, fig. 3a, b.

Proteoina difflugiformis (Brady) Phleger and Parker, 1955 p. 2, pl. 1 fig. 4, 5.

Saccammina difflugiformis (Brady) Todd and Bröniman, 1957, p. 52, pl. 1, fig. 15.

Distribution: Restricted to inner shelf assemblages 3 and 4.

Saccammina sphaerica Brady

## Plate 1 Fig. 4

Saccammina sphaerica Brady, 1871, v. 7, p. 187; Barker, 1960, p. 36, pl. 123, fig. 11-15.

Distribution: Rare, not observed live and restricted to the inner shelf.

AMMODISCIDAE	Reuss 1862
AMMODISCINAE	Reuss 1862
AMMODISCUS	Reuss 1862

Ammodiscus catinus Hoglund

## Plate 1 Fig. 5

Ammodiscus catinus Hoglund, 1947, p. 122, pl. 8, fig. 1, 7; Gregory, 1971, p. 166, pl. 1, fig. 7.

Distribution: Not an important component in any assemblage, occurs mainly on sandy substrates.

LITUOLACEA	de Blainville 1825
HORMOSINIDAE	Haeckel 1894
HORMOSININAE	Haeckel 1894
REOPHAX	Montfort 1808

Reophax arctica (Brady)

## Plate 1 Fig. 8

Reophax arctica Brady, p. 405, pl. 21, fig. 2a-b; Parker, 1952, p. 85, pl. 13, fig. 3; Gregory, 1971, p. 168, pl. 2, fig. 3; Cole, 1981, p. 22, pl. 16, fig. 22.

Distribution: Occurs mainly in assemblage 1. Restricted mainly to inner shelf areas on sandy and silty substrates.

Reophax scotti Chaster

Plate 1 Fig. 11

Reophax arctica Chaster, 1890, p. 57, pl. 1, fig. 1; Gregory, 1971, p. 172, pl. 11, fig. 7; Cole, 1981, p. 26, pl. 2, fig. 17.

Distribution: Similar in occurrence to R. arctica although less abundant and more fragile.

Reophax nodulosa Brady

Plate 1. Fig. 14

Reophax nodulosa Brady, 1879, v. 19, p. 52, pl. 4, fig. 7, 8; Cole, 1981, p. 25, pl. 4, fig. 7,8.

Distribution: Found mainly in central basin assemblage 3. Distinctive but easily broken.

Reophax guttifera (Brady)

Plate 1 Fig. 10.

Lituola guttifera Brady, 1881, v. 21, p. 49.

Reophax guttifera (Brady) Brady, 1884, p. 295, pl. 31, fig. 10-15; Cole, 1981, p. 25, pl. 2, fig. 11.

Distribution: Mainly associated with central shelf basin assemblages.

Reophax scorpiurus(Montfort)

Plate 1 Fig. 9

Reophax scorpiurus Montfort, 1808, p. 330; Loeblich and Tappan, 1953, p. 24, pl. 2, fig. 7-10; Leslie, 1965, p. 169, pl. 1, fig. 6, 7.

Distribution: Occurs on the periphery of the basin assemblage as a sub-dominant component of assemblage 4.

LITUOLIDAE de Blainville 1825

HAPLOPHRAGMOIDINAE Maync 1925

HAPLOPHRAGMOIDES Cushman

Haplophragmoides canariensis (d'Orbigny)

Plate 1 Fig. 13

Nonionina canariensis d'Orbigny, 1839, p. 128, pl. 2, fig. 33.

Haplophragmoides canariensis (d'Orbigny) Cushman, 1910, v. 1, p. 101.

Distribution: Found mainly in central basin assemblage 3; apertural characters enable its distinction from C. jeffreysi.

ADERCOTRYMA Loeblich and Tappan 1952

Adercotryma glomerata (Brady)

Plate 1 Fig. 16

Lituola glomerata Brady, 1878, v. 1, p. 433, pl. 20, fig. 1a-c.

Adercotryma glomerata (Brady) Loeblich and Tappan, 1953, v. 121, p. 26, pl. 8, fig. 1-4; Vilks, 1969, p. 44, pl. 1, fig. 15; Gregory, 1971, p. 173, pl. 11, fig. 9-11; Cole, 1981, p. 29, pl. 4, fig. 5-6.

Distribution: Most abundant in the north east of the study area.

Dominant component of shelf assemblage 1, and occurs in basins and on banks.

RECURVOOIDES Earland 1934

Recurvoides turbinatus (Brady)

Plate 1 Fig. 17

Haplophragmium turbinatus Brady, 1881, p. 50.

Recurvoides turbinatus (Brady) Parker, 1952, p. 402, pl. 2, fig. 23, 24; Vilks, 1969, p. 45, pl. 1, fig. 19; Gregory, 1971, p. 176, pl. 3, fig. 3, 4; Cole, 1981, p. 32, pl. 6, fig. 7, 8.

Distribution: Essentially the same as A. glomerata.

CRIBROSTOMOIDES Cushman 1910

Cribrostomoides crassimargo (Norman)

Plate 1 Fig. 6, 7.

Haplophragmium crassimargo Norman, 1892, p. 17.

Labrospira crassimargo (Norman) Hoglund, 1947, p. 11, fig. 1, text-fig. 121-125.

Cribrostomoides crassimargo (Norman) Leslie, 1965, p. 158, pl. 2, fig. 2a, b; Barbieri and Medioli, 1969, p. 855, fig. 4; Gregory, 1971, p. 176, pl. 3, fig. 1, 2.

Distribution: Occurs mainly with A. glomerata in assemblage 1.

Cribrostomoides jeffreysi (Williamson)

Plate 1 Fig. 12

Nonionina jeffreysi Williamson, 1858, p. 34, pl. 3, fig. 72, 73

Cribrostomoides jeffreysi (Williamson) Barbieri and Medioli, 1969, p. 855, fig. 4; Vilks, 1969, p. 45, pl. 1, fig. 17a, b; Cole, 1981, p. 30, pl. 6, fig. 6.

Distribution: Associated mainly with assemblages 3 and 4, and sandy/silty substrate.

Cribrostomoides subglobosum (Sars)

Plate 1 Fig. 15

Lituola subglobosum Sars, 1871, p. 253.Cribrostomoides subglobosum (Sars) Cole, 1981, p. 31, pl. 4, fig. 13.

Distribution: Rare and confined to assemblage 1 and 4.

CYCLAMMINAE Marie 1941

ALVEOPHRAGMIUM Tschendrina 1936

Alveophragmium scitulum (Brady)

Plate 2 Fig. 6

Alveophragmium scitulum (Brady) Parker, 1952, p. 487; Barker, 1960, p. 70, pl. 34, fig. 11-13.

Distribution: Maximum occurrences in basin assemblages 3 and 4. Never a dominant component.

CYCLAMMINA Brady 1879

Cyclammina cancellata Brady

Plate 2 Fig. 10

Cyclammina cancellata Brady, 1879, v. 19, p. 62; Cole, 1981, p. 12, pl. 5, fig2.

Distribution: Very rare, restricted to slope regions.

LITUOLINAE de Blainville 1825

AMMOBACULITES Cushman 1910

Ammobaculites dilatatus Cushman

## Plate 2 Fig. 1

Ammobaculites dilatatus Cushman. Cushman and Bronnimann, 1948, p. 39, pl. 7, fig. 10-11; Schnitker, 1971, p. 193, pl. 1, fig. 7

Distribution: Restricted to basins and slope. Not very abundant.

## AMMOTIUM Loeblich and Tappan

Ammotium cassis (Parker)

## Plate 2 Fig. 2

Lituola cassis Parker in Dawson, 1870, p. 177, fig. 3.

Ammobaculites cassis (Parker) Cushman, 1920, p. 63, pl. 12, fig. 5.

Ammotium cassis (Parker) Loeblich and Tappan, 1953, p. 33, pl. 2, fig. 12-18; Gregory, 1970, p. 176, pl. 3, fig. 5.

Distribution: Restricted to inner shelf near-shore areas of silts and fine sands. Common in assemblage 6.

## TEXTULARIDAE Ehrenberg, 1838

## SPIROPLECTAMMINAE Cushman, 1927

## SPIROPLECTAMMINA Cushman, 1927

Spiroplectammina biformis (Parker and Jones)

## Plate 2 Fig. 5

Textularia agglutinans d'Orbigny var. biformis Parker and Jones, 1865, p. 370, pl. 15, fig. 23, 24.

Spiroplectammina biformis (Parker and Jones) Cushman, 1927, p. 23, pl. 5, fig. 1; Vilks, 1969, p. 45, pl. 1, fig. 20a, b; Gregory, 1970, p. 178, pl. 3, fig. 6; Cole, 1981, p. 34, pl. 6, fig. 3-4.

Distribution: Abundant in shelf assemblages 1 and 6. Present on both banks and basins, sands and silts.

TEXTULARINAE Ehrenberg 1838

TEXTULARIA Defrance in de Blainville 1824

Textularia torquata Parker

Plate 2 Fig. 4.

Textularia torquata Parker, 1952, p. 403, pl. 3, fig. 9-11; Vilks, 1968, p. 18, pl. 1, fig. 10, 11; Gregory, 1971, p. 179, pl. 4, fig. 1, 2.

Distribution: Extremely small specimens, low in abundance. Restricted to arenaceous assemblages.

TROCHAMMINIDAE Schwager 1877

TROCHAMMININAE Schwager 1877

TROCHAMMINA Parker and Jones 1859

Trochammina inflata (Montagu)

Plate 2. Fig. 12, 13.

Nautilus inflata Montagu, 1808, p. 81, pl. 18, fig. 3.

Trochammina inflata (Montagu) Parker and Jones, 1859, p. 347; Schafer and Cole, 1978, p. 29, fig. 1; Scott and Medioli, 1980, p. 44, pl. 3, fig. 12, 14.

Distribution: Restricted to northern periphery of central basins, not dominant.

The form of T. inflata as seen in this study differs from the marsh forms described in Scott and Medioli (1980); the shelf forms lack an inner organic lining.

Trochammina lobata Cushman

Plate 2. Fig. 15

Trochammina lobata Cushman, 1944, p. 18, pl. 2, fig. 10; Parker, 1952, p. 408, pl. 4, fig. 8a,b; Schnitker, 1971, p. 212, pl. 1, fig. 18; Cole and Furgerson, 1975, p. 14, pl. 4, fig. 5,6.

Distribution: Not an important component in any assemblage.

Trochammina squamata Parker and Jones

Plate 2. Fig. 8, 9.

Trochammina squamata Parker and Jones, p. 407, pl. 15, fig. 30-31 a-c; Schafer and Cole, 1978, p. 29, pl. 5, fig.1; Scott and Medioli, 1980, p. 45, pl. 4, fig. 6,7.

Distribution: Restricted to shallow banks and in-shore areas.

Trochammina globigeriniformis (Parker and Jones)

Plate 2. Fig. 14.

Lituola nautiloidea globigeriniformis Parker and Jones, 1865,p. 407, pl. 15, fig. 46,47.

Trochammina globigeriniformis Cushman, 1910, pl. 24, text-fig. 193-195; Cole, 1981, p.37, pl. 7, fig. 2-3;

Distribution: Same as T. squamata although in lesser abundance.

TRITAXIS Schubert, 1921

Tritaxis conica Parker and Jones

Plate 2. Fig. 11

Valvulina triangularis conica Parker and Jones, p. 406, pl. 15, fig. 27.

Tritaxis conica (Parker and Jones) Barker, 1960, p. 100, pl. 49, fig. 15,16; Cole, 1981, p. 41, pl. 17, fig. 19,20.

Distribution: Restricted to assemblage 4.

ATAXOPHRAGMIIDAE Schwager 1877

GLOBOTEXTULARINAE Cushman 1927

EGGERELLA Cushman 1933

Eggerella advena (Cushman)

Plate 2. Fig. 3.

Verneuilina advena Cushman, 1922, p. 141.

Eggerella advena (Cushman) Cushman, 1937, p.51, pl. 5, fig. 12-15;

Schnitker, 1971, pl.1, fig. 19; Cole, 1981, p. 42, pl. 6, fig. 2.

Distribution: Predominant component of assemblage 6, rare outside these limits.

KARRERIELLA Cushman 1933

Karreriella bradyi (Cushman)

Plate 2. Fig. 7.

Gaudryina bradyi Cushman, 1911, p. 67, fig. 107.

Karreriella bradyi (Cushman) Cushman, 1937, p. 135; Cole, 1981, p. 44, pl. 6, fig. 5.

Distribution: Occurs in many assemblages, never more than 5%.

MILIOACEA Ehrenberg 1839

MILIOIDAE Ehrenberg 1839

QUINQUELOCULININAE Cushman 1917

QUINQUELOCULINA d'Orbigny 1826

Quinqueloculina seminulum (Linne)

Plate 3. Fig. 1

Serpula seminulum Linne, 1758, p. 786.Miliolina seminulum (Linne) Brady, 1884, p. 157, pl. 5, fig. 6.Quinqueloculina seminulum (Linne) Cushman, 1917, p. 44, pl. 11, fig. 2; Todd and Bronnimann, 1957, p. 27, pl. 3, fig. 9-10; Barbieri and Medioli, 1969, p. 855, pl. 62, fig. 1a, b; Gregory, 1971, p. 187, pl. 6, fig. 1.

Distribution: Restricted to near-shore open marine areas. Not an important component in any assemblage.

PATEORIS Loeblich and Tappan 1953

Pateoris hauerinoides (Rhumbler)

Plate 3. Fig. 2

Quinqueloculina subrotunda (Montagu) forma hauerinoides Rhumbler, 1936, p. 206, 217, 226, fig. 208-212.Pateoris hauerinoides (Rhumbler) Loeblich and Tappan, 1953, p. 42, pl. 6, fig. 8-12; Gregory, 1971, p. 188, pl. 6, fig. 3-4.

Distribution: Very rare occurrence, only once seen greater than 3 %.

Found attached to sea-weed.

SIGMOIOPSIS Findley 1947

Sigmoilopsis schlumbergeri (Silvestri)

Plate 3. Fig. 4

Sigmoilina schlumbergeri Silvestri, 1904, v. 22, p. 267.Sigmoilopsis schlumbergeri (Silvestri) Findley, 1947, v. 28, p. 270;

Barker, 1960, p. 16; pl. 8, fig. 1-4.

Distribution: Not an important component. Found only in deep slope waters.

TRILOCULINA d'Orbigny 1826

Triloculina trihedra Loeblich and Tappan

Plate 3. Fig. 3

Triloculina trihedra Loeblich and Tappan, 1953, p. 45, pl. 4, fig. 10;  
Schnitker, p. 212, pl. 3, fig. 10.

Distribution: Rare. Similar to Q. seminulum.

NODOSARIACEA

NODOSARIDAE Ehrenberg 1838

NODOSARININAE Ehrenberg 1838

ASTACOLUS de Montfort 1808

Astacolus hyalacrulus Loeblich and Tappan

Plate 3. Fig. 5

Astacolus hyalacrulus Loeblich and Tappan, 1953, p. 52, pl. 9, fig. 1-4;  
Cole, p. 58, pl. 17, fig. 48.

Distribution: Probably relict, restricted to the upper slope.

LAGENA Walker and Jacob

Lagena mollis (Cushman)

Plate 3. Fig. 6

Lagena gracillima (Seguenza) var. mollis Cushman, 1944, p. 21., pl. 3,

fig. 3.

Lagena mollis (Cushman) Vilks, 1969, p. 47, pl. 2, fig. 20; Gregory, 1971, p. 195, pl. 7, fig. 4; Cole, 1981, p. 64, pl. 18, fig. 20.

Distribution: Rare occurrences on the upper slope.

GLANDULINIDAE Reuss 1860

GLANDULININAE

GLANDULINA d'Orbigny

Glandulina laevigata (d'Orbigny)

Plate 3. Fig. 7

Nodosaria laevigata d'Orbigny, 1826, p. 252, pl. 10, fig. 1-3.

Glandulina laevigata (d'Orbigny) d'Orbigny, 1846, p. 29, pl. 1, fig. 45; Loeblich and Tappan, 1953, p. 18, pl. 16, fig. 2-5; Gregory, 1971, p. 198, pl. 9, fig. 4; Cole, 1981, p. 73, pl. 19, fig. 2.

Distribution: Restricted to upper slope or central basins.

OOLINIDAE Loeblich and Tappan 1961

FISSURINA Reuss 1850

Fissurina marginata (Montagu)

Plate 3. Fig. 8

Vermiculum marginata Montagu, 1803, p. 524.

Fissurina marginata (Montagu) Loeblich and Tappan, 1953, p. 77, pl. 14, fig. 6-9; Vilks, 1969, p. 48, pl. 2, fig. 2a,b; Gregory, 1971, p. 207, pl. 10, fig. 1; Cole, 1981, p. 81, pl. 19, fig. 25.

Distribution: Rare on shelf and slope.

BULIMINACEA Jones 1875

BOLIVINITIDAE Cushman 1927

BOLIVINA d'Orbigny 1839

Bolivina subspinescens Cushman

Plate 3. Fig. 10

Bolivina subspinescens Cushman, 1922, p. 48, pl. 17, fig. 5; Barker, 1960, p. 108, pl. 52, fig. 24; Cole, 1981, p. 88, pl. 10, fig. 4.

Distribution: A rare species, found mainly in the central basins and upper slope.

Bolivina spathulata (Williamson)

Plate 3 . Fig. 9

Textularia variabilis Williamson var. spathuleata, Williamson, 1858, p. 76, pl. 6, fig. 164-165.

Bolivina spathulata (Williamson) Barker, 1960, p. 106, pl. 52, fig. 20-21; Schnitker, 1971, p. 194, pl. 4, fig. 24.

Distribution: In association with T. angulosa this species forms one of the main components of the upper slope / shelf edge assemblages

Bolivina subaenariensis Cushman

Plate 3. Fig. 11

Bolivina subaenariensis Cushman, 1922, p. 46, pl. 7, fig. 6; Phleger and Parker, 1951, p. 15, pl. 7, fig. 8-10; Cole and Fergerson, 1975, p. 33.

Distribution: A dominant component of the central basin areas. Not so common on the slope. Associated with G. auriculata.

ISLANDIELLIDAE Loeblich and Tappan 1964

## ISLANDIELLA Norvang 1958

Islandiella teretis (Tappan)

Plate 3. Fig. 19

Cassidulina laevigata d'Orbigny. Brady, 1884, p. 428, pl. 54, fig. 1-3.Cassidulina teretis Tappan, 1951, p. 7, pl. 1, fig. 30a-c.Islandiella teretis (Tappan) Vilks, 1969, p. 49, pl. 3, fig. 5.

Distribution: Restricted occurrence, found mainly in central basin and shelf edge assemblages. This species may in fact have been confused with I. norcrossi (Cushman) as figured in Rodrigues et al 1980.

Islandiella islandica (Norvang)

Plate 5. Fig. 18

Cassidulina islandica Norvang, 1945, p. 41, fig. 7, 8.; Leslie, 1965, p. 57, pl. 10, fig. 4a-c.Islandiella islandica Norvang. Vilks; 1969, p. 49, pl. 3, fig. 3; Barbieri and Medioli, 1969, p. 857, pl. 62, fig. 3a-c; Gregory, 1971, p. 213, pl. 11, fig. 1.

Distribution: Restricted to outer sandy banks where it is predominant in the absence of C. lobatulus. Occurs also but in less abundance in nearshore samples.

BULIMINIDAE Jones 1875

BULIMININAE Jones 1875

BULIMINA d'Orbigny 1826

Bulimina marginata d'Orbigny

Plate 3. Fig. 13

Bulimina marginata d'Orbigny, 1826, p. 269, pl. 12, fig. 10-12; Phleger

and Parker, 1951, p. 16, pl. 7, fig. 27-28; Feyling-Hansen, 1964, p. 303, pl. 14, fig. 2-3; Schnitker, 1971, p. 194, pl. 5, fig. 5.

Distribution: Not present in large numbers, restricted to central basin assemblages.

Bulimina aculeata d'Orbigny 1826

Plate 3. Fig. 12

Bulimina aculeata d'Orbigny, 1826, v. 7, p. 269; Cushman, 1944, p. 27, pl. 3, fig. 47; Phleger and Parker, 1951, p. 15, pl. 7, fig. 23; Schnitker, 1971, p. 194, pl. 5, fig. 4.

Distribution: Occurs mainly in the central basins and their southern approaches. Although often occurring in large numbers it is never found living.

Bulimina exilis Brady 1884

Plate 3. Fig. 15

Bulimina elegans var. exilis Brady, 1884, p. 399, pl. 50, fig. 5-6.  
Bulimina exilis Brady. Cushman and Parker, 1940, p. 11, pl. 2, fig. 18-21; Cole, 1981, p. 89, pl. 19, fig. 43.

Distribution: Dominant component of the upper slope assemblage (400--1000m). Also found in the central basins, but of lesser importance.

Bulimina striata d'Orbigny

Plate 3. Fig. 14

Bulimina striata d'Orbigny; Barker, 1960, p. 104, pl. 51, fig. 10-12; Cushman and Parker, 1947, p. 119.

Distribution: Not an important component in any assemblage. Restricted

to lower slope areas and as such is a good depth indicator.

GLOBOBULIMINA Cushman 1927

Globobulimina auriculata (Bailey) 1851

Plate 3. Fig. 17

Bulimina auriculata Bailey, 1851, p. 12, pl. 1, fig. 25-27.

Globobulimina auriculata (Bailey) var. arctica, Hoglund, 1947, p. 254, text-fig. 266, 267, 270, 271.

Globobulimina auriculata arctica (Hoglund) Leslie, 1965, p. 161, pl. 9, fig. 6a-c.

Globobulimina auriculata Cole, 1981, p. 90.

Distribution: One of the most pre-dominant species in the central basins; occurs less frequently in the upper slope areas.

Globobulimina pacifica Cushman

Plate 3. FFig. 16

Globobulimina pacifica Cushman. Galloway and Wissler, 1927, p. 74; Loeblich and Tappan, 1964, p. 560, fig. 442, no. 4.

Distribution: Rare, occurs mainly on the upper slope.

STAINFORTHIA Hofker 1956

Stainforthia concava (Hoglund)

Plate 3. Fig. 18

Virgulina concava Hoglund, 1947, p.257, pl. 23, fig. 3, 4.

Stainforthia concava (Hoglund) Loeblich and Tappan, 1964, c561; Cole, 1981, p. 91, pl. 9, fig. 5.

Distribution: Rare, mainly slope occurrence.

UVIGERINIDAE Haeckel 1894

UVIGERINA d'Orbigny 1826

Uvigerina peregrina Cushman 1923

Plate 4. Fig. 3

Uvigerina peregrina Cushman, 1923, p. 166, pl. 42, fig. 7-10; Schnitker, 1971, p. 212, pl. 5, fig. 11; Cole, 1981, p. 92, pl. 10, fig. 9.

Distribution: Restricted to depths greater than 800 m. A dominant component at these depths.

TRIFARINA Cushman 1923

Trifarina angulosa (Williamson)

PLATE 4. FIG. 1

Uvigerina angulosa Williamson, 1858, p. 67, pl. 5, fig. 140.

Angulogerina angulosa (Williamson) Leslie, 1965, p. 155, pl. 18, fig. 13a-c.

Trifarina angulosa (Williamson) Barbieri and Medioli, 1969, p. 857, pl. 65, fig. 1-5; Cole, 1981, p. 92, pl. 19, fig. 46.

Distribution: Dominant shelf edge component.

Trifarina cf. occidentalis Cushman

Plate 4. Fig. 2

Trifarina cf. occidentalis (Cushman) Todd, 1948, p. 291, pl. 36, fig. 4; Todd and Bronnimann, 1957, p. 36, pl. 9, fig. 5-6.

Distribution: Found in only a few upper slope samples

Most probably a megalospheric form of T. angulosa.

DISCORBACEA

DISCORBIDAE Ehrenberg 1838

DISCORBINAЕ Ehrenberg 1838

BUCELLA Anderson 1952

Buccella frigida (Cushman)

Plate 4. Fig. 4, 5

Pulvinulina frigida Cushman, 1922, p. 12.

Eponides frigidus (Cushman), Cushman, 1941, p. 37, pl. 9, fig. 16-17;

Phleger, 1952, p. 419, pl. 6, fig. 12a-b.

Buccella frigida (Cushman) Anderson, 1952, p. 144, fig. 4, 5, 6; Vilks, 1969, p. 49, pl. 3, fig. 7; Gregory, 1971, p. 220, pl. 12, fig. 1-3.

Distribution: Found only on the shelf but is common in several assemblages. Often found associated with C. lobatulus.

BAGGININAE Cushman 1927

VALVULINERIA Cushman 1927

Valvularia laevigata Phleger and Parker

Plate 4. Fig. 6, 7

Valvularia laevigata Phleger and Parker, 1951, p. 25, pl. 13, fig. 11-12; Schnitker, 1971, p. 212, pl. 6, fig. 6; Cole, 1981, p. 98, pl. 11, fig. 1.

Distribution: Rare occurrence, restricted to the shelf.

ROTALIACEA Ehrenberg 1839

ELPHIDIIDAE Galloway 1933

ELPHIDIUM de Montfort 1808

Elphidium excavatum (Terquem) forma clavata Cushman 1930

Plate 5. Fig. 9

Elphidium incertum (Williamson) var. clavatum Cushman, 1930, p. 20, pl. 7, fig. 10.

Elphidium clavatum Cushman. Buzas, 1965, p. 23, pl. 3, fig. 3-4; SenGupta, 1971, p. 89, pl. 2, fig. 28,29.

Elphidium excavatum forma clavata Cushman. Miller, 1979, p. 27, pl. 1, fig. 3-8,pl. 3, fig. 2-8.

Distribution: Forms a dominant component of the possibly relict assemblage on the shelf and slope. For a complete discussion of this variable form see Miller *et al.* 1982.

Elphidium bartletti (Cushman)

Plate 5. Fig. 10

Elphidium bartletti Cushman, 1933, p. 4, pl. 1, fig. 9; Leslie, 1965, p. 160, pl. 8, fig. 10a-b; Vilks, 1969, p. 50, pl. 3, fig. 12; Cole, 1981, p. 100, pl. 20, fig. 1.

Distribution: Common on the shelf edge and in the central basin area.

Elphidium subarcticum Cushman

Plate . Fig.

Elphidium subarcticum Cushman 1944, p. 27, pl. 3, fig. 34,35; Schnitker, 1971, p. 198, pl. 7, fig. 3; Cole 1981, p. 101, pl. 20, fig. 4.

Distribution: Rare, occurs mainly on the shelf.

## ELPHIDIELLA Cushman 1936

Elphidiella arctica(Parker and Jones)

## Plate 5. Fig. 7

Polystomella arctica Parker and Jones, p. 471, pl. 48, fig. 18Elphidiella arctica (Parker and Jones) Cushman, 1939, p. 65, pl. 18, fig. 11-14; Phleger, 1952, p. 83, pl. 14, fig. 14; Leslie, 1965, p. 159, pl. 8, fig. 12a-b.Distribution: Commonly occurs associated with C. lobatulus on shallow banks and nearshore.

## ORBITOIDACEA

## CIBICIDIDAE Cushman 1927

## PLANULININAE Bermudez 1952

## PLANULINA d'Orbigny 1826

Planulina wuellerstorfi (Schwager)

## Plate 4. Fig. 8, 9

Anomalina wuellerstorfi Schwager, 1886, p. 258, pl. 7, fig. 107, 105.Planulina wuellerstorfi (Schwager) Cushman, 1929, p. 104, pl. 15, fig. 1-2; Barker, 1960, pl. 93, fig. 9; Cole, p. 103, pl. 12, fig. 9

Distribution: Not a dominant species. Only found at deep (+ 2000m )slope depths.

CIBICIDINAE Cushman 1927

CIBICIDES de Montfort 1808

Cibicides lobatulus (Walker and Jacob)

Plate 4. Fig. 10, 11

Nautilus lobatulus Walker and Jacob, 1798, p. 642, pl. 14, fig. 36

Cibicides lobatulus (Walker and Jacob) Parker, 1952, p. 427, pl. 6, fig. 26; Feyling-Hansen, 1964, p. 339, pl. 19, fig. 1-3; Vilks, 1969, p. 50, pl. 3, fig. 17; Cole, 1981, p. 104, pl. 12, fig. 4.

Distribution: Predominant component of near-shore assemblages and shallow banks. Its variable morphology is a result of its attached nature.

Cibicides pseudoungerianus (Cushman)

Plate 4. Fig. 17, 18

Truncatulina pseudoungerianus Cushman, 1922, p. 97, pl. 20, fig. 9.

Cibicides pseudoungerianus (Cushman) Feyling-Hansen, 1964, p. 340, pl. 19, fig. 4-6; Schafer and Cole, 1978, p. 22; Schnitker 1971, p. 196, pl. 9, fig. 7.

Distribution: Not a dominant species, main occurrence is in the central basin area and upper slope.

CASSIDULINACEA

CAUCAUSINIDAE Bykova 1959

FURSENKOININAE Loeblich and Tappan 1961

FURSENKOINA Loeblich and Tappan 1961

Furstenkoina fusiformis (Williamson)

## Plate 5. Fig. 1

Bulimina pupoides d'Orbigny var. fusiformis Williamson, 1858, p. 63, pl. 5, fig. 129, 130.

Virgulina fusiformis (Williamson) Feyling-Hanssen, 1964, p. 307, pl. 14, fig. 15-18.

Furstenkoina fusiformis (Williamson) Schnitker, 1971, p. 200, pl. 10, fig. 1; Cole, 1981, p. 106, pl. 14, fig. 1.

Distribution: Found mainly on the upper slope and central basin area.

NONIONIDAE Schultz 1854

NONIONINAE Schultz 1854

ASTRONONION Cushman and Edwards 1937

Astrononion gallowayi Loeblich and Tappan

Plate 4. Fig. 12

Astrononion gallowayi Loeblich and Tappan, 1953, p. 90, pl. 17, fig. 4-7; Vilks, 1969, p. 51, pl. 3, fig. 19; Cole, 1981, p. 109, pl. 13, fig. 6.

Distribution: Not a common species, it is usually found in shelf edge or upper slope environments.

NONIONELLINA Voloshinova 1958

Nonionellina labradorica (Dawson)

Plate 4. Fig. 14, 15

Nonionina labradorica Dawson, 1860, p. 191, fig. 4.

Nonion labradorica (Dawson) Cushman, 1927, p. 148, pl. 2, fig. 7-8.

Nonionellina labradorica (Dawson) Schafer and Cole, 1978, p. 28, pl. 9,

fig. 4; Cole, 1981, p. 110, pl. 13, fig. 20.

Distribution: A dominant component of the central basin area.

NONIONELLA Cushman 1926

Nonionella turgida (Williamson)

Plate 4. Fig. 13

Rotalina turgida Williamson, 1858, p. 50, pl. 4, fig. 95-97.

Nonionella turgida (Williamson) Cushman, 1930, p. 15, pl. 6, fig. 1-4;  
Cole, 1981, p. 110, pl. 13, fig. 3.

Distribution: Occurs mainly on the upper slope areas.

PULLENIA Parker and Jones 1862

Pullenia quinqueloba (Reuss)

Plate 4. Fig. 16

Nonionina quinqueloba Reuss, 1851, p. 71, pl. 5, fig. 31.

Pullenia quinqueloba (Reuss) Schnitker, 1971, p. 206, pl. 10, fig. 11;  
Cole, 1981, p. 111, pl. 14, fig. 6.

Distribution: Occurs mainly in the central basins but is also found in  
slope samples.

ALABAMINIDAE Hofker 1951

GYROIDINA d'Orbigny 1826

Gyroidina soldani d'Orbigny

Plate 5. Fig. 5, 6

Rotalia soldani d'Orbigny, p. 278, fig. 36.

Gyroidina soldani d'Orbigny var. altiformis Stewart. Phleger and Parker, 1952, p. 22, pl. 11, fig. 15-16.

Gyroidina soldani d'Orbigny. Cole, 1981, p. 112, pl. 14, fig. 7.

Distribution: Shelf edge and upper slope environments.

ORIDORSALIS Andersen 1961

Oridorsalis umbonatus (Reuss)

Plate 5. Fig. 3, 4

Rosalina umbonata Reuss, 1857, p. 75, pl. 5, fig. 35.

Eponides umbonatus (Reuss) Parker, 1952, p. 419, pl. 6, fig. 13.

Eponides tener (Brady) Cushman. Todd and Postma, 1954, p. 359.

Oridorsalis umbonatus (Reuss) Todd, 1965, p. 23, pl. 6, fig. 2; Cole, 1981, p. 113, pl. 14, fig. 8.

Distribution: Relatively rare occurrence in central basin and deeper slope areas.

OSANGULARIIIDAE Loeblich and Tappan 1964

OSANGULARIA Brotzen 1940

Osangularia rugosa (Phleger and Parker)

Plate 5. Fig. 13, 14

Pseudoparrella rugosa Phleger and Parker, 1951, p. 28, pl. 15, fig. 8-9.

Osangularia rugosa (Phleger and Parker) Pfleiderer and Frerichs, 1976, pl. 7, fig. 2-4; Cole, 1981, p. 114, pl. 20, fig. 12, 13.

Distribution: Restricted to rare deep water occurrences.

ANOMALINIDAE Cushman 1927

DISCANOMALINA Asano 1951

Discanomalina semipunctata (Bailey)

Plate 5. Fig. 15

Rotalina semipunctata Bailey, 1851, p. 1-15, pl. 11, fig. 17-19.

Anomalina semipunctata (Bailey) Cushman, 1931, p. 106, pl. 18, fig. 1-2.

Discanomalina coronata (Parker and Jones) Asano, 1951, p. 13, fig. 1-2.

Discanomalina semipunctata (Bailey) Medioli and Scott, 1979, p. 298, pl. 1,2.

Distribution: Restricted to s.w. areas of the shelf, may be locally very abundant.

MELONIS de Montfort 1808

Melonis pompilioides (Fichel and Moll)

Plate 5. Fig. 8

Nautilus pompilioides Fichel and Moll, 1798, p. 31, pl. 2, fig. a-c.

Nonionina pompilioides (Fichel and Moll) Parker, Jones and Brady, 1865, p. 18, pl. 3, fig. 98.

Melonis pompilioides (Fichel and Moll) Voloshinova, 1958, p. 117-119; Cole 1981, p. 114, pl. 14, fig. 9.

Distribution: Restricted to deep waters, rare.

CASSIDULINA d'Orbigny 1826

Cassidulina neocarinata Thalmann

Plate 5. Fig. 17.

Cassidulina laevigata d'Orbigny.var. carinata Cushman, 1922, p. 124, pl. 25, fig. 6,7.

Cassidulina neocarinata Thalmann, 1950, p. 44, nom. nov. for C. laevigata d'Orbigny var. carinata Cushman, 1922. Rodrigues et al 1980, p. 58, pl. 5, fig. 2, 5, 8, pl. 6, fig. 3,4.

Distribution: Rare, found mainly in the central basins.

ROBERTINACEA Reuss 1850

CERATOBULIMINIDAE Cushman 1927

EPISTOMININAE Wedekind 1937

HOEGLUNDINA Brotzen 1948

Hoeglundina elegans (d'Orbigny)

Plate 5. Fig. 11, 12

Rotalia elegans d'Orbigny, 1826, p. 276, no. 54.

Epistomina elegans (d'Orbigny) Cushman, 1931, p. 65, pl. 13, fig. 6.

Hoeglundina elegans (d'Orbigny) Cole, 1981, p. 116, pl. 14, fig. 10.

Distribution: Not a common species, found in deep slope environments (+2000m) and contrary to the literature at shelf depths in the central basins.

## Listing of rare species.

Ammobaculites foliaceus (Brady) Cushman, 1920, v. 104, p. 64.

Astrorhiza arenaria Carpenter. Cole, 1981, p. 3, pl. 2, fig. 6.

Astrononion stellatum Cushman and Edwards, 1937, p. 32, pl. 3, fig.

9-11.

Bottellina labyrinthica Brady. Cole, 1981, p. 11, pl. 3, fig. 7.

Cibicides refulgens Montfort. Barker, 1960, p. 190, pl. 92, figs. 7-9.

Cystammina pauciloculata (Brady) Barker, 1960, p. 84, pl. 41, fig. 1-2.

Cassidella complanata (Egger) Parker, 1952, p. 417, pl. 6, fig. 1-2

Dentalina frobisherensis Loeblich and Tappan, 1953, p. 55, pl. 10, fig. 1-9.

Dentalina ittai Loeblich and Tappan, 1953, p. 56, pl. 10, fig. 10-12.

Fissurina serrata (Schlumberger) Cole, 1981, p. 83, pl. 19, fig. 31.

Fissurina submarginata (Broomgaart) Barker, 1960, p. 122, pl. 59, fig.

21-22.

Glomospira charoides (Jones and Parker) Barker, 1960, p. 78, pl. 38, fig. 10-16.

Islandiella norcrossi (Cushman) Loeblich and Tappan, 1953, p. 20, pl. 24, fig. 2.

Lagena gracillima (Seguenza) Loeblich and Tappan, 1953, p. 60, pl. 11, fig. 1-4.

Lagena apliopleura Loeblich and Tappan, 1953, p. 59, pl. 10, fig. 14, 15.

Lagena distoma Parker and Jones. Barker 1960, p. 119, pl. 58, fig. 11-15.

Lagena elongata (Ehrenberg) Barker, 1960, p. 116, pl. 56, fig. 27-29.

Lagena laevis Loeblich and Tappan, 1953, p. 61, pl. 11, fig. 5-8.

Liebusella soldani (Jones and Parker) Barker, 1960, p. 66, pl. 32, fig.

14-18.

Lenticulina peregrina (Schwager) Barker, 1960, p. 144, pl. 68, fig.

11-16.

Miliammina fusca (Brady) Scott and Medioli, 1980, p. 40, pl. 2, fig.

1-3.

Marsipella cylindrica Brady. Barker, 1960, p. 48, pl. 24, fig. 20-22.

Nodosaria pyrula d'Orbigny. Chapman and Parr, p. 60.

Oolina melo d'Orbigny. Loeblich and Tappan, 1953, p. 73, pl. 13, fig.

9-10.

Oolina borealis Loeblich and Tappan, 1953, p. 68, pl. 13, fig. 4-6

Marginulina obesa Cushman. Barker, 1960, p. 136, pl. 65, fig. 5-6.

Pyrgo murrhyna (Schwager) Barker, 1960, p. 4, pl. 2, fig. 10, 11, 15.

Patellina corrugata Williamson. Barker, 1960, p. 178, pl. 86, fig. 1-7.

Quinqueloculina agglutinata Cushman. Loeblich and Tappan, 1953, p. 39, pl. 15, fig. 1-4.

Rhabdammina abyssorum Carpenter. Cole, 1981, pl. 16, fig. 3.

Rhizammina indivisa (Brady) Cole, 1981, pl. 2, fig. 16.

Robulus thalmanni Hessland. Barker, 1960, p. 144, pl. 69, fig. 13.

Rosalina coumbiensis (Cushman) Uchio, 1960, p. 66, pl. 8, fig. 1-2.

Robertinoides charlottensis (Cushman) Parker, 1952, p. 416, pl. 5, fig. 3a-c.

Saccorhiza ramosa (Brady) Cole, 1981, p. 11, pl. 2, fig. 7.

Saracenaria italicica Defrance. Barker, 1960, p. 142, pl. 67, fig. 9.

Technitella legumen Norman. Barker, 1960, p. 50, pl. 25, fig. 8-12.

Trochammina nana (Brady) Barker, 1960, p. 72, pl. 35, fig. 6-8.

Trochammina nitida Brady. Barker, p. 84, pl. 41, fig. 5-6.

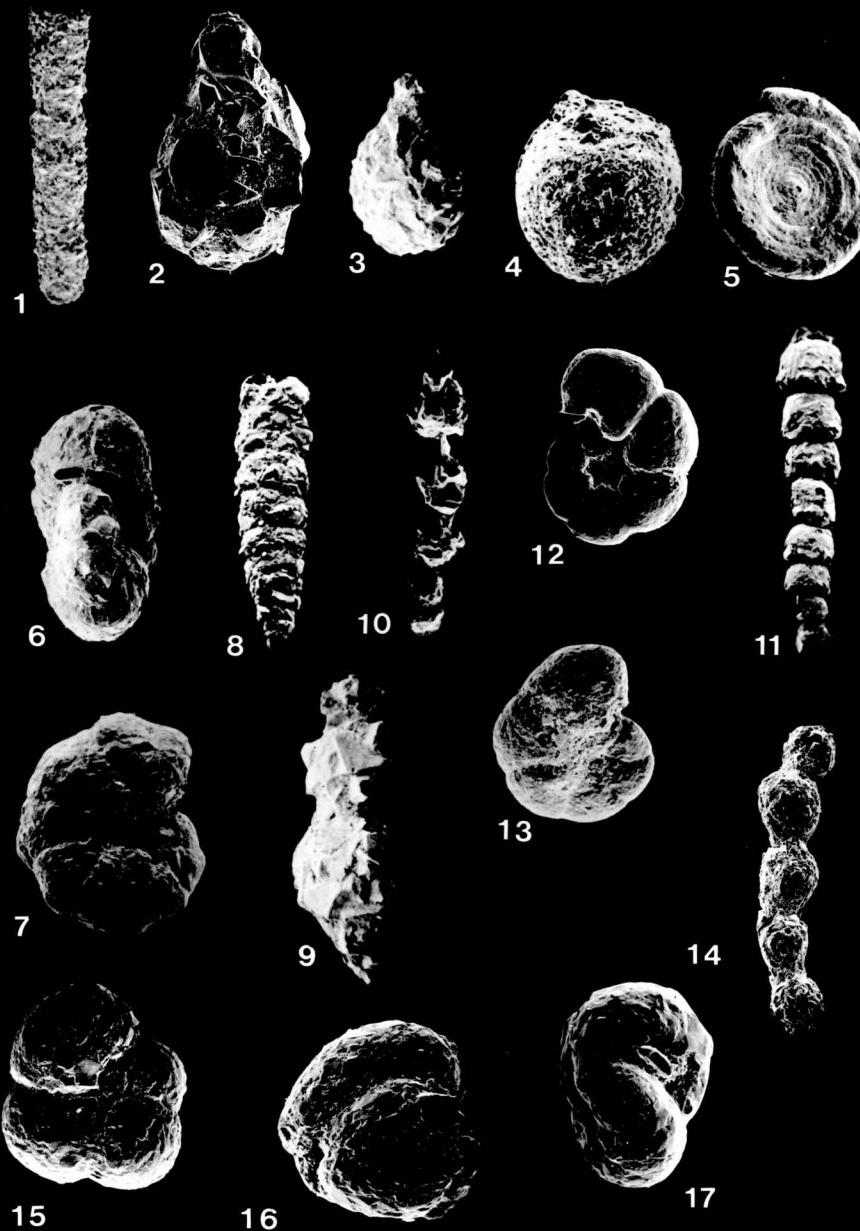
Tolypammina schandini Rhumbler. Barker, 1960, p. 48, pl. 24, fig. 6-9.

Vaginulina spinigera Brady. Barker, 1960, p. 42, pl. 65, fig. 13, 14.

## PLATE 1.

- Figure 1      Hyperammina elongata Brady. x 21
- Figure 2      Saccammina atlantica (Cushman) x 74
- Figure 3      Saccammina difflugiformis (Brady) x53
- Figure 4      Saccammina sphaerica Brady. x113
- Figure 5      Ammodiscus catinus Hoglund. x43.5
- Figure 6      Cribrostomoides crassimargo (Norman) x36
- Figure 7      Cribrostomoides crassimargo (Norman) x50
- Figure 8      Reophax arctica (Brady) x145
- Figure 9      Reophax scorpiurus (Montfort) x47
- Figure 10      Reophax guttifer (Brady) x69.5
- Figure 11      Reophax scotti Chaster. x180.5
- Figure 12      Cribrostomoides jeffreysi (Williamson) x103
- Figure 13      Haplophragmoides canariensis (d'Orbigny) x75
- Figure 14      Reophax nodulosa Brady. x41.2
- Figure 15      Cribrostomoides subglobosum (Sars) x 57
- Figure 16      Adercotryma glomerata (Brady) x146
- Figure 17      Recurvooides turbinatus (Brady) x146

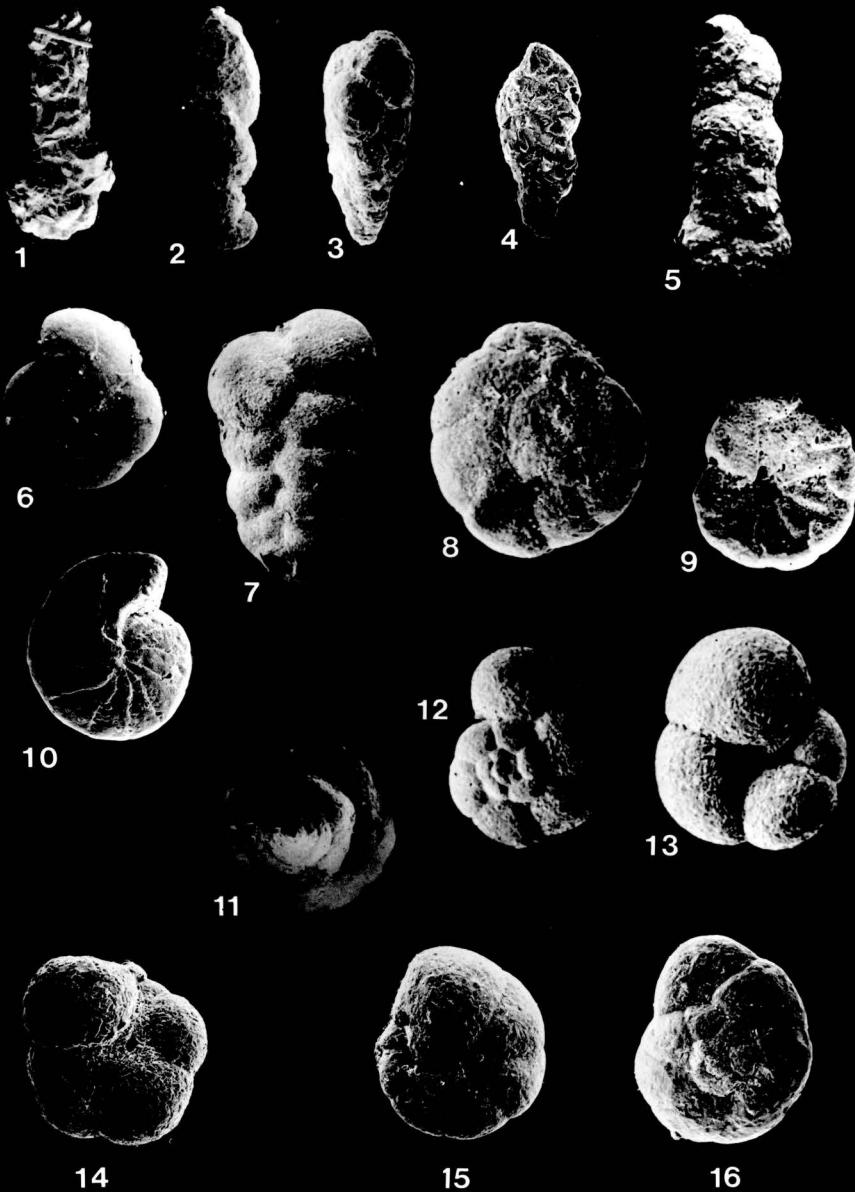
PLATE I



## PLATE 2.

- Figure 1      Ammobaculites dilatatus Cushman. x 36.5
- Figure 2      Ammotium cassis (Parker) x 25
- Figure 3      Eggerella advena (Cushman) x 138
- Figure 4      Textularia torquata Parker. x 136
- Figure 5      Spiroplectammina biformis (Parker and Jones) x 133
- Figure 6      Alveophragmium scitulum (Brady) x 135
- Figure 7      KarerIELLA bradyi (Cushman) x 73.5
- Figure 8      Trochammina squamata Parker and Jones. x 171
- Figure 9      Trochammina squamata Parker and Jones. x 120
- Figure 10      Cyclammina cancellata Brady. x 20
- Figure 11      Tritaxis conica Parker and Jones. x 46.5
- Figure 12      Trochammina inflata (Montagu) x 52
- Figure 13      Trochammina inflata (Montagu) x 146
- Figure 14      Trochammina globigeriniformis (Parker and Jones) x 96
- Figure 15      Trochammina lobata Cushman. x 102
- Figure 16      Trochammina lobata Cushman. x 93

PLATE 2



## PLATE 3.

- Figure 1            Quinqueloculina seminulum (Linne) x 50.5
- Figure 2            Pateroris hauerinoides (Rhumbler) x 67.5
- Figure 3            Triloculina trihedra Loeblich and Tappan. x 89
- Figure 4            Sigmoilopsis schlumbergeri (Silvestri) x 42.5
- Figure 5            Astracolus hyalancrulus Loeblich and Tappan. x 36
- Figure 6            Lagena mollis (Cushman) x 24.5
- Figure 7            Glandulina laevigata (d'Orbigny) x 60
- Figure 8            Fissurina marginata (Montagu) x 170
- Figure 9            Bolivina spathulata (Williamson) x 133
- Figure 10          Bolivina subspinescens Cushman. x 163
- Figure 11          Bolivina subaenariensis Cushman. x 43
- Figure 12          Bulimina aculeata d'Orbigny. x 74
- Figure 13          Bulimina marginata d'Orbigny. x 102.5
- Figure 14          Bulimina striata d'Orbigny. x 105.5
- Figure 15          Bulimina exilis Brady. x 44.5
- Figure 16          Globobulimina pacifica Cushman. x 53
- Figure 17          Globobulimina auriculata (Bailey) x 121.5
- Figure 18          Stainforthia concava (Hoglund) x 110.5
- Figure 19          Islandiella teretis (Tappan) x 70

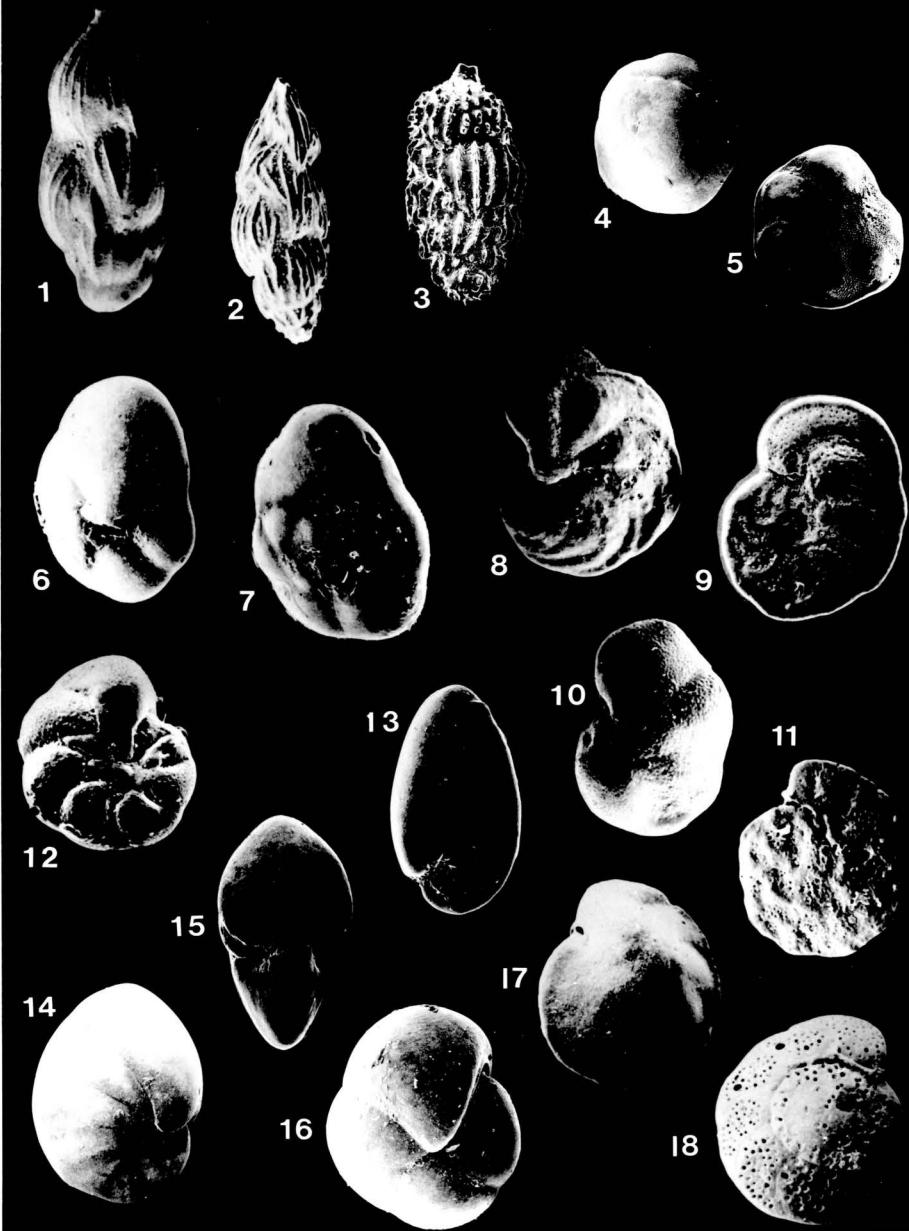
PLATE 3



## PLATE 4.

- Figure 1      Trifarina angulosa (Williamson) x 91
- Figure 2      Trifarina occidentalis Cushman. x 106
- Figure 3      Uvigerina peregrina Cushman. x 58
- Figure 4      Buccella frigida (Cushman) x 60
- Figure 5      Buccella frigida (Cushman) x 60
- Figure 6      Valvularia laevigata Phleger and Parker. x 135
- Figure 7      Valvularia laevigata Phleger and Parker. x 136
- Figure 8      Planulina wuellerstorfi (Schwager) x 63
- Figure 9      Planulina wuellerstorfi (Schwager) x 57.5
- Figure 10      Cibicides lobatulus (Walker and Jacob) x 58
- Figure 11      Cibicides lobatulus (Walker and Jacob) x 121.5
- Figure 12      Astrononion gallowavi Loeblich and Tappan. x 109
- Figure 13      Nonionella turgida (Williamson) x 128.5
- Figure 14      Nonionellina labradorica (Dawson) x 54.5
- Figure 15      Nonionellina labradorica (Dawson) x 60.5
- Figure 16      Pullenia quinqueloba (Reuss) x 82.5
- Figure 17      Cibicides pseudoungerianus (Cushman) x 56.5
- Figure 18      Cibicides pseudoungerianus (Cushman) x 81

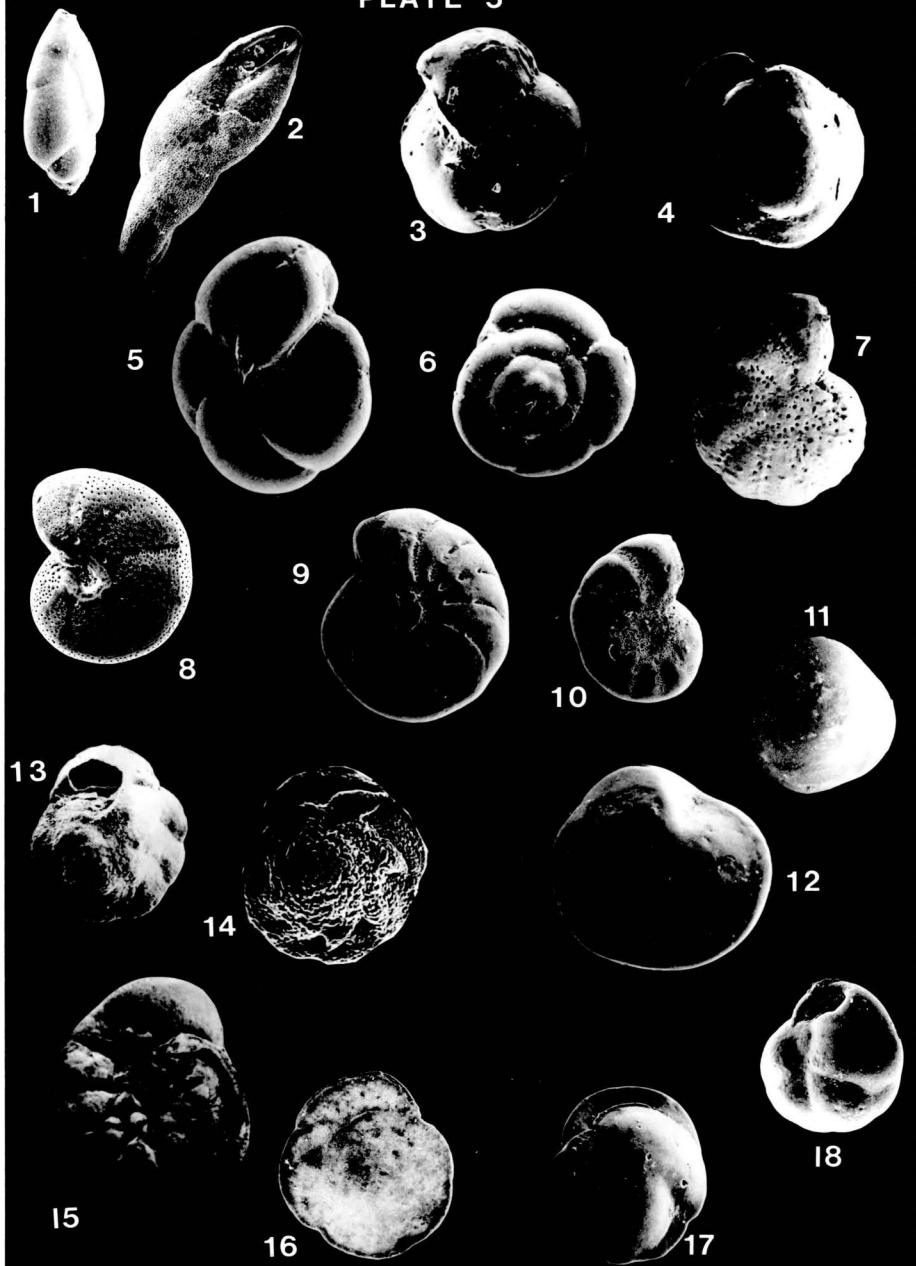
PLATE 4



## PLATE 5

- Figure 1 Furstenkoina fusiformis (Williamson) x 115.5  
Figure 2 Furstenkoina loeblichii (Williamson) x 170  
Figure 3 Oridorsalis umbonatus (Reuss) x 96  
Figure 4 Oridorsalis umbonatus (Reuss) x 114  
Figure 5 Gyroidina soldani d'Orbigny. x 117.5  
Figure 6 Gyroidina soldani d'Orbigny. x 157  
Figure 7 Elphidiella arctica (Parker and Jones) x 31  
Figure 8 Melonis pompilioides (Fichel and Moll) x 71  
Figure 9 Elphidium excavatum f. clavatum Cushman. x 65  
Figure 10 Elphidium bartletti (Cushman) x 77.5  
Figure 11 Hoeglundina elegans (d'Orbigny) x 60  
Figure 12 Hoeglundina elegans (d'Orbigny) x 47.5  
Figure 13 Osangularia rugosa (Phleger and Parker) x 156  
Figure 14 Osangularia rugosa (Phleger and Parker) x 109  
Figure 15 Discanomalina semipunctata (Bailey) x 36  
Figure 16 Eponides spp. x 130.5  
Figure 17 Cassidulina laevigata d'Orbigny. x 118.5  
Figure 18 Islandiella islandica (Norvang) x 62.5

PLATE 5



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## APPENDIX I

### SAMPLE LOCATION AND DEPTH

STATION LATS. AND LONGS. AS USED IN FACTOR ANALYSIS OF FORAM. DATA.  
SAMLOC.

SAMPLE NO.	LATITUDE	LONGITUDE	DEPTH
7902	42 46.5	63 25.0	602
7903	42 45.6	63 25.0	700
7904	42 43.3	63 24.5	798
7906	42 41.4	63 25.0	1000
7907	42 39.0	63 24.5	1200
7908	42 37.6	63 25.0	1400
7909	42 32.0	63 25.0	1600
7910	42 26.2	63 25.0	1800
7911	42 09.0	63 25.0	2400
7912	42 13.5	63 25.0	2200
7913	42 19.9	63 25.0	2000
7914	43 45.9	63 04.8	271
7917	43 46.3	62 54.0	251
7918	43 46.0	62 51.7	240
7920	43 46.4	62 46.7	220
7947	42 47.6	61 53.5	1200
7948	42 40.2	61 53.5	1680
7949	42 37.0	61 53.5	1800
7950	42 34.2	61 35.5	2000
7951	42 32.2	61 54.0	2200
7952	42 29.8	61 54.0	2400
7953	42 51.1	63 17.7	170
7955	42 58.7	63 21.9	156
7956	43 04.7	63 29 1	166
7957	43 09.5	63 32.0	140
7958	43 13.5	63 34.6	175
7959	43 20.4	63 33.2	190
7960	43 24.3	63 42.7	220
0802	43 22.5	61 27.5	18
0817	45 21.8	60 48.5	110
0818	45 21.0	60 44.0	70
0819	45 19.5	60 37.5	113
0820	45 18.5	60 30.5	163
0821	45 18.5	60 13.5	119
0822	45 13.0	59 59.3	137
0824	45 09.1	59 38.1	102
0825	45 51.3	60 08.6	16
0830	45 50.3	60 02.5	59
0832	45 48.7	60 01.3	29
0833	45 25.6	61 02.1	31
0835	45 25.2	61 07.7	51
0838	45 24.5	61 16.6	27
0839	45 24.1	61 18.8	33
0841	45 23.3	61 24.4	33
0843	45 19.0	60 36.5	73
0844	45 10.6	60 46.0	174
0846	45 52.8	60 59.0	238
0847	44 44.0	61 07.5	82
0848	44 34.8	61 11.3	104
0849	44 27.9	61 17.5	128
0850	44 20.5	61 20.6	196

0851	44 15.0	56 12.3	200
0852	44 08.5	61 34.3	108
0856	43 40.0	62 12.5	101
0857	43 28.0	62 22.5	86
8002	43 47.5	62 45.1	214
8004	43 24.8	63 20.0	206
8005	43 24.0	63 20.0	203
8006	43 23.0	63 20.0	197
8007	43 22.0	63 20.0	190
8010	43 22.5	63 20.0	195
8011	43 22.5	63 20.0	195
8012	43 22.5	63 20.0	195
8013	42 05.0	63 30.3	2500
8014	41 48.5	63 28.7	3050
8017	42 50.2	63 27.0	180
8018	42 48.7	63 28.0	195
8027	42 50.2	63 22.2	185
8035	45 02.6	60 42.1	135
8036	45 04.5	60 44.0	114
8037	45 06.4	60 45.8	154
8038	45 06.6	60 47.6	167
8039	45 09.8	60 49.6	156
8040	45 10.9	60 50.4	135
8042	45 12.7	60 52.2	103
8043	45 13.5	60 54.0	60
8044	45 14.0	60 53.0	76
8060	43 12.3	60 59.5	492
8061	43 10.8	61 00.4	648
8062	43 09.8	61 00.5	775
8063	43 07.6	61 00.9	990
8064	43 04.2	61 00.7	1708
8065	43 02.3	61 05.8	2120
8066	43 09.5	61 09.7	1350
8067	43 11.0	61 09.0	1050
8068	43 12.3	61 09.5	781
7802	44 08.4	58 06.1	834
6003	44 03.6	63 30.5	156
6005	43 45.0	63 30.0	175
6006	43 30.1	63 30.2	160
6007	43 20.5	63 29.9	167
6018	42 48.5	63 27.5	230
6032	42 43.5	63 22.1	750
6033	42 42.5	63 22.5	970
6034	42 42.5	63 22.6	1045
6036	42 46.9	63 21.8	450
6041	42 47.8	63 26.2	280
6042	42 47.5	63 26.5	280
6050	42 47.7	63 27.5	350
6051	42 47.3	63 25.5	380
6052	44 10.7	62 50.2	170
6053	44 21.4	62 51.5	152
6055	44 24.5	62 52.0	100
6056	44 28.5	62 52.0	115
6057	44 31.5	62 52.0	88

6058	44 34.3	62 50.3	75
6062	43 54.2	62 06.6	165
6063	43 44.1	64 11.5	185
6064	43 39.1	64 14.5	180
6065	43 34.3	64 12.2	190
6066	43 30.9	64 07.8	125
6068	43 22.7	64 06.5	97
6069	43 13.8	64 06.3	87
6070	43 01.4	64 11.8	95
6071	42 55.0	64 14.4	116
6072	42 48.1	64 18.9	105
6073	42 40.9	64 20.5	140
6074	42 37.3	64 22.8	280
6075	42 37.3	64 22.8	300
6076	42 31.2	64 26.6	770
6077	42 31.3	64 27.3	650
6078	42 33.4	64 28.3	400
6079	42 39.8	64 30.3	130
6080	42 45.7	64 33.4	110
6081	42 55.8	64 38.8	100
6082	43 04.8	64 42.6	94
6083	43 09.7	64 54.7	123
6084	43 17.4	64 57.9	140
6085	43 17.8	64 58.4	160
6086	43 19.2	64 58.4	155
6087	43 19.2	64 55.1	142
6088	43 22.1	64 53.3	128
6089	43 23.6	64 49.6	110
6090	43 26.4	64 43.7	56
6092	43 16.2	64 59.6	161
6094	43 16.7	64 59.4	155
6096	43 19.4	65 10.0	150
6097	43 36.7	64 36.4	141
6098	43 31.8	64 29.0	144
6099	44 03.2	64 12.8	113
6100	44 03.0	64 12.8	112
6101	44 02.2	64 13.7	110
6102	44 04.5	64 12.2	103
6106	44 15.2	64 10.6	66
6108	43 57.1	64 09.2	150
6109	43 57.5	64 09.2	149
6111	43 51.6	64 00.0	205
6112	43 49.4	64 03.9	228
6113	43 50.5	64 04.0	170
6114	43 52.2	64 05.8	168
6115	43 54.7	64 08.5	155
6116	43 55.1	64 08.9	150
6117	43 55.7	64 09.5	155
6110	43 57.3	64 09.5	148
1008	42 56.0	62 15.0	360
1009	43 00.0	62 07.0	130
1010	43 07.0	61 56.5	100
1011	43 22.0	61 40.0	77
1015	43 44.0	60 04.0	60

1016	43 44.0	59 27.0	140
1017	43 39.0	59 21.2	200
1018	43 38.6	59 20.0	280
1020	43 33.5	59 18.1	320
1021	43 32.7	59 19.5	440
1022	43 43.0	59 19.0	120
1026	44 11.0	59 19.0	260
1027	44 14.4	59 19.1	235
1028	44 20.0	50 20.0	215
1029	44 23.0	59 20.5	120
1034	44 13.0	58 12.0	129
1035	44 13.0	58 10.5	200
1036	44 11.9	58 08.0	440
1037	44 11.5	58 07.5	625
1038	44 10.7	58 06.8	775
1039	44 09.9	58 05.6	1000
1040	44 07.2	57 58.7	1600
1043	44 41.0	57 23.0	63
1044	44 44.0	57 20.0	70
1045	44 44.3	57 19.0	160
1046	44 44.2	57 18.6	240
1047	44 44.3	57 16.7	300
1048	44 41.1	57 15.5	380
1051	45 04.5	57 46.0	120
1053	45 07.0	57 52.5	60
1054	45 10.0	57 54.0	72
1055	45 13.0	57 55.5	175
1056	45 15.5	58 02.0	84
1057	45 20.0	58 04.0	220
1058	45 23.5	58 09.0	160
1059	45 27.0	58 12.0	100
1060	45 29.5	58 14.5	180
1062	45 32.0	58 21.5	100
1063	45 34.5	58 24.0	113
1064	45 35.5	58 24.0	340
1065	45 40.0	58 30.0	360
1066	45 51.5	58 34.5	80
1067	45 46.0	58 43.0	270
1068	45 51.0	58 52.0	200
1069	45 54.5	58 52.0	140
1070	45 56.0	59 02.0	105
1071	45 57.5	59 08.0	70
1072	45 54.0	59 25.0	55
1073	45 47.5	59 23.0	120
1074	45 42.0	59 13.0	160
1075	45 39.0	59 03.5	275
1076	45 28.0	59 02.0	145
1077	45 22.5	58 56.0	80
1078	45 18.0	58 51.0	80
1080	45 06.0	58 43.5	75
1081	45 00.0	58 39.6	145
1082	44 54.1	58 29.8	91
1083	44 48.0	58 38.2	200
1084	44 43.0	58 41.0	100
1085	44 43.0	58 49.0	85

1086	44 48.0	58 51.0	260
1087	44 50.5	58 52.0	200
1088	44 52.5	58 55.0	300
1089	44 48.0	59 01.0	310
1090	44 43.0	59 05.0	280
1093	44 52.5	59 10.0	209
1095	44 53.5	59 18.5	100
1096	44 52.0	59 23.9	166
1097	44 48.5	59 29.5	240
1099	44 42.0	59 44.0	100
1100	44 40.0	59 47.0	155
1101	44 43.5	59 53.0	80
1102	44 48.0	59 55.0	234
1103	44 50.9	59 51.6	130
1104	44 58.0	59 55.0	150
1105	44 54.0	60 00.0	210
1106	44 52.5	60 02.5	170
1107	44 42.0	60 02.0	210
1109	44 38.0	59 56.0	90
1110	44 34.0	59 53.0	110
1111	44 24.0	59 49.0	140
1112	44 20.5	59 53.0	140
1113	44 13.0	59 59.0	113
1114	44 09.0	59 49.0	135
1115	44 11.0	59 56.0	130
1116	44 11.5	60 10.0	220
1119	44 13.9	60 37.0	140
1120	44 17.0	60 46.0	160
1124	44 18.0	61 11.0	83
1125	44 34.0	61 16.0	105
1126	44 44.0	61 19.0	112
1127	44 48.0	61 22.0	92
1128	44 53.0	61 30.0	130
1130	44 53.0	61 40.0	105
1131	44 45.3	61 38.0	118
1132	44 35.0	61 37.5	165
1133	44 25.0	61 37.0	171
1134	44 21.0	61 51.0	160
1135	44 17.0	61 57.0	170
1136	44 27.5	61 56.0	170
1137	44 39.0	61 56.0	135
1138	44 43.5	61 09.5	115
1139	44 38.5	62 18.5	155
1140	44 22.0	62 22.0	230
1141	44 19.5	62 19.0	160
1142	44 18.0	62 18.4	80
1143	44 14.0	62 23.0	92
1118	44 12.0	60 27.5	125
4003	43 53.5	64 33.5	77
4004	43 36.5	64 59.5	55
4005	43 21.0	65 23.2	55
4006	43 14.3	65 49.0	44
4007	43 14.6	66 24.6	83

## APPENDIX II

SHELF VARIMAX FACTOR MATRIX AND FACTOR SCORES

## SPECIES ABBREVIATIONS AS USED IN THE ANALYSES

ALVESCIT	<u>Alveophragmium scitulum</u>
AMMOBACU	<u>Ammobaculites dilatatus</u>
AMMOCATI	<u>Ammodiscus catinus</u>
AMMOTCAS	<u>Ammotium cassis</u>
ASTRGALO	<u>Astrononion gallowayi</u>
BOLISUBS	<u>Bolivina subspinescens</u>
BULIEXIL	<u>Bulimina exilis</u>
BULIACLE	<u>Bulimina aculeata</u>
BULISTRI	<u>Bulimina striata</u>
BUCCFRIG	<u>Buccella frigida</u>
BRIZSPAT	<u>Bolivina spathulata</u>
BRIZSUBA	<u>Bolivina subaenariensis</u>
CIBILOBA	<u>Cibicides lobatulus</u>
CIBIPSEU	<u>Cibicides pseudoungerianus</u>
CIBIWUEL	<u>Planulina wuellerstorfi</u>
CASSLAEV	<u>Cassidulina laevigata</u>
CRIBCRAS	<u>Cribrostomoides crassimargo</u>
CRIBJEFF	<u>Cribrostomoides jeffreysi</u>
CRIBSUBG	<u>Cribrostomoides subglobosum</u>
CYCLCANC	<u>Cyclammina cancellata</u>
ELPHEXCL	<u>Elphidium excavatum f. clavata</u>
ELPHBART	<u>Elphidium bartletti</u>
ELPHSUBA	<u>Elphidium subarcticum</u>
EGGLADVA	<u>Eggerella advena</u>
EPONUMBO	<u>Eponides sp.</u>
FISSERAT	<u>Fissurina serrata</u>

FURSLOEB	<u>Furstenkoina loeblichii</u>
FURSFUSI	<u>Furstenkoina fusiformis</u>
GLANLAEV	<u>Glandulina laevigata</u>
GYROSOLD	<u>Gyroidina soldani</u>
GLOBAURC	<u>Globobulimina auriculata</u>
GLOBTURG	<u>Globobulimina pacifica</u>
HOEGELGS	<u>Hoeglundina elegans</u>
HYPEELON	<u>Hyperammina elongata</u>
HAPLACAR	<u>Haplophragmoides canariensis</u>
ISLAISLA	<u>Islandiella islandica</u>
ISLATERE	<u>Islandiella teretis</u>
KAREINFL	<u>Karreriella bradyi</u>
LAGEMOLL	<u>Lagena mollis</u>
LENTSPAA	<u>Astocolus hyalacrulus</u>
NONILABR	<u>Nonionellina labradorica</u>
NONITURG	<u>Nonionella turgida</u>
MELOPOMP	<u>Melonis pompilioides</u>
PATEHAUR	<u>Pateroris hauerinoides</u>
PULEQUIN	<u>Pullenia quinqueloba</u>
QUICSEMI	<u>Quinqueloculina seminulum</u>
REOPSCOR	<u>Reophax scorpiurus</u>
REOPGUTT	<u>Reophax guttifera</u>
REOPNODU	<u>Reophax nodulosus</u>
REOPARCT	<u>Reophax arctica</u>
REOPSCOT	<u>Reophax scotti</u>
PAROSEMI	<u>Discamolina semipunctata</u>
SACCATLA	<u>Saccammina atlantica</u>
SACCSPHA	<u>Saccammina sphaerica</u>

SACCDIFF	<u>Saccammina difflugiformis</u>
STAINAPT	<u>Stainforthia concava</u>
TROCHINF	<u>Trochammina inflata</u>
TROCHLOB	<u>Trochammina lobata</u>
TROCHGLO	<u>Trochammina globigeriniformis</u>
TRIFANGU	<u>Trifarina angulosa</u>
TRITCONI	<u>Tritaxis conica</u>
TEXTTORQ	<u>Textularia torquata</u>
UVIGPERE	<u>Uvigerina peregrina</u>
ORIDTENR	<u>Oridorsalis umbonatus</u>
VALVLAEV	<u>Valvulineria laevigata</u>
RECUTURB	<u>Recurvoides turbinatus</u>
ADERGLOM	<u>Adercotryma glomerata</u>
SPIROBIF	<u>Spiroplectammina biformis</u> ✓
TRILOCSP	<u>Triloculina trihedra</u>
ELPHACTT	<u>Elphidiella arctica</u>
SIGMSCHL	<u>Sigmoilopsis schlumbergeri</u>
OSANGSPA	<u>Osangularia rugosa</u>
TRIFOCCI	<u>Trifarina occidentalis</u>
TROCHSQA	<u>Trochammina squamata</u>
MBULIMAR	<u>Bulimina marginata</u>

## VARIMAX FACTOR MATRIX

COMM.	1	2	3	4	5	6	7	8
7914	.549	-.012	.062	.321	.033	.151	-.032	-.022
7917	.719	-.006	.176	.526	-.009	-.015	-.014	-.039
7918	.767	-.002	.172	.417	-.012	-.043	-.014	-.739
7920	.593	-.004	.105	.588	-.067	-.072	-.002	-.469
7953	.964	.001	.470	-.001	-.005	.727	-.025	.217
7955	.939	-.003	.401	-.061	.002	.656	-.029	.112
7956	.891	-.003	.293	-.065	.014	.606	-.033	.052
7957	.891	-.000	.146	-.045	.035	.439	-.028	.014
7958	.862	-.000	.373	.198	.066	.220	-.029	.066
7959	.920	.001	.076	.637	.072	.156	-.010	.775
7960	.827	-.013	.066	.859	.061	.036	-.008	.566
8002	.936	.056	.020	-.008	-.003	-.003	.965	.015
8017	.889	.719	.026	.012	-.103	-.009	.511	-.021
8019	.996	-.006	.987	.007	.000	.110	-.003	-.099
8019	.953	.972	.043	.011	-.010	-.007	.014	-.020
8020	.951	.974	.012	.018	-.007	-.007	-.037	-.012
8021	.909	.626	.571	.077	-.025	.418	-.010	.074
8022	.973	.981	.027	.010	.047	-.006	-.058	-.000
8024	.963	.977	.010	.013	.072	-.006	-.020	-.006
8025	.927	.044	.018	-.006	.008	.004	.957	.021
8030	.964	.114	.018	-.001	.245	-.001	.940	.027
8032	.970	-.007	.936	-.024	.014	.144	.043	-.092
8033	.846	.392	.029	.016	-.020	-.007	.731	-.015
8035	.872	.214	.021	-.003	-.104	-.006	.879	-.003
8038	.968	.091	.020	.011	.186	-.003	.960	-.020
8039	.815	.141	.017	.032	.111	-.005	.832	.010
8041	.738	.259	.020	.010	.141	-.004	.767	-.002
8043	.983	.009	.980	.001	-.000	.111	-.003	-.095
8044	.917	.890	.018	.022	.348	-.003	.013	-.006
8046	.916	.951	.011	.019	.095	-.006	-.049	-.006
8047	.980	.446	.871	.016	.001	.118	-.015	.079
8048	.937	.839	.123	.041	.423	.167	-.005	.009
8049	.873	.057	.039	.407	.817	.053	.078	.039
8050	.607	-.009	.012	.741	.078	-.028	.013	-.018
8051	.904	-.010	.005	.940	-.056	-.009	-.004	-.026
8052	.590	.005	.019	.756	.035	.100	-.009	-.025
8055	.595	.049	.172	.127	.178	.540	.011	.382
8057	.979	.046	.100	.031	.184	.109	.063	.943
8002	.351	-.004	-.028	.280	-.029	.274	.010	-.425
8004	.784	-.006	.035	.521	-.055	-.027	.036	-.007
8005	.912	-.003	.067	.523	-.036	.035	.019	-.001
8006	.923	.021	.012	.374	-.022	.027	-.027	.008
8007	.930	-.009	.007	.375	-.023	.222	-.001	.018
8010	.915	-.004	.049	.192	-.038	.142	.011	.014
8011	.908	-.006	.017	.252	-.015	.120	.015	.021
8012	.895	-.001	.013	.272	-.028	.064	.035	.029
8017	.960	.003	.284	-.003	-.017	.891	-.012	.195
8019	.953	.001	.279	-.023	-.014	.806	-.015	.188
8027	.974	.002	.357	-.004	-.015	.883	-.011	.180
8035	.819	.841	.017	.046	.169	-.009	-.000	-.015
8036	.967	.002	.403	.019	-.003	.862	-.010	.179

8037	.955	.975	.010	.018	-.035	-.008	-.043	-.011	-.007-
8039	.956	.967	.007	.011	-.045	-.005	-.077	-.007	-.000
8039	.973	.966	.009	.018	.177	-.006	-.049	-.000	-.005
8040	.947	.971	.013	.014	-.011	-.002	-.048	-.014	-.007
8042	.928	.923	.015	.006	-.065	-.010	.254	-.011	-.007
8043	.851	.901	.020	.022	-.042	-.002	.095	-.022	-.008
8044	.964	-.003	.965	.013	-.003	.122	-.016	-.022	-.083
6003	.857	.004	.008	.874	-.033	-.010	.004	-.021	-.298
6005	.960	.730	.058	.161	.604	.039	-.005	.018	.011
6006	.919	.008	.175	.123	.005	.482	.018	.021	.800
6007	.972	-.010	.095	.161	.013	.276	-.019	-.002	.868
6019	.787	-.003	.414	.009	-.010	.753	-.006	-.015	.213
6041	.895	.004	.172	.153	.031	.886	-.008	.029	.214
6042	.893	.000	.348	.034	.005	.846	-.002	.101	.206
6052	.662	-.006	.003	.604	-.037	-.039	.058	.010	.446
6053	.884	-.009	.004	.929	.006	-.006	.005	-.023	.132
6055	.910	.016	.848	.042	.012	.403	-.011	.131	.051
6056	.699	.011	.001	.742	.194	.058	-.006	-.011	.011
6057	.868	.891	.016	.018	-.002	-.002	.189	-.014	.009
6058	.736	.099	.630	.129	-.023	.536	.114	-.020	.087
6062	.816	.067	.044	.891	-.029	.045	-.012	-.024	.099
6063	.650	.189	.024	.113	.400	-.004	.073	-.017	.028
6064	.910	-.013	.003	.939	.025	.004	-.006	-.018	.104
6065	.664	.016	.011	.777	.129	.040	-.003	.016	.002
6066	.782	.002	.639	.139	.066	.424	-.013	-.049	.408
6068	.995	.000	.628	.040	-.036	.076	.002	.770	.002
6069	.997	.025	.411	.017	-.045	.018	.041	.907	.001
6070	.994	.034	.738	.005	-.025	.195	-.001	.638	.034
6071	.974	.038	.572	-.016	.030	.629	-.009	.469	.169
6072	.967	-.001	.472	-.025	-.009	.845	-.013	-.004	.172
6073	.877	.005	.175	-.001	-.016	.918	-.000	-.045	.021
6074	.983	.004	.076	-.011	-.011	.939	-.006	.063	.302
6075	.789	.056	-.021	.065	.069	.477	.030	.008	.641
6079	.929	.014	.418	-.001	.012	.865	.003	.025	.060
6080	.966	.004	.760	-.015	.013	.573	-.017	-.040	.239
6081	.991	.010	.722	.009	.012	.031	-.023	.684	.028
6082	.903	.013	.862	.041	.043	.201	.098	.008	.310
6083	.906	.124	.671	.085	.296	.413	-.002	.010	.418
6084	.958	.585	.020	.029	.761	.000	.078	.032	.006
6085	.845	.062	.023	.841	.269	.016	.014	.233	.015
6086	.949	.331	.004	.041	.908	-.001	.067	.037	.007
6087	.967	.638	.007	.032	.728	-.004	.090	.031	.005
6088	.923	.690	.028	.014	.573	.000	.291	.026	.004
6089	.995	.018	.868	.012	-.023	.045	.017	.489	.009
6090	.995	-.002	.728	.011	-.031	.016	-.025	.680	.014
6092	.881	.197	-.000	.646	.638	.060	.042	.018	.001
6094	.937	.381	.004	.036	.878	.002	.073	.035	.012
6096	.824	.229	.003	.691	.529	.114	.025	.019	.001
6097	.814	.100	.010	.084	.886	.002	.083	.031	.005
6098	.769	.260	.008	.087	.830	.004	.065	.033	.002
6100	.897	.940	.013	.014	.039	-.009	.084	-.011	-.009
	.976	.978	.010	.014	.127	-.008	.047	-.003	-.006

6101	.756	.854	.006	.017	-.023	-.005	.137	-.006	-.014
6102	.948	.958	.014	.016	-.022	-.008	.124	-.013	-.010
6106	.886	.003	.675	.038	-.018	.270	.047	.411	-.429
6108	.745	.309	.020	.104	.707	-.011	.069	.008	-.009
6109	.953	.076	.005	.089	.960	.002	.086	.037	-.007
6110	.567	.357	.015	.120	.559	.003	.049	.010	-.002
6111	.878	.047	.048	.921	-.005	-.007	.012	-.016	-.144
6112	.717	.179	.013	.570	.529	.050	.016	.011	.018
6113	.883	.028	.040	.882	.011	.004	-.000	.284	-.139
6114	.814	.604	.011	.332	.574	-.042	.036	.018	-.008
6115	.788	.406	.010	.108	.764	-.004	.018	.022	-.005
6116	.974	.175	.002	.066	.965	.002	.078	.040	-.007
6117	.780	.078	.025	.825	.284	.071	.024	-.006	-.057
1009	.991	.023	.247	.020	-.028	.202	.008	.942	-.026
1010	.926	.013	.031	.013	.061	.954	.016	.061	.075
1011	.995	.012	.195	-.018	-.016	-.050	.054	.975	-.006
1015	.923	.055	.019	-.007	-.046	-.002	.956	.016	-.002
1016	.825	.175	.006	.079	.145	.073	.776	.050	-.197
1017	.917	.047	-.059	.232	.331	.724	.395	.089	-.036
1018	.584	.126	.139	.384	.241	.256	.170	.031	-.246
1020	.822	.046	.321	.283	.614	.211	.135	.298	-.232
1021	.626	.019	.252	.578	.387	.138	.062	.002	-.183
1022	.942	.258	.019	.003	.166	.002	.914	.026	-.003
1026	.990	.007	.926	.046	.008	.324	.152	.046	-.013
1027	.903	.465	.019	.018	.449	.009	.688	.028	-.006
1028	.172	.001	-.000	.350	-.052	.032	.013	-.013	-.040
1029	.988	.417	.227	.038	.193	.115	.197	.808	-.030
1034	.932	.081	.167	.013	-.059	.312	.807	.363	-.013
1035	.813	.022	.121	.064	.062	.783	.370	.053	-.007
1043	.992	-.006	.985	.003	-.007	.092	.063	-.090	-.027
1044	.994	-.011	.992	.015	-.004	.092	-.019	.020	.025
1045	.982	-.002	.795	.010	-.030	.365	.003	.462	.055
1046	.927	.004	.484	.119	-.010	.804	.121	.003	.052
1047	.647	-.006	.280	.709	-.054	.098	.003	-.039	-.138
1051	.974	.176	.189	.051	-.037	.531	.062	.784	.049
1053	.997	-.011	.992	.001	-.003	.107	-.017	-.009	.018
1054	.992	-.008	.991	.011	-.007	.090	-.018	.001	.022
1055	.838	.788	.009	.021	.363	.002	.291	.015	-.007
1056	.973	.047	.859	.011	-.025	.117	.291	.366	.021
1057	.867	.410	-.042	.360	.516	.516	.291	.366	.076
1058	.922	.946	.010	.020	.155	.005	.053	-.000	-.005
1059	.970	.094	.676	.024	-.033	.066	.154	.687	.011
1060	.886	.934	.018	.013	.007	-.010	-.023	-.018	-.010
1062	.852	.918	.009	.025	-.014	.012	-.042	-.002	.008
1063	.972	.894	.058	.012	-.078	-.029	.083	.391	-.006
1064	.930	.920	.013	.039	.253	-.006	-.015	-.004	-.016
1065	.806	.724	.015	.394	-.025	.204	-.055	.263	.041
1066	.984	.129	.705	.013	-.036	.118	-.031	.672	-.053
1067	.969	.972	.007	.008	.061	-.007	-.070	-.002	-.002
1068	.956	.962	.005	.012	.057	.003	-.066	.001	.001
1069	.983	.886	.006	.024	.414	.001	-.017	.016	-.003
1070	.914	.738	.101	.060	-.056	.231	-.001	.547	-.011

1071	.994	-.011	.991	.002	-.003	.102	-.017	.026	-.007
1072	.992	-.004	.991	.002	-.004	.091	.024	-.005	.012
1073	.931	.949	.018	.013	.025	-.008	.121	-.015	-.012
1074	.974	.972	.007	.015	.154	-.003	.016	.002	-.004
1075	.744	.575	-.009	.603	.114	.170	-.024	.004	.065
1076	.946	.943	.014	.025	.209	-.009	.064	-.005	-.013
1077	.991	-.004	.805	-.001	-.025	.125	-.026	.566	-.070
1078	.994	.003	.740	.021	-.032	.098	.004	.660	-.004
1080	.994	-.008	.945	.011	-.017	.062	-.024	.308	-.019
1081	.904	.945	.007	.021	.098	-.008	.015	-.002	-.008
1082	.972	.268	.503	.065	.242	.194	-.009	.736	-.015
1083	.944	.967	.008	.014	.046	-.006	.078	-.003	-.006
1084	.862	.908	.048	.033	.102	.063	.037	.121	-.014
1085	.992	.154	.906	.020	-.010	.115	-.017	.366	-.005
1086	.980	.987	.012	.024	.061	-.002	.009	-.006	-.002
1087	.970	.976	.013	.021	.113	-.003	.034	-.009	-.009
1088	.972	.980	.010	.020	.086	-.010	.046	-.007	-.009
1089	.957	.975	.010	.007	-.060	-.010	.042	-.013	-.007
1090	.966	.981	.011	.005	.000	-.009	.003	-.008	-.005
1093	.950	.956	.016	.035	.107	-.005	.026	-.012	-.014
1095	.693	.791	.078	.044	.051	-.005	.070	.081	-.001
1098	.685	.812	.004	.008	-.058	-.002	.146	-.003	-.005
1097	.958	.971	.009	.019	.111	-.009	.053	-.004	-.008
1099	.932	.957	.011	.024	.054	-.003	-.047	-.000	.006
1100	.950	.940	.015	.023	.240	-.010	.061	-.003	-.012
1101	.921	.550	.558	.032	.076	.136	.200	.485	-.034
1102	.964	.955	.014	.022	.199	-.008	.073	-.005	-.013
1103	.984	.866	.008	.029	.474	-.003	.001	.016	-.004
1104	.984	.949	.006	.016	.251	-.002	-.048	.007	-.003
1105	.981	.988	.009	.011	-.031	-.008	-.041	-.009	-.004
1106	.967	.954	.007	.020	.226	-.005	.002	.006	-.004
1107	.943	.969	.015	.013	.036	-.002	-.001	-.013	-.008
1109	.915	.830	.036	.029	.377	.006	.226	.018	-.006
1110	.865	.855	.015	.022	-.036	.006	.344	.006	-.009
1111	.951	.717	.030	.041	.633	.005	.138	.024	-.005
1112	.913	.885	.024	.028	.307	.066	-.011	.035	-.001
1113	.622	.707	.018	.039	.334	.009	.057	.005	-.003
1114	.910	.756	.009	.079	.509	.043	.237	.024	-.008
1115	.955	.837	.009	.032	.468	.006	.112	.021	-.002
1116	.930	.955	.012	.030	.082	.005	.075	.002	-.003
1117	.933	.748	.026	.048	.602	-.004	.055	.013	-.017
1120	.971	.911	.018	.020	.230	-.005	.295	.005	-.006
1124	.948	.310	.017	.011	.378	.003	.827	.034	-.004
1125	.975	.761	.004	.029	.591	.040	.085	.031	-.005
1126	.961	.950	.011	.016	.176	.000	.164	.002	-.008
1127	.909	.948	.027	.018	.054	-.002	.076	-.005	-.001
1129	.867	.926	.010	.023	.051	-.009	-.003	-.007	-.004
1130	.922	.940	.146	.040	.048	.064	-.021	.064	-.038
1131	.877	.688	.014	.036	.622	.006	.020	.021	-.001
1132	.801	.205	.004	.080	.835	-.006	.063	.021	-.023
1133	.925	.010	.088	.799	.041	.020	.026	.027	-.133
1134	.904	.028	.027	.879	.329	-.003	.032	-.009	-.043
1134 <sup>c</sup>	.912	.058	.060	.864	.120	-.018	.022	.276	-.250

1136	.943	.115	.005	.835	.438	.014	.039	.012	-.142
1137	.954	.500	.010	.065	.835	-.003	.035	.028	-.008
1138	.952	.803	.009	.025	.551	-.006	.041	.015	-.012
1139	.883	.508	.008	.232	.753	.020	.036	.022	-.033
1140	.563	.011	-.008	.556	-.039	-.007	.073	-.000	-.444
1141	.409	.450	.060	.369	.103	.086	-.006	.027	-.217
1142	.977	.046	.863	.020	-.011	.165	-.011	.445	-.034
1143	.893	.291	.700	.042	.072	.454	-.017	.285	-.051
1119	.973	.947	.010	.019	.203	.001	.148	.009	.001
4003	.983	.062	.604	.030	-.022	.134	.011	.768	-.065
4004	.952	.048	.920	.006	-.020	.063	.042	.312	.001
4005	.994	.046	.881	.005	.075	.128	-.006	.439	-.041
4006	.993	-.008	.980	-.009	.008	.142	-.000	.089	-.064
4007	.992	-.010	.973	-.011	.013	.204	-.019	-.014	-.058
VARIANCE	29.428	14.935	9.018	8.425	7.783	5.499	5.998	6.118	
CUM. VAR	29.428	44.363	53.381	61.806	69.589	75.088	81.086	87.204	

## VARIMAX FACTOR SCORES

VAR.	1	2	3	4	5	6	7	8
ALVFSCLT	-.049	.013	.253	.565	.014	-.080	-.025	.010
AMMBACU	.005	.000	-.000	.006	-.000	-.005	-.001	.001
AMMCATI	.050	-.005	.021	.009	.004	.047	-.009	.004
AMMOTCAS	.044	.056	-.038	.123	-.050	.548	.075	-.057
ASTPGALD	-.011	.117	.134	-.005	.441	.280	.099	.084
BOLISURS	-.017	-.018	-.037	.079	.008	.004	.039	-.154
BULIFXTL	.050	-.006	.307	-.084	.057	-.035	.014	.085
BULIACLF	.013	.093	3.157	-.771	-1.115	.102	-.110	-2.697
RUCCFRIG	.036	.230	.205	-.007	.619	.099	.492	-1.866
BRI7SPAT	-.023	-.224	.053	-.038	1.300	.061	.134	-.646
BRI7ZSURA	-.004	-.002	2.237	-.423	.027	.076	.128	-.156
CIBTLBRA	-.009	8.070	-.018	.015	.767	-.129	.936	.197
CIRIPSFJU	.005	-.048	.019	-.025	.049	.004	.001	-.193
CASSLAFLV	.012	-.056	-.016	-.015	.389	-.024	.037	-.271
CRICRACAS	1.251	.178	.137	-.329	-.012	.876	.256	-.156
CRIBJFFF	.275	.189	.541	1.098	.034	-.037	.027	.260
CRIBSURG	-.007	.019	.086	.104	-.021	-.001	.005	-.028
ELPHFXCL	-.036	.090	-.939	.224	1.129	-.293	.090	-6.837
FLPHBART	-.000	.087	-.257	.022	.854	-.122	.161	-.910
FLPHSURB	-.002	.125	-.027	.005	-.052	-.012	.025	-.094
EGGLADVA	.255	.155	-.083	-.532	-.015	7.809	.130	.026
EPONIJMAR	.001	-.003	.031	.030	-.008	-.007	-.003	-.035
FISSEFRAT	-.003	.034	-.003	.013	-.002	.001	.004	-.019
FURSLDNR	.033	.017	.399	-.066	.029	-.026	.019	.054
FURSFUST	-.074	-.100	.713	-.046	-.287	.596	.177	-2.465
GLANLAFLV	.003	-.012	.011	.012	.101	-.009	.003	.021
GYROSOLU	-.001	-.035	-.020	.006	.198	-.015	.005	-.046
GLORAIIRC	-.087	-.072	5.884	-.163	.358	-.127	.072	.866
GLORTBURG	.000	-.002	-.002	-.001	.016	-.002	.003	-.000
HDFGFLGS	.003	-.024	.009	-.011	.036	-.003	.006	-.085
HYPEELDN	.034	.032	.089	.080	-.011	.013	-.020	-.008
HAPLCARE	.014	.108	.321	.222	.103	-.026	-.042	.110
TSLLISLD	.059	.984	.132	-.341	-.488	-.164	.068	.020
TSLLITEPF	.021	.108	1.646	-.047	.801	-.099	.149	.038
KARFINFL	-.015	.020	.071	.044	-.002	.002	.019	.007
LAGFMOLL	.000	-.005	.087	-.025	-.025	.009	.013	-.095
LENTSPCA	-.002	-.058	.037	-.014	.260	-.017	-.004	-.024

NONILARP	- .131	- .013	3 .347	.159	.211	.003	- .016	- .244
NONITURP	- .019	.002	.070	.055	.033	.004	- .032	- .080
PATEHAUP	- .001	.017	.003	.001	.001	.002	- .005	- .114
PULFOUNTN	.001	- .025	- .020	.006	.039	.004	- .000	- .020
QUICCSFMI	.000	.093	.013	- .001	.078	.119	- .007	- .140
REEDPGUTT	.181	.042	.707	1 .747	.054	.008	- .049	- .237
REEDPNODU	.027	- .009	.271	.223	.028	.028	- .024	- .204
REEDPARCT	.000	- .007	.013	.042	.028	.005	- .017	- .006
REEDPSCTT	.812	- .044	.044	- .495	.024	.022	- .028	- .065
PARNFSMTA	.077	.021	.090	.054	.028	.016	- .045	- .053
SACCATLA	.001	- .026	.001	- .011	.024	.007	- .024	- .073
SACCDIFF	.558	.057	.180	7 .612	.070	.007	- .065	- .043
SACCSOPHA	- .134	.002	.657	1 .143	.056	.001	- .004	- .061
STATNAPT	- .009	- .003	.126	.031	.056	.001	- .025	- .157
TR0CHINF	- .004	- .003	.109	- .002	.036	.037	- .021	- .015
TKOCHLNR	- .038	.042	.100	.424	.371	.427	- .127	- .120
TKOCHGLD	.499	- .036	.166	.317	.317	.068	- .034	- .091
TRIFANGU	.004	- .013	.012	- .000	.031	.074	- .582	- .827
TRITCONTI	.067	- .740	.073	- .186	.693	.023	- .083	- .255
TEXTTORG	.090	.099	- .044	.617	.138	.115	- .032	- .025
DRIDTENR	.308	- .012	.019	.330	.020	.007	- .041	- .125
VALVLAJV	- .009	.022	.045	.023	.063	.002	- .001	- .074
RECUTURB	- .008	- .008	.002	.036	.024	.051	- .120	- .078
ADFPGLOM	1 .666	.087	- .076	.018	.089	.780	- .038	- .008
SPIRORIF	7 .727	.043	.049	- .427	.046	.502	- .105	- .061
TRILNCSP	1 .169	.026	.026	- .617	.038	.003	- .003	- .014
FLPHACTT	- .001	- .010	.005	- .001	.052	.026	- .049	- .004
TR0C4SOA	- .000	.077	- .001	.002	.026	.003	- .009	- .067
MBULIMAR	.012	- .041	.022	- .038	.154	.213	- .009	- .002
		- .029	.205	- .056	.045	.011		- .286

## APPENDIX III

SLOPE VARIMAX FACTOR MATRIX AND FACTOR SCORES

CMM.	1	2	3	4
7902	.794	.174	.851	-.174
7903	.735	.184	.822	-.150
7904	.773	.378	.682	-.394
7905	.818	.604	.415	-.522
7907	.945	.841	.262	-.401
7908	.980	.909	.209	-.320
7909	.991	.944	.135	-.275
7910	.990	.963	.129	-.199
7911	.954	.943	.153	-.462
7912	.960	.853	.153	-.450
7913	.980	.910	.137	-.356
7948	.939	.958	.078	.103
7949	.942	.963	.053	.091
7950	.969	.979	.060	-.021
7951	.874	.888	.012	.290
7952	.843	.905	.093	-.106
8013	.923	.770	.171	-.542
8014	.008	.051	.069	.012
8060	.560	.024	.743	-.079
8061	.774	.122	.829	-.237
8062	.543	.121	.616	-.377
8063	.546	.318	.464	-.444
8064	.891	.930	.084	.128
8065	.91C1	.922	.174	-.149
8066	.900	.635	.353	-.595
8067	.754	.298	.753	-.278
8068	.336	.062	.546	-.173
7802	.402	.085	.595	-.196
6032	.880	.074	.180	-.286
6033	.775	.424	.585	-.452
6034	.794	.403	.307	-.478
6036	.811	.475	.445	-.613
6050	.547	.124	.207	-.695
6051	.277	.075	.139	-.499
6076	.895	.033	.075	-.102
6077	.742	.421	.380	-.646
6078	.532	.099	.568	-.411
1008	.518	.041	.598	-.390
1036	.841	.030	.904	.089
1037	.851	.104	.859	-.010
1038	.920	-.018	.106	-.000
1039	.686	.249	.534	-.205
1040	.604	.069	.528	-.151
VARIANCE	33.513	21.427	9.039	12.079
CMM. VAR	33.513	54.940	63.979	76.058

## VARIMAX FACTOR SCORES

	VAD.	1	2	3	4
ALVFSCT	-.048	.380	.661	.176	
AMMOMBACU	.014	.335	-.162	.028	
AMMOCATI	-.003	.011	-.007	-.034	
AMMOTCAS	-.014	.025	-.016	-.042	
ASTRALIN	-.077	.466	.171	.239	
BOLISURS	.011	.333	.663	.090	
BULIFEXTL	-.204	5.421	.687	2.486	
BULISTPI	.506	.002	.142	.155	
BULIACLE	-.286	3.557	.118	.593	
BUCCFRIG	-.415	1.526	.867	-.109	
BRI7SPAT	-.846	-.296	.407	-3.213	
BRI7SURA	-.033	-.081	.324	-.062	
CIRIPLOPA	-.215	1.745	-.137	-.215	
CIRIPSFU	.203	.377	-.039	-.576	
CIRTWUEL	.022	.086	.198	.092	
CASSLAEV	-.127	.591	.602	-.183	
CRIBCRAS	-.009	.131	-.010	-.037	
CRIBJFFF	-.008	.029	-.055	-.201	
CRIBSLURG	-.004	.037	-.025	-.028	
CYCLCANC	.020	.069	.014	-.065	
ELPHFXCL	5.116	1.494	.667	-5.645	
ELPHART	-.178	1.177	-.281	.150	
FLPHSURA	-.033	.245	-.090	.040	
EGGLADVA	.039	.163	.115	-.026	
EPONIJMRD	.049	.020	.118	.055	
FISSEPAT	.016	.135	-.000	.075	
FURSELOER	.255	.002	-.017	.107	
FURSEUST	-.334	1.049	.720	-.015	
GLANLAEV	-.022	.218	-.027	-.055	
GYROSCLO	.022	.580	.137	.084	
GLOBAIRC	-.101	.741	-.165	-.468	
GLOBSTURG	.492	.264	.173	.365	
HOEGFLGS	.387	-.035	-.020	.146	
HYPFFLUN	.236	.057	-.098	-.208	
HAPLCARE	-.003	.134	-.056	-.006	
ISLIISLD	-.014	.051	-.038	.115	
ISLITFRE	-.347	1.775	-.845	.522	
KARETNFL	-.033	.471	-.001	-.158	
LAGEMOLL	-.018	.107	.064	-.096	
LENTSPCA	-.045	.013	.102	-.132	
NONILARR	-.204	1.808	-.378	.132	
NONITURG	-.062	.965	1.534	.104	
MFLOPOMP	.020	.147	-.061	.041	
PATEHAUR	-.001	.047	-.012	-.004	
PULFOUTN	.106	.498	.416	-.138	
REOPNSCOR	-.138	1.510	.271	.399	
REOPGUTT	-.62	.449	-.036	-.026	
RFOPNODH	-.003	.135	-.013	-.085	
RFUPARCT	-.007	.053	-.025	.009	
REOPSCOT	-.001	.075	-.038	-.032	
PAROSEMTI	-.001	.032	-.006	-.061	
SACCATLA	-.304	1.164	.502	-.467	
SACCDIFF	.068	.732	-.014	-.093	
SACCSPHA	.032	.360	.256	.174	
STAINAPT	.057	.015	-.039	-.191	
TR0CHINF	-.015	-.046	.276	-.060	
TR0CHLNB	-.035	.123	.033	-.036	
TR0CHGLQ	-.003	.301	-.142	-.166	
TRIFANGU	-.671	1.345	-.595	-2.027	
TRITCONI	-.052	.335	-.025	.052	
UVIGPFRF	6.458	-.324	.150	3.815	
ORIDTFNR	.070	.063	-.040	-.075	
VALVLAEV	-.099	.720	.435	-.484	
RECUTURB	.078	.089	-.026	.093	
ADERGLOM	-.006	.009	.000	-.015	
SPIROBI	-.003	.026	-.018	.002	
SIGMSCHL	.119	.043	-.025	.058	
OSANGSPA	-.058	-.045	-.010	-.226	
TRIFGCC	-.488	-.944	7.898	.243	
TR0CHSOA	.002	.329	.314	.422	
MBULIMAR	-.161	.692	.028	-.010	

#### APPENDIX IV

LIVE SPECIES VERSUS FACTOR SCORE TABLES (1--8)

## COMPARISON OF FACTOR 1 WITH OBSERVED LIVING DATA

STATION	FACTOR LOADING	MAIN SPECIES FROM THE FACTOR SCORE MATRIX (FSM)				OTHER LIVE SPECIES NOT DOMINANT IN FSM
		ADERGLOM 7.1	RECUTURB 1.6	SPIROBIF 1.1	CRIBCRAS 1.2	
0819	.97	74	14		11	
0820	.97	100				
0822	.98	83				
8037	.96	76				
8038	.96	50	13			
8039	.97	43	13	2	17	
1058	.94	63	9			
1062	.91	41	5	1	6	
1086	.98	50	6			
1087	.97	85	9			
1088	.98	87	12			
1089	.97	60	40			
1090	.98	68	33			
1093	.95	62	19	2		
1097	.97	70	23			
1100	.94	71	20			
1104	.95	56	11	7		
1105	.98	63	9	9		
1106	.95	67	8			
1127	.94	50	20		9	
1128	.92	50	20	4	2	
1130	.94	72	6	3	8	
		OBSERVED LIVE				

## COMPARISON OF FACTOR 2 WITH OBSERVED LIVING DATA

### COMPARISON OF FACTOR 3 WITH OBSERVED LIVING DATA

## COMPARISON OF FACTOR 4 WITH OBSERVED LIVING DATA

## COMPARISON OF FACTOR 5 WITH OBSERVED LIVING DATA

## COMPARISON OF FACTOR 6 WITH OBSERVED LIVING DATA

## COMPARISON OF FACTOR 7 WITH OBSERVED LIVING DATA

## COMPARISON OF FACTOR 8 WITH OBSERVED LIVING DATA

## APPENDIX V

## GENERAL FAUNAL DATA

$H(s)$  = Shannon-Weiner diversity value

E = Equitability

$\alpha$  = Fisher index

S = Species number

L = Total living

L/D = Live/Dead

T = Total, live + dead

STATION	H(S)	E	$\alpha$	S	L	L/D	T
7914	1.6	0.23	4.0	17.0	0.0	0.0	972.0
7917	2.2	0.52	4.5	18.0	0.0	0.0	368.0
7818	2.1	0.43	3.75	18.0	4.0	0.007	512.0
7920	1.8	0.48	1.75	12.0	0.0	0.0	963.0
7955	2.9	0.60	4.5	24.0	112.0	0.093	1200.0
7956	2.2	0.37	3.4	20.0	40.0	0.016	2496.0
7957	1.9	0.35	5.5	30.0	120.0	0.090	1320.0
7958	2.0	0.23	7.85	35.0	27.0	0.030	927.0
7959	2.7	0.42	5.5	30.0	105.0	0.070	1560.0
7960	2.9	0.58	3.7	16.0	10.0	0.010	570.0
0802	2.1	0.52	4.0	7.0	38.0	0.075	506.0
0817	1.8	0.70	1.5	8.0	0.0	0.0	221.0
0818	1.1	0.17	2.0	16.0	384.0	0.03	10688.0
0819	1.6	0.45	2.5	11.0	27.0	0.13	201.0
0820	1.4	0.51	1.0	8.0	6.0	0.08	71.0
0821	2.1	0.61	2.2	14.0	108.0	0.08	1208.0
0822	1.2	0.37	2.5	9.0	6.0	0.03	164.0
0824	1.5	0.48	2.0	9.0	17.0	0.05	285.0
0825	0.6	0.45	1.0	9.0	40.0	0.17	231.0
0830	1.3	0.40	1.8	9.0	22.0	0.04	532.0
0832	1.4	0.27	3.2	15.0	128.0	0.11	1168.0
0833	1.9	0.60	2.1	10.0	60.0	0.17	350.0
0835	1.7	0.35	2.2	10.0	49.0	0.09	516.0
0838	1.3	0.70	2.5	11.0	35.0	0.12	295.0
0839	1.6	0.70	1.1	7.0	16.0	0.02	1016.0
0841	1.9	0.74	1.5	9.0	108.0	0.22	486.0
0843	1.4	0.20	5.5	20.0	352.0	0.157	2240.0
0844	1.56	0.40	2.75	13.0	34.0	0.062	542.0
0846	1.3	0.45	2.0	8.0	28.0	0.172	162.0
0847	2.0	0.35	4.75	20.0	55.0	0.19	278.0
0848	2.2	0.46	4.75	20.0	51.0	0.26	191.0
0849	1.8	0.11	3.75	17.0	22.0	0.04	546.0
0850	1.8	0.60	1.25	10.0	0.0	0.0	8616.0
0851	1.8	0.74	1.50	8.0	0.0	0.0	1795.0
0852	2.4	0.48	3.75	22.0	152.0	0.112	1352.0
0856	2.5	0.52	5.0	24.0	20.0	0.032	618.0
0857	1.7	0.35	3.25	16.0	46.0	0.08	566.0
8002	2.8	0.60	4.75	27.0	68.0	0.005	11136.0
8005	2.7	0.60	5.5	25.0	34.0	0.05	574.0
8006	2.3	0.41	4.25	20.0	120.0	0.07	1734.0
8007	2.5	0.61	3.5	21.0	24.0	0.01	1728.0
8010	2.5	0.48	3.8	25.0	108.0	0.03	2704.0
8011	2.5	0.60	3.25	20.0	90.0	0.04	1818.0
8012	2.6	0.60	4.25	25.0	150.0	0.08	1776.0
8017	2.3	0.38	3.75	26.0	96.0	0.018	5088.0
8018	2.4	0.51	2.75	21.0	144.0	0.028	5120.0
8027	2.4	0.43	4.5	24.0	36.0	0.027	1324.0
8035	1.5	0.20	1.75	7.0	16.0	0.105	152.0
8036	2.3	0.36	4.5	27.0	104.0	0.026	3927.0
8037	1.1	0.13	2.0	8.0	42.0	0.078	534.0
8038	1.1	0.11	2.05	12.0	6.0	0.027	221.0
8039	1.5	0.38	2.0	12.0	92.0	0.142	646.0
8040	1.4	0.33	2.75	12.0	47.0	0.206	228.0

8042	1.8	0.57	1.30	11.0	384.0	0.162	2360.0
8043	1.8	0.68	2.0	9.0	25.0	0.13	192.0
8044	1.7	0.27	2.3	21.0	96.0	0.008	11296.0
6003	2.1	0.38	3.5	21.0	0.0	0.0	1744.0
6005	1.9	0.40	3.5	15.0	0.0	0.0	245.0
6007	2.3	0.52	3.75	19.0	0.0	0.0	1662.0
6018	1.9	0.51	1.5	13.0	0.0	0.0	2544.0
6041	2.3	0.47	6.0	21.0	0.0	0.0	189.0
6042	1.9	0.40	3.0	19.0	0.0	0.0	215.0
6052	2.3	0.53	3.2	18.	0.0	0.0	1248.0
6053	2.1	0.51	3.0	16.0	0.0	0.0	1012.0
6055	1.7	0.34	2.2	16.0	0.0	0.0	2816.0
6056	2.0	0.41	4.5	18.0	0.0	0.0	888.0
6057	2.0	0.67	2.0	11.0	0.0	0.0	542.0
6058	2.1	0.30	4.5	24.0	0.0	0.0	1028.0
6062	2.4	0.62	3.2	19.0	0.0	0.0	2196.0
6063	1.9	0.79	2.1	9.0	0.0	0.0	153.0
6064	2.0	0.47	4.0	17.0	0.0	0.0	261.0
6065	2.4	0.59	5.3	20.0	0.0	0.0	203.0
6066	2.9	0.62	5.2	28.0	0.0	0.0	570.0
6068	1.3	0.26	2.2	15.0	112.0	0.04	2496.0
6069	1.1	0.30	1.5	10.0	0.0	0.0	2960.0
6070	1.7	0.30	2.5	17.0	272.0	0.04	7088.0
6071	2.0	0.44	3.0	18.0	0.0	0.0	3736.0
6072	1.7	0.36	1.0	15.0	0.0	0.0	5664.0
6073	1.7	0.34	2.2	16.0	0.0	0.0	2608.0
6074	1.9	0.40	3.7	18.0	0.0	0.0	1520.0
6075	2.9	0.68	8.2	27.0	0.0	0.0	242.0
6079	2.0	0.54	3.0	15.0	0.0	0.0	164.0
6080	1.6	0.25	2.75	19.0	0.0	0.0	6128.0
6081	1.5	0.32	1.8	14.0	0.0	0.0	5724.0
6082	2.0	0.45	3.1	17.0	0.0	0.0	790.0
6083	2.6	0.52	3.75	25.0	0.0	0.0	3348.0
6084	1.6	0.41	3.2	12.0	9.0	0.05	185.0
6085	2.3	0.56	2.8	17.0	200.0	0.20	1008.0
6086	1.8	0.43	3.5	14.0	0.0	0.0	282.0
6087	1.6	0.45	2.0	11.0	0.0	0.0	282.0
6088	1.7	0.43	3.5	13.0	0.0	0.0	251.0
6089	1.3	0.30	1.4	12.0	200.0	0.04	5408.0
6090	0.7	0.70	0.9	3.0	0.0	0.0	12928.0
6092	2.3	0.52	3.5	19.0	12.0	0.10	960.0
6094	1.7	0.46	2.2	12.0	0.0	0.0	294.0
6096	2.4	0.55	3.7	19.0	108.0	0.012	1404.0
6097	1.8	0.50	2.0	13.0	50.0	0.08	580.0
6098	2.0	0.40	4.0	19.0	84.0	0.013	639.0
6099	1.8	0.56	1.8	11.0	34.0	0.070	466.0
6100	1.6	0.49	2.0	10.0	35.0	0.124	282.0
6101	1.5	0.42	2.0	11.0	30.0	0.13	222.0
6102	1.9	0.60	2.2	12.0	42.0	0.159	264.0
6106	2.0	0.47	2.2	17.0	0.0	0.0	3702.0
6108	2.3	0.70	7.8	25.0	0.0	0.0	245.0
6109	1.8	0.63	2.0	10.0	0.0	0.0	241.0
6111	2.3	0.44	4.0	22.0	37.0	0.117	315.0
6112	2.5	0.57	5.1	21.0	0.0	0.0	162.0

6113	2.1	0.40	4.0	22.0	0.0	0.0	1706.0
6114	2.6	0.57	6.0	24.0	0.0	0.0	213.0
6115	1.9	0.70	1.75	9.0	0.0	0.0	488.0
6116	1.7	0.45	1.7	12.0	0.0	0.0	768.0
6117	2.5	0.56	3.7	23.0	152.0	0.135	1124.0
6110	2.2	0.80	2.8	12.0	0.0	0.0	197.0
1009	1.5	0.32	2.0	14.0	0.0	0.0	1468.0
1010	1.5	0.28	3.2	16.0	8.0	0.01	554.0
1011	0.8	0.24	1.7	10.0	59.0	0.09	604.0
1015	0.5	0.21	1.5	8.0	159.0	0.71	223.0
1016	2.0	0.44	3.0	16.0	84.0	0.11	734.0
1017	2.0	0.58	2.7	13.0	4.0	0.006	572.0
1018	2.6	0.58	4.5	22.0	38.0	0.07	490.0
1020	2.6	0.49	6.5	29.0	76.0	0.12	602.0
1021	2.9	0.78	4.5	22.0	0.0	0.0	536.0
1022	1.5	0.38	2.1	11.0	80.0	0.26	306.0
1026	1.9	0.39	5.1	18.0	2.0	0.008	232.0
1027	1.8	0.75	5.0	18.0	0.0	0.0	149.0
1028	1.8	0.75	2.9	14.0	10.0	0.028	245.0
1029	1.9	0.42	3.2	17.0	18.0	0.046	391.0
1034	1.9	0.40	3.5	15.0	18.0	0.05	307.0
1035	2.1	0.67	2.5	13.0	28.0	0.06	434.0
1043	1.1	0.41	1.2	7.0	58.0	0.26	219.0
1044	0.7	0.40	1.0	5.0	10.0	0.56	178.0
1045	0.8	0.33	1.2	7.0	40.0	0.04	1100.0
1046	2.4	0.51	3.1	17.0	28.0	0.47	602.0
1047	2.4	0.50	4.0	20.0	40.0	0.06	600.0
1051	2.1	0.47	3.5	18.0	36.0	0.05	632.0
1053	0.6	0.27	1.0	7.0	2880.0	0.21	13344.0
1054	1.0	0.38	1.0	7.0	97.0	0.57	170.0
1055	1.9	0.79	2.0	9.0	5.0	0.03	157.0
1056	1.9	0.51	2.0	9.0	34.0	0.128	264.0
1057	2.3	0.64	3.8	15.0	24.0	0.10	219.0
1058	1.7	0.60	2.2	10.0	5.0	0.02	233.0
1059	2.0	0.47	4.0	16.0	24.0	0.07	322.0
1060	1.4	0.41	2.2	10.0	49.0	0.18	261.0
1062	1.8	0.58	2.5	10.0	72.0	0.38	185.0
1063	1.3	0.33	2.2	11.0	231.0	0.70	330.0
1064	1.6	0.48	1.7	10.0	128.0	0.24	522.0
1065	2.2	0.68	3.0	14.0	96.0	0.31	309.0
1066	1.8	0.38	2.5	16.0	248.0	0.10	2328.0
1067	0.8	0.55	0.9	4.0	241.0	0.69	348.0
1068	1.3	0.42	1.7	9.0	66.0	0.10	624.0
1069	1.6	0.38	2.7	13.0	356.0	0.51	698.0
1070	2.1	0.52	3.0	16.0	122.0	0.14	864.0
1071	1.0	0.30	0.9	9.0	792.0	0.03	20544.0
1072	1.2	0.33	1.3	10.0	290.0	0.29	996.0
1073	1.6	0.47	2.0	11.0	47.0	0.15	303.0
1074	1.5	0.53	2.0	9.0	48.0	0.19	247.0
1075	2.1	0.51	3.0	17.0	486.0	0.25	1936.0
1076	1.8	0.61	1.8	10.0	241.0	0.50	482.0
1077	1.3	0.54	0.9	7.0	0.0	0.0	5120.0

1078	1.4	0.33	2.2	12.0	0.0	0.0	376.0
1080	0.9	0.49	0.5	5.0	0.0	0.0	4614.0
1081	1.8	0.61	1.8	11.0	62.0	0.14	442.0
1082	2.1	0.51	2.5	17.0	0.0	0.0	614.0
1083	1.8	0.61	2.0	10.0	50.	0.14	339.0
1084	2.1	0.54	3.5	15.0	35.0	0.11	298.0
1085	1.7	0.37	2.3	15.0	344.0	0.12	2835.0
1086	1.8	0.45	3.5	14.0	68.0	0.21	318.0
1087	1.6	0.48	2.0	10.0	0.0	0.0	226.0
1088	1.7	0.53	1.5	10.0	80.0	0.08	939.0
1089	1.3	0.48	1.8	8.0	080.0	0.07	852.0
1090	1.4	0.58	1.3	7.0	76.0	0.16	466.0
1093	1.8	0.47	2.2	13.0	186.0	0.29	628.0
1095	1.8	0.62	2.2	10.0	0.0	0.0	171.0
1096	1.7	0.62	2.2	9.0	16.0	0.09	163.0
1097	1.5	0.61	1.75	8.0	140.0	0.23	608.0
1099	1.4	0.40	2.0	10.0	0.0	0.0	307.0
1100	1.8	0.52	2.3	12.0	49.0	0.17	286.0
1101	2.1	0.48	4.0	17.0	63.0	0.13	481.0
1102	1.8	0.65	2.0	10.0	26.0	0.09	282.0
1103	1.8	0.52	2.5	11.0	91.0	0.33	274.0
1104	1.5	0.47	1.8	10.0	52.0	0.11	498.0
1105	1.3	0.35	1.8	11.0	46.0	0.06	730.0
1106	1.7	0.49	2.2	11.0	40.0	0.13	290.0
1107	1.6	0.62	2.2	10.0	37.0	0.15	241.0
1109	1.9	0.66	2.2	10.0	43.0	0.21	237.0
1110	1.5	0.62	1.5	7.0	0:0	0.00	151.0
1111	1.9	0.55	2.0	12.0	59.0	0.12	506.0
1112	2.4	0.56	1.9	20.0	51.0	0.04	1284.0
1113	1.9	0.62	1.9	11.0	103.0	0.11	994.0
1114	2.3	0.52	3.5	20.0	36.0	0.29	1260.0
1115	2.0	0.58	2.3	13.0	128.0	0.18	702.0
1116	2.0	0.59	3.0	14.0	55.0	0.22	285.0
1119	1.6	0.53	2.25	10.0	60.0	0.24	250.0
1120	1.8	0.65	1.5	10.0	192.0	0.16	1184.0
1124	1.6	0.52	2.0	10.0	65.0	0.21	299.0
1125	1.6	0.59	1.75	9.0	50.0	0.09	527.0
2619	1.4	0.17	1.1	11.0	58.0	0.06	945.0
1127	1.9	0.58	2.1	12.0	34.0	0.12	277.0
1128	1.9	0.72	1.5	10.0	264.0	0.19	1332.0
1130	2.2	0.56	3.5	17.0	168.0	0.20	816.0
1131	2.0	0.65	3.1	12.0	198.0	0.27	714.0
1132	1.6	0.38	2.0	14.0	106.0	0.09	1041.0
1133	2.3	0.54	3.9	19.0	38.0	0.05	664.0
1134	2.3	0.67	3.7	16.0	0.0	0.0	252.0
1135	2.6	0.58	4.7	23.0	0.0	0.0	474.0
1136	2.5	0.60	4.0	20.0	6.0	0.01	530.0
1137	2.0	0.58	2.2	13.0	76.0	0.10	702.0
1138	1.9	0.46	2.6	14.0	96.0	0.18	520.0
1139	2.2	0.43	4.2	20.0	0.0	0.0	508.0
1140	2.6	0.47	5.0	29.0	180.0	0.15	1200.0
1141	2.6	0.55	4.0	26.0	288.0	0.15	1812.0

1142	1.8	0.40	2.75	15.0	58.0	0.07	768.0
1143	2.1	0.61	5.0	18.0	62.0	0.28	217.0
1118	1.9	0.59	1.7	13.0	240.0	0.13	1240.0
4003	1.9	0.43	1.2	15.0	0.0	0.0	5184.0
4004	1.6	0.43	1.1	12.0	1184.0	0.20	5728.0
4005	1.2	0.28	1.5	14.0	0.0	0.0	10464.0
4006	1.5	0.31	0.7	7.0	0.0	0.0	2656.0
4007	1.0	0.37	0.5	6.0	0.0	0.0	3216.0
7902	3.1	0.60	7.5	38.0	12.0	0.032	387.0
7903	2.9	0.54	11.5	34.0	0.0	0.0	235.0
7904	2.8	0.51	16.0	34.0	1.0	0.0004	207.0
7906	3.1	0.52	14.0	40.0	2.0	0.0008	230.0
7907	2.4	0.42	6.0	27.0	0.0	0.0	536.0
7908	2.3	0.36	5.75	27.0	0.0	0.0	618.0
7909	1.6	0.26	3.75	19.0	0.0	0.0	570.0
7910	1.6	0.23	4.50	21.0	0.0	0.0	614.0
7911	1.45	0.23	3.75	18.0	0.0	0.0	588.0
7912	1.39	0.25	2.0	16.0	48.0	0.006	7392.0
7913	1.37	0.26	2.75	15.0	0.0	0.0	550.0
7948	1.24	0.16	4.5	21.0	0.0	0.0	1680.0
7949	1.45	0.28	2.25	15.0	0.0	0.0	2016.0
7950	1.4	0.30	2.25	14.0	0.0	0.0	1690.0
7951	1.2	0.37	2.0	12.0	0.0	0.0	1140.0
7952	1.8	0.35	3.5	17.0	0.0	0.0	480.0
8013	1.9	0.37	4.0	19.0	0.0	0.0	410.0
8014	2.9	0.72	6.2	20.0	0.0	0.0	400.0
8060	2.8	0.52	10.0	30.0	0.0	0.0	202.0
8061	2.7	0.50	7.2	30.0	0.0	0.0	576.0
8062	3.0	0.66	9.0	30.0	0.0	0.0	460.0
8063	3.1	0.63	9.5	36.0	0.0	0.0	454.0
8064	1.7	0.16	7.0	33.0	0.0	0.0	786.0
8065	2.1	0.37	4.0	22.0	0.0	0.0	840.0
8066	2.8	0.60	7.8	30.0	0.0	0.0	452.0
8067	2.7	0.62	9.7	30.0	0.0	0.0	267.0
8068	1.6	0.21	9.0	23.0	0.0	0.0	158.0
6033	2.0	0.39	4.0	20.0	0.0	0.0	864.0
6034	2.0	0.39	4.0	20.0	0.0	0.0	864.0
6036	2.9	0.70	6.0	23.0	0.0	0.0	215.0
6050	2.5	0.45	4.0	22.0	0.0	0.0	964.0
6051	2.1	0.47	3.5	18.0	0.0	0.0	630.0
6076	2.4	0.43	4.7	25.0	33.0	0.03	1008.0
6077	2.0	0.43	7.0	29.0	0.0	0.0	680.0
6078	3.0	0.63	10.0	31.0	0.0	0.0	275.0
1008	2.73	0.69	4.5	22.0	0.0	0.0	570.0
1036	2.7	0.53	6.1	28.0	0.0	0.0	592.0
1037	2.8	0.59	5.5	28.0	0.0	0.0	1212.0
1038	2.2	0.32	5.5	30.0	15.0	0.01	1116.0
1039	2.9	0.79	5.0	23.0	4.0	0.007	520.0
1040	2.4	0.48	5.0	24.0	0.0	0.0	610.0

## APPENDIX VI

MEASUREMENTS OF DEPTH, TEMPERATURE, SALINITY  
AND SUBSTRATE CHARACTER

SAMPLE	DEPTH	T°C	S%	%G	%S	%M
7914	271.0	10.0	34.8	0.0	4.0	95.0
7917	251.0	10.0	34.8	0.0	4.0	95.0
7918	240.0	10.0	34.8	0.0	4.0	95.0
7920	220.0	10.0	34.8	0.0	4.0	95.0
7953	160.0	9.5	35.0	8.0	90.0	2.0
7955	156.0	9.5	35.0	8.0	90.0	2.0
7956	166.0	9.5	35.0	18.0	78.0	4.0
7957	140.0	9.5	35.0	12.0	88.0	0.0
7958	175.0	9.5	35.0	12.0	88.0	0.0
7959	190.0	9.5	35.0	12.0	88.0	0.0
7960	220.0	9.5	35.0	0.0	4.0	95.0
0802	18.0	4.0	31.8	2.0	67.0	31.0
0817	110.0	1.9	32.3	5.0	60.0	35.0
0818	70.0	1.9	32.3	22.0	78.0	0.0
0819	113.0	1.9	32.3	20.0	80.0	0.0
0820	163.0	1.9	32.3	0.0	58.0	42.0
0821	119.0	1.9	32.3	0.0	68.0	32.0
0822	137.0	1.9	32.3	0.0	20.0	80.0
0824	102.0	1.9	32.3	10.0	90.0	0.0
0825	16.0	4.0	31.8	0.0	56.0	44.0
0830	59.0	4.0	31.8	20.0	80.0	0.0
0832	29.0	4.0	31.8	0.0	60.0	40.0
0833	31.0	4.0	31.8	20.0	76.0	4.0
0835	51.0	4.0	31.8	0.0	60.0	40.0
0838	27.0	4.0	31.8	20.0	76.0	4.0
0839	33.0	4.0	31.8	20.0	76.0	4.0
0841	33.0	4.0	31.8	0.0	60.0	40.0
0843	73.0	4.0	31.8	10.0	76.0	12.0
0844	174.0	1.9	32.3	0.0	66.0	34.0
0846	238.0	1.9	32.3	0.0	52.0	48.0
0847	82.0	4.3	33.3	2.0	67.0	31.0
0848	104.0	4.3	33.3	0.0	95.0	5.0
0849	128.0	4.3	33.3	0.0	95.0	5.0
0850	196.0	4.3	33.3	0.0	0.0	100.0
0851	200.0	8.5	34.4	0.0	0.0	100.0
0852	108.0	7.1	34.1	0.0	80.0	20.0
0856	101.0	8.5	34.3	20.0	80.0	0.0
0857	86.0	8.5	34.3	10.0	90.0	0.0
8002	214.0	8.5	34.3	0.0	4.0	95.0
8004	206.0	10.0	34.8	0.0	30.0	70.0
8005	203.0	9.0	34.9	0.0	30.0	70.0
8006	197.0	10.0	34.8	0.0	30.0	70.0
8007	190.0	10.0	34.8	0.0	30.0	70.0
8010	195.0	10.0	34.8	0.0	30.0	70.0
8011	195.0	10.0	34.8	0.0	30.0	70.0
8012	195.0	10.0	34.8	0.0	30.0	70.0
8017	180.0	10.3	35.0	0.0	88.0	12.0
8018	195.0	10.3	35.0	0.0	88.0	12.0
8027	185.0	10.3	35.0	0.0	88.0	12.0

8035	135.0	1.9	32.3	2.0	58.0	40.0
8036	114.0	1.9	32.3	12.0	78.0	10.0
8037	154.0	1.9	32.3	2.0	58.0	40.0
8038	167.0	1.9	32.3	0.0	40.0	60.0
8039	156.0	1.9	32.3	0.0	52.0	48.0
8040	135.0	1.9	32.3	0.0	52.0	48.0
8042	103.0	4.0	31.8	22.0	68.0	10.0
8043	60.0	4.0	31.8	10.0	77.0	13.0
8044	76.0	4.0	31.8	20.0	80.0	0.0
6003	156.0	4.0	31.8	0.0	4.0	96.0
6005	175.0	9.5	35.0	0.0	1.0	99.0
6006	160.0	9.5	35.0	0.0	1.0	99.0
6007	167.0	9.5	35.0	9.0	88.0	3.0
6018	230.0	9.5	35.0	0.0	78.0	22.0
6041	280.0	9.5	35.0	0.0	69.0	31.0
6042	280.0	9.5	35.0	0.0	72.0	28.0
6052	170.0	3.1	32.8	0.0	12.0	88.0
6053	152.0	3.1	32.8	0.0	8.0	92.0
6055	100.0	3.1	32.8	0.0	72.0	28.0
6056	115.0	3.1	32.8	0.0	80.0	20.0
6057	88.0	3.1	32.8	10.0	90.0	0.0
6058	75.0	3.1	32.8	9.0	91.0	0.0
6062	165.0	8.5	34.3	0.0	1.0	99.0
6063	185.0	8.5	34.3	0.0	1.0	99.0
6064	180.0	8.5	34.3	0.0	0.0	100.0
6065	190.0	8.5	34.3	0.0	1.0	99.0
6066	125.0	8.5	34.3	0.0	21.0	79.0
6068	97.0	5.6	33.5	96.0	4.0	0.0
6069	87.0	5.6	33.5	1.0	97.0	2.0
6070	95.0	5.6	33.5	99.0	1.0	0.0
6071	116.0	5.6	33.5	46.0	50.0	3.0
6072	105.0	5.6	33.5	12.6	88.0	0.0
6073	140.0	5.6	33.5	11.0	89.0	0.0
6074	280.0	10.3	35.0	0.0	75.0	25.0
6075	300.0	10.3	35.0	0.0	80.0	20.0
6079	130.0	5.6	33.5	10.0	90.0	0.0
6080	110.0	5.6	33.5	8.0	92.0	0.0
6081	100.0	5.6	33.5	7.0	93.0	0.0
6082	94.0	5.6	33.5	11.0	89.0	0.0
6083	123.0	5.0	33.5	0.0	12.0	88.0
6084	140.0	5.0	33.0	0.0	8.0	92.0
6085	160.0	5.0	33.0	0.0	25.0	75.0
6086	155.0	5.0	33.0	0.0	8.0	92.0
6087	142.0	5.0	33.0	0.0	8.0	92.0
6088	128.0	6.0	32.0	18.0	82.0	0.0
6089	110.0	6.0	32.0	24.0	76.0	6.0
6090	56.0	6.0	32.0	14.0	86.0	0.0
6092	161.0	6.0	32.0	0.0	12.0	88.0
6094	155.0	6.0	32.0	0.0	10.0	90.0
6096	150.0	6.0	32.0	0.0	0.0	100.0
6097	141.0	6.0	32.0	0.0	20.0	80.0
6098	144.0	6.0	32.0	0.0	18.0	82.0
6099	113.0	6.0	32.0	0.0	56.0	44.0

6100	112.0	6.0	32.0	0.0	65.0	35.0
6101	110.0	6.0	32.0	0.0	60.0	40.0
6102	103.0	6.0	32.0	0.0	59.0	41.0
6106	66.0	6.0	32.0	64.0	36.0	0.0
6108	150.0	6.0	32.0	0.0	36.0	54.0
6109	149.0	6.0	32.0	0.0	41.0	59.0
6111	205.0	6.0	32.0	0.0	30.0	70.0
6112	228.0	6.0	32.0	0.0	28.0	72.0
6113	170.0	6.0	32.0	0.0	38.0	62.0
6114	168.0	6.0	32.0	0.0	38.0	62.0
6115	155.0	6.0	32.0	0.0	38.0	62.0
6116	150.0	6.0	32.0	0.0	38.0	62.0
6117	155.0	6.0	32.0	0.0	38.0	62.0
6110	148.0	6.0	32.0	0.0	38.0	62.0
1009	130.0	6.0	34.3	2.0	79.0	19.0
1010	100.0	8.5	34.3	1.0	99.0	0.0
1011	77.0	8.5	34.3	2.0	98.0	0.0
1015	60.0	8.5	34.3	6.0	94.0	0.0
1016	140.0	9.5	35.0	0.0	99.0	1.0
1017	200.0	7.3	34.5	0.0	90.0	10.0
1018	280.0	9.0	35.0	0.0	90.0	10.0
1020	320.0	9.0	35.0	0.0	90.0	10.0
1021	440.0	9.0	35.0	0.0	87.0	13.0
1022	120.0	9.0	35.0	0.0	86.0	14.0
1026	260.0	6.0	32.0	0.0	83.0	17.0
1027	235.0	6.0	32.0	0.0	20.0	80.0
1028	215.0	6.0	32.0	0.0	0.0	100.0
1029	120.0	5.0	33.7	0.0	0.0	100.0
1034	129.0	7.0	34.5	10.0	90.0	0.0
1035	200.0	7.0	34.5	10.0	90.0	0.0
1043	63.0	2.1	32.1	2.0	98.0	0.0
1044	70.0	1.9	32.3	5.0	95.0	0.0
1045	160.0	1.9	32.3	5.0	95.0	0.0
1046	240.0	1.9	32.3	2.0	80.0	18.0
1047	300.0	1.9	32.3	0.0	78.0	22.0
1051	120.0	1.4	32.4	0.0	97.0	3.0
1053	60.0	1.4	32.4	18.0	82.0	0.0
1054	72.0	1.2	32.3	9.0	91.0	0.0
1055	175.0	1.9	32.3	0.0	97.0	3.0
1056	84.0	1.3	32.3	0.0	97.0	3.0
1057	220.0	1.9	32.3	0.0	70.0	30.0
1058	160.0	0.8	32.3	0.0	86.0	14.0
1059	100.0	1.5	32.8	0.0	83.0	17.0
1060	180.0	2.5	33.1	0.0	18.0	82.0
1062	100.0	1.9	32.3	10.0	60.0	30.0
1063	113.0	1.9	32.3	10.0	53.0	37.0
1064	340.0	1.9	32.3	0.0	0.0	100.0
1065	360.0	1.9	32.3	0.0	3.0	97.0
1066	80.0	1.9	32.3	0.0	97.0	3.0
1067	270.0	1.9	32.3	0.0	2.0	98.0
1068	200.0	1.9	32.3	0.0	8.0	92.0
1069	140.0	1.9	32.3	0.0	78.0	22.0

1070	105.0	1.9	32.3	30.0	70.0	0.0
1071	70.0	1.9	32.3	10.0	90.0	0.0
1072	55.0	1.9	32.3	10.0	90.0	0.0
1073	120.0	1.9	32.3	0.0	56.0	44.0
1074	160.0	1.9	32.3	0.0	75.0	25.0
1075	275.0	1.9	32.3	0.0	18.0	82.0
1076	145.0	1.9	32.3	10.0	90.0	0.0
1077	80.0	1.9	32.3	69.0	31.0	0.0
1078	80.0	1.9	32.3	35.0	65.0	0.0
1080	75.0	1.9	32.3	67.0	33.0	0.0
1081	145.0	1.9	32.3	5.0	90.0	5.0
1082	91.0	2.3	32.4	1.0	99.0	0.0
1083	200.0	2.3	32.4	0.0	27.0	73.0
1084	100.0	1.1	32.5	0.0	42.0	58.0
1085	85.0	1.9	32.3	7.0	63.0	0.0
1086	260.0	1.9	32.3	0.0	26.0	74.0
1087	200.0	1.9	32.3	0.0	20.0	80.0
1088	300.0	1.9	32.3	0.0	24.0	76.0
1089	310.0	1.9	32.5	0.0	18.0	82.0
1090	280.0	1.4	32.3	0.0	22.0	78.0
1093	209.0	1.9	32.3	0.0	31.0	69.0
1095	100.0	1.9	32.3	2.0	79.0	19.0
1096	166.0	1.9	32.3	0.0	32.0	68.0
1097	240.0	1.9	32.3	0.0	10.0	90.0
1099	100.0	2.2	32.7	10.0	90.0	0.0
1100	155.0	1.9	32.3	0.0	47.0	53.0
1101	80.0	1.9	32.3	8.0	92.0	0.0
1102	234.0	3.7	33.0	0.0	7.0	96.0
1103	130.0	1.9	32.3	0.0	39.0	61.0
1104	150.0	1.9	32.3	0.0	35.0	65.0
1105	210.0	1.9	32.3	0.0	20.0	80.0
1106	170.0	1.9	32.3	0.0	32.0	68.0
1107	210.0	1.9	32.3	0.0	0.0	100.0
1109	90.0	1.9	32.3	2.0	98.0	0.0
1110	110.0	1.9	32.3	8.0	91.0	0.0
1111	140.0	1.9	32.3	0.0	92.0	8.0
1112	140.0	1.9	32.3	0.0	47.0	53.0
1113	113.0	2.3	33.3	0.0	72.0	28.0
1114	135.0	1.9	32.3	0.0	87.0	13.0
1115	130.0	1.9	32.3	41.0	59.0	0.0
1116	220.0	2.2	32.7	0.0	39.0	61.0
1119	140.0	1.9	32.3	0.0	39.0	61.0
1120	160.0	1.9	32.3	0.0	39.0	61.0
1124	83.0	1.9	32.3	1.0	99.0	0.0
1125	105.0	1.9	32.3	4.0	66.0	30.0
1126	112.0	1.9	32.3	2.0	55.0	43.0
1127	92.0	2.9	32.0	10.0	90.0	0.0
1128	130.0	2.9	32.0	10.0	90.0	0.0
1130	105.0	2.9	32.0	88.0	12.0	0.0
1131	118.0	2.9	32.0	0.0	40.0	60.0
1132	165.0	2.9	32.0	0.0	15.0	75.0
1133	171.0	8.5	34.3	0.0	15.0	75.0

1134	160.0	8.5	34.3	0.0	15.0	85.0
1135	170.0	8.5	34.3	0.0	15.0	85.0
1136	170.0	8.5	34.3	0.0	15.0	85.0
1137	135.0	8.5	34.3	0.0	54.0	46.0
1138	115.0	8.5	34.3	0.0	54.0	46.0
1139	155.0	8.5	34.3	0.0	15.0	85.0
1140	230.0	10.0	35.0	0.0	44.0	55.0
1141	160.0	10.0	35.0	0.0	40.0	60.0
1142	80.0	10.0	35.0	9.0	89.0	2.0
1143	92.0	6.1	33.6	8.0	92.0	0.0
1118	125.0	1.9	32.3	0.0	39.0	61.0
4003	77.0	4.0	31.8	99.0	1.0	0.0
4004	55.0	4.0	31.8	94.0	6.0	0.0
4005	55.0	4.0	31.8	97.0	3.0	0.0
4006	44.0	4.0	31.8	76.0	24.0	0.0
4007	83.0	4.0	31.8	99.0	1.0	0.0
4007	83.0	4.0	31.8	99.0	1.0	0.0

7902	602.0	9.0	35.5
7903	700.0	9.0	35.5
7904	798.0	9.0	35.5
7906	1000.0	6.0	35.0
7907	1200.0	4.0	35.0
7908	1400.0	4.0	35.0
7909	1600.0	4.0	35.5
7910	1800.0	4.0	35.5
7911	2400.0	4.0	35.0
7912	2200.0	4.0	35.0
7913	2000.0	4.0	35.0
7948	1680.0	4.0	35.0
7949	1800.0	4.0	35.0
7950	2000.0	4.0	35.0
7951	2200.0	4.0	35.0
7952	2400.0	4.0	35.0
8013	2500.0	4.0	35.0
8014	3050.0	4.0	35.0
8060	492.0	9.0	35.5
8061	592.0	9.0	35.5
8062	648.0	9.0	35.5
8063	990.0	9.0	35.5
8064	1708.0	4.0	35.0
8065	2120.0	4.0	35.0
8066	1350.0	4.0	35.0
8067	1050.0	6.0	35.5
8068	781.0	9.0	35.5
7802	834.0	9.0	35.5
6032	750.0	9.0	35.5
6033	970.0	9.0	35.5
6034	1045.0	6.0	35.5
6036	450.0	9.0	35.5
6050	350.0	9.0	35.5
6051	380.0	9.0	35.5
6076	777.0	9.0	35.5
6077	650.0	9.0	35.5
6078	400.0	9.0	35.5
1008	360.0	4.5	34.5
1036	440.0	4.5	34.5
1037	625.0	4.5	34.5
1038	775.0	4.5	34.5
1039	1000.0	4.5	34.5
1040	1600.0	4.5	34.5

APPENDIX VII

RAW ABUNDANCE DATA

The following list consists of the raw data for each species entered into the analysis. The column to the extreme left indicates each station number. Values in each column to the right of this are for % total occurrence (live+dead) of each species shown at the top of the column(in code),any value to the immediate right of each column represents % live occurrence of that species at that station.

## SPECIES ABBREVIATIONS AS USED IN THE ANALYSES

ALVESCIT	<u>Alveophragmium scitulum</u>
AMMOBACU	<u>Ammobaculites dilatatus</u>
AMMOCATI	<u>Ammodiscus catinus</u>
AMMOTCAS	<u>Ammotium cassis</u>
ASTRGALO	<u>Astronion gallowayi</u>
BOLISUBS	<u>Bolivina subspinescens</u>
BULIEXIL	<u>Bulimina exilis</u>
BULIACLE	<u>Bulimina aculeata</u>
BULISTRI	<u>Bulimina striata</u>
BUCCFRIG	<u>Buccella frigida</u>
BRIZSPAT	<u>Bolivina spathulata</u>
BRIZSUBA	<u>Bolivina subaenariensis</u>
CIBILOBA	<u>Cibicides lobatulus</u>
CIBIPSEU	<u>Cibicides pseudoungerianus</u>
CIBIWUEL	<u>Planulina wuellerstorfi</u>
CASSLAEV	<u>Cassidulina laevigata</u>
CRIBCRAS	<u>Cribrostomoides crassimargo</u>
CRIBJEFF	<u>Cribrostomoides jeffreysi</u>
CRIBSUBG	<u>Cribrostomoides subglobosum</u>
CYCLCANC	<u>Cyclammina cancellata</u>
ELPHEXCL	<u>Elphidium excavatum f. clavata</u>
ELPHBART	<u>Elphidium bartletti</u>
ELPHSUBA	<u>Elphidium subarcticum</u>
EGGLADVA	<u>Eggerella advena</u>
EPONUMBO	<u>Eponides sp.</u>
FISSERAT	<u>Fissurina serrata</u>

FURSLOEB	<u>Furstenkoina loeblichii</u>
FURSFUSI	<u>Furstenkoina fusiformis</u>
GLANLAEV	<u>Glandulina laevigata</u>
GYROSOLD	<u>Gyroidina soldani</u>
GLOBAURC	<u>Globobulimina auriculata</u>
GLOBTURG	<u>Globobulimina pacifica</u>
HOEGELGS	<u>Hoeglundina elegans</u>
HYPEEOLN	<u>Hyperammina elongata</u>
HAPLACAR	<u>Haplophragmoides canariensis</u>
ISLAISLA	<u>Islandiella islandica</u>
ISLATERE	<u>Islandiella teretis</u>
KAREINFL	<u>Karreriella bradyi</u>
LAGEMOLL	<u>Lagena mollis</u>
LENTSPAA	<u>Astocolus hyalacrulus</u>
NONILABR	<u>Nonionellina labradorica</u>
NONITURG	<u>Nonionella turgida</u>
MELOPOMP	<u>Melonis pompilioides</u>
PATEHAUR	<u>Pateroris hauerinoides</u>
PULEQUIN	<u>Pullenia quinqueloba</u>
QUICSEMI	<u>Quinqueloculina seminulum</u>
REOPSCOR	<u>Reophax scorpiurus</u>
REOPGUTT	<u>Reophax guttifera</u>
REOPNODU	<u>Reophax nodulosus</u>
REOPARCT	<u>Reophax arctica</u>
REOPSCOT	<u>Reophax scotti</u>
PAROSEMI	<u>Discamolina semipunctata</u>
SACCATLA	<u>Saccammina atlantica</u>
SACCSPHA	<u>Saccammina sphaerica</u>

SACCDIFF	<u>Saccammina difflugiformis</u>
STAINAPT	<u>Stainforthia concava</u>
TROCHINF	<u>Trochammina inflata</u>
TROCHLOB	<u>Trochammina lobata</u>
TROCHGLO	<u>Trochammina globigeriniformis</u>
TRIFANGU	<u>Trifarina angulosa</u>
TRITCONI	<u>Tritaxis conica</u>
TEXTTORQ	<u>Textularia torquata</u>
UVIGPERE	<u>Uvigerina peregrina</u>
ORIDTENR	<u>Oridorsalis umbonatus</u>
VALVLAEV	<u>Valvulineria laevigata</u>
RECUTURB	<u>Recurvoides turbinatus</u>
ADERGLOM	<u>Adercotryma glomerata</u>
SPIROBIF	<u>Spiroplectammina biformis</u>
TRILOCSP	<u>Triloculina trihedra</u>
ELPHACTT	<u>Elphidiella arctica</u>
SIGMSCHL	<u>Sigmoilopsis schlumbergeri</u>
OSANGSPA	<u>Osangularia rugosa</u>
TRIFOCCI	<u>Trifarina occidentalis</u>
TROCHSQA	<u>Trochammina squamata</u>
MBULIMAR	<u>Bulimina marginata</u>

## PERCENT FOR NS GRABS AND CORE TOPS

% OF VARIABLES = 75 NUMBER OF SAMPLES = 258  
 PERCENT FOR NS GRABS AND CORE TOPS

## RAW DATA

SAMPLE	NAME	1	2	3	4	5	6	7	8	9	10
		ALVESKIT	AMMOMBACU	AMMOCATI	AMMOTCAS	ASTRGALO	BOLISUBS	BULIEXIL	BULISTRI	BULIACLE	BUCCFRIG
7902		1.000	0.000	0.000	.300	0.000	0.000	19.000	0.000	3.500	5.500
7903		4.400	4.300	0.000	0.000	0.000	0.000	10.200	0.000	13.600	1.700
7904		1.400	2.400	0.000	0.000	0.000	0.000	7.800	0.000	0.800	0.000
7906		.800	1.300	0.000	0.000	0.000	4.000	3.000	0.000	4.000	4.000
7907		0.000	0.000	0.000	0.000	0.000	.800	.800	1.800	3.000	0.000
7908		0.000	1.300	0.000	0.000	0.000	.600	.600	3.400	2.900	0.300
7909		0.000	.700	0.000	0.000	0.000	0.000	0.000	2.800	1.000	0.000
7910		0.000	0.000	0.000	0.000	0.000	.700	1.300	0.000	0.000	0.000
7911		0.000	0.000	0.000	0.000	0.000	0.000	0.000	2.700	0.000	0.000
7912		0.000	0.000	0.000	0.000	0.000	0.000	0.000	2.900	0.000	0.000
7913		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.700	0.000	0.000
7914		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7917		.700	0.000	0.000	0.000	0.000	0.000	1.000	0.000	32.900	0.000
7916		0.000	0.000	0.000	0.000	0.000	0.000	1.800	0.000	30.800	0.000
7948		0.000	0.000	0.000	0.000	0.000	1.000	1.700	5.000	0.000	0.000
7949		0.000	0.000	0.000	0.000	0.000	.800	0.000	4.300	0.000	0.000
7950		0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.700	0.000	0.000
7952		0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.700	0.000	0.000
7953		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	.600	2.500 9.0
7955		0.000	0.700	0.000	0.000	0.000	.700	3.6	0.000	0.000	2.000
7956		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.600	.500
7957		0.000	0.000	0.000	0.000	0.000	.300	3.3	0.000	0.000	11.300 26
7958		1.300	0.000	0.000	0.000	0.000	.600	0.000	0.000	4.200	.500
7959		0.300 22	0.000	0.000	0.000	0.000	0.000	0.000	0.000	10.000	4.000
7960		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	13.000	1.400
7947		0.000	0.020	0.000	0.000	0.000	0.000	0.000	1.800	7.400	5.000
0816		0.000	0.000	0.000	14.000	0.000	0.000	0.000	0.000	0.000	0.000
0817		0.000	0.000	0.000	1.000	0.000	0.000	0.000	0.000	0.000	0.000
0818		0.000	0.000	0.000	0.000	0.000	1.200	0.000	0.000	0.000	0.000
0819		0.000	0.000	0.000	0.000	.500	0.000	0.000	0.000	0.000	3.000 17
0820		0.000	0.000	0.000	0.000	5.600	0.000	0.000	0.000	0.000	0.000
0821		0.000	0.000	0.000	0.000	5.000	0.000	0.000	0.000	0.000	3.200
0822		0.000	0.000	0.000	0.000	6.000	0.000	0.000	0.000	0.000	0.000
0824		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0825		0.000	0.000	0.000	0.000	.800	0.000	0.000	0.000	0.000	0.000
0830		0.000	0.000	0.000	0.000	3.70	9	0.000	0.000	0.000	0.000
0832		0.000	0.000	0.000	0.000	3.200	1	0.000	0.000	0.000	0.000
0833		0.000	0.000	0.000	0.000	4.00	7	0.000	0.000	0.000	0.000
0835		0.000	0.000	0.000	.400	11.60	19	0.000	0.000	0.000	0.000
0838		0.000	0.000	0.000	.300	.600	0.000	0.000	0.000	0.000	0.000
0839		0.000	0.000	0.000	0.000	2.000	0.000	0.000	0.000	0.000	0.000
0841		0.000	0.000	0.000	0.000	12.300	0.000	0.000	0.000	0.000	0.000
0843		0.000	0.000	0.000	0.000	0.000	1.400	0.000	0.000	0.000	1.700 5
0844		0.000	0.000	1.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0846		0.000	.600	0.000	3.800	0.000	0.000	0.000	0.000	0.000	0.000
0847		0.000	0.000	0.000	0.000	4.300	11	0.000	0.000	0.000	1.000
0848		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.000
0849		1.500	0.000	0.000	0.000	0.000	0.000	0.400	0.000	1.800	0.000

0851	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	23.400	0.000
0852	2.600	0.000	0.000	0.000	0.300	0.000	0.000	0.000	23.500	0.900
0856	0.000	0.000	0.000	0.000	1.600	0.000	0.000	0.000	2.900	0.000
0857	0.000	0.000	0.000	0.000	1.400	0.000	0.000	0.000	0.000	0.000
9002	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.600	0.000
9004	0.000	0.000	0.000	0.000	0.800	0.000	0.000	0.000	30.000	2.300
9005	0.700	0.000	0.000	0.000	0.200	0.000	0.000	0.000	22.000	10.800
9006	0.000	0.000	0.000	0.000	0.000	1.000	0.000	0.000	19.000	8.600
9007	0.300	0.000	0.000	0.000	0.000	0.600	3.300	0.000	7.500	13.000
9010	0.000	0.000	0.000	0.000	0.000	0.300	0.000	0.000	11.600	12.600
9011	0.000	0.000	0.000	0.000	0.000	0.030	0.000	0.000	10.800	16.800
9012	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	12.800	11.800
9013	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
9018	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
9027	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.300	4.000
9035	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
9036	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
9037	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
9039	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
9039	0.000	1.200	0.000	2.500	16	0.000	0.000	0.000	0.000	0.000
9040	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
9042	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
9043	0.000	0.000	0.000	0.000	1.300	2	0.000	0.000	0.000	0.000
9044	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
9060	0.500	0.000	0.000	0.000	0.000	4.200	67	0.000	0.000	10.700
8061	0.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000	17.300	10.400
8062	3.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	10.700	0.000
8063	2.600	0.000	0.000	0.000	0.000	1.300	0.000	0.000	5.200	2.000
8064	0.900	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.300	1.700
8065	0.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
f057	1.100	10	0.000	0.000	0.000	0.000	0.000	0.000	3.500	0.000
6003	1.900	0.000	0.000	0.000	0.000	0.000	0.000	0.000	21.500	0.000
6005	2.500	0.000	1.300	0.000	0.000	0.000	0.000	0.000	6.800	12.800
6006	0.000	0.000	0.000	0.000	0.000	2.100	0.000	0.000	2.800	6.500
6007	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.600	6.600
6019	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6032	1.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000	9.600	4.300
6031	0.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000	5.800	0.000
6034	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	7.000	0.000
6036	2.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	10.600	10.600
6041	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	5.000	2.400
6042	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	5.700
6050	0.000	0.000	0.000	0.000	0.000	1.200	0.000	0.000	0.000	7.000
6051	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	5.000
6052	1.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	17.000	7.000
6054	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	17.000	7.000
6055	0.000	0.000	0.000	0.000	0.000	1.100	0.000	0.000	0.000	16.800
6057	0.000	0.000	0.000	0.000	0.000	0.900	0.000	0.000	1.400	0.000
6058	0.000	0.000	0.000	0.000	0.000	3.000	0.000	0.000	0.000	0.000
6062	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	14.600	4.000
6063	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6064	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	14.600	4.000
6065	1.400	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.900	0.000
6074	2.700	0.000	0.000	0.000	0.000	1.700	0.000	0.000	1.700	5.500
6075	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	2.700
6076	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	4.000
6077	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.000
6079	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6072	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.000
6073	0.000	0.000	0.000	0.500	0.000	0.000	0.000	0.000	0.000	1.000
6076	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.000
6075	2.400	0.000	0.000	0.000	0.000	3.700	0.000	1.200	1.700	0.000





1134	0.000	0.000	1.200	0.000	0.030	0.000	0.000	0.000	10.400	0.000
1135	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	13.500	4.200
1136	2.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000	10.900	3.000
1137	2.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000	10.000	0.000
1138	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1139	1.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.500	0.050
1140	1.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	20.300	1.000
1141	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.000	3.900
1142	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	8.200	2
1143	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
4003	0.000	0.000	0.000	0.000	0.000	0.900	0.000	0.000	2.000	0.000
4004	0.000	0.000	0.000	0.000	0.000	0.500	0.000	0.000	0.000	2.200
4005	0.000	0.000	0.000	0.000	0.000	0.300	0.000	0.000	0.000	2.400
4006	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	.900
4007	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000

SAMPLE	NAME	11	12	13	14	15	16	17	18	19	20
	BRIZSPAT	BRI7SUBA	CIB1LCRA	CIB1PSFU	CIB1WIUEL	CASSLAEV	CP1BCRAS	CRIBJFFF	CRISUBG	CYCLCANC	
7902	3.500	0.000	4.500	1.000	0.000	2.000	0.000	0.000	0.000	0.000	0.000
7903	0.000	0.000	3.000	3.000	0.000	2.000	0.000	0.000	0.000	0.000	1.000
7904	0.000	0.000	2.000	1.400	0.000	2.000	0.000	0.000	0.000	0.000	0.000
7907	0.000	0.000	4.000	2.600	0.000	3.000	1.300	0.000	0.000	0.000	0.000
7908	0.000	0.000	5.000	5.500	0.000	4.000	0.000	0.000	0.000	0.000	0.000
7909	0.000	0.000	0.000	1.800	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7910	0.000	0.000	0.000	1.300	0.000	1.000	0.000	0.000	0.000	0.000	0.700
7911	0.400	0.000	0.000	.700	0.000	0.000	0.000	0.400	0.000	0.000	0.000
7912	0.000	0.000	0.000	0.000	0.300	0.000	0.000	0.000	0.000	0.000	0.000
7913	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7914	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7917	0.000	1.000	6.800	0.000	0.000	2.000	0.000	0.000	2.900	0.500	0.000
7934	0.000	6.400	7.000	0.000	0.000	0.000	0.000	0.000	1.600	0.000	0.000
7942	0.000	0.000	1.070	1.70	0.000	0.000	0.000	0.000	0.300	0.000	0.000
7949	0.000	0.000	0.000	1.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7950	0.000	0.000	0.000	0.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7952	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7953	6.000	0.300	18.000	18	0.000	2.000	0.000	0.000	0.000	0.000	0.000
7955	4.000	11	0.000	14.700	4	0.000	0.000	2.000	0.000	0.000	0.000
7956	0.000	0.000	15.200	0.000	0.000	2.600	6	0.000	0.000	0.000	0.000
7957	3.000	10	0.000	7.000	0.000	2.400	0.000	0.000	0.000	0.000	0.000
7958	0.000	24.000	12.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7959	0.300	1.600	1.300	0.000	0.000	1.200	5	0.000	0.000	0.000	0.000
7960	0.000	3.500	2.500	0.000	0.000	0.000	0.000	0.000	0.700	0.000	0.000
7947	0.000	0.000	0.000	19.000	0.000	0.000	0.600	1.200	0.000	0.000	1.200
0F02	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0F17	0.000	0.000	0.000	0.000	0.000	0.000	0.000	22.600	0.000	0.000	0.000
0F18	0.000	0.000	77.000	83	0.000	0.000	0.000	1.500	0.000	0.000	0.000
0819	0.000	0.000	1.500	0.000	0.000	0.000	0.000	12.900	11	500	0.000
0820	0.000	0.000	0.000	0.000	0.000	0.000	0.000	11.200	0	0.000	0.000
0821	0.000	0.000	23.000	0.000	0.000	0.000	0.000	11.000	16	0.000	0.000
0822	0.000	0.000	0.200	0.000	0.000	0.000	0.000	0.700	0.000	0.000	0.000
0824	0.000	0.000	0.000	0.000	0.000	0.000	0.000	6.700	0.000	0.000	0.000
0825	0.000	0.000	0.300	0.000	0.000	0.000	0.000	6.000	0.000	0.000	0.000
0830	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.100	0.000	0.000	0.000
0832	0.000	0.000	60.000	67	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0833	0.000	0.000	0.000	0.000	0.000	0.000	0.000	25.500	42	0.000	0.000
0835	0.000	0.000	0.000	0.000	0.000	0.000	0.000	2.700	8	0.000	0.000
0838	0.000	0.000	0.000	0.000	0.000	0.000	0.000	2.700	7	0.000	0.000
0839	0.000	0.000	0.000	0.000	0.000	0.000	0.000	6.300	0.000	0.000	0.000
0841	0.000	0.000	0.000	0.000	0.000	0.000	0.000	14.800	0.000	0.000	0.000
0843	0.000	0.000	0.000	0.000	0.000	0.000	0.000	6.200	16	0.000	0.000
0844	0.000	0.000	0.400	0.000	0.000	0.000	0.000	15.800	1.100	0.000	0.000
0846	0.000	0.000	0.000	0.000	0.000	0.000	0.000	14.800	0.000	0.000	0.000

0847	0.000	0.000	41,400	27	0.000	0.000	0.000	2,900	8	0.700	0.000	0.000
0848	0.000	0.000	5,200	12	0.000	0.000	0.000	8,900		0.000	0.000	0.000
0849	0.000	7,300	2,200		0.000	0.000	0.000	0.000		1,500	18	
0851	0.000	15,700	0,000		0.000	0.000	0.000	0.000		0.000	0.000	0.000
0852	0.000	11,200	9,300		0.000	0.000	0.000	0.000		0.000	0.000	0.000
0856	.300	1,500	5,100		0.000	0.000	0.000	0.000		0.000	0.000	0.000
0857	0.000	0,500	4,900		0.000	0.000	0.000	0.000		0.000	0.000	0.000
P002	11,200	9,400	2,000		10,300	0.000	0.000	5,400		0.000	0.000	0.000
P004	2,400	2,000	1,200		0.000	0.000	0.000	0.000		0.000	0.000	0.000
P005	3,500	7,000	1,300	18	0.000	0.000	0.000	0.000		0.000	0.000	0.000
P006	0.000	4,000	0,000		2,700	0.000	0.000	0.000		0.000	0.000	0.000
P007	4,400	6,800	0,000		0.000	0.000	0.000	0.000		0.000	0.000	0.000
P010	9,300	5,000	1,300		0.000	0.000	0.000	0.000		0.000	0.000	0.000
P011	4,000	5,000	0,000		0.000	0.000	0.000	0.000		0.000	0.000	0.000
P012	1,700	8	3,000		0.000	0.000	0.000	0.000		0.000	0.000	0.000
P013	0,000	0,000	0,000		2,400	1,400	0.000	0.000		0.300	0.000	0.000
P017	1,900	33	13,900		0.300	0.000	0.000	0.000		0.000	0.000	0.000
P018	6,800	44	3,000	11	11,600	0.000	0.000	5,900	33	0.000	0.000	0.000
P027	9,500	77	0,000	16,000	0.000	0.000	0.000	1,200		0.000	0.000	0.000
P035	0.000	0,000	0,000		0.000	0.000	0.000	0.000		2,600	50	0.000
P036	11,200	92	0,200	19,300	0.000	0.000	0.000	1,400		0.000	0.000	0.000
P037	0.000	0,000	0,000		0.000	0.000	0.000	0.000		8,600	0.000	0.000
P038	0.000	0,000	0,000		0.000	0.000	0.000	0.000		0.000	0.000	0.000
P039	0.000	0,000	0,000		0.000	0.000	0.000	0.000		0.000	0.000	0.000
P040	0.000	0,000	0,000		0.000	0.000	0.000	0.000		18,800	17	0.000
P042	0.000	0,000	0,000		0.000	0.000	0.000	0.000		15,800	8	1,000
P043	0.000	0,000	0,000		0.000	0.000	0.000	0.000		16,200	16	0.000
P044	0.000	0,000	0,000		0.000	0.000	0.000	0.000		300	0.000	0.000
P060	0.000	0,000	0,000		0.000	0.000	0.000	0.000		0.000	0.000	0.000
P061	2,000	0,000	4,100		1,000	0.000	0.000	0.000		0.000	0.000	0.000
P062	0,000	0,000	0,000		4,300	0.000	0.000	1,700		0.000	0.000	0.400
P063	0,000	0,000	0,000		7,400	0.000	0.000	1,700		0.000	0.000	1,300
P064	0,000	0,000	0,000		1,000	0.000	0.000	0.000		0.000	0.000	0.000
P065	0,000	0,000	0,000		0,000	0.000	0.000	0,000		0.000	0.000	0.000
P066	0,000	0,000	0,000		0,000	0.000	0.000	0,000		0.000	0.000	0.000
P067	0,000	0,000	0,000		0,000	0.000	0.000	0,000		0.000	0.000	0.000
P003	0,000	0,000	1,300		0,000	0.000	0,000	0,000		0,000	0.000	0.000
P005	0,000	5,500	0,000		0,000	0.000	0,000	0,000		0,000	0.000	0.000
P006	1,500	0,000	6,000		0,000	0.000	0,000	4,300		0,000	0.000	0.000
P007	6,100	0,000	3,600		0,000	0.000	0,000	2,100		0,000	0.000	0.000
P018	2,500	0,000	20,700		0,000	0.000	0,000	2,800		0,000	0.000	0.000
P032	0,100	0,000	0,000		1,000	0.000	0,000	0,000		1,000	1,400	0,000
P033	4,000	0,000	0,000		0,000	0.000	0,000	5,400		0,000	0,000	0,700
P034	3,700	0,000	2,700		0,000	0.000	0,000	5,000		0,000	0,000	0,000
P036	0,000	0,000	4,100		0,000	0.000	0,000	0,000		0,000	0.000	0,000
P041	2,100	0,000	9,000		0,000	0.000	0,000	2,600		0,000	0,000	0,000
P042	16,200	0,000	16,800		0,000	0.000	0,000	1,100		0,000	0.000	0,000
P043	0,000	0,000	0,000		0,000	0.000	0,000	3,300		0,000	0,000	0,000
P051	41,000	0,000	8,700		0,000	0.000	0,000	2,700		0,000	0,000	0,000
P052	0,000	12,500	0,900		0,000	0.000	0,000	0,300		600	1,900	5
P053	0,000	19,000	0,300		0,000	0.000	0,000	0,000		0,000	0,000	0,000
P055	0,000	40,000	0,000		0,000	0.000	0,000	0,000		0,000	0,000	0,000
P056	0,000	2,700	0,000		0,000	0.000	0,000	0,000		0,000	0,000	2,200
P057	0,000	0,000	0,000		0,000	0.000	0,000	0,000		14,700	1,800	0,000
P058	0,000	22,000	0,000		0,000	0.000	0,000	1,100		0,000	0,000	0,000
P062	6,500	0,000	1,800		0,000	0.000	0,000	400		0,000	0,000	0,000
P063	0,700	0,000	0,000		0,000	0.000	0,000	2,800		7,800	9,000	0,000
P065	0,000	2,000	0,000	6	0,000	0.000	0,000	0,000		1,900	6	1,500
P066	0,400	0,000	18,000		0,000	0.000	0,000	3,800		0,000	0,000	0,000
P069	0,000	0,300	32,200	21	0,000	0.000	0,000	0,000		0,000	0,000	0,000
P070	0,000	0,000	20,600	25	0,000	0.000	0,000	0,000		0,000	0,000	0,000
P071	1,300	0,000	23,000		0,000	0.000	0,000	0,000		0,000	0,000	0,000
P072	1,400	0,000	27,400		0,000	0.030	0,800	0,000		0,000	0,000	0,000



1053	£.000	0.000	88.400	100	0.000	0.000	0.000	0.000	0.000	3.000	0.000
1054	£.000	0.000	74.130	85	0.000	0.000	7.500	0.000	0.000	3.000	0.000
1055	£.000	0.000	4.400	11	0.000	0.000	1.700	0.000	0.000	0.000	0.000
1056	£.000	1.500	5.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1057	£.000	0.000	25.000	0.000	0.000	0.000	12.100	6	0.000	0.000	0.000
1058	£.000	0.000	0.000	0.000	0.000	0.000	20.200	13	0.000	0.000	0.000
1059	£.000	0.000	0.000	0.000	0.000	0.000	3.200	0.000	0.000	0.000	0.000
1060	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1061	£.000	0.000	0.000	0.000	0.000	0.000	1.200	0.000	0.000	0.000	0.000
1062	£.000	0.000	0.000	0.000	0.000	0.000	2.200	1	0.000	0.000	0.000
1063	£.000	0.000	0.000	0.000	0.000	0.000	3.400	6	0.000	0.000	0.000
1064	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1065	£.000	0.000	0.000	0.000	0.000	0.000	14.000	11	0.000	0.000	0.000
1066	£.000	0.000	30.000	10	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1067	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1068	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1069	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1070	£.000	0.000	1.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1071	£.000	0.000	74.100	71	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1072	£.000	0.000	70.000	86	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1073	£.000	0.000	0.000	0.000	0.000	0.000	17.000	8	0.000	0.000	0.000
1074	£.000	0.000	0.000	0.000	0.000	0.000	3.600	0.000	0.000	0.000	0.000
1075	£.000	0.000	0.000	0.000	0.000	0.000	2.900	3	0.000	0.000	0.000
1076	£.000	0.000	0.000	0.000	0.000	0.000	2.200	7	0.000	0.000	0.000
1077	£.000	0.000	0.000	42.100	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1078	£.000	0.000	0.000	38.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1080	£.000	0.000	0.000	64.000	0.000	0.000	0.000	1.400	0.000	0.000	0.000
1082	£.000	0.000	0.000	0.000	0.000	0.000	1.900	0.000	0.000	0.000	0.000
1083	£.000	0.000	0.000	0.000	0.000	0.000	0.900	0.000	0.000	0.000	0.000
1084	£.000	0.000	0.000	1.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1085	£.000	0.000	0.000	43.400	62	0.000	0.000	0.000	0.000	0.000	0.000
1086	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1087	£.000	0.000	0.000	0.000	0.000	0.000	13.700	0.000	0.000	0.000	0.000
1088	£.000	0.000	0.000	0.000	0.000	0.000	5.400	0.000	0.000	0.000	0.000
1089	£.000	0.000	0.000	0.000	0.000	0.000	4.200	0.000	0.000	0.000	0.000
1090	£.000	0.000	0.000	0.000	0.000	0.000	4.200	0.000	0.000	0.000	0.000
1093	£.000	0.000	0.000	0.300	0.000	0.000	0.300	11.700	0.000	0.000	0.000
1095	£.000	0.000	0.000	1.700	0.000	0.000	0.000	23.300	20.400	29	0.000
1098	£.000	0.000	0.000	0.200	0.000	0.000	2.400	3.900	0.000	0.000	0.000
1099	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	13.600	0.000	0.000
1100	£.000	0.000	0.000	13.500	25	0.000	0.000	0.000	0.000	0.000	0.000
1101	£.000	0.000	0.000	0.000	0.000	0.000	1.000	0.000	0.000	0.000	0.000
1102	£.000	0.000	0.000	0.000	0.000	0.000	10.200	0.000	0.000	0.000	0.000
1103	£.000	0.000	0.000	0.000	0.000	0.000	1.800	0.000	0.000	0.000	0.000
1104	£.000	0.000	0.000	0.000	0.000	0.000	3.200	0.000	0.000	0.000	0.000
1105	£.000	0.000	0.000	0.000	0.000	0.000	3.000	0.000	0.000	0.000	0.000
1106	£.000	0.000	0.000	0.000	0.000	0.000	1.000	0.000	0.000	0.000	0.000
1107	£.000	0.000	0.000	0.000	0.000	0.000	15.700	0.000	0.000	0.000	0.000
1109	£.000	0.000	0.000	1.000	0.000	0.000	2.400	2	0.000	0.000	0.000
1110	£.000	0.000	0.000	0.900	2	0.000	0.000	0.000	0.000	0.000	0.000
1111	£.000	0.000	0.000	0.800	2	0.000	0.000	0.000	0.000	0.000	0.000
1112	£.000	0.000	0.000	0.000	0.000	0.000	4.300	2	0.000	0.000	0.000
1114	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1115	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1116	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1117	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1119	£.000	0.000	0.000	0.800	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1120	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1124	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1125	£.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1126	£.000	0.000	0.000	0.300	0.000	0.000	2.200	9.200	22	0.000	0.000
1127	£.000	0.000	0.000	0.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1128	£.000	0.000	0.000	0.000	0.000	0.000	2.700	2.700	8.000	8	0.000
1130	£.000	0.000	0.000	5.300	0.000	0.000	2.100	1	8.000	8	0.000

1131	0.000	0.000	0.000	0.000	0.000	0.000	1.100	12.800	27	0.000	0.000
1132	0.000	0.000	0.000	0.000	0.000	0.000	0.000	.500	10	0.000	0.000
1133	0.000	12.500	26	1.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1134	0.000	10.400	0.800	0.000	0.000	0.000	1.800	3.100	0.000	0.000	0.000
1135	0.000	4.000	0.800	0.000	0.000	0.000	3.600	1.200	0.000	0.000	0.000
1136	0.000	10.500	100	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1137	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1138	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1139	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1140	7.500	0.000	0.000	2.600	0.000	0.000	1.200	2.200	0.000	0.000	0.000
1142	0.000	0.000	41.000	47	0.000	0.000	0.000	0.000	0.000	1.200	4
1143	0.000	0.000	23.200	32	0.000	0.000	0.000	0.000	0.000	6.400	62
1116	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	4.100	3
4003	0.000	0.000	25.000	0.000	0.000	0.000	0.000	3.000	0.000	0.000	0.000
4004	0.000	0.000	45.800	48	0.000	0.000	0.000	0.000	0.000	0.800	0.000
4005	0.000	0.000	48.300	0.000	0.000	0.000	0.000	2.100	0.000	0.000	0.000
4006	0.000	0.000	66.600	100	0.000	0.000	0.000	0.000	0.000	0.900	0.000
4007	0.000	0.000	73.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
SAMPLE NAME	21	22	23	24	25	26	27	28	29	30	
	ELPHEXCL	FLPHBART	FLPHSURA	EGGLADVA	FPONUMBER	FISSERAT	FURSLNEB	FURSFUSI	GLANLAEV	GYROSOLD	
7902	11.000	1.700	1.000	.300	0.000	0.000	0.000	1.000	11	1.400	1.000
7903	9.500	2.600	0.000	.900	0.000	0.000	0.000	0.000	0.000	0.900	2.600
7904	18.600	0.000	0.000	2.900	0.000	0.000	0.000	0.000	0.000	0.000	1.000
7904	24.700	.800	0.000	4.400	0.000	0.000	0.000	0.000	0.000	.800	1.800
7905	35.800	0.000	0.000	1.500	0.000	0.000	0.000	0.000	0.000	0.300	1.000
7906	38.500	0.000	0.000	.600	2.000	0.000	0.000	0.000	0.000	0.000	0.000
7909	47.500	0.000	0.000	.400	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7910	42.300	0.000	0.000	1.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7912	60.600	0.000	0.000	.400	0.000	0.000	0.000	0.000	0.000	0.000	0.400
7913	63.100	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.600
7914	43.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.300
7917	19.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.200
7918	26.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7948	27.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7949	28.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7950	37.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.700
7952	29.400	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.600
7953	18.200	4.700	1.400	0.000	0.000	0.000	0.000	0.000	0.000	.300	1.200
7955	30.700	3.700	1.700	4	0.000	0.000	0.000	0.000	0.000	1.000	0.300
7956	37.500	6.600	1.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7957	40.500	2.200	10	0.600	0.000	0.000	0.000	0.000	0.000	0.900	0.000
7958	30.400	2.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7959	13.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.200	0.000
7960	6.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7947	25.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0802	0.000	0.000	0.000	63.600	92	0.000	0.000	0.000	0.000	0.000	0.000
0817	0.000	0.000	0.000	18.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0818	1.700	1.000	4.300	.400	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0819	0.000	0.000	0.000	1.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0820	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0821	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0824	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0825	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0832	20.000	3.400	0.000	82.200	100	0.000	0.000	0.000	0.000	0.000	0.000
0833	0.000	0.000	0.000	52.000	91	0.000	0.000	0.000	0.000	0.000	0.000
0834	0.000	0.000	0.000	36.800	38	0.000	0.000	0.000	0.000	0.000	0.000
0839	0.000	0.000	0.000	56.000	71	0.000	0.000	0.000	0.000	0.000	0.000
0841	0.000	0.000	0.000	35.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000



6070	2.700	5.000	0.000	1.500	0.000	0.000	0.000	.300	10	0.000	0.000
6071	2.600	5.300	0.000	.400	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6072	13.800	7.900	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	.600
6073	4.500	1.800	0.000	0.000	0.000	0.000	0.000	0.000	0.000	2.400	1.300
6074	17.600	3.100	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6075	17.500	.600	0.000	0.000	0.000	1.200	0.000	0.000	4.500	.800	0.000
6076	7.500	0.000	0.000	0.000	0.000	2.400	0.000	0.000	1.500	0.000	0.000
6077	25.600	1.700	0.000	0.000	0.000	0.000	0.000	0.000	1.200	0.000	.800
6078	9.300	1.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.800	0.000
6079	6.700	5.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6080	13.600	10.700	0.000	0.000	0.000	0.000	0.000	0.000	3.400	0.000	0.000
6081	2.100	4.100	0.000	0.000	0.000	0.000	0.000	0.000	1.400	0.000	0.000
6082	10.000	0.000	0.000	3.700	0.000	0.000	0.000	11.600	0.000	0.000	0.000
6083	15.000	2.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6084	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.000
6085	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6086	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6087	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6088	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6089	1.200	1.400	0.000	2.800	0.000	0.000	0.000	0.000	0.900	0.000	0.000
6090	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6091	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.200	0.000	0.000
6092	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.100	0.000	0.000
6093	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6094	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6095	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6096	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6097	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6098	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6099	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6100	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6101	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6102	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6103	15.000	2.000	0.000	2.500	0.000	0.000	0.000	0.000	3.200	0.000	0.000
6104	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6105	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6106	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6107	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6108	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6109	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6110	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.800	6.600	0.000
6111	0.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.200	0.000	0.000
6112	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.200	1.200	0.000
6113	0.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6114	0.000	0.000	0.000	0.000	1.000	0.000	0.000	0.000	0.000	1.500	0.000
6115	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6116	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6117	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.400	5.300	0.000
1006	7.400	1.600	0.000	0.000	0.000	0.000	0.000	0.000	10.000	0.000	.700
1009	2.100	1.400	0.000	1.400	0.000	0.000	0.000	0.000	3.800	0.000	0.000
1010	0.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.800	4.700	0.000
1011	0.000	0.000	0.000	0.000	5.600	0.000	0.000	0.000	0.000	1.000	0.000
1015	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1016	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1017	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1020	3.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1023	1.400	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1024	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1026	3.000	2.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1028	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1034	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1035	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1036	5.500	4.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.300
1037	8.500	3.600	0.000	1.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1038	3.600	0.000	0.000	0.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1039	1.300	0.000	0.000	0.000	0.000	1.900	0.000	0.000	0.000	1.500	0.000
1040	1.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1043	0.000	0.000	0.000	0.000	6.300	0.000	0.000	0.000	0.000	0.000	0.000
1044	0.000	0.000	0.000	0.000	1.400	0.000	0.000	0.000	0.000	0.000	0.000
1045	0.000	0.000	0.000	0.000	4.600	0.000	0.000	0.000	0.000	0.000	0.000
1046	.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.000	0.000	0.000

047	0.000	3.300	0.000	0.000	0.000	0.000	5.000	10	0.000
1048	12.000	2.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1051	0.000	2.500	0.000	0.000	3.700	12	0.000	0.000	0.000
1053	0.000	0.000	.700	0.000	0.000	0.000	0.000	0.000	0.000
1054	0.000	0.000	5.100	0.000	0.000	0.000	0.000	0.000	0.000
1055	0.000	0.000	0.000	0.000	8.200	20	0.000	0.000	0.000
1056	0.000	0.000	0.000	0.000	17.200	59	0.000	0.000	0.000
1057	0.000	0.000	0.000	0.000	2.000	0.000	0.000	0.000	0.000
1058	0.000	0.000	0.000	0.000	1.600	0.000	0.000	0.000	0.000
1059	0.000	1.000	0.000	0.000	6.800	0.000	0.000	0.000	0.000
1060	0.000	0.000	0.000	0.000	.700	0.000	0.000	0.000	0.000
1062	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1063	0.000	0.000	0.000	0.000	8.000	7	0.000	0.000	0.000
1064	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1065	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1066	0.000	0.000	2.700	0.000	0.000	0.000	0.000	0.000	0.000
1067	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1068	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1069	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1070	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1071	1.500	0.000	4.400	0.000	0.000	0.000	0.000	0.000	0.000
1072	0.000	0.000	3.900	0.000	2.000	0.000	0.000	0.000	0.000
1073	0.000	0.000	0.000	0.000	4.600	0.000	0.000	0.000	0.000
1074	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1075	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1076	0.000	0.000	0.000	0.000	2.900	0.000	0.000	0.000	0.000
1077	0.000	0.000	4.600	0.000	0.000	0.000	0.000	0.000	0.000
1078	0.000	0.000	.500	1.300	0.000	0.000	0.000	0.000	0.000
1080	0.000	0.000	1.600	0.000	0.000	0.000	0.000	0.000	0.000
1081	0.000	0.000	0.000	1.800	3	0.000	0.000	0.000	0.000
10F2	0.000	0.000	1.200	0.000	0.000	0.000	0.300	0.000	0.000
1083	0.000	0.000	0.000	4.400	0.000	0.000	0.000	0.000	0.000
1084	0.000	1.500	0.000	1.300	0.000	0.000	0.000	0.000	0.000
1084	0.000	2.500	0.000	.800	1.800	0.000	0.000	0.000	0.000
1087	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
10P8	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1089	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1090	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1093	0.000	0.000	0.000	0.000	1.300	0.000	0.000	0.000	0.000
1095	0.000	0.000	0.000	2.500	0.000	0.000	0.000	0.000	0.000
1097	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1099	0.000	0.000	0.000	0.000	1.900	0.000	0.000	0.000	0.000
11C0	0.000	0.000	0.000	0.000	3.400	0.000	0.000	0.000	0.000
11C1	0.000	0.000	.800	6.600	8	0.000	0.000	0.000	0.000
11C2	0.000	0.000	0.000	3.100	0.000	0.000	0.000	0.000	0.000
11C3	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
11C4	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
11C5	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
11C6	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
11C7	0.000	0.000	0.000	0.000	1.200	0.000	0.000	0.000	0.000
11C9	0.000	0.000	0.000	0.000	11.800	11	0.000	0.000	0.000
1110	0.000	0.000	0.000	72.500	0.000	0.000	0.000	0.000	0.000
1111	0.000	0.000	0.000	6.300	8	0.000	0.000	0.000	0.000
1112	0.000	0.000	0.000	1.000	0.000	0.000	0.000	0.000	0.000
1113	0.000	0.000	0.000	0.200	2	0.000	0.000	0.000	0.000
1114	0.000	0.000	0.000	4.500	0.000	0.000	0.000	0.000	0.000
1115	0.000	0.000	0.000	5.000	3	0.000	0.000	0.000	0.000
1116	0.000	0.000	0.000	6.600	4	0.000	0.000	0.000	0.000
1120	0.000	0.000	0.000	14.500	16	0.000	0.000	0.000	0.000
1124	0.000	0.000	0.000	4.600	47	0.000	0.000	0.000	0.000
1125	0.000	0.000	0.000	3.900	0.000	0.000	0.000	0.000	0.000
1126	0.000	0.000	0.000	6.800	0.000	0.000	0.000	0.000	0.000

1127	0.000	0.000	0.000	3.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1128	0.000	0.000	0.000	0.700	1	0.000	0.000	0.000	0.000	0.000	0.000
1130	0.000	0.100	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1131	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1132	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1133	0.000	0.000	0.000	0.600	1.000	0.000	0.000	1.200	0.5100	0.000	0.000
1134	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.100	0.5100	0.000	0.000
1135	2.400	0.000	0.000	0.000	0.000	0.000	0.000	3.800	6.700	0.000	0.000
1136	0.700	0.000	0.000	0.000	0.000	2.200	0.000	0.000	4.500	0.000	0.000
1137	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1138	0.000	0.000	0.000	0.800	1.000	0.000	0.000	0.000	0.000	0.000	0.000
1139	0.000	0.000	0.000	0.000	0.000	2.700	0.000	1.500	1.900	0.000	0.000
1140	2.000	0.000	0.000	0.500	1.000	0.000	0.000	1.600	23.300	77	0.000
1142	0.000	5.800	2	0.000	0.000	0.000	0.000	0.000	5.900	7	0.000
1143	0.000	6.400	0.000	0.000	0.000	0.000	0.000	1.800	4.100	5	0.000
111 <sup>a</sup>	0.000	0.000	0.000	0.000	7.700	0.000	0.000	0.000	0.000	0.000	0.000
4003	3.000	0.000	1.200	1.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000
4004	3.000	0.000	0.000	3.600	5	0.000	2.800	0.000	0.000	0.000	0.000
4005	3.900	0.900	1.500	1.600	0.000	0.000	0.600	0.000	0.000	0.000	0.000
4006	7.400	0.000	0.900	1.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000
4007	8.100	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
SAMPLE NAME	31	32	33	34	35	36	37	38	39	40	
	GLOURAC	GLOBTURG	HNEGELGS	HYPEFLON	HAPLCARE	ISLIISLD	ISLITERE	KARFINFL	LAGEMOLL	LFTNTSPCA	
7202	4.500	10	0.000	0.000	1.000	1.000	22	1.000	2.400	1.700	0.000
7203	0.000	0.000	0.000	0.000	0.900	1.400	0.000	2.000	5.600	3.800	0.000
7204	0.200	0.000	0.000	0.500	1.400	0.000	0.000	3.000	2.900	0.000	0.000
7206	0.000	1.200	0.000	0.000	1.800	0.000	0.000	0.000	1.200	0.000	0.000
7207	0.000	0.000	0.000	0.000	4.800	0.000	0.000	0.000	3.000	0.000	0.300
7208	0.000	0.600	0.000	0.000	2.200	0.000	0.000	0.000	1.500	0.000	0.000
7209	1.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.300	0.000	0.000
7210	0.020	2.900	0.400	1.300	0.000	0.000	0.000	0.000	0.700	0.000	0.400
7211	0.000	0.300	6.000	0.000	0.000	0.000	0.200	0.000	0.000	0.000	0.000
7212	0.000	2.000	1.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7213	0.000	1.400	0.700	2.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7214	12.300	0.000	0.400	0.000	0.000	0.400	0.000	3.200	0.000	0.000	0.000
7217	9.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.000	0.000
7218	7.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7248	0.000	1.000	0.000	0.300	0.300	0.000	0.000	0.000	0.700	0.000	0.000
7249	0.000	1.200	0.000	2.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7250	0.000	2.400	0.000	1.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7252	0.000	5.000	17.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7253	0.300	0.000	0.600	0.000	0.000	0.000	0.000	5.700	0.000	0.000	0.000
7254	0.000	0.300	0.000	0.000	0.000	0.000	0.000	5.700	4	0.000	0.300
7255	0.300	0.000	0.000	0.000	0.000	0.000	0.000	2.700	4	0.000	0.000
7257	2.400	0.000	0.000	0.000	0.000	0.300	0.000	2.700	0.000	0.000	0.300
7258	7.400	0.000	0.300	0.600	11	0.000	3.100	0.000	0.600	7	0.000
7259	11.400	24	0.000	1.000	0.000	0.000	3.100	0.000	2.500	11	0.000
7260	21.700	33	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7264	0.200	0.000	0.000	5.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0812	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0817	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0819	0.000	0.000	0.000	0.000	0.000	0.000	1.200	0.000	0.000	0.000	0.000
0820	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0820	0.000	0.000	0.000	0.000	0.000	4.200	0.000	0.000	0.000	0.000	0.000
0821	0.300	8	0.000	0.000	0.000	0.000	0.500	5.100	9.000	0.000	0.000
0822	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.600	17	0.000	0.000
0824	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0825	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0830	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0832	0.000	0.000	0.000	0.000	0.000	0.000	1.200	5.100	9	0.000	0.000
0833	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0835	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000



6065	2.400	0.000	0.000	0.000	2.700	0.000	3.000	0.000	0.000	0.000
6059	.600	0.000	0.000	0.000	.300	50.000	14	0.000	0.000	0.000
6069	0.000	0.000	0.000	0.000	.800	64.500	44	0.000	0.000	0.000
6070	0.000	.300	0.000	0.000	.600	37.700	30	0.000	0.000	0.000
6071	0.000	0.000	0.000	0.000	.600	20.000	0.000	0.000	0.000	0.000
6072	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6073	0.000	0.000	0.000	0.000	.900	.500	0.000	0.000	0.000	0.000
6074	0.000	0.000	0.000	0.000	0.000	.800	4.700	0.000	0.000	1.500
6075	0.000	0.000	.800	0.000	0.000	0.000	3.700	0.000	0.000	0.800
6076	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6077	13.200	0.000	0.000	0.000	0.000	1.800	10.000	0.000	0.000	0.000
6078	0.000	0.000	0.000	0.000	4.700	0.000	5.500	0.000	0.000	0.000
6079	0.000	0.000	0.000	0.000	4.800	1.200	1.200	0.000	0.000	3.600
6080	0.000	0.000	0.000	0.000	0.800	45.500	1.800	0.000	0.000	1.200
6081	0.000	0.300	0.000	0.300	.800	5.300	3.300	0.000	0.000	0.000
6082	1.000	0.000	0.000	0.000	0.000	0.000	2.000	4.800	0.000	0.000
6083	1.100	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6084	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6085	30.600	70	0.000	0.000	.800	1.100	6	9.100	0.000	0.000
6086	0.000	0.000	0.000	0.000	0.000	1.200	0.000	0.000	0.000	0.000
6087	0.000	0.000	0.000	0.000	0.000	0.400	0.000	0.000	0.000	0.000
6088	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6089	0.000	0.000	0.000	0.000	0.000	0.900	36.400	0.000	0.000	0.000
6090	0.000	0.000	0.000	0.000	0.000	0.000	53.700	77	0.000	0.000
6091	17.300	100	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6092	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6093	18.200	33	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6094	0.000	0.000	0.000	0.000	0.000	3.700	0.000	0.000	0.000	0.000
6095	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6096	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6097	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6098	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6099	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6100	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6101	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6102	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6103	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6104	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6105	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6106	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6107	0.000	0.000	0.000	0.000	0.000	1.200	0.000	0.000	0.000	0.000
6108	0.000	0.000	0.000	0.000	0.000	0.500	0.000	0.000	0.000	0.000
6109	0.000	0.000	0.000	0.000	0.000	1.200	0.000	0.000	0.000	0.000
6110	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6111	30.700	0.000	0.000	0.000	.300	0.000	0.000	0.000	0.000	1.600
6112	15.400	45	0.000	0.000	1.200	0.000	2.400	0.000	0.000	0.000
6113	20.000	50	0.000	0.000	.200	1.700	9.000	0.000	0.000	0.000
6114	7.500	0.000	0.000	0.000	1.500	3.200	0.000	0.000	0.000	0.000
6115	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6116	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6117	17.400	25	0.000	0.000	1.700	9.600	23	0.000	12.500	0.000
1004	1.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1009	0.000	0.000	0.000	0.000	0.000	0.000	57.800	0.000	0.000	0.000
1010	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.000
1011	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1012	0.000	0.000	0.000	0.000	0.000	0.000	78.400	0.000	0.000	0.000
1013	0.000	0.000	0.000	0.000	1.300	0.000	0.000	0.000	0.000	0.000
1014	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1015	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1016	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1017	4.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1020	4.600	20	0.000	0.000	1.200	1.200	7.300	0.000	0.000	0.600
1021	13.400	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1022	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1023	0.800	0.000	0.000	0.000	0.000	0.000	6.900	1.700	0.000	0.000
1024	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1025	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1026	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1027	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1028	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1029	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1030	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1031	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1032	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1033	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1034	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1035	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1036	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1037	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1038	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1039	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1040	0.000	4.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1041	0.000	0.000	0.000	0.000	1.800	0.000	1.800	3	0.000	0.000

1044	0.000	0.000	0.000	0.000	1.600	11.000	20	5.600	0.000
1044	0.000	0.000	0.000	0.000	.700	30.900	20	1.400	0.000
1047	1.320	0.000	0.000	0.000	.900	0.000	5.000	4.000	0.000
1048	10.360	30	25	12	0.200	12.600	1	1.400	0.000
1048	0.100	0.000	0.000	0.000	2.500	33.200	1	1.500	0.000
1053	0.000	0.000	0.000	0.000	0.000	8.600	3	2.200	0.000
1056	0.000	0.000	0.000	0.000	0.300	0.000	2	0.000	0.000
1057	0.200	0.000	0.000	0.000	0.000	23.900	8	3.000	0.000
1058	12.300	15	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1060	0.000	0.000	0.000	0.000	0.000	31.500	33	0.000	0.000
1062	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1063	0.000	0.000	0.000	0.000	0.000	21.000	27	0.000	0.000
1064	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1065	11.300	6	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1066	0.000	0.000	0.000	0.000	0.000	35.400	48	0.000	0.000
1067	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1068	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1070	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1071	0.000	0.000	0.000	0.000	0.000	13.200	5	4.400	0.000
1072	0.000	0.000	0.000	0.000	2.400	1	10.200	10	0.000
1073	0.000	0.000	0.000	0.000	0.000	7.500	4	0.000	0.000
1074	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1075	12.800	27	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1076	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1077	0.000	0.000	0.000	0.000	0.000	37.100	0.000	0.000	0.000
1078	0.000	0.000	0.000	0.000	0.000	42.200	0.000	0.000	0.000
1081	0.000	0.000	0.000	0.000	0.000	29.500	0.000	0.000	0.000
1082	0.000	0.000	0.000	0.000	0.000	0.000	35.500	0.000	0.000
1083	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1084	0.000	0.000	0.000	0.000	0.000	0.000	5.000	0.000	0.000
1085	0.000	0.000	0.000	0.000	0.000	23.400	19	0.000	0.000
1086	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1087	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1088	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1089	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1090	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1091	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1092	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1093	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1094	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1095	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1096	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1097	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1098	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1099	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1100	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1101	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1102	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1103	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1104	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1105	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1106	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1107	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1108	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1109	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1110	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1111	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1112	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1113	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1114	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1115	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1116	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1117	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1118	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1119	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1120	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
					16				
						14.500	15		

1124	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1125	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1126	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1127	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1128	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1129	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1130	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1131	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1132	14.700	26	0.000	0.000	.300	9.000	14.600	47	0.000	0.000	0.000
1133	24.200	10	0.000	0.000	.100	2.300	0.000	0.000	0.000	0.000	0.000
1134	20.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1135	18.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1136	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1137	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1138	3.900	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1139	10.300	8	0.000	0.000	0.000	0.000	24.500	18	0.000	0.000	0.000
1140	0.000	0.000	0.000	0.000	0.000	0.000	13.500	11	0.000	0.000	0.000
1141	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
4003	0.000	0.000	0.000	0.000	0.000	0.000	40.100	0.000	0.000	0.000	0.000
4004	0.000	0.000	0.000	0.000	0.000	0.000	21.700	0.000	0.000	0.000	0.000
4005	0.000	0.000	0.000	0.000	0.000	0.000	30.900	0.000	0.000	0.000	0.000
4006	0.000	0.000	0.000	0.000	0.000	0.000	13.600	0.000	0.000	0.000	0.000
4007	0.000	0.000	0.000	0.000	0.000	0.000	5.900	0.000	0.000	0.000	0.000

SAMPLE NAME	41	42	43	44	45	46	47	48	49	50	
	MONTHLY	MONTHLY	MFLYDOPMP	PATEHAUR	PULFQUTN	QUICSFMI	RFOPSCDR	RFOPGUTT	REOPNNDU	REOPPARCT	
7902	6.500	33	1.000	22	0.000	.700	0.000	0.000	3.500	1.000	0.000
7903	4.300	0.000	0.000	0.000	0.000	2.600	0.000	5.100	0.000	.400	0.000
7904	6.800	0.000	0.000	0.000	0.000	3.900	0.000	7.400	.500	.500	0.000
7905	0.000	.800	0.000	0.000	0.000	4.300	0.000	3.000	0.000	0.000	0.000
7907	.700	.700	0.000	0.000	0.000	4.500	0.000	3.700	0.000	0.000	0.000
7908	0.000	.900	0.000	0.000	0.000	1.900	0.000	2.200	0.000	0.000	0.000
7909	0.000	0.000	0.000	0.000	0.000	.400	0.000	.700	0.000	0.000	0.000
7910	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7911	0.000	0.000	0.000	.700	0.000	.700	0.000	0.000	0.000	0.000	0.000
7912	0.000	0.000	0.000	0.000	0.000	.500	0.000	0.000	0.000	0.000	0.000
7913	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	.400	0.000
7914	.900	0.000	0.000	0.000	0.000	0.000	0.000	0.000	.400	0.000	0.000
7915	0.000	0.000	0.000	0.000	0.000	0.000	0.000	4.700	.300	0.000	0.000
7916	3.900	0.000	0.000	0.000	0.000	0.000	0.000	5.200	0.000	0.000	0.000
7944	0.000	0.000	0.000	0.000	0.000	1.000	0.000	1.000	0.000	0.000	0.000
7947	0.000	0.000	0.000	0.000	0.000	1.000	0.000	1.000	0.000	0.000	0.000
7950	0.000	.700	.400	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7953	0.000	0.000	0.000	0.000	0.000	2.500	0.000	0.000	0.000	0.000	0.000
7954	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.300	0.000	0.000	0.000
7955	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7956	.900	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7957	.900	0.000	0.000	0.000	0.000	0.000	0.000	1.000	.900	.600	0.000
7958	4.500	11	0.000	0.000	0.000	0.000	0.000	0.000	1.300	0.000	0.000
7959	10.200	14	0.000	0.000	2.000	0.000	0.000	0.000	1.800	5.800	14
7960	1.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7964	1.200	0.000	0.000	0.000	0.000	1.800	0.000	1.800	0.000	1.200	0.000
0802	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.100	0.000
0817	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	4.900
0818	.500	0.000	0.000	.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0819	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	.500	0.000	0.000
0820	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	4.200	0.000	0.000
0821	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0822	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0824	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.400	0.000	0.000	0.000
0826	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0827	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000







1114	0.000	0.000	0.000	0.000	0.000	0.000	4.200	5	0.000	0.000	2.800
1117	0.000	0.000	0.000	0.000	0.000	0.000	15.600	25	0.000	0.000	0.000
1120	0.000	0.000	0.000	0.000	0.000	0.000	2.700		0.000	1.600	0.000
1124	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
1125	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
1126	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	6.600
1127	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	8.300
1128	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	13.500
1129	1.000	0.000	0.000	0.000	0.000	0.000	1.200	3	0.000	0.000	6.300
1131	0.000	0.000	0.000	0.000	0.000	0.000	30.000	36	0.000	0.000	1.800
1132	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
1133	9.400	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
1134	9.100	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
1135	12.200	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
1136	12.300	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
1137	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
1138	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
1139	2.300	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
1140	2.400	3	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
1142	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
1143	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
1116	0.600	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	5.400
4003	0.000	0.000	0.000	0.000	0.000	0.000	2.000		0.000	0.000	1.200
4004	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
4005	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
4006	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000
4007	0.000	0.000	0.000	0.000	0.000	0.000	0.000		0.000	0.000	0.000

SAMPLE NAME	51	52	53	54	55	56	57	58	59	60	TRIFANGU
RFOPSCT	PARTSEM1	SACCATLA	SACCOIFF	SACCPH4	STAINAPT	TRCHINE	TRCHLOB	TRCHGLO	TRIFANGU		
7902	0.000	0.000	0.000	3.500	.300	0.000	0.000	0.000	0.000	0.000	3.400
7903	0.000	0.000	0.000	7.700	.400	0.000	0.000	0.000	0.000	0.000	9.000
7904	0.000	0.000	3.400	6.300	1.400	0.000	0.000	0.000	0.000	0.000	1.000
7906	0.000	0.000	4.000	4.000	3.400	0.000	0.000	0.000	0.000	0.000	4.800
7907	0.000	0.000	0.000	3.800	.800	0.000	0.000	0.000	0.000	0.000	1.300
7908	0.000	0.000	0.000	1.300	0.000	0.000	0.000	0.000	0.000	0.000	0.300
7909	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7910	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7911	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7912	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7913	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7914	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7917	0.000	0.000	0.000	0.000	5.700	0.000	1.100	0.000	0.000	0.000	0.000
7918	0.000	0.000	0.000	0.000	3.500	0.000	0.000	0.000	0.000	0.000	1.500
7949	0.000	0.000	0.000	0.300	0.300	0.000	0.000	1.000	0.000	0.000	0.300
7940	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7950	0.000	0.000	0.400	0.000	0.000	0.000	1.300	0.000	0.000	0.000	0.000
7952	0.000	0.000	0.000	0.000	0.000	0.000	1.200	0.000	0.000	0.000	0.000
7953	0.000	0.000	2.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	21.800
7955	0.000	0.000	0.000	0.000	0.000	0.000	0.300	0.000	0.000	0.000	22.700
7956	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	14
7957	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	11.800
7958	0.000	0.000	1.900	22	0.000	0.000	0.000	0.000	1.300	0.000	1.900
7959	0.000	0.000	2.400	14	1.200	0.000	0.000	0.000	1.300	0.000	2.400
7960	0.300	0.000	0.700	6.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7947	0.000	0.000	1.200	1.400	1.200	0.000	0.000	0.000	0.000	0.000	0.000
0802	0.000	0.000	4.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0817	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0818	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
CB20	4.400	0.000	3.400	0.000	0.000	0.000	0.300	0.000	0.000	0.000	0.000
CB21	4.800	0.000	1.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000	14.000
CB22	1.200	0.000	8.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000

0824	1.400	C.000	R.400	0.000	0.000	0.000	0.000	0.000	0.700	0.000	0.000
0827	0.000	0.000	4.000	0.000	0.000	0.000	0.000	0.000	10.800	0.000	0.000
0830	0.000	0.000	21.000	0.000	0.000	0.000	0.000	0.000	1.500	0.000	0.000
0832	0.000	0.000	2.000	0.000	0.000	0.000	0.000	0.000	1.300	0.000	0.000
0833	0.000	0.000	2.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0836	0.000	0.000	12.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0839	0.000	0.000	6.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0841	0.000	0.000	12.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0843	0.000	0.000	6.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0844	0.000	0.000	23.200	0.700	0.000	0.000	0.000	0.000	0.000	1.500	0.000
0846	0.000	0.000	23.300	21	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0847	0.000	0.000	1.200	0.000	0.000	0.000	0.000	0.000	4.100	0.000	0.000
0849	0.000	0.000	20.400	6	0.000	0.000	0.500	0.000	4.000	0.000	0.000
0851	0.000	0.000	44.000	45	0.000	0.000	.700	0.000	4.000	0.000	1.100
0852	0.000	0.000	3.000	5.000	0.000	0.000	0.000	0.000	0.000	0.000	0.300
0856	0.000	0.000	7.100	0.000	0.000	0.000	0.000	0.000	1.300	0.000	0.000
0857	0.000	0.000	12.700	20	0.000	0.000	0.000	0.000	12.000	9	0.000
0902	0.000	0.000	0.000	0.000	0.600	0.000	0.000	0.000	0.000	0.000	0.000
0904	0.000	0.000	0.000	0.400	0.000	0.000	2.000	0.000	0.000	1.200	0.000
0905	0.000	0.000	0.700	1.800	0.000	1.000	0.000	0.000	0.000	0.000	0.000
0906	0.000	0.000	0.000	0.000	0.000	1.000	0.000	0.000	0.000	0.000	0.000
0907	0.000	0.000	0.000	0.000	0.000	1.300	0.000	0.000	0.000	1.200	0.000
0910	0.000	0.000	1.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0911	0.000	0.000	2.000	0.000	0.000	0.000	2.000	0.000	0.000	1.000	0.000
0912	0.000	0.000	1.500	0.000	0.000	0.000	2.700	0.000	0.000	2.300	0.000
0913	0.000	0.000	0.000	0.000	1.000	0.000	4.900	0.000	0.000	0.000	0.000
0917	0.000	0.000	0.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000	34.000
0921	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0922	0.000	0.000	1.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0925	0.000	0.000	0.000	0.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0934	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0937	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0939	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0940	0.000	0.000	15.000	3.000	0.000	0.000	0.000	0.000	12.000	6	0.000
0942	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.300
0943	0.200	0.000	1.700	0.000	0.000	0.000	0.000	0.000	4.100	0.000	0.000
0944	0.000	0.000	1.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0946	0.000	0.000	2.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.600
0951	0.000	0.000	7.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	2.000
0952	0.000	0.000	2.600	0.000	0.1730	0.000	0.000	0.000	0.000	3.400	15.200
0953	0.000	0.000	6.400	2.200	1.300	1.300	0.000	0.000	0.000	3.000	15.200
0954	0.000	0.000	1.200	1.400	0.000	0.000	0.000	0.000	0.000	1.900	1.100
0955	0.000	0.000	0.000	0.400	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0956	0.000	0.000	1.700	0.000	0.000	0.000	1.700	0.000	0.000	0.000	2.000
0957	0.000	0.000	7.300	4.000	0.000	0.000	0.000	0.000	0.000	0.000	5.400
5003	0.000	0.000	0.500	1.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000
5005	0.000	0.000	30.000	0.000	0.000	0.000	0.000	0.000	2.000	0.000	1.300
6006	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.000	0.000	11.000
6007	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.200
6014	0.000	0.000	0.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	25.500
6032	0.000	0.000	0.400	0.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6033	0.000	0.000	3.200	0.000	3.200	0.000	0.000	0.700	0.000	0.000	0.000
6034	0.000	0.000	0.000	3.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6034	0.000	0.000	5.000	0.000	0.000	0.000	0.000	1.400	0.000	0.000	0.000
6041	0.000	0.000	1.000	0.000	1.000	0.000	0.000	0.000	0.000	0.000	0.000
6042	0.000	0.000	.900	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6050	0.000	0.000	.800	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6051	0.000	0.000	.300	0.000	1.200	0.000	0.000	1.200	0.000	0.000	1.100
6052	0.000	0.000	1.000	4.800	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6053	0.000	0.000	1.000	2.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6055	0.000	0.000	3.100	23.000	0.000	0.000	0.000	0.000	0.000	0.000	14.200



1036	C.000	0.000	1.000	0.000	0.000	0.000	0.000	0.300	0.000	1.300
1037	C.000	0.000	2.000	0.000	0.000	0.000	0.000	0.700	0.000	1.000
1038	C.000	0.000	2.100	0.000	0.000	0.000	0.000	1.000	0.000	.700
1039	C.000	0.000	3.800	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1040	C.000	0.000	1.300	0.500	5.500	0.000	0.000	0.000	0.000	0.000
1043	C.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1044	C.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1045	C.000	0.000	0.900	0.000	0.000	0.000	0.000	1.300	0.000	20.000
1046	C.000	0.000	1.600	0.000	0.000	0.000	0.000	0.000	0.000	2.000
1047	C.000	0.000	1.400	0.300	3.000	0.000	0.000	0.000	0.000	31.500
1048	C.000	0.000	0.600	0.000	0.000	0.000	0.000	0.000	0.000	1.100
1051	C.000	0.000	0.000	0.000	0.000	0.000	0.000	4.700	0.000	25.000
1053	C.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.600
1054	C.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1055	C.000	0.000	19.700	0.000	0.000	0.000	0.000	5.700	0.000	3.700
1056	C.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	5
1057	C.000	0.000	20.000	0.000	0.000	0.000	0.000	11.400	25	0.000
1058	C.000	0.000	11.500	0.000	0.000	0.000	0.000	11.100	9	0.000
1059	C.000	0.000	0.000	0.000	0.000	0.000	0.000	3.700	0.000	2.000
1060	C.000	0.000	2.400	0.000	0.000	0.000	0.000	0.500	0.000	0.000
1062	C.000	0.000	0.000	0.000	0.000	0.000	0.000	14.500	0.000	0.000
1063	C.000	0.000	3.500	0.000	0.000	0.000	0.000	1.200	0.000	0.000
1064	C.000	0.000	13.000	3	0.000	0.000	0.000	3.500	0.000	7.700
1065	C.000	0.000	1.300	0.000	0.000	0.000	0.000	2.000	0.000	3
1066	C.000	0.000	0.000	0.000	0.000	0.000	0.000	2.000	0.000	0.000
1067	C.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1068	C.000	0.000	19.000	7	0.000	0.000	0.000	12.500	12	0.000
1069	C.000	0.000	2.000	0.000	0.000	0.000	0.000	3.400	5	0.000
1070	C.000	0.000	24.300	14	0.000	0.000	0.000	12.500	3	0.000
1071	C.000	0.000	1.100	0.000	0.020	0.000	0.000	3.400	5	0.000
1072	C.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1073	C.000	0.000	2.200	0.000	0.020	0.000	0.000	0.000	0.000	0.000
1074	C.000	0.000	12.900	8	0.000	0.000	0.000	4.800	0.000	0.000
1075	C.000	0.000	12.700	5	0.000	0.000	0.000	7.400	3	0.000
1076	C.000	0.000	12.000	5	0.000	0.000	0.000	1.600	0.000	3.300
1077	C.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	4.600
1078	C.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	4.200
1080	C.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1081	C.000	0.000	5.100	7	0.000	0.000	0.000	0.000	0.000	0.000
1082	C.000	0.000	14.300	0.000	0.000	0.000	0.000	0.000	0.000	0.100
1083	C.000	0.000	5.500	0.000	0.000	0.000	0.000	2.500	0.000	0.000
1084	C.000	0.000	6.700	0.000	0.000	0.000	0.000	0.000	0.000	3.000
1085	C.000	0.000	6.000	4	0.000	0.000	0.000	6.300	0.000	2.800
1087	C.000	0.000	6.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1088	C.000	0.000	6.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1089	C.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1090	C.000	0.000	3.800	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1093	C.000	0.000	6.700	3	0.000	0.000	0.000	0.000	0.000	0.000
1095	C.000	0.000	2.900	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1098	C.000	0.000	1.800	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1097	C.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1099	C.000	0.000	2.200	0.000	0.000	0.000	0.000	3.000	0.000	0.000
1100	C.000	0.000	13.900	8	0.000	0.000	0.000	0.000	0.000	0.000
1101	C.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1102	C.000	0.000	2.060	0.000	0.070	0.000	0.000	4.700	3	0.000
1103	C.000	0.000	24.200	0.000	0.000	0.000	0.000	1.700	0.000	3.000
1104	C.000	0.000	12.500	9	0.000	0.000	0.000	2.992	2	0.000
1105	C.000	0.000	1.600	0.000	0.000	0.000	0.000	5.600	9	0.000
1106	C.000	0.000	15.500	5	0.000	0.000	0.000	1.700	0.000	0.000
1107	C.000	0.000	14.900	0.000	0.000	0.000	0.000	7.000	10	0.000
1109	C.000	0.000	19.300	5	0.000	0.000	0.000	7.300	11	0.000
1111	C.000	0.000	0.000	0.000	0.000	0.000	0.000	13.200	0.000	0.000
1112	C.000	0.000	30.800	16	7.900	5	0.000	3.500	7	0.000

1113	0.000	0.000	12,600	2	9,800	0.000	0.000	0.000	4,200	4	0.000	0.000
1114	0.000	0.000	20,300	3	0.000	0.000	0.000	0.000	7,900	1	200	0.000
1115	0.000	0.000	22,300	18	0.000	0.000	0.000	0.000	10,700	3	5	0.000
1116	0.000	0.000	4,900	0	0.000	0.000	0.000	0.000	2,700	0	000	0.000
1117	0.000	0.000	30,000	28	0.000	0.000	0.000	0.000	3,900	4	000	0.000
1118	0.000	0.000	24,000	21	0.000	0.000	0.000	0.000	8,000	0	000	0.000
1119	0.000	0.000	32,900	20	0.000	0.000	0.000	0.000	4,400	5	000	0.000
1120	0.000	0.000	12,300	10	0.000	0.000	0.000	0.000	4,300	2	000	0.000
1121	0.000	0.000	5,400	9	0.000	0.000	0.000	0.000	3,100	0	000	0.000
1122	0.000	0.000	4,200	6	2,200	0.000	0.000	0.000	0.000	0	000	0.000
1123	0.000	0.000	3,600	2	0.000	0.000	0.000	0.000	0.000	0	000	0.000
1124	0.000	0.000	28,200	22	0.000	0.000	0.000	0.000	0.000	0	000	0.000
1125	0.000	0.000	34,500	42	3,400	1,000	0.000	0.000	0.000	0	000	0.000
1126	0.000	0.000	3,900	2	2,700	0.000	0.000	0.000	0.000	0	000	0.000
1127	1,000	0.000	11,100	0	7,200	0.000	4,000	0.000	0.000	0	000	0.000
1128	0.000	0.000	5,900	0	0.000	1,200	4,500	0.000	0.000	0	000	0.000
1129	0.000	0.000	16,200	0	0.000	0.700	0.000	0.000	0.000	0	000	0.000
1130	0.000	0.000	33,600	39	5,400	0.000	0.000	0.000	7,100	0	000	0.000
1131	0.000	0.000	29,800	14	0.000	0.000	0.000	0.000	0.000	0	000	0.000
1132	0.000	0.000	27,100	0	0.000	0.000	0.000	0.000	4,300	0	000	0.000
1133	0.000	0.000	0.000	0	2,600	0.000	0.000	0.000	1,200	0	000	0.000
1134	0.000	0.000	0.000	0	0.000	0.000	0.000	0.000	0.000	0	000	0.000
1135	0.000	0.000	0.000	0	0.000	0.000	0.000	0.000	0.000	0	000	0.000
1136	0.000	0.000	0.000	0	0.000	0.000	0.000	0.000	0.000	0	000	0.000
1137	0.000	0.000	0.000	0	0.000	0.000	0.000	0.000	0.000	0	000	0.000
1138	0.000	0.000	0.000	0	0.000	0.000	0.000	0.000	0.000	0	000	0.000
1139	0.000	0.000	0.000	0	0.000	0.000	0.000	0.000	0.000	0	000	0.000
1140	0.000	0.000	0.000	0	0.000	0.000	0.000	0.000	0.000	0	000	0.000
1141	0.000	0.000	0.000	0	0.000	0.000	0.000	0.000	0.000	0	000	0.000
1142	0.000	0.000	0.000	0	0.000	0.000	0.000	0.000	0.000	0	000	0.000
1143	0.000	0.000	2,700	2	0.000	0.000	0.000	0.000	4,600	3	0	000
1144	0.000	0.000	12,900	0	0.000	0.000	0.000	0.000	5,700	0	000	0.000
4003	0.000	0.000	900	0	0.000	0.000	0.000	0.000	0.000	0	000	0.000
4004	0.000	0.000	0.000	0	0.000	0.000	0.000	0.000	0.000	0	000	0.000
4005	0.000	0.000	6,000	0	0.000	0.000	0.000	0.000	1,000	0	000	0.000
4006	0.000	0.000	1,900	0	0.000	0.000	0.000	0.000	0.000	0	000	0.000
4007	0.000	0.000	1,200	0	0.000	0.000	0.000	0.000	0.000	0	000	0.000

SAMPLE	NAME	61	62	63	64	65	66	67	68	69	70
		TRITCONI	TEXTDRO	UVIGPERF	ORIOTENR	VALVLAEV	RECUTURB	ADEPGLOM	SPIROBIF	TRILOCSP	ELPHACTT
7902		2.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7903		2.100	0.000	0.300	0.000	0.000	1.300	0.000	0.000	0.000	0.000
7904		2.000	0.000	0.000	0.500	0.000	0.000	0.000	0.000	0.000	0.000
7905		0.800	0.000	3.900	1.300	3.000	0.400	0.000	0.000	0.000	0.000
7906		0.000	0.000	17.500	0.000	1.000	0.000	0.000	0.000	0.000	0.000
7907		0.000	0.000	24.900	0.000	0.600	0.000	0.000	0.000	0.000	0.000
7908		0.000	0.000	35.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7910		0.000	0.000	37.000	3.600	4.000	0.700	0.000	0.000	0.000	0.000
7911		0.000	0.000	23.400	1.000	0.000	0.000	0.000	0.000	0.000	0.000
7912		0.000	0.000	25.600	0.600	0.000	0.000	0.000	0.000	0.000	0.000
7913		0.000	0.000	32.300	0.400	0.000	0.000	0.000	0.000	0.000	0.000
7914		0.000	0.000	0.000	0.400	0.000	0.000	0.000	0.000	0.000	0.000
7917		0.000	0.000	0.000	1.000	0.000	0.000	0.000	0.000	0.000	0.000
7918		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7948		0.000	0.000	53.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7949		0.000	0.000	54.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7950		0.000	0.000	52.200	0.000	0.000	0.700	0.000	0.000	0.000	0.000
7952		0.000	0.000	27.800	0.400	0.000	0.000	0.000	0.000	0.000	0.000
7953		0.000	0.000	0.000	0.300	0.000	0.000	0.000	0.000	0.000	0.000
7955		0.000	0.000	0.000	1.700	0.000	0.000	0.000	0.000	0.000	0.000
7956		0.000	0.000	0.300	1.900	0.000	0.000	0.000	0.000	0.000	0.000
7957		0.300	0.000	0.000	1.500	0.000	0.000	0.000	0.000	0.000	0.000
7958		1.600	0.300	0.000	1.900	0.000	0.000	0.000	0.000	0.000	0.600
7959		0.000	0.000	0.000	1.000	0.000	0.000	0.000	0.000	0.000	0.000
7960		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
7947		0.600	0.000	12.400	0.000	0.000	0.600	0.000	0.000	0.000	0.000
0802		2.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0817		0.000	0.000	0.000	0.000	0.000	10.000	22.600	17.100	0.000	0.000
0818		0.000	0.000	0.000	0.000	0.000	0.900	0.000	0.000	0.000	0.000
0819		0.000	0.000	0.000	0.000	0.000	15.000	14.46.200	74.12.400	0.000	0.000

0820	0.000	0.000	0.000	0.000	0.000	4.200	59.000	100	8.400	0.000	0.000
0821	0.000	0.000	0.000	0.000	0.000	4.800	22.700	50	4.500	0.000	0.000
0822	0.000	0.000	0.000	0.000	0.000	8.000	67.600	830	4.700	0.000	0.000
0824	0.000	0.000	0.000	0.000	0.000	8.100	55.800	720	7.700	0.000	0.000
0825	0.000	0.000	0.000	0.000	0.000	0.000	1.100	0.000	2.600	6.300	0.000
0827	0.000	0.000	0.000	0.000	0.000	0.000	0.600	0.000	0.000	0.000	0.000
0829	0.000	0.000	0.000	0.000	0.000	0.000	21	7.300	7	11.300	8
0830	0.000	0.000	0.000	0.000	0.000	0.000	3.800	0.000	24.800	0.000	0.000
0832	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0835	0.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0839	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	12.100	11
0841	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.200	10.600	100
0843	0.000	0.000	0.000	0.000	0.000	0.000	0.700	0.000	26.700	0.000	0.000
0844	0.000	4.800	17	0.000	0.000	0.000	0.000	2.200	12	40.300	41
0846	1.300	0.000	0.000	0.000	0.000	0.000	0.000	61.000	60	20.900	32
0847	0.000	0.000	0.000	0.000	0.000	0.000	0.000	9	0.000	0.000	0.400
0848	0.030	2.000	0.000	0.000	0.000	0.000	4.100	31	0.000	0.000	0.000
0849	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0851	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0852	0.300	0.000	0.000	0.000	0.000	0.000	0.000	1.200	0.000	0.000	0.000
0856	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0857	0.000	13	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.000
8002	0.600	0.000	0.000	0.000	0.000	1.200	0.000	0.000	0.000	0.000	0.000
8004	0.400	0.000	0.000	0.000	0.000	2.000	0.000	0.000	0.000	0.000	0.000
8005	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
8006	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
8007	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
P010	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
P011	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
P012	0.000	1.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
P013	0.600	0.000	12.000	2.400	0.000	0.000	0.000	0.000	0.000	0.000	0.000
P017	0.000	0.000	0.000	12.000	2.400	0.000	0.000	0.000	0.000	0.000	0.000
P019	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
P027	0.000	0.000	0.000	0.000	0.000	0.300	0.000	0.000	0.000	0.000	0.000
P035	0.000	0.000	0.000	0.000	0.000	0.000	0.000	4.600	12	40.000	31
P036	0.200	0.000	0.000	0.000	0.000	0.600	0.000	0.000	0.000	0.000	0.000
P037	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
P039	0.000	0.000	0.000	0.000	0.000	0.000	0.000	6.700	67.400	76	4.500
P040	0.000	0.000	0.000	0.000	0.000	0.000	0.000	8.100	72.000	50	6.300
P042	0.000	0.000	0.000	0.000	0.000	0.000	0.000	9.200	13	64.000	43
P043	0.000	0.000	0.000	0.000	0.000	0.000	0.000	10.500	9	57.000	53
P044	0.000	0.000	0.000	0.000	0.000	0.000	0.000	14.200	10	36.000	44
P046	1.500	0.000	0.000	0.000	0.000	0.000	0.000	9.300	24	14.600	16
P052	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
P053	0.900	0.000	1.300	0.000	0.000	8.300	100	0.000	0.000	0.000	0.000
P054	0.000	0.000	58.000	0.000	1.600	0.000	0.000	0.000	0.000	0.000	0.000
P055	0.000	0.000	28.600	0.000	1.000	0.000	0.000	0.000	0.000	0.000	0.000
P056	0.000	0.000	3.500	0.000	4.400	0.000	0.000	0.000	0.000	0.000	0.000
P067	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6003	0.000	1.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6005	0.500	0.000	0.000	0.000	0.000	0.000	0.000	10.000	30	0.000	0.000
6006	1.500	0.000	0.000	0.000	0.000	0.000	0.000	0.600	0.000	0.000	0.000
6007	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6010	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6032	0.400	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6033	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.200	0.000	0.000	0.000
6034	0.000	0.000	0.000	0.000	0.000	0.000	0.000	4.600	0.000	0.000	0.000
6036	0.900	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6041	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	3.500
6042	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6051	1.500	0.000	0.000	0.000	0.000	0.000	0.000	3.000	0.000	0.000	0.000
6052	0.200	0.000	0.000	0.000	0.000	0.000	0.000	0.600	0.000	0.000	0.000
	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000

6051	0.000	.400	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6055	.900	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6058	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6057	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6059	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6062	.700	0.000	0.000	0.000	0.000	0.000	1.500	2.300	3.000	0.000	0.000
6063	7.100	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6064	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6065	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6066	9.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6068	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6069	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6070	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6071	1.700	0.000	0.000	0.000	0.000	0.000	0.000	.400	1.200	0.000	0.000
6072	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6073	1.800	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6074	1.300	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6075	1.200	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6076	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6077	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6079	1.000	0.000	0.000	0.000	1.400	0.000	0.000	0.700	0.700	0.000	0.000
6080	.500	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6081	1.400	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6082	1.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6083	7.400	0.000	0.000	0.000	1.100	0.000	0.000	.900	3.400	0.000	0.000
6084	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6085	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6087	.700	3.100	0.000	0.000	0.000	0.000	0.000	9.500	13.000	1.400	0.000
6088	.800	4.600	0.000	0.000	0.000	0.000	0.000	10.200	29.200	0.000	0.000
6089	0.000	0.000	0.000	0.000	0.000	0.000	0.000	14.700	27.000	.800	0.000
6090	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.400	0.000	0.000
6092	1.600	4.800	0.000	0.000	0.000	0.000	0.000	4.400	4.400	0.000	0.000
6094	6.800	.700	0.000	0.000	0.000	0.000	0.000	8.500	15.600	1.000	0.000
6097	1.100	.900	0.000	0.000	0.000	0.000	0.000	1.700	6.000	0.000	0.000
6098	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.000	2.300	0.000
6099	1.000	12.600	0.000	0.000	0.000	0.000	0.000	0.000	6.500	7	0.000
6100	0.000	0.000	0.000	0.000	0.000	0.000	0.000	7.700	4.500	600	21.400
6101	0.000	0.000	0.000	0.000	0.000	0.000	0.000	13.200	18	30	10.400
6102	0.000	0.000	0.000	0.000	0.000	0.000	0.000	15.700	16	31	20.200
6103	0.000	3.000	0.000	0.000	0.000	0.000	0.000	3.000	33	35	13.600
6104	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6105	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6106	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6107	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6108	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6109	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6110	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6111	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6112	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6113	0.000	1.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
6114	5.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	18.300	1.400	0.000
6115	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	15.500	0.000	0.000
6116	0.000	3.500	0.000	0.000	0.000	0.000	0.000	0.000	5.400	0.000	0.000
6117	1.000	1.400	0.000	0.000	0.000	0.000	0.000	0.000	2.100	0.000	0.000
1008	0.000	0.000	0.000	0.000	0.700	0.000	0.000	0.000	0.000	0.000	0.000
1009	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1010	1.400	0.000	0.000	0.000	.700	0.000	0.000	0.000	0.000	0.000	0.000
1011	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.000	0.000
1012	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	1.800	2	1.800
1013	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	6.200	0.000	0.000
1014	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1015	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1016	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1017	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1020	1.900	0.000	0.000	0.000	1.000	1.200	0.000	0.000	0.000	0.000	0.000
1021	0.000	0.000	0.000	0.000	2.200	0.000	0.000	0.000	0.000	0.000	0.000
1022	0.000	0.000	0.000	0.000	2.600	0.000	0.000	0.000	0.000	0.000	0.000
1023	.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000	12.400	10	1.900
1024	2.600	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
1025	6.700	0.000	0.000	0.000	0.000	0.000	0.000	0.000	16.700	0.000	0.000
1027	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000



SAMPLE	NAME	71	72	73	74	75
7902		0.000	0.000	0.000	0.000	4.000
7903		0.000	0.000	0.000	0.000	0.000
7904		0.000	0.000	0.000	0.000	1.000
7906		0.000	0.000	0.000	0.000	0.000
7907		0.000	0.000	0.000	0.000	0.000
7908		0.000	0.000	0.000	0.000	0.000
7909		0.000	0.000	0.000	0.000	0.000
7910		0.000	0.000	0.000	0.000	0.000
7911		0.000	0.000	0.000	0.000	0.000
7912		0.000	0.000	0.000	0.000	0.000
7913		0.000	0.000	0.000	0.000	0.000
7914		0.000	0.000	0.000	0.000	2.000
7917		0.000	0.000	0.000	0.000	0.000
7918		0.000	0.000	0.000	0.000	0.000
7920		0.000	0.000	0.000	0.000	0.000
7921		0.000	0.000	0.000	0.000	0.000
7922		0.000	0.000	0.000	0.000	0.000
7923		0.000	0.000	0.000	0.000	0.000
7925		0.000	0.000	0.000	0.000	0.000
7927		0.000	0.000	0.000	0.000	0.000
7928		0.000	0.000	0.000	0.000	0.000
7929		0.000	0.000	0.000	0.000	0.000
7930		0.000	0.000	0.000	0.000	0.000
7931		0.000	0.000	0.000	0.000	0.000
7932		0.000	0.000	0.000	0.000	0.000
7933		0.000	0.000	0.000	0.000	0.000
7935		0.000	0.000	0.000	0.000	0.000
7937		0.000	0.000	0.000	0.000	0.000
7938		0.000	0.000	0.000	0.000	0.000
7939		0.000	0.000	0.000	0.000	0.000
7940		0.000	0.000	0.000	0.000	0.000
7947		0.000	0.000	0.000	0.000	0.000
C002		0.000	0.000	0.000	1.000	0.000

0817	0.000	0.000	0.000	0.000	0.000
0418	0.000	0.000	0.000	0.000	0.000
0319	0.000	0.000	0.000	0.000	0.000
0420	0.000	0.000	0.000	0.000	0.000
0421	0.000	0.000	0.000	0.000	0.000
0822	0.000	0.000	0.000	0.000	0.000
0824	0.000	0.000	0.000	0.000	0.000
0825	0.000	0.000	0.000	0.000	0.000
0930	0.000	0.000	0.000	0.000	0.000
0832	0.000	0.000	0.000	0.000	0.000
0833	0.000	0.000	0.000	1.700	0.000
0835	0.000	0.000	0.000	1.400	0.000
0838	0.000	0.000	0.000	0.000	0.000
0839	0.000	0.000	0.000	0.000	0.000
0841	0.000	0.000	0.000	0.000	0.000
0843	0.000	0.000	0.000	0.000	0.000
0844	0.000	0.000	0.000	0.000	0.000
0845	2.000	0.000	0.000	0.000	0.000
0847	0.000	0.000	0.000	0.000	0.000
0848	0.000	0.000	0.000	0.000	0.000
0849	0.000	0.000	0.000	0.000	0.000
0851	0.000	0.000	0.000	0.000	0.400
0852	0.000	0.000	0.000	0.000	0.000
0854	0.000	0.000	0.000	0.000	0.000
0857	0.000	0.000	0.000	0.000	0.000
6002	0.000	0.000	0.000	0.000	1.100
6004	0.000	0.000	0.000	0.000	0.000
6005	0.000	0.000	0.000	0.000	0.000
6006	0.000	0.000	0.000	0.000	0.000
6007	0.000	0.000	0.000	2.000	0.000
8011	0.000	0.000	0.000	0.000	0.000
8012	0.000	0.000	0.000	0.000	0.000
8013	3.900	0.000	0.000	0.000	0.000
5017	0.000	0.000	0.000	0.000	0.000
8015	0.000	0.000	0.000	0.000	0.000
8027	0.000	0.000	0.000	0.000	0.000
8035	0.000	0.000	0.000	0.000	0.000
8036	0.000	0.000	0.000	0.000	0.000
8037	0.000	0.000	0.000	0.000	0.000
8028	0.000	0.000	0.000	0.000	0.000
8029	0.000	0.000	0.000	0.000	0.000
8040	0.000	0.000	0.000	0.000	0.000
8042	0.000	0.000	0.000	0.000	0.000
8043	0.000	0.000	0.000	0.000	0.000
8044	0.000	0.000	0.000	0.000	0.000
8060	0.000	0.000	0.000	0.000	7.400
8061	0.000	0.000	0.000	0.000	1.700
8062	0.000	0.000	0.000	0.000	0.000
8064	0.000	0.000	0.000	0.000	0.000
8065	0.000	0.000	0.000	0.000	0.000
8066	0.000	0.000	0.000	0.000	0.000
8067	0.000	0.000	0.000	0.000	0.000
6003	0.000	0.000	0.000	0.000	0.000
6005	0.000	0.000	0.000	0.000	0.000
6006	0.000	0.000	0.000	0.000	0.000
6007	0.000	0.000	0.000	0.000	0.000
6014	0.000	0.000	0.000	0.000	0.000
6032	0.000	0.000	30.700	0.000	0.000
6033	0.000	0.000	2.100	0.000	0.000
6034	0.000	0.000	12.000	0.000	0.000
6036	0.000	0.000	0.000	0.000	1.800
6041	0.000	11.600	0.000	0.000	0.000
6042	0.000	11.900	0.000	0.000	0.000

6050	0.000	1.600	0.000	0.000	
6051	0.000	4.700	0.000	0.000	
6052	0.000	4.000	0.000	0.000	
6053	0.000	0.000	0.000	0.000	
6055	0.000	0.000	0.000	0.000	
6056	0.000	0.000	0.000	0.000	
6057	0.000	0.000	0.000	0.000	
6058	0.000	0.000	0.000	0.000	
6059	0.000	0.000	0.000	0.000	
6060	0.000	0.000	0.000	0.000	
6061	0.000	0.000	0.000	0.000	
6062	0.000	0.000	0.000	0.000	
6063	0.000	0.000	0.000	0.000	
6064	0.000	0.000	0.000	0.000	
6065	0.000	0.000	0.000	0.000	
6066	0.000	0.000	0.000	0.000	
6067	0.000	0.000	0.000	0.000	
6068	0.000	0.000	0.000	0.000	
6069	0.000	0.000	0.000	0.000	
6070	0.000	0.000	0.000	0.000	
6071	0.000	0.000	0.000	0.000	
6072	0.000	0.000	0.000	0.000	
6073	0.000	0.000	0.000	0.000	
6074	0.000	0.000	0.000	0.000	
6075	0.000	0.000	0.000	0.000	
6076	0.000	0.000	0.000	0.000	
6077	0.000	0.000	0.000	0.000	
6078	0.000	0.000	0.000	0.000	
6079	0.000	0.000	0.000	0.000	
6080	0.000	0.000	0.000	0.000	
6081	0.000	0.000	0.000	0.000	
6082	0.000	0.000	0.000	0.000	
6083	0.000	0.000	0.000	0.000	
6084	0.000	0.000	0.000	0.000	
6085	0.000	0.000	0.000	0.000	
6086	0.000	0.000	0.000	0.000	
6087	0.000	0.000	0.000	0.000	
6088	0.000	0.000	0.000	0.000	
6089	0.000	0.000	0.000	0.000	
6090	0.000	0.000	0.000	0.000	
6091	0.000	0.000	0.000	0.000	
6092	0.000	0.000	0.000	0.000	
6093	0.000	0.000	0.000	0.000	
6094	0.000	0.000	0.000	0.000	
6095	0.000	0.000	0.000	0.000	
6096	0.000	0.000	0.000	0.000	
6097	0.000	0.000	0.000	0.000	
6098	0.000	0.000	0.000	0.000	
6099	0.000	0.000	0.000	0.000	
6100	0.000	0.000	0.000	0.000	
6101	0.000	0.000	0.000	0.000	
6102	0.000	0.000	0.000	0.000	
6103	0.000	0.000	0.000	0.000	
6104	0.000	0.000	0.000	0.000	
6105	0.000	0.000	0.000	0.000	
6106	0.000	0.000	0.000	0.000	
6107	0.000	0.000	0.000	0.000	
6108	0.000	0.000	0.000	0.000	
6109	0.000	0.000	0.000	0.000	
6110	0.000	0.000	0.000	0.000	
6111	0.000	0.000	0.000	0.000	
6112	0.000	0.000	0.000	0.000	
6113	0.000	0.000	0.000	0.000	
6114	0.000	0.000	0.000	0.000	
6115	0.000	0.000	0.000	0.000	
6116	0.000	0.000	0.000	0.000	
6117	0.000	0.000	0.000	0.000	
1008	0.000	0.000	0.000	0.000	
1009	0.000	0.000	0.000	0.000	
1010	0.000	0.000	0.000	0.000	
1011	0.000	0.000	0.000	0.000	
1015	0.000	0.000	0.000	0.000	
1016	0.000	0.000	0.000	0.000	
1017	0.000	0.000	0.000	0.000	
1020	0.000	0.000	0.000	0.000	
1021	0.000	0.000	0.000	0.000	

1022	0.000	0.000	0.000	0.000	0.000
1027	0.000	0.000	0.000	0.000	0.000
1029	0.000	0.000	0.000	0.000	0.000
1034	0.000	0.000	0.000	6.500	0.000
1035	0.000	0.000	0.000	10.500	0.000
1036	0.600	0.000	0.000	2.600	0.000
1037	0.000	0.000	4.200	2.600	0.000
1038	0.000	0.000	38.300	1.400	2.100
1039	0.000	0.000	10.700	0.000	2.600
1040	0.000	0.000	6.200	3.700	0.000
1044	0.000	0.000	0.000	0.000	0.000
1045	0.000	0.000	0.000	0.000	0.000
1046	0.000	0.000	0.000	0.000	0.000
1048	0.000	0.000	0.000	0.000	0.000
1051	0.000	0.000	0.000	0.000	0.000
1053	0.000	0.000	0.000	0.000	0.000
1054	0.000	0.000	0.000	0.000	0.000
1055	0.000	0.000	0.000	0.000	0.000
1056	0.000	0.000	0.000	0.000	0.000
1057	0.000	0.000	0.000	0.000	0.000
1058	0.000	0.000	0.000	0.000	0.000
1060	0.000	0.000	0.000	0.000	0.000
1062	0.000	0.000	0.000	2.700	0.000
1063	0.000	0.000	0.000	0.000	0.000
1064	0.000	0.000	0.000	0.000	0.000
1065	0.000	0.000	0.000	0.000	0.000
1066	0.000	0.000	0.000	0.000	0.000
1067	0.000	0.000	0.000	0.000	0.000
1068	0.000	0.000	0.000	0.000	0.000
1069	0.000	0.000	0.000	0.000	0.000
1070	0.000	0.000	0.000	0.000	0.000
1071	0.000	0.000	0.000	0.000	0.000
1072	0.000	0.000	0.000	0.000	0.000
1073	0.000	0.000	0.000	0.000	0.000
1074	0.000	0.000	0.000	0.000	0.000
1075	0.000	0.000	0.000	0.000	2.400
1076	0.000	0.000	0.000	0.000	0.000
1077	0.000	0.000	0.000	0.000	0.000
1078	0.000	0.000	0.000	0.000	0.000
1080	0.000	0.000	0.000	0.000	0.000
1081	0.000	0.000	0.000	0.900	0.000
1082	0.000	0.000	0.000	0.000	0.000
1083	0.000	0.000	0.000	0.000	0.000
1084	0.000	0.000	0.000	0.000	0.000
1085	0.000	0.000	0.000	0.000	0.000
1086	0.000	0.000	0.000	0.900	0.000
1087	0.000	0.000	0.000	0.000	0.000
1088	0.000	0.000	0.000	0.000	0.000
1089	0.000	0.000	0.000	0.000	0.000
1090	0.000	0.000	0.000	0.000	0.000
1095	0.000	0.000	0.000	0.000	0.000
1098	0.000	0.000	0.000	0.000	0.000
1097	0.000	0.000	0.000	0.000	0.000
1099	0.000	0.000	0.000	0.000	0.000
1100	0.000	0.000	0.000	0.000	0.000
1101	0.000	0.000	0.000	0.000	0.000
1102	0.000	0.000	0.000	0.000	0.000
1103	0.000	0.000	0.000	0.000	0.000
1104	0.000	0.000	0.000	0.000	0.000
1105	0.000	0.000	0.000	0.000	0.000

38.300 45

1106	0.000	0.000	0.000	0.000	0.000
1107	0.000	0.000	0.000	0.000	0.000
1109	0.000	0.000	0.000	0.000	0.000
1110	0.000	0.000	0.000	0.000	0.000
1111	0.000	0.000	0.000	0.000	0.000
1112	0.000	0.000	0.000	0.600	0.000
1113	0.000	0.000	0.000	0.800	0.000
1114	0.000	0.000	0.000	0.300	0.000
1115	0.000	0.000	0.000	0.500	0.000
1116	0.000	0.000	0.000	0.800	0.000
1119	0.000	0.000	0.000	0.000	0.000
1120	0.000	0.000	0.000	0.000	0.000
1124	0.000	0.000	0.000	2.600	0.000
1125	0.000	0.000	0.000	0.000	0.000
1126	0.000	0.000	0.000	0.000	0.000
1127	0.000	0.000	0.000	0.000	0.000
1128	0.000	0.000	0.000	1.500	0.000
1130	0.000	0.000	0.000	0.000	0.000
1131	0.000	0.000	0.000	0.000	0.000
1132	0.000	0.000	0.000	0.000	0.000
1133	0.000	0.000	0.000	0.000	0.000
1134	0.000	0.000	0.000	0.000	0.000
1135	0.000	0.000	0.000	0.000	0.000
1136	0.000	0.000	0.000	0.000	0.000
1137	0.000	0.000	0.000	0.900	0.000
1138	0.000	0.000	0.000	0.000	0.000
1140	0.000	0.000	0.000	0.000	0.000
1142	0.000	0.000	0.000	0.000	0.000
1143	0.000	0.000	0.000	0.000	0.000
1148	0.000	0.000	0.000	0.000	0.000
4003	0.000	0.000	0.000	0.000	0.000
4004	0.000	0.000	0.000	0.000	0.000
4005	0.000	0.000	0.000	0.000	0.000
4006	0.000	0.000	0.000	0.000	0.000
4007	0.000	0.000	0.000	0.000	0.000

## APPENDIX VIII

Regression statistics produced through applying a 0.001 cut off. (refer to Chapter 6). SEE = Standard Error of Estimate.

VARIABLE NUMBER	VARIABLE NAME	DEPTH REGRESSION COEFFICIENT	STD. ERROR OF REG. COEFF.	COMPUTED t VALUE
8	F2	-59.90625	22.40832	-2.673
12	F6	-72.93839	24.04646	-3.033
11	F5	79.36601	20.54404	3.860
9	F3	79.02030	25.22223	3.126
13	F7	-21.67884	19.49142	-1.112
14	F8	-58.46799	23.55363	-2.482
7	F1	36.88112	22.65556	1.628
10	F4	12.55029	18.02700	.696
SEE = 54.2m				

TEMPERATURE				
14	F8	-5.60670	.69694	-8.046
7	F1	-3.77164	.43887	-8.594
8	F2	-3.16055	.50773	-6.232
11	F5	1.66009	.57536	2.885
10	F4	1.44440	.56352	2.563
SEE = 1.8 C				

SALINITY				
14	F8	-2.48216	.27198	-9.126
11	F5	1.48253	.25270	5.867
9	F3	.51991	.27033	1.997
13	F7	.79693	.25758	3.094
SEE = .77				

GRAVEL				
9	F2	22.89394	4.21771	5.428
13	F7	17.83077	5.52203	3.229
SEE = 16				

SAND				
9	F3	-14.27837	11.63006	-1.228
11	F5	52.30125	9.96432	5.249
12	F6	52.04661	10.57108	4.923
8	F2	26.75690	9.07305	2.949
SEE = 28				

MUD				
8	F2	-48.74423	9.24458	-5.273
9	F3	23.07817	11.42211	2.020
12	F6	-51.72531	10.22297	-5.060
11	F5	-41.47570	9.82860	-4.220
13	F7	-31.18985	8.93054	-3.492
7	F1	-15.51767	8.60288	-1.804
SEE = 26				