COOLING AND UPLIFT PATTERN OF THE LIBERTY HILL PLUTON, SOUTH CAROLINA, USA

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Abstract

The Liberty Hill pluton is one of a series of post-metamorphic plutons emplaced in the central and eastern Piedmont province of the Appalachians in the late Paleozoic. An investigation of the post-magmatic cooling and uplift pattern of the pluton utilized the 40 Ar/ 39 Ar and fission track dating techniques. Similar 40 Ar/ 39 Ar ages of 314 - 302 Ma shown by hornblende, muscovite and biotite separates as well as a previously calculated Rb/Sr whole rock date indicate the pluton cooled rapidly after emplacement in the low metamorphic grade rocks of the Carolina Slate Belt. An 40 Ar/ 39 Ar microcline age of 260 ma as well as fission track ages of 123, 97.0, 92.4, and 82.1 Ma indicate the pluton cooled slowly after reaching ambient country rock temperatures. A crystallization pressure of 4.6 \pm 1 Kb from a hornblende geobarometer and an estimate of the paleogeothermal gradient permitted the construction of a depth-time curve consistent with a three stage uplift history for the Liberty Hill pluton. Chapter 1

1.0 Introduction

The Appalachian orogen was constructed along eastern North America by a series of collisions which spanned much of the Paleozoic. Taconic (Ordovician) and Acadian (Devonian) collisions involved the accretion of island arcs while the Alleghanian (Pennsylvanian-Permian) collision of Africa and North America caused considerable crustal thickening and orogenic development (Secor et al. 1986; Hatcher, 1987). Current attempts to understand the constructional histories of this and other mountain belts throughout the world have relied on a wide variety of geological and geophysical techniques. One method of investigation involves tracing the uplift history of plutons contained within orogenic belts. Isotopic dating of coexisting minerals with different closure temperatures combined with geobarometric data provides a convenient way to reconstruct the pressuretemperature-time (P-T-t) history of a pluton. This study will use this method to investigate the cooling and uplift history of the Liberty Hill pluton, emplaced in the Southern Appalachians during the Alleghanian orogeny.

1.1 <u>Regional Tectonic Setting</u>

The Liberty Hill pluton is located in the South Carolina section of the Carolina Slate Belt. This belt forms the eastern section of the Appalachian Piedmont province which is bounded on the west by the hills of the Blue Ridge and on the east by the sedimentary rocks of the Coastal Plain (Figure 1). Hatcher (1978) proposed that the Carolina Slate Belt formed part of a late Precambrian volcanic island arc which collided with North America during the Acadian orogeny (~400 - 350 Ma) (Hatcher, 1987). The Alleghanian collision of



Figure 1: Generalized geologic map of the central and southern Appalachians showing location of Carolina Slate Belt. LH denotes location of Liberty Hill pluton (after: Glover et al., 1983).

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North America with Africa (~300 - 245 Ma) produced a string of granitic plutons which intruded into the low-grade rocks of the Carolina Slate Belt (Figure 2). The plutons all yield emplacement ages of 325 to 265 Ma and only show signs of deformation adjacent to regional ductile deformation zones (Glover et al, 1983).



Figure 2: Generalized geologic map showing 325 to 265 Ma (Pennsylvanian-Permian) granitic plutons (stippled pattern) in the central and eastern Piedmont of North Carolina, South Carolina and Georgia. LH denotes the Liberty Hill pluton (after: Fullagar and Butler, 1979).

1.2 Geology of the Liberty Hill Region

The Carolina Slate Belt is predominantly late Precambrian to early Paleozoic pyroclastic and epiclastic volcanic rocks with a small percentage of lava flows and pelitic rocks (Fullagar and Butler, 1979).

The Liberty Hill pluton was emplaced in metamorphosed hypabyassal, pyroclastic and epiclastic rocks (Speer, 1981). Pelitic rocks on the eastern margin of the pluton have preserved a well-developed contact aureole, that-superimposed on the regional Precambrian greenschist facies metamorphic assemblage of minerals--indicates ambient country rock temperatures of 300°C -400°C.

The Liberty Hill Pluton is a roughly elliptical, steep-sided intrusion with a single root (Bell and Popenoe, 1976). It covers approximately 360 km² in Kershaw, Lancaster and Fairfield Counties in north-central South Carolina. The pluton consists of three texturally and mineralogically distinct facies (Figure 3). The predominant facies is a central coarse-grained amphibolebiotite granite which grades to a coarse-grained biotite granite at the margins. A fine-to-medium-grained biotite granite co-exists with the coarsegrained granites particularly in the west-central regional. Compositional similarity and the manner in which one granite grades into another have ruled out multiple intrusions (Speer, 1987).

The pluton was first isotopically dated by Overstreet and Bell (1965) using a zircon separate which yielded a lead-alpha age of 245 ma. A subsequent Rb-Sr whole-rock age of 299+ 8 Ma was determined by Fullagar (1971).

Spear (1981) estimated a contact metamorphic temperature of 725° C for the central granitoids, 647° C for the marginal rocks and an emplacement pressure





Figure 3: Geologic map of the Liberty Hill pluton showing sample location (after: Speer, 1987).

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of 4.5 to 5.1 Kb. after examining the contact aureole. This pressure estimate is dependent on the choice of the aluminosilicate triple point; the point suggested by Richardson et al. (1969) gives a pressure of 4.5 to 5.1 Kb while that of Holdaway (1971) would yield a pressure estimate of 2.0 - 3.5 Kb. Dr. Casey Ravenhurst collected samples of the coarse-grained amphibole-biotite granite (LH-1, -2, and -5) and the fine-grained biotite granite (LH-3, -4) in December 1986 (location map, Figure 3). 40Ar/39Ar and fission track dating of minerals from these samples as well as the use of a hornblende geobarometer allow the calculation of the emplacement and cooling history of the pluton. With an estimate of the paleogeothermal gradient, the temperature-time history can be converted into a depth-time history allowing estimation of the uplift of the Liberty Hill pluton.

Chapter 2: 40 Ar/39 Ar Dating Technique

2.0 <u>Introduction</u>

The Rb/Sr whole rock date of 299 ± 8 (Fullagar, 1979) constrains the crystallization age of the Liberty Hill pluton. Dating the subsequent cooling history of the pluton requires isotopic systems with lower closure temperatures, such as 40Ar/39Ar. The 40Ar/39Ar dating technique is a modification of the K-Ar technique. It is based on the radioactive decay of K to Ar. At high temperatures any Ar produced by this decay is lost to the environment. However, at the closure temperature Ar becomes immobile and begins to accumulate within a mineral. The temperature at which Ar becomes immobile varies from mineral to mineral; therefore, dating a set of coexisting minerals will reveal a time-temperature history. Current experimental difficulties in quantifying the parameters that affect closure temperatures have led to large errors in their estimates. In this chapter the theoretical basis of 40Ar/39Ar dating, laboratory procedure and results from the Liberty Hill samples will be examined.

2.1 ⁴⁰<u>Ar/</u>³⁹<u>Ar Dating Theory</u>

The ${}^{40}\text{Ar}/{}^{39}\text{Ar}$ dating technique is a variant of K-Ar dating. Both systems depend on the accumulation of radiogenic ${}^{40}\text{Ar}$ (${}^{40}\text{Ar}*$) from decay of ${}^{40}\text{K}$. In the K-Ar method ${}^{40}\text{Ar}*$ and ${}^{40}\text{K}$ are measured on separate aliquots of a sample, the ratio of the two yielding the age. The ${}^{40}\text{Ar}/{}^{39}\text{Ar}$ technique involves conversion of ${}^{39}\text{K}$ atoms to ${}^{39}\text{Ar}$ by irradiation with fast neutrons. The amount of ${}^{39}\text{Ar}$ produced via this reaction is proportional to the total amount of ${}^{40}\text{K}$, since the ratio of ${}^{40}\text{K}$ to ${}^{39}\text{K}$ is essentially constant. Thus the ratio of

 $40 \star Ar/^{39}$ Ar gives a measure of the $40 Ar \star / 40 K$ ratio and the age of the sample.

The amount of 39 Ar produced via the nuclear reaction 39 K (n,p) 39 Ar is given by the equation:

$${}^{39}\text{Ar} = {}^{39}\text{K}\,\Delta \top \,\mathcal{S}\,\phi(\mathcal{E})\,\sigma(\mathcal{E})\,\partial(\mathcal{E},\ldots,\mathcal{C})$$

where, 39 K - is the number of 39 K isotope atoms

 $\Delta \, T$ - is the length of the irradiation in seconds

 $\emptyset(\mathcal{E})$ is the neutron flux with energy \mathcal{E}

 $\mathcal{K}(\mathcal{E})$ - cross section for the reaction ³⁹K (n,p)

³⁹Ar at energy €

The number of radiogenic 40 Ar atoms due to decay of 40 K is given by the equation:

$${}^{40}\text{Ar}* = \frac{\lambda c}{\lambda} {}^{40}\text{K} (e^{\lambda t} 1) \dots (4)$$

where; λ = the total decay constant for ${}^{40}\text{K}$

 λe = the decay constant of 40K for electron capture

t = the age of the sample

After irradiation of a sample the ratio of 40Ar*/ 39 Ar is given by the equation:

The value of $S \varphi(\epsilon) \nabla (\epsilon) \delta \epsilon$ is difficult to evaluate accurately. This has led to a system of age calculation in which samples of known age (flux monitors) are interspersed with unknown samples in an irradiation canister. Then, using the J parameter which is defined as:

$$J = \frac{\lambda}{\lambda e} \frac{39 k}{40 k} \frac{1}{\Delta \tau} \frac{e^{\lambda t} - 1}{Sp(c)\sigma(c) dc} \dots \dots$$

and substituting this into equation 5 we obtain:

The J parameter for a canister uses the known age of the standard samples and their ${}^{40}\text{Ar}/{}^{39}\text{Ar}$ ratio measured in the mass spectrometer. A simple graph of J value versus canister position allows interpolation of a J value for the unknowns. The age of the unknowns is calculated using the equation:

$$t = \frac{1}{2} lm \left[\left(\frac{40}{34} \frac{1}{\beta_{T}} \right) J + 1 \right] \dots (8)$$

The estimated analytical error of the age \pm is given by:

$$\sigma_t \simeq \left[\frac{1}{2^2 F^2 \left[\sigma_F^2 + \sigma_3^2 \right]} \dots \left(q \right) \right]$$

where; $F = \frac{40}{\text{Ar}^{*}}/\frac{39}{\text{Ar}}$

 $\mathcal{O}_{\mathcal{F}}^{2} \sim \partial_{\mathcal{O}_{\mathcal{J}}}^{2}$ - are variance of F and J expressed in percent (Faure, 1977)

2.2 ⁴⁰Ar/³⁹Ar Lab Procedure

The samples used in the 40 Ar/ 39 Ar dating technique were crushed using a hydraulic press and a tungsten-carbide shatterbox, sieved and washed to obtain the 70 to 120 mesh size fraction. A Frantz magnetic separator sorted the samples according to their magnetic susceptibilities. The fractions containing large quantities of biotite or muscovite were paper panned and handpicked to obtain very pure samples. The magnetic fractions containing hornblende were placed in a methylene iodide heavy liquid solution yielding a hornblende separate. This hornblende separate was handpicked to eliminate biotite contamination. The potassium feldspar rich fraction was purified in a sodium polytungstate solution (density $^2.6 \text{ g/cm}^3$).

Aliquots of approximately 100 mg of biotite and muscovite, 200 mg of potassium feldspar and 500 mg of hornblende, wrapped in aluminum foil inside

an irradiation canister were irradiated at McMaster University nuclear reactor for approximately 10 hours. After being returned to Dalhousie and allowed to "cool", samples from the canister were placed in a quartz tube in a Lindberg extraction furnace. Gradual heating in a step-wise fashion extracted the argon gas.

A typical heating step consisted of:

- Heating the furnace up by a 50°C interval and allowing the sample to outgas for one hour. (Extraction system shown in Figure 4.)
- Lowering the charcoal finger temperature to -196°C by immersion in liquid nitrogen, resulting in absorption of the gas onto the charcoal.
- 3) Isolating the system by closing the valve to the furnace.
- 4) Cleaning the surface of the titanium "getter" by heating it to 800°C for 10 minutes while it is being pumped.
- 5) Cleaning the sample by allowing it to interact with the titanium "getter" for twenty minutes. Volatile gases bind to the titanium surface leaving a relatively pure sample of argon gas.



Figure 4: Argon extraction system. The Ti "getter" chemically buries reactive impurities in its surface. The freeze-down fingers are immersed in liquid nitrogen. One finger contains activated charcoal to condense argon, the other condenses out gases not absorbed by the Ti getter (after: Grist, 1986).

- 6) The empty freeze-down finger was then immersed in liquid nitrogen in order to condense out the gases not absorbed by the titanium "getter".
- 7) The argon gas was slowly admitted to the MS-10 mass spectrometer where it was analyzed for quantities of the various isotopes of argon. Measured isotope ratios were corrected for:
 - 1) mass spectrometer discrimination
 - 2) atmospheric argon contamination
 - interfering isotopes of argon (³⁶Ar, ³⁹Ar) produced by irradiation of calcium-containing minerals.

The apparent age corresponding to each heating step, calculated from the corrected argon isotope ratios and plotted as a function of the percentage of 39 Ar released, produced an age spectrum for each mineral.

2.3 <u>Results</u>

The age spectra obtained from the biotite, muscovite and microcline of sample LH-4 and hornblende and biotite of LH-2 are shown in Figure 5. A crystal description of each of these phases is shown in Table 1.

Three of the spectra show relatively concordant ages, in which over 80% of the gas fractions released have approximately the same age.

The LH-2 hornblende spectrum shows a distinct six-step plateau representing 85% of the gas released. A weighted average of these six steps, based on the percentage of the spectrum each represents, gives an age of 303 ± 3 Ma. The LH-2 biotite spectrum also has a well-defined plateau consisting of seven heating steps representing 80% of the gas released. The weighted average of this plateau yields an age of 302 Ma \pm 3. The LH-4 biotite spectrum has a plateau which represents nearly 90% of the gas released. The

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SAMPLE	MINERAL	GRAIN SIZE	COMMENTS
LH-2	Hornblende	Up to 4mm	 Euhedral to subhedral Pleochroic - yellow green to dark green Some intergrown with biotite Little alteration
LH-2	Biotite	Up to 5mm	 Euhedral to subhedral Pleochroic - yellow to green-black Little alteration
LH-4	Biotite	Up to 0.5mm Fine grained	- As above
LH-4	Muscovite	Up to 0.25mm	 Subhedral to anhedral grains Intergrown with biotite, but predominantly as discrete grains
LH-4	Alkali– Feldspar (Microcline)	Up to 1.5mm	- Sub-to anhedral grains - Shows slight alteration to white mica (less than 10%)
LH-4	Orthoclase	Up to 1.5mm	 Shows 10-90% alteration to white mica Microprobe revealed perthitic lamallea although not generally visible in thin section

crystalline nature of samples used in $^{40}{\rm Ar}/^{39}{\rm Ar}$ dating

small step with the anomalously high age represents 5% of the total gas released and is therefore significant. This plateau yields an age of 309 Ma \pm 3.

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(See text for details.)

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The final two spectra are discordant, that is, they fail to form an age plateau representing a large percentage of the gas released. At some point in the history of these samples they have been partially outgassed. The LH-4 muscovite produced an unusual spectrum in that it yielded insignificantly small quantities of gas above 850° C. The low temperature steps show the staircase pattern typical of a partially outgassed sample. The best estimate of an age for this spectrum is 314 ± 3 Ma as represented by the small plateau of the three high temperature steps.

The microcline separates were shown to be intermediate to maximum microcline, using the Wright (1986) method X-ray diffraction "3" peak measurement (diffraction spectra in back pocket). The age spectrum of the microcline shows a steplike pattern of increasing age with increasing temperature, leveling off to a plateau for the high temperature steps. This common shape for microcline spectra (Harrison and Be, 1983) results from either a partial loss of gas due to a reheating event, or from slow cooling through the closure temperature, allowing some escape of gas. Lack of geological evidence for a reheating event in this region, as constrained by the lower greenschist facies nature of the country rock, suggests the sample shape resulted from gas lost during passage of the feldspar through the closure temperature. In this case, the 260±3 Ma age was calculated by means of a weighted average of the total gas released.

<u>Chapter 3: Theoretical Background and Analytical Methods</u> <u>Fission Track Analysis</u>

3.0 Introduction

The ${}^{40}\text{Ar}/{}^{39}\text{Ar}$ technique provides temperature-time (T-t) data spanning the range of 550° to 150°C. The fission track dating technique allows T-t constraints below this level. The apparent fission track age as well as the distribution of track lengths within crystals reveal important aspects of the thermal history. This chapter examines the theory, lab procedure and results of fission track analysis of apatite crystals from the Liberty Hill pluton.

3.1 Fission Track Dating Theory

The fission track dating technique depends upon the natural spontaneous fission of 238 U. As fission products travel through a crystal lattice they have a damage track. These damage tracks are typically a few $\overset{\circ}{A}$ wide and 10-20 long (Figure 6).



Figure 6: Photomicrograph of an etched apatite crystal showing fossil fission tracks. The scale bar is 10 (after: Gleadow, 1983).

Treatment fo the host crystal with a chemical etchant allows the fission tracks to be observed with a petrographic microscope. Fission tracks are stable for long periods of time at room temperature; however, at elevated temperatures the damage zones gradually anneal rendering the tracks invisible. Fission tracks are completely retained in apatite crystals when formed below the annealing (closure) temperature range of 70°-120°C (Harrison and McDougall, 1980).

The use of fission tracks as a dating technique is based on the assumption that in a system where there is no loss or gain of 238 U and the amount of time the system has been cooler than the closure temperature.

This study uses the "zeta calibration" age calculation technique outlined by Hurford and Green (1982). The age, t, is determined from the equation:

$$t = \frac{1}{\lambda D} \ln \left[1 + \lambda D Z \left(\frac{P_s}{P_i} \right) g P_D \right] \dots \dots$$

where $Z(zeta) = \frac{e \times p (\lambda D t S T D)'}{\lambda D (PS/Pi) S T D S PD}$... L2)

 ∂_D = total decay constant of ²³⁸U (1.555125 x 10⁻¹⁰/yr) ρ_5 = spontaneous track density

 \mathcal{P}_{L}^{\cdot} = induced track density

 $P_{\rm D}$ = track density of glass dosimeters

 t_{5TD} = age of standard (Fish Canyon apatite, 27.8 0.2 Ma)

A = geometry factor (0.5)

The zeta parameter is evaluated using glass dosimeters and age standards. This allows the calculation of an apparent age for unknowns without having to measure the thermal neutron flux in the reactor, as well as other poorly known parameters. Individual determinations of zeta may contain large unexplained errors (Hurford and Green, 1983). It is therefore more precise to use an average value of zeta calibrated from many capsule runs. The zeta used in this study is the laboratory average described in Section 3.2.

3.2 Fission Track Laboratory Procedure

The samples used for fission track dating were crushed using a hydraulic press and a tungsten-carbide shatterbox. The samples were washed and sieved to obtain the 70 to 120 mesh size fraction. The separation of apatite grains was initiated by removing the magnetic fraction using a Frantz-magnetic separator. The sample was further purified using a sodium polytungstate solution (density 3.0 g/cm^3) and a methylene iodide solution (density 3.3 g/cm^3). Approximately 150 apatite crystals, hand-picked from each separate, were mounted in araldite epoxy resin.

All grain mounts were polished using a Polimet polisher, starting with a 600 grit paper and then continuing with a 15, 9, 3, and 1/4 diamond paste.

Spontaneous tracks were revealed by etching the grain mounts in 7% HNO₃ for 40 seconds. The grain mounts were covered with a freshly cleaved flake of low uranium mica (Brazilian Ruby Clear) to record induced tracks produced in the reactor. Apatite grain mounts of known age (Fish Canyon Standards) and SRM-614 uranium dosimeter glasses (from the National Bureau of Standards) were also covered with mica detectors. The grain mounts and dosimeters were placed in an irradiation canister, sent to McMaster University nuclear reactor and irradiated in site 9D for 555 seconds.

Upon return of the capsule the mica detectors and glass dosimeters were removed and etched in 48% HF for 30 seconds. The micas were carefully washed

in distilled water and heated on a hot plate eliminating remaining HF. The number of tracks per grid was counted at 400 X magnification with a microscope in which a 5 x 5 graticule in one ocular served as a grid. Approximately 250 grid fields were counted for each dosimeter. The mica detectors for the standards and unknowns were etched in 48% HF for 13 minutes and placed in bilateral symmetry with the grain mount on a glass slide. The density of spontaneous tracks, counted in the grain mounts, and of induced tracks, counted in the portion of the mica detector directly above the grain, was counted at 1000 x magnification. The glass dosimeter track density and the ratio of spontaneous to induced track density from the standards were used to calculate a zeta parameter using equation 2 (Section 3.1). This parameter, averaged with zetas from previous capsule runs, gives a value of 11.4+0.4 x 10^3 . The ratio of spontaneous to induced track density, determined for fewer than 50 grains for each unknown, is independent of the concentration of uranium and should therefore show similar values, subject to the variation allowed by a Poissonian distribution. However, there are several sources of error which can conspire to make the distribution larger than Poissonian. According to the statistical method developed by Green (1981) if the $ho_{f s/
ho i}$ data pass a chi-square statistical distribution it is analyzed in a conventional Poissonian manner (the ratio $= \mathbf{z} \rho_{\mathbf{z}} | \mathbf{z} \rho_{\mathbf{z}}$). However, if the data fail the chi-square distribution test a mean track density ratio $(\mathbf{z} \mathcal{P}_{5} | \mathcal{P}_{L})$ is calculated and a standard error of the mean assigned to that quantity.

3.3 Track Length Distribution

The variation of the length of confined fission tracks yields important information about thermal histories. As figure 7 shows, confined fission

tracks do not intersect the polished surface of a crystal.



Figure 7: Diagram showing the etching of confined fission tracks used to measure track lengths (after: Gleadow et al., 1983).

Instead they intersect other features such as cracks or surface tracks which act as conduits of etchant to the confined tracks. When freshly formed, fission tracks in apatite have a fairly constant mean length of about 16.3 ± 0.9 (Gleadow et al., 1983). New tracks are continuously formed throughout the history of a sample. Therefore each track will spend a different amount of time in the annealing temperature range and show a different degree of shortening. The overall track length distribution therefore represents a compilation of the temperature history of each track. A narrow distribution of long tracks (Figure 8a) represents a relatively short residence time within the partial annealing zone, whereas a broad distribution of tracks with a shorter mean length (Figure 8b) indicates a longer time in the annealing zone.



Figure 8: Variation in the distribution of track length. (See text for details)

3.4 Track Length Laboratory Procedure

The measurement of track lengths, restricted to confined tracks horizontal or sub-horizontal to the viewing surface, ensured that a real, not an apparent, length was measured. These tracks were measured with a digitizing tablet linked to a microcomputer. The tablet cursor was fitted with a light emitting diode (LED) visible through the microscope via a drawing tube. The track lengths, measured by placing the LED image at either end of a selected track and energizing the cursor, were computed with reference to a previously entered calibrated distance presented as a histogram.

3.5 <u>Results</u>

The results of the analysis of apatite mounts from four samples of the pluton are shown in Table 2. Samples LH-1, -2, and -5 show similar ages (82.1, 97.0, 92.4 Ma), while the age of LH-3 seems anomalously old (123 Ma). The reason for this discrepancy is not fully understood, although it was noted that the apatite grains from sample LH-3 were small and anhedral compared to the larger euhedral crystals obtained from the other three samples. There also exists an elevation range of 100m between the highest (LH-3) and lowest (LH-5) samples. At an uplift rate of .05mm/year this would pass LH-3 through the annealing zone 2 x 10^6 years before LH-5.

The distribution of confined fission track lengths is shown in Figure 9. The pattern of this length distribution is diagnostic of the cooling history of a sample (Hurford, 1986). A steady slow cooling history will produce a broad, asymmetric, unimodal distribution of track lengths (Hurford, 1986). Hurford (1986) also found that 150 slowly cooled plutonic rocks had mean track lengths that lay within the range of 12-14 . The samples LH-5, -2, and -1

SAMPLE	# GRAINS				U ppm	AGE <u>(Ma)</u>
 דע_1	24	6 29	7 13	1 94	12	97.0 + 7.5
LH-1 LH-2	19	3.31	4.03	1.94	7	92.4 ± 8.9
LH-3	28	6.47	5.86	1.98	10	123 ± 9.4
LH-5	30	5.53	7.42	1.94	13	82.1 <u>+</u> 6.0

TABLE 2

Ages calculated using a laboratory zeta factor of 1142.0 ± 380.0 ; neutron flux monitored using standard glass SRM 614; error calculated using the Poissonian Statistical Method (Green, 1981) and reported to 10.

show unimodal distributions, with mean track lengths between 12.8 and 13.5 . Sample LH-3 has a mean track length within the slow cooling range; however, the frequency distribution shows a slightly bimodal population. Bimodal distributions usually result from a subsequent heating event after initial cooling. The proximity of this sample to the perimeter of the pluton may have left it susceptible to reheating from hydrothermal fluids. However, mineral alteration within the samples does not provide conclusive evidence for this assertion. The overall pattern of the track length distributions and mean lengths suggests slow steady cooling.



Chapter 4: Hornblende Geobarometer

4.0 <u>Introduction</u>

Calculating the uplift history of a pluton necessitates an estimate of the depth at which it crystallized and thus stopped moving upward relative to the host country rock. Geobarometers are available which allow a calculation of the ambient pressure at the time of crystallization. This study used a hornblende geobarometer to calculate crystallization pressure and thus yield the starting point on the depth-time profile of the Liberty Hill pluton.

4.1 <u>Hornblende Geobarometer Theory</u>

Hammerstrom and Zen (1986) proposed an empirical correlation between the estimated pressures of crystallization of calc-alkaline plutons and the total Al content of hornblende. In common rocks, there are usually too many thermodynamic degrees of freedom to constrain amphibole-bearing systems enough to use compositional parameters to determine pressure. However, in calcalkaline magma systems there are sufficient constraints to correlate the Al content of hornblende to the pressure of solidification (Hollister et al, 1987).

According to the phase rule, at a specific temperature and pressure the number of degrees of freedom of a system is determined by the number of phases subtracted from the minimum number of chemical components (F = C-P+2; "Phase Rule"). A calc-alkaline rock may be described by the 10 major oxides as well as seven solid phases. Assuming that melt and an H₂O-bearing vapour phase are present at the end of crystallization, the system has a variance of 3: P, T and one compositional degree of freedom. This variance is lowered to one by

the observation that in most calc-alkaline plutons the rim composition of plagioclase is nearly uniform and the temperature range of solidification of hornblende is only approximately 100°C for pressures above 2 Kb (Hollister et al., 1987). The final degree of freedom is accounted for by variation in the Al content of hornblende.

To use the geobarometer, the following four conditions must be met:

- The phases quartz, plagioclase, hornblende, biotite, orthoclase, titanite and magnatite must have crystallized together form the melt.
- 2) Only rim compositions of the hornblende should be analyzed.
- 3) The pressure should be above 2 Kb.
- 4) The rim plagioclase composition should be between An_{23} and An_{35} (Hollister et al., 1987).

If these conditions are met then the pressure P is given by the equation: P $(\pm 1 \text{ Kb}) = -4.76 + 5.64 \text{ Al}_{t} \dots (10)$

where; Al_t is the total number of cations of Al per formula unit based on 23 oxygens. The Liberty Hill pluton meets these conditions.

4.2 <u>Results</u>

Eight hornblende crystals from samples LH-5 and LH-2 were analyzed on the JEOL-733 Microprobe operating at 15KV and 5 nanno amps. Both samples contained euhedral to subhedral dark yellow-green hornblende crystals up to 4 mm in size. As suggested by Hollister et al. (1987) only crystal rim analyses were used in pressure calculations. A list of the total number of aluminum cations per formula unit based on 23 oxygens and the calculated pressures are presented in Table 3.

TABLE 3

SAMPLE	#	AL	PRESSURE (Kb ± 1)
TH-2	1	1 6948	4 80
	2	1 6670	4,64
	3	1,6750	4,69
	"	1.7195	4.94
	4	1.6808	4.72
1 21 28 3 3 1 20 1 20 1 20 1 20 1 20 1 20 1 20			
LH-5	5	1.9122	6.02
	n	1.6017	4.27
	6	1.4784	3.58
	7	1.6348	4.46
	"	1.6372	4.47
	8	1.55546	4.01
	"	1.5574	4.14

The average pressure of these results if 4.6 ± 1 Kb (standard error estimate for this technique). This is comparable to values of 3.1 to 4.7 Kb estimated by Speer (1987) using the same geobarometer.

5.0 <u>Temperature-Time</u>

In principle, different cooling ages measured on coexisting minerals should allow construction of a cooling history for the Liberty Hill pluton. However, in practice there are limitations on the development of a temperature-time history because of variability of closure temperatures. Closure temperatures are sensitive to a number of variables, among which cooling rate and effective grain size are the most important (Cliff, 1985). At present, difficulty in determining the precise effective grain size of diffusive gas loss has led to closure temperatures with significant errors. The closure temperatures chosen for this study and the ages obtained are presented in Table 4.

MINERAL	ISOTOPIC SYSTEM	CLOSURE TEMPERATURE	AGE (MA)
Hornblende	40 _{Ar/} 39 _{Ar}	530 ± 40°C*	303 <u>+</u> 3
Muscovite	40 _{Ar/} 39 _{Ar}	350 <u>+</u> 25 ⁰ C**	314 <u>+</u> 3
Biotite	40 _{Ar/} 39 _{Ar}	280 ± 40°C*	302 ± 3 309 ± 3
Microcline	40 _{Ar/} 39 _{Ar}	150 ± 30°C*	206 ± 3
Apatite	Fission Track	105 <u>+</u> 10 ^o C*	$ \begin{array}{r} 123 \pm 9.4 \\ 97.0 \pm 7.5 \\ 92.4 \pm 8.9 \\ 82.1 \pm 6.0 \end{array} $

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Table 4

* (Harrison and McDougall, 1980)

** (Dallmeyer et al., 1981)

The data in Table 4 is open to a number of interpretations. With additional information it may be possible to further constrain the post-magmatic cooling pattern of the Liberty Hill pluton. The simplest interpretation of current data follows:

- A. ~300 Ma The Liberty Hill pluton was emplaced at a depth of approximately 15km, corresponding to 4.6 Kb of pressure, with a central temperature of approximately 725°C (Speer, 1987). Similar cooling ages for hornblende, muscovite and biotite suggest the pluton cooled rapidly through the closure temperature of these minerals to the ambient temperature of the host Carolina Slate Belt, after which both followed similar cooling histories.
- B. ~260 Ma The pluton cooled through the temperature required for retention of argon in microcline (~150°C).
- C. ~100 Ma to present The pluton cooled through ~150°C allowing fission tracks to accumulate in apatite, after which the temperature stayed below this level.

These data are presented in the T - t (temperature - time) graph shown in Figure 10. The initial portion of the curve shows rapid cooling from 700° C to 300° C.

An attempt was made to estimate the amount of time a body of magma of this size would take to cool by conduction from 700°C to 300°C. Jaeger (1967) developed simple cooling models of magma suddenly emplaced at a uniform temperature and cooled by conduction into country rocks with the same thermal properties as the magma. This cooling model ignores the affects of latent heat, convection and differentiation within the magma, and thus gives only an order of magnitude estimates of the cooling time. Jaeger (1967) determined



Figure 10: Simplest interpretation of the cooling history of the Liberty Hill pluton, plotted from measured mineral ages against preferred system closure temperatures.
that the sizes, positions and conductivities of cooling bodies is expressed in a family of curves which relate Eand Twhere:

 $\xi = \frac{\gamma}{\alpha}$; x is a position co-ordinate and a is the linear dimension (radius or half thickness) and

 $\gamma = \frac{Kt}{\alpha^2}$ t is time and k is thermal diffusivity. Using the k value of granite (0.01 cm²/sec) γ equals:

∽ _ 31.5t/a²

where: t is time in years and is a linear dimension in meters.

At this stage determination of the time at which the temperature at the center of an intrusion reaches a specified value K of the initial temperature is possible. In this case:

 $K = 300^{\circ}C/700^{\circ}C = .0428$

Approximating the pluton as a cubic intrusion, the time for this much cooling is given by the function:

 $Q(0, \gamma) = K^{1/3} (K^{1/3} = .754)$

From Table 2 (Jaeger, 1967) for K = .754, \Im is .04. Therefore for the Liberty Hill pluton, which has a half width of 13,000m the cooling time is given by:

 $t = (13,000)^2 (0.4)/31.5 = 2.2 \times 10^6$ years

A cooling time of less than 5 x 10^6 years is within the error of the ages obtained for this portion of the curve.

5.1 <u>Depth-Time</u>

There are three time points to constrain in order to interpret a depthtime path; the 300 Ma age of emplacement, the 260 Ma age given by the microline, and the 98 Ma average of the fission track ages.

Two independent pressure indicators, the hornblende geobarometer and the metamorphic mineral assemblage of the contact aureoloe, agree that the emplacement pressure was ~4.5 kb giving a depth of about 15 km. The Rb/Sr whole rock the hornblende, muscovite and biotite ${}^{40}\text{Ar}/{}^{39}\text{Ar}$ ages all of 300 Ma indicate that the pluton cooled rapidly to a temperature of about 300°C. This is consistent with the metamorphic mineral assemblage of the surrounding country rock. The presence of metamorphic vermiculite in this assemblage suggests a maximum ambient temperature of 400°C (Veld, 1978). Thus at 15 km depth the upper temperature limit of the host rock if 400°C, while the lower limit is 300°C, as supplied by the cooling history. This yields a geothermal gradient between 26°C/km and 20°C/km. Assuming the paleogeothermal gradient decayed in a linear fashion to the present day Liberty Hill geothermal gradient of 14.9 01°C/km (Costain et al., 1986) the depth of the 150°C + 30°C microcline temperature is between 5 and 9 km. The 105 10°C closure temperature of the fission track dates yields a depth range of 5 to 7 km. The final point of the curve is the surface where the Liberty Hill pluton is exposed today (summarized in Figure 11).

The initial rapid uplift of the pluton (~0.2 mm/year) between 300 Ma and 280 Ma corresponds well to uplift rates determined elsewhere in the Piedmont Belt during this period of time (Dallmeyer et al., 1986). From 280 Ma to 100 Ma the uplift rate of the pluton appears slower (0.01mm/year). Possible regional events which affected this uplift are late Triassic-early Jurassic (205 to 190 Ma) systems of basins which opened up along the length of the Appalachians in response to the opening of the Atlantic Ocean (Swanson, 1986). During this time the Crowburg and Wadesboro basins developed north of the Liberty Hill pluton. The large fault that activated these basins is in

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time (Ma)





contact with the northwestern edge of the pluton, dipping in along the pluton edge at an angle of approximately 40° (Bell and Popenoe, 1976). The pluton lies on the uplifted side of the fault. If significant upward movement occurred along the fault the assumption of a continuous slow uplift may be invalid. However, as there is such a narrow closure temperature range between the microcline and apatite (50°C), had significant uplift occurred the fission tracks would record an age much closer to that of the microcline. The rate of uplift from 100 Ma until the present appears to increase (~.05 mm/year). A study of the nature of the sediments of the nearby Coastal Plain (Owens, 1970) suggests that the Southern Appalachians has undergone a period of gentle uplift since Cretaceous times, except for a mid-Tertiary hiatus. The gentle uplift trend of the Southern Appalachians is consistent with the uplift data from the Liberty Hill pluton. The main conclusions of this study are as follows:

- 1) The LH-4 hornblende, muscovite and biotite plateau ages are 303 ± 3 Ma, and 309 ± 3 Ma respectively.
- 2) The LH-2 biotite yields a plateau age of 302 ± 3 Ma while the LH-2 microcline total gas age is 260 ± 3 Ma.
- 3) The fission track ages for apatites from samples LH-1, -2, -3 and -5 are 97.0 7.5, 92.4 8.9, 123 9.4, and 82.1 6.0 respectively.
- 4) The hornblende geobarometer gives a crystallization pressure of 4.6 \pm 1 kb.
- 5) The shape of the track length distributions of unconfined tracks suggests slow steady uplift through the apatite annealing temperature range.
- 6) Conversion of the T-t data to depth-time data reveals a composite uplift pattern, involving initial rapid uplift of .02 mm/year (during the Alleghanian orogeny). The uplift rate slows to ~0.01 mm/year between 280 and 100 Ma, and speeds up again to ~0.5 mm/year from 100 Ma until the present.

Further work involving forward modeling of the fission track distribution may help constrain the most recent portion of the uplift history of the Liberty Hill pluton. Also fission track dates from zircon crystals would be useful to obtain as these have approximately the same closure temperature as 40Ar/39Ar dating the microcline.

The depth and time of emplacement is known for relatively few plutons in the central and southern Appalachians; the subsequent cooling and uplift history for even fewer. The variety of uplift patterns for the studied

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plutons indicates a complex constructional variability across the Appalachians. Investigating cooling histories of other plutons in this region will aid in our understanding of the uplift history of the southern-central Appalachians.

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Temp Step (^o C)	Age	(Ma)	MV ³⁹ Ar
200 - 500	0		.05
500 - 550	236.25	.654	.023
550 - 600	246.23	.480	.050
600 - 650	252.27	.217	.071
650 - 700	261.28	.184	.087
700 - 750	266.16	.246	.064
750 - 800	265.81	.481	.036
800 - 850	267.50	.780	.034
850 - 900	268.33	1.14	.017
900 - 950	268.90	.558	.043
950 - 1000	271.17	2.51	.009
1000 - 1050	268.81	.951	.027
1050 - 1100	266.87	1.70	.015
1100 - 1140	269.54	1.65	.015

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Temp Step (^o C)	Age (Ma	a)	_ع 39 _{Ar}	
200 - 650	194.50	7.90	.009	
650 - 725	278.97	7.67	.010	
725 - 800	275.94	11.13	.008	
800 - 850	279.71	6.46	.007	
850 - 900	300.88	1.18	.030	
900 - 925	301.04	. 571	.091	
925 - 950	304.38	. 309	.074	
950 - 975	304.68	. 399	.056	
975 - 1000	299.16	1.36	.023	
1000 - 1050	302.92	4.49	.010	
1050 - 1125	240.37 2	23.71	.015	

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Temp Step (^o C)	Age (Ma)	MV ³⁹ Ar
200 - 500	254.17 1.12	.016
500 - 600	283.80 .758	.039
600 - 650	287.27 1.92	.012
650 - 700	296.87 .240	.072
700 - 750	323.47 8.16	.014
750 - 800	309.92 1.41	.019
800 - 850	309.55 .936	. 029
850 - 900	320.01 1.66	.031
900 - 950		
950 - 1000		

1000 - 1050

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<u>LH-2 Biotite</u>

Temp	St	tep (^o C)	Age	(Ma)	MV ³⁹ Ar
200	-	500	145.40	5.22	.013
500	-	600	296.58	1.05	.025
600	-	650	302.91	. 539	.056
650	-	700	302.91	. 303	.101
700	-	750	302.12	. 908	.023
750	-	800	301.54	.451	.079
800	-	850	300.58	1.34	.020
850	-	900	301.84	.660	.033
900	-	950	301.48	.668	.030
950	-	1000	295.30	1.13	.028
1000	-	1050	272.33	6.36	.009

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LH-	4	В	i	ο	t	i	t	е
		_	_	_	_	_		

Temp	St	ep (^o C)	Age	(Ma)	MV ³⁹ Ar
200	-	550	249.89	.716	.045
550	-	600	306.62	.567	.041
600	-	650	302.58	.536	.070
650	-	700	303.50	. 579	.053
700	-	750	302.03	1.11	.028
750	-	800	303.25	2.05	.020
800	-	850	323.19	2.52	.019
850	-	900	307.05	.624	.061
900	-	950	298.91	1.02	.037
950	- :	1000	278.58	4.45	.014

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	Glass dosimeter Nr counted	= =	SRN611 1225	98.0 _	5.4 MA
-	RHOP D for class	Ξ	1.950E+01 ± 0.560E+00		
-	Area of 1 QUAD	Ξ	1.248-06		
	Fooled NS/NI	=	0.880 / 0.020	MEAN	AGE
-	Mean NS/N1	Ξ	0.889 ± 0.989		
				08.7 :	10.8 MA

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APPENDIX 4

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Raw data from track length distributions

FT LENGTH DATA POINTS IN MICRONS

2 	.4.51	14.34	14.87	14.01	15.41	11.11	· · ·	• • • • •	11.34
1.37			10.23		11.14	11.11		- 1	11.53

SUMMARY OF MEASURED STATISTICS

		20	FΤ	LENGTHS	MEASURED
THE	MEAN IS	13.	495		
THE	STD DEV IS	0.9	979		
THE	STD ERROR IS	0.2	225		
THE	SKEWNESS IS	-0.0	024		
THE	KURTOSIS IS	-0.8	877		

MEASURED FREQUENCY TABLE

INTERVAL RANGE	INT. MID.	FREQUENCY	CUMULATIVE	%
FROM TO	POINT	VALUE	FREQUENCY	
0.000- 1.000	0.500	0	0	0.0
1.000- 2.000	1.500	0	Ō	0.0
2.000- 3.000	2.500	0	0	0.0
3.000-4.000	3.500	0	Ο.	0.0
4.000- 5.000	4.500	0	0	0.0
5.000- 6.000	5.500	О	0	0.0
6.000- 7.000	6.500	0	0	0.0
7.000- 8.000	7.500	0	. 0	0.0
8.000- 9.000	8.500	0	0	0.0
9.000-10.000	9.500	0	0	0.0
10.000-11.000	10.500	0	0	0.0
11.000-12.000	11.500	2	2	10.0
12.000-13.000	12.500	5	ī	25.0
13.000-14.000	13.500	6	13	30.0
14.000-15.000	14.500	6	19	30.0
15.000-16.000	15.500	1	20	5.0
16.000-17.000	16.500	0	20	0.0
17.000-18.000	17.500	0	20	0.0
18.000-19.000	18.500	0	20	0.0
19.000-20.000	19.500	0	20	0.0



FT LENGTH DATA POINTS IN MICRONS

16.12 14.64 14.61 14.61 14.61 14.61 14.61 14.61 14.66 16.65 10.11 11.44

SUMMARY OF CORRECTED STATISTICS

		24.70	FT	LENGTHS	MEASURED
THE	MEAN IS	13.363			
THE	STD DEV IS	1.101			
THE	STD ERROR IS	0.226			
THE	SKEWNESS IS	-0.072			
THE	KURTOSIS IS	-1.017			

CORRECTED FREQUENCY TABLE

INTERVAL RANGE	INT. MID.	FREQUENCY	CUMULATIVE	%
FROM TO	POINT	VALUE	FREQUENCY	
0.000- 1.000	0.500	0	0	0.0
1.000- 2.000	1.500	Ο	0	0.0
2.000- 3.000	2.500	0	0	0.0
3.000- 4.000	3.500	0	0 ·	0.0
4.000- 5.000	4.500	0	0	0.0
5.000- 6.000	5.500	0	0	0.0
6.000- 7.000	6.500	0	0	0.0
7.000- 8.000	7.500	0	0	0.0
8.000- 9.000	8.500	0	0	0.0
9.000-10.000	9.500	0	0	0.0
10.000-11.000	10.500	0	0	0.0
11.000-12.000	11.500	3	3	11.6
12.000-13.000	12.500	1	9	26.7
13.000-14.000	13.500	\overline{t}	. 17	29.7
14.000-15.000	14.500	ī	24	27.6
15.000-16.000	15.500	1	25	4.3
16.000-17.000	16.500	0	25	0.0
17.000-18.000	17.500	0	25	0.0
18.000-19.000	18.500	0	25	0.0
19.000-20.000	19.500	0	25	0.0



FT LENGTH DATA POINTS IN MICRONS

SUMMARY OF MEASURED STATISTICS

		66 FT	LENGTHS	MEASURED
THE	MEAN IS	13.565		
THE	STD DEV IS	1.210		
THE	STD ERROR IS	0.150		
THE	SKEWNESS IS	-0.117		
THE	KURTOSIS IS	0.434		

MEASURED FREQUENCY TABLE

INTERVAL RANGE	INT. MID.	FREQUENCY	CUMULATIVE	%
FROM TO	POINT	VALUE	FREQUENCY	
· .				
0.000- 1.000	0.500	0	0	0.0
1.000- 2.000	1.500	0	0	0.0
2.000- 3.000	2.500	0	0	0.0
3.000- 4.000	3.500	0	0	0.0
4.000- 5.000	4.500	0	0	0.0
5.000- 6.000	5.500	0	0	0.0
6.000- 7.000	6.500	0	0	0.0
7.000- 8.000	7.500	0	0	0.0
8.000- 9.000	8.500	0	0	0.0
9.000-10.000	9.500	U	0	0.0
10.000-11.000	10.500	3	3	4.5
11.000-12.000	11.500	3	6	4.5
12.000-13.000	12.500	14	20	21.2
13.000-14.000	13.500	19	. 39	28.8
14.000-15.000	14.500	21	60	31.8
15.000-16.000	15.500	5	65	7.6
16.000-17.000	16.500	0	65	0.0
17.000-18.000	17.500	1	66	1.5
18.000-19.000	18.500	0	66	0.0
19.000-20.000	19.500	0	66	0.0



FT LENGTH DATA POINTS IN MICRONS

 	 	14.14			

SUMMARY OF CORRECTED STATISTICS

				80.88	FΤ	LENGTHS	MEASURED
THE	MEAN IS		1	3.464			
THE	STD DEV I	S		1.316			
THE	STD ERROR	2 IS	(0.147			
THE	SKEWNESS	IS	- (0.208			
THE	KURTOSIS	IS	(0.354			

CORRECTED FREQUENCY TABLE

INTERVAL	RANGE	INT.	MID.	FREQU	ENCY	CUMULA	TIVE	%
FROM	ТО	POINI	•	VALUE		FREQUE	NCY	
0.000- 1	.000	0.50	0	0		0		0.0
1.000- 2	.000	1.50	0	0		0		0.0
2.000- 3	.000	2.50	0	0		()		0.0
3.000- 4	.000	3.50	0	0		0		0.0
4.000- 5	.000	4.50	0	0	,	0		0.0
5.000- 6	.000	5.50	0	0		0		0.0
6.000- 7	.000	6.50	0	0		υ		0.0
7.000- 8	.000	7.50	0	0		0		0.0
8.000- 9	.000	8.50	0	0		0		0.0
9.000-10	.000	9.50	0	0		0		0.0
10.000-11	.000	10.50	0	5		ō		5.8
11.000-12	.000	11.50	0	-1		9		5.3
12.000-13	.000	12.50	0	18		27		22.8
13.000-14	.000	13.50	Q	23		51		28.7
14.000-15	.000	14.50	0	24		75		29.5
15.000-16	.000	15.50	0	5		80		6.6
16.000-17	.000	16.50	()	0		30		0.0
17.000-18	.000	17.50	0	1		81		1.2
18.000-19	.000	18.50	0	0		81		0.0
19.000-20	.000	19.50	0 .	0		81		0.0





SAMPLE LOCATION (OR STD)...LIBERTY HILL, SOUTH CAROLINA DATE OF ANALYSIS......01/15/88 LAB IDENTIFICATION CODE....LIBH-003 IRRADIATION CODE.....A: DISKETTE.....A: DATAFILE.....LIBHJ003.A1L MINERAL ANALYZED.....APATITE ANALYSIS BY.....JM

FT LENGTH DATA POINTS IN MICRONS

	12.10 14.21 14.51		14.22			
11.38	12.67.	2.20	12.46	11.95	 	

SUMMARY OF MEASURED STATISTICS

		36	FΤ	LENGTHS	MEASURED
THE	MEAN IS	12.	779		
THE	STD DEV IS	1.	097		
THE	STD ERROR IS	Ο.	185		
THE	SKEWNESS IS	0.	362		
THE	KURTOSIS IS	-0.	476		

MEASURED FREQUENCY TABLE

INTERVAL RANGE	INT. MID.	FREQUENCY	CUMULATIVE	%
FROM TO	POINT	VALUE	FREQUENCY	
0.000- 1.000	0.500	0	0	0.0
1.000- 2.000	1.500	0	0	0.0
2.000- 3.000	2.500	0	0	0.0
3.000- 4.000	3.500	0	0	0.0
4.000- 5.000	4.500	0	0	0.0
5.000- 6.000	5.500	0	Q	0.0
6.000- 7.000	6.500	0	0	0.0
7.000- 8.000	7.500	0	0	0.0
8.000- 9.000	8.500	0	0	0.0
9.000-10.000	9.500	0	0	0.0
10.000-11.000	10.500	1	1	2.8
11.000-12.000	11.500	ī	3	19.4
12.000-13.000	12.500	16	24	44.4
13.000-14.000	13.500	-1	28	11.1
14.000-15.000	14.500	ī	35	19.4
15.000-16.000	15.500	1	36	2.8
16.000-17.000	16.500	0	36	0.0
17.000-18.000	17.500	0	3.6	0.0
18.000-19.000	18.500	0	36	0.0
19.000-20.000	19.500	0	36	0.0

SAMPLE LOCATION (OR STD)...LIBERTY HILL.SOUTH CAROLINA DATE OF ANALYSIS......01/15/88 LAB IDENTIFICATION CODE....LIBH-003 IRRADIATION CODE.....A: DISKETTE.....A: DATAFILE.....LIBHJ003.A1L MINERAL ANALYZED.....APATITE ANALYSIS BY.....JM

FT LENGTH DATA POINTS IN MICRONS

13.59 122.20 122.00	11.30 14.12 12.64	18.06 11.08 12.53		13.09 14.09 12.09		
i La Color	 تشاهد	1	 			

SUMMARY OF CORRECTED STATISTICS

		46.65	FΤ	LENGTHS	MEASURED
THE	MEAN IS	12.732			
THE	STD DEV IS	1.148			
THE	STD ERROR IS	0.170			
THE	SKEWNESS IS	0.458			
THE	KURTOSIS IS	-0.478			

CORRECTED FREQUENCY TABLE

INTERVAL	RANGE	INT. MID.	FREQUENCY	CUMULATIVE	%
FROM	ТО	POINT	VALUE	FREQUENCY	
0 000	1 000	0 500	0	0	0 0
1.000-	2.000	0.500	0	0	0.0
1.000- 4	2.000	1.500	U	0	0.0
2.000- 3	3.000.	2.500	0	0	0.0
3.000	1.000	3.500	0	0	0.0
4.000- 3	5.000	4.500	0	0 ·	0.0
5.000- (5.000	5.500	0	0	0.0
6.000- 7	7.000	6.500	0	0	0.0
7.000- 8	3.000	7.500	0	Ŭ	0.0
8.000- 9	9.000	8.500	0	· 0	0.0
9.000-10	.000 .	9.500	0	0	0.0
10.000-11	.000	10.500	2	2	3.4
11.000-12	2.000	11.500	10	12	21.5
12.000-13	3.000	12.500	21	33	45.3
13.000-14	.000	13.500	5	38	10.5
14.000-15	5.000	14.500	8	46	17.1
15.000-16	.000	15.500	1	47	2.3
16.000-17	.000	16.500	()	47	0.0
17.000-18	.000	17.500	()	-17	0.0
18.000-19	.000	18.500	0	47	0.0
19.000-20	.000	19.500	0	47	0.0



A:LIBHJ003.A1L

SAMPLE LOCATION (OR STD)...LIBERTY HILL, SOUTH CAROLINA DATE OF ANALYSIS......01/14/88 LAB IDENTIFICATION CODE...LIBH-005 IRRADIATION CODE.....A: DISKETTE.....A: PATAFILE.....LIBHJ005.A2L MINERAL ANALYZED.....APATITE ANALYSIS BY.....JM

FT LENGTH DATA POINTS IN MICRONS

	· 4 / 1 ·	· (1 ·			· • · ;			• • •	· · · ·
			20105						
					· / 57		14 42	:	
11111		14.14			1	11111			
		- C - C - C				· · / /			
4.44	20.00		101.4	14	20100				
,		•• • /		•••		• • • • •			
	:			÷7 ÷ =		1111			-
			14.11					. .	
					••	· · · · ·			-
			1.1.1.1					· · · · .	

SUMMARY OF MEASURED STATISTICS

		70	FΤ	LENGTHS	MEASURED
THE	MEAN IS	13.	508		
THE	STD DEV IS	0.	954		
THE	STD ERROR IS	0.	115		
THE	SKEWNESS IS	-0.	368		
THE	EURTOSIS IS	Ο.	374		

MEASURED FREQUENCY TABLE

INTERVAL RANGE	INT. MID.	FREQUENCY	CUMULATIVE	0/ /0
FROM TO	POINT	VALUE	FREQUENCY	
0.000- 1.000	0.500	O	Ġ	0.0
1.000-2.000	1.500	Û	0	÷.0
2.000- 3.000	2.500	С	()	0.0
3.000- 4.000	3.500	0	0	0.0
4.000- 5.000	4.500	0	0	0.0
5.000- 6.000	5.500	0	0	0.0
6.000- 7.000	6.500	0	0	<u>.</u>
7.000- 8.000	7.500	(¹	0	14.C
8.000- 9.000	8.300	0	:)	(
9.000-10.000	9.500	0	12	2.0
10.000-11.000	10.500	1	•	
11.000-12.000	11.300	3	4	. .3
12.000-13.000	12.500	13	17	18.6
13.000-14.000	13.500	32	4.9	15.7
14.000-15.000	14.500	18	67	25.7
15.000-16.000	15.500	; ;	7.0	4.3
16.000-17.000	16.500	Û	7.0	0.0
17.000-18.000	17.500	0	7.0	6.0
18.000-19.000	18.500	Ú	7()	0.0
19.000-20.000	19.500	0	$\overline{i}(0)$	9.0



SAMPLE LOCATION (OR STD)...LIBERTY HILL, SOUTH CAROLINA DATE OF ANALYSIS......01/14/88 LAB IDENTIFICATION CODE....LIBH-005 IRRADIATION CODE.....A: DISKETTE.....A: DATAFILE.....LIBHJ005.A2L MINERAL ANALYZED.....APATITE ANALYSIS BY.....JM

FT LENGTH DATA POINTS IN MICRONS

	11.41								11.16
		12 12		14 5.9				-	
10 15				14 14				14.132	
		17 EX	19.05	7 5					14.32
10.05	15 79 .		10.20	10.00			11 11	17 21	17.19
10.02	10.10	10.01	10.12		10.00	10 / 0			
-2-32		10.00	10.10	10.00					
11111	1-1-2	16.00	19,00						

SUMMARY OF CORRECTED STATISTICS

		85.82	FT	LENGTHS	MEASURED
THE	MEAN IS	13.458			
THE	STD DEV IS	0.981			
THE	STD ERROR IS	0.107			
THE	SKEWNESS IS	-0.465			
THE	KURTOSIS IS	0.494			

CORRECTED FREQUENCY TABLE

INTERVAL RANGE	INT. MID.	FREQUENCY	CUMULATIVE	%
FROM TO	POINT	VALUE	FREQUENCY	
0.000- 1.000	0.500	0	0	0.0
1.000- 2.000	1.500	0	Û	0.0
2.000- 3.000	2.500	Q	0	0.0
3.000- 4.000	3.500	0	0	0.0
4.000- 5.000	4.500	0	0	0.0
5.000- 6.000	5.500	0	0	0.0
6.000- 7.000	6.500	0	0	0.0
7.000- 8.000	7.500	0	. 0	0.0
8.000- 9.000	8.500	0	0	0.0
9.000-10.000	9.500	0	0	0.0
10.000-11.000	10.500	2	2	1.8
11.000-12.000	11.500		6	5.0
12.000-13.000	12:500	17	23	20.0
13.000-14.000	13.500	39	G 2	45.6
14.000-15.000	14.500	20	83	23.9
15.000-16.000	15.300	3	86	3.7
16.000-17.000	16.500	0	86	0.0
17.000-18.000	17.500	0	86	0.0
18.000-19.000	18.500	0	86	0.0
19.000-20.000	19.500	0	86	0.0



JМ

APPENDIX 5

Microprobe result for hornblende geobarometer

JN SCHEDULE ZZZ 3-NOV-87 14:15:34

WDS:SI WDS:TI WDS:AL WDS:CR WDS:FE WDS:NI WDS:MN WDS:MG W WDS:NA WDS:K 4F CORRECTION 15.00 KV 40.00 Deg 0. Of Iterations 0 --- K [Z] [A] [F] [ZAF] ATOM.% WT.%

							• / • · · · • ·	•
-K	0.150	0.998	1.343	0.998	1.340	16.83	20.18	
$-\kappa$	0.006	1.134	1.033	0.981	1.151	0.34	0.70	
K	0.032	1.030	1.483	0.991	1.515	4.32	5.00	÷
₹-K	0.000	1.139	1.014	0.954	1.102	0.02	0.03	
E-K	0.149	1.146	1.004	1.000	1.152	7.19	17.23	
-K	0.000	1.138	1.013	1.000	1.153	0.00	0.00	
(-K	0.004	1.163	1.008	1.000	1.173	0.21	0.50	
i-K	0.027	0.985	1.721	0.994	1.688	4.56	4.68	÷
,-K	0.078	1.029	1.044	0.992	1.067	4.88	8.35	
-κ	0.003	1.023	2.250	0.996	2.295	1.45	1.43	×
-к	0.010	1.051	1.065	0.981	1.099	0.66	1.11	
-к	0.186	0.954	2.298	0.999	2,192	59.54	40.79	D *
					· ·	TOTAL=	100.00	

High Absorbance

IDE	RESULTS
	WT%
82	43.23
02	1.17
<i>.</i>	9.44
3	0.05
	22.15
0	0.00
0	0.65
0	7.81
:0	11.69
20	1.93
20	1.34
TAL=	99.44

OF OXY = 23.00 #16 (NEAR XTL RIM)

SI	6.5884
ΤI	0.1337
AL	1.6948
CR	0.0030
FE	2.8239
NI	0.0000
MN	0.0840
MG	1.7730
CA	1.9088
NA	0.5692
r N	0.2597

14:18:39

 $kb = 4.80 \pm 1$

N_SCHEDULĘ ZZŻ 1-87

21:28

WDS:TI WDS:AL WDS:CR WDS:FE WDS:NI WDS:MN WDS:MG W F.S WDS:SI :CA WDS:NA WDS:K

F CORRECTION 15.00 KV 40.00 Dec

. Of Iterations 0

	ĸ	[2]	[A]	[F]	[ZAF]	ATOM	.% WT.:	×.
-K	0.147	0.998	1.339	0.998	1.336	16.38	19.75	
-K	0.007	1.135	1.033	0.981	1.151	0.42	0.86	
-K	0.031	1.030	1.480	0.991	1.513	4.14	4.81	×
-K	0.000	1.139	1.014	0.954	1.104	0.01	0.02	
-K	0.147	1.147	1.004	1.000	1.152	7.07	17.04	
-K	0.000	1.139	1.013	1.000	1.154	0.00	0.00	
-K	0.004	1.164	1.008	1.000	1.174	0.22	0.52	
-K	0.025	0.986	1.725	0.994	1.693	4.22	4.36	×
-K	0.077	1.030	1.043	0.992	1.067	4.81	8.28	
-K	0.003	1.023	2.260	0.996	2.306	1.40	-1.39	×
-K	0.010	1.052	1.063	0.981	1.098	0.71	1.18	
-K	0.190	0.955	2.296	0.999	2.191	60.64	41.79	D *
						TOTAL=	100.01	

- High Absorbance

T	RESULTS	
	WT%	
J2	.42.33	
02	1.44	
203	9.09	
203	0.03	
Э	21.91	
Э	0.00	
5	0.67	
5	7.27	
7	11.30	
20	1.87	
20	1,43	
-AI =	97.62	
		#20 (NEAR XTL RIM)
		 11 1 4 1

OF 0XY = 23.00 KD = 4.64

,

 $\mathbb{S}I$ 6.5870 ΤI 0.1684 1.6670 AL 0.0034 CR FE 2.8520 0.0000 I 0.0877 --1N МG 1.6859 CA 1.9333 0.5646 NA 0.2834 К
N SCHEDULE ZZZ -NOV-87 10:43: 2

WDS:SI WDS:TI WDS:AL WDS:CR WDS:FE WDS:NI WDS:MN WDS:MG W WDS:NA WDS:K

F CORRECTION 15.00 KV 40.00 Deg

. 0.	f Itera	tions	0					
	к	[]]	[A]	[F]	[ZAF]	ATOM	.%. WT.;	<i>.</i>
·К	0.147	0.998	1.341	0.998	1.338	16.35	19.70	
·Κ	0.007	1.135	1.033	0.981	1.151	0.40	0.83	
К	0.031	1.030	1.483	0.991	1.516	4.17	4.85	÷
К	0.000	1.139	1.014	0.954	1.103	0.01	0.02	
К	0.148	1.147	1.004	1.000	1.152	7.12	17.16	
К	0.000	1.139	1.013	1.000	1.154	0.00	0.00	
К	0.004	1.164	1.008	1.000	1.174	0.20	0.48	
К	0.026	0.986	1.725	0.994	1.692	4.41	4.56	×
К	0.078	1.030	1.043	0.992	1.067	4.89	8.42	
\times	0.005	1.023	2.259	0.996	2.306	1.39	1.37	×
K	0.009	1.052	1.064	0.981	1.098	0.30	1.01	
\langle	0.189	0.955	2.294	0.999	2.190	60.44	41.60	D *
						TOTAL=	100.00	

- High Absorbance

ΣC	RESULTS	
	WT%	
3	42.21	
2	1.39	
-	9.16	
	. 0.04	
	22.06	
	0.00	
	0.62	
	7.59	
	11.79	
	1.85	
	1.22	
1 =	97.91	

≇ OF OXY = 23.00

31 AREUNG & 🖯	6.5489 0.1621 1.6750 0.0043 2.8634 0.0000 0.0810 1.7555 1.9595 0.5564	# 39 (NEAR RIMOFYTL) Kb= 4.69 ±1
46:	0.2415 7	

4 SCHEDULE ZZZ -NOV-87 0:28:33 - ^{- - -} -WDS:TI WDS:AL WDS:CR WDS:FE WDS:NI WDS:MN WDS:MG W WDS:SI WDS:NA WDS:K CORRECTION 15.00 KV 40.00 Dec Of Iterations 0 [ZAF] ATOM.% WT.% []] [A] [F] К 0.998 1.342 0.998 1.338 16.54 19.87 К 0.148 0.007 1.134 1.033 0.981 1.150 0.44 0.91 К 0.031 1.030 1.485 0.991 1.517 4.13 4.78 ÷ К К 0.000 1.139 1.014 0.954 1.102 0.00 0.00 17.38 К 1.804 1.000 1.152 7.24 0.150 1.146 К 0.000 1.138 1.013 1.000 1.153 0.00 0.00 K 0.49 0.004 1.163 1.008 1.000 1.173 0.21 \langle 0.985 1.728 0.994 1.695 4.45 4.58 0.027 ÷ \langle 8.28 0.077 1.029 1.044 0.992 1.066 4.82 < 0.996 2.309 1.43 0.006 1.023 2.263 1.45 ÷ 0.981 1.098 0.65 1.09 \leq 1.051 1.064 0.009 0.954 2.299 0.999 2.193 60.06 41.22 D * < 0.187 TOTAL= 100.03 - High Absorbance DE RESULTS WT% 2 :42.57 3 1.52 1: 9.03 0.00 22.35 0.00 0.63 7.64 11.59 1.93 1.31 98.56 #26 (NEAR XTL RIM) 23.00 # OF OXY = $kb = 4.50 \pm 1$ ΞI 6.5656 ΤI 0.1761 ΆL 1.6413 0R 0.0000 ΞE 2.8838 11 0.0000 1N0.0824 1G 1.7550 CΑ 1.9147 0.5781 0.2585 31:38

N SCHEDULE ZZZ -NOV-87 14:27: 5

F WDS:SI WDS:TI WDS:AL WDS:CR WDS:FE WDS:NI WDS:MN WDS:MG W WDS:NA WDS:K

1.067

4.98

÷

÷

8.53

F CORRECTION 15.00 KV 40.00 Dec . Of Iterations 0 ATOM.% WT.% [Z] [A] [F] [ZAF] _ _ К -К 0.998 1.343 0.998 0.149 1.339 16.67 19.98 1.033 -К 0.004 1.134 0.981 1.151 0.26 0.54 -K 0.033 1.030 1.482 0.991 1.514 4.35 5.02 0.000 1.139 1.014 0.953 1.101 0.00 0.00 -К 7.37 17.66 -к 0.153 1.146 1.004 1.000 1.151 -К 1.000 1.153 0.00 0.000 1.138 1.013 0.00 0.48 -К 0.004 1.163 1.008 1.000 1.173 0.20 -К 0.023 0.985 1.726 0.994 1.692 4.35 4.47

1.044

1.24 -K 1.023 2.266 0.996 2.311 1.26 0.005 ÷ 1.04 -K 0.009 1.051 1.065 0.981 1.099 0.62 2.290 0.999 59.93 41.04 D * -К 0.187 0.954 2.185 TOTAL= 99.99

0.992

- High Absorbance

1.029

0.079

-К

IDE	RESULTS
	WT%
02	42.80
02	0.90
2	9.49
د(0.00
5 [°]	22.70
С	0.00
Э	0.62
ā i	7.46
ā	11.94
20	1.67
20	1.25
TAL=	98.83

OF OXY = 23.00

36 (NEFR XTL RIM) KE= 4.94 ±1

SI 6.5820 0.1040 ΤI 1.7195 AL CR 0.0000 2.9205 FΕ NI 0.0000 0.0811 MN MG 1.7085 1.9664 CA '4 0.4980 0.2454

14:30:11

N - CHEDULE ZZZ - J-87 : 9:12

F.S WDS:SI WDS:TI WDS:AL WDS:CR WDS:FE WDS:NI WDS:MN WDS:MG W :CA WDS:NA WDS:K

ΗL	ORREUTI	UN	15.00	KV	40.0	10 Deg		
. 0)f Itera	tions	1					
	К	[2]	(A)	[F]	[ZAF]	ATOM	.% WT.:	×.
-К	0.145	0.998	1.348	0.998	1.344	16.24	19.50	
-K	0.005	1.134	1.033	0.981	1.150	0.29	0.59	
-К	0.036	1.030	1.480	0.991	1.513	4.79	5.55	÷
-K	0.000	1.138	1.013	0.952	1.100	0.01	0.03	
-K	0.153	1.146	1.004	1.000	1.151	7.38	17.72	
-K	0.000	1.138	1.013	1.000	1.153	0.00	0.00	
-K	0.003	1.163	1.008	1.000	1.172	0.17	0.41	
-K	0.024	0.985	1.731	0.994	1.698	4.04	4.16	÷
-К	0.077	1.029	1.044	0.992	1.067	4.80	8.23	
-К	0.003	1.023	2.269	0.996	2.314	1.45	1.43	×
-K	0.010	1.051	1.064	0.981	1.098	0.66	1.11	
-K	0.190	0.954	2.272	0.999	2.167	60.15	41.27	D *
						TOTAL=	100.00	

- Hígh Absorbance

}	RESULTS	
32	41.79	
02	0.99	
203	10.48	
203	0.04	
D	22.78	
0	0.00	
0	0.53	
Э	6.93	
Э	11.52	
20	1.93	
20	1.34	
FAL=	98.33	

OF OXY = 23

23.00 ++ /

#46 (NEAR EDGE AGRINST PLAG)

Kb= 6.02 ±1 , SΙ 6.4706 0.1147 ΤI 1.9122 AL 0.0052 CR ΞĒ 2.9504 1 0.0000 MN 0.0698 1.5995 MG CA 1.9111 0.5791 NA 0.2645 К

N SCHEDULE ZZZ -/~'-87 7:13

F.S WDS:SI WDS:TI WDS:AL WDS:CR WDS:FE WDS:NI WDS:MN WDS:MG W :CA WDS:NA WDS:K

F CORRECTION 15.00 KV 40.00 Deg

2 . Of Iterations ĸ [2] [A] [F] [ZAF] ATOM.X WT.X 1.340 -К 0.148 0.998 0.998 1.337 16.56 19.90 -K 0.006 1.137 1.031 0.980 1.150 0.38 0.79 0.030 1.488 0.991 1.521 4.01 4.65 -K 1.030 ÷ 0.951 0.04 -K 0.000 1.141 1.013 1.101 0.02 7.24 17.40 0.151 1.147 1.004 1.000 1.152 -K -K 0.000 1.141 1.013 1.000 1.156 0.00 0.00 -К 0.003 1.166 1.007 1.000 1.175 0.19 0.45 0.026 0.986 1.731 0.994 1.698 4.44 4.57 -К × 4.92 0.079 0.992 1.066 8.44 -K 1.030 1.043 -К 0.005 1.023 2:266 0.996 2.312 1.45 1.43 ÷ 0.979 0.58 0.97 1.056 1.090 -K 0.008 1.054 60.21 41.35 D * -K 0.192 0.956 2.246 0.999 2.147 TOTAL= 99.98

- High Absorbance

I	RESULTS		
32	· 42.65		
02	1.31		
203	8.79		
203	0.05	· .	
)	22.37		
)	0.00		
0	0.58		
)	7.62		
]	11.82		
30	1.93		
20	1.17		
"AL=	98.28		

23.00 # 4c (NEAR XTL RIM) # OF 0XY =

SI TI	3.5961 0.1524	,	КЬ= 4.27± !
AL	1.6017		-
CR	0.0033		
FE	2.8946		·
۲	0.0000		
1 and	0.0764		
۰G	1.7557		
CA	1.9588		
NA	0.5775		
к	Ø.2314		

F.S WDS:SI WDS:TI WDS:AL WDS:CR WDS:FE WDS:NI WDS:MN WDS:MG W :CA WDS:NA WDS:K

F CORRECTION 15.00 KV 40.00 Deg

. 0	f Itera	tions	2					
	ĸ	[]]	[A]	[F]	[ZAF]	ATOM.	.% WT.>	
-K	0.153	0.999	1.335	0.998	1.333	16.90	20.41	
-K	0.008	1.138	1.030	0.981	1.151	0.47	0.98	
-K	0.028	1.031	1.486	0.991	1.520	3.71	4.33	÷
-K	0.000	1.142	1.013	0.953	1.104	0.01	0.03	
-K	0.144	1.148	1.004	1.000	1.153	6.91	16.69	
-K	0.000	1.142	1.012	1.000	1.157	0.00	0.00	
-K	0.004	1.167	1.007	1.000	1.176	0.22	0.53	
-K	0.028	0.986	1.721	0.994	1.690	4.68	4.84	×
-K	0.076	1.031	1.043	0.992	1.067	4.73	8.16	
-K	0.006	1.024	2.246	0.996	2.293	1.52	1.51	÷
-K	0.007	1.054	1.057	0.979	1.092	0.52	0.87	
-K	0.194	0.957	2.239	0.999	2.142	60.33	41.64	Ð¥
					-	TOTAL=	99.98	

- High Absorbance

I 52 02 203 203	RESULTS WT% 43.74 1.63 8.17 0.04	·	•.	
	21.45 0.00 0.68 8.07 11.42			
20 20 7AL=	2.04 1.05 : 98.30			
Ħ	OF OXY = 23.	+ 00	5a (chose to	XTL RIM) (NEAR GTZ)
			KD= 3.58 21.	
SI TI	6.7123 0.1886	,		
AL	1.4784			
CR FF	0.0050 2 7540			
Ţ	0.0000			
	0.0884			
CA	1.8783			·
NA	0.6068			
к	0.2030		-	

1:38:35

F.S WDS:SI WDS:TI WDS:AL WDS:CR WDS:FE WDS:NI WDS:MN WDS:MG W :CA WDS:NA WDS:K

F CORRECTION 15.00 KV 40.00 Deg

۰,

. (]f Itera	tions	0					
	к	[]]	[A]	[F]	[ZAF]	ATOM	.% WT.;	
-К	0.149	0.998	1.341	0.998	1.337	16.70	20.05	
-К	0.006	1.134	1.033	0.981	1.151	0.37	0.76	
-К	0.031	1.030	1.483	0.991	1.516	4.12	4.77	÷
-K	0.000	1.139	1.014	0.954	1.102	0.01	0.02	
-K	0.150	1.146	1.004	1.000	1.152	7.25	17.40	
-к	0.000	1.138	1.013	1.000	1.153	0.00	0.00	
-K	0.004	1.163	1.008	1.000	1.173	0.21	0.49	
-K	0.026	0.985	1.728	0.994	1.695	4.37	4.50	÷
-К	0.078	1.029	1.043	0.992	1.066	4.90	8.40	
-К	0.003	1.023	2.262	0.996	2.307	1.49	1.47	÷
-K	0.009	1.051	1.064	0.981	1.098	0.60	1.01	
-K	0.187	0.954	2.295	0.999	2.190	59.98	41.15	D *
						TOTAL=	100.01	

- High Absorbance

I	RESULTS
	WT%
	• 42.96
02	1.27
203	9.02
203	0.02
Э	22.37
Э	0.00
3	0.63
Э	7.50
Э С	11.77
30	1.98
20	1.21
"AL=	98.72

Ħ	OF	ΟXY	=	23.00	#66	(NEF.F	GRAIN	FIRA
					КЬ	= 4.46	+1	

81	6.608/
ΤI	0.1469
AL	1.6348
CR	0.0030
FE	2.8780
े र	0.0000
1	0.0816
1G	1.7184
CA	1.9391
NA	0.5906
к	0.2380

JN SCHEDULE ZZZ 3-⁄**`V−87 .26:40 ;

F.S WDS:SI WDS:TI WDS:AL WDS:CR WDS:FE WDS:NI WDS:MN WDS:MG W B:CA WDS:NA WDS:K

+F CORRECTION 15.00 KV 40.00 Dea

o. O	f Itera	tions	0					
	к	[2]	[A]	[F]	[ZAF]	AT OM.	.% WT.;	
-K	0.151	0.997	1.342	0.998	1.338	16.99	20.30	
-к	0.005	1.133	1.033	0.981	1.149	0.30	0.61	
K	0.031	1.029	1.485	0.991	1.516	4.18	4.81	×
:-K	0.000	1.138	1.014	0.952	1.099	0.01	0.02	
:-к	0.156	1.145	1.004	1.000	1.151	7.52	17.96	
-K	0.000	1.137	1.013	1.000	1.153	0.00	0.00	
1-K	0.004	1.162	1.008	1.000	1.172	0.21	0.49	
i-K	0.026	0.985	1.732	0.994	1.397	4.31	4.42	÷
1-K	0.078	1.029	1.044	0.992	1.066	4.88	8.32	
η-K	0.005	1.022	2.272	0.996	2.316	1.39	1.37	×
-к	0.009	1.051	1.065	0.981	1.099	0.62	1.04	
-K	0.186	0.954	2.287	0.999	2.180	59.59	40.67	D *
						TOTAL=	100.00	

- High Absorbance

	RESULTS	
	WT%	
UΖ	43.49	
02	1.02	
203	9.09	
203	0.03	
0	23.09	
0	0.00	
0	0.63	
0	7.36	
0	11.65	
20	1.84	
20	1.25	
TAL=	99.45	

OF OXY = 23.00 #6c (FRA COGE OR YTL)

kb-4.47=1

51	0.0400
ΤI	0.1175
AL	1.6372
CR	0.0030
FE	2.9520
, Ι	0.0000
M PM	0.0814
1G	1.6766
CA	1.9079
NA	0.5461
К	0.2435

F.S WDS:SI WDS:TI WDS:AL WDS:CR WDS:FE WDS:NI WDS:MN WDS:MG W :CA WDS:NA WDS:K

= _CORRECTION 15.00 KV 40.00 Deg

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	'f Itera	tions	8					
	К	[2]	[A]	[F]	[ZAF]	ATOM.	.%. WT.>	
-K	0.154	0.999	1.336	0.998	1.334	17.11	20.36	
-K	0.003	1.136	1.033	0.981	1.152	0.21	0.44	
-K	0.030	1.031	1.479	0.991	1.513	3.92	4.56	÷
-K	0.000	1.140	1.013	0.953	1.102	0.01	0.01	
-К	0.148	1.148	1.004	1.000	1.153	7.10	17.15	
-K	0.000	1.140	1.012	1.000	1.155	0.00	0.00	
-К	0.003	1.165	1.007	1.000	1.174	0.20	0.46	
-К	0.028	0.987	1.716	0.994	1.684	4.58	4.74	÷
-K	0.078	1.031	1.043	0.992	1.068	4.86	8.39	
-K	0.005	1.024	2.251	0.996	2,299	1.18	1.17	÷
·К	0.007	1.053	1.065	0.981	1.101	0.47	0.79	
-К	0.192	0.955	2.265	0.999	2.163	60.37	41.66	D ÷
						TOTAL=	100.03	

- High Absorbance

DE RESULTS WT% 2 44.28 0.73 2 8.62 03 03 0.02 22.05 0.00 1 0.30 7.90 . 11.74 Ū 1.58 0.95 0 · AL= 98.46

OF OXY = 23.00 # 76 (NEAR RIM, NEXT TO BIOTITE)

 $kb = 4.01 \pm 1$

C T	4 7700
51	0.//33
ΤI	0.0836
AL	1.5546
CR	0.0023
FE	2.8210 -
NI	0.0000
MN	0.0773
MG	1.8003
CA	1.9247
NA	0.4689
К	0.1850

3.42:48

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JN SCHEDULE ZZZ 3-NOV-87 13:45:41

EF ∦i	WDS:SI WDS:NA	WDS:T WDS:K	I WDS	:AL (WDS:CR	WDS:FE	WDS:NI	WDS:MN	WDS:MG	ω
ŧF CO	RRECTION	4	15.00	KV	40	.00 Deg				
 . 0f . 1 . 4 . 4<td>Iterati K 0.151 0 0.007 1 0.030 1 0.030 1 0.046 1 0.003 1 0.029 0 0.078 1 0.029 1 0.029 0 0.078 1 0.088 1 0.088 1 0.187 0</td><td>ions 0 [Z] 3.999 1.135 1.031 1.140 1.147 1.139 1.164 1.986 1.030 1.986 1.024 1.955 1.955</td><td>[A] 1.339 1.033 1.484 1.014 1.014 1.012 1.008 1.716 1.043 2.243 1.045 2.298</td><td>[F] 0.990 0.990 0.995 1.000 1.000 1.000 0.992 0.992 0.992 0.993 0.993</td><td>[ZA 8 1.33 2 1.15 1 1.51 5 1.10 3 1.15 3 1.17 4 1.68 3 1.15 3 1.15 3 1.17 4 1.68 3 1.15 3 1.15 3 1.17 4 1.28 3 1.10 4 1.10 5 1.1</td><td>FJ ATOM 7 16.86 2 0.41 8 4.00 4 0.00 3 7.01 4 0.00 4 0.19 4 4.84 7 4.85 3 1.40 3 0.54 4 59.89 TOTAL=</td><td>.% WT.% 20.31 0.86 4.65 0.00 16.90 0.44 5.00 8.35 1.39 0.91 41.24 100.04</td><td>: * * D *</td><td></td><td></td>	Iterati K 0.151 0 0.007 1 0.030 1 0.030 1 0.046 1 0.003 1 0.029 0 0.078 1 0.029 1 0.029 0 0.078 1 0.088 1 0.088 1 0.187 0	ions 0 [Z] 3.999 1.135 1.031 1.140 1.147 1.139 1.164 1.986 1.030 1.986 1.024 1.955 1.955	[A] 1.339 1.033 1.484 1.014 1.014 1.012 1.008 1.716 1.043 2.243 1.045 2.298	[F] 0.990 0.990 0.995 1.000 1.000 1.000 0.992 0.992 0.992 0.993 0.993	[ZA 8 1.33 2 1.15 1 1.51 5 1.10 3 1.15 3 1.17 4 1.68 3 1.15 3 1.15 3 1.17 4 1.68 3 1.15 3 1.15 3 1.17 4 1.28 3 1.10 4 1.10 5 1.1	FJ ATOM 7 16.86 2 0.41 8 4.00 4 0.00 3 7.01 4 0.00 4 0.19 4 4.84 7 4.85 3 1.40 3 0.54 4 59.89 TOTAL=	.% WT.% 20.31 0.86 4.65 0.00 16.90 0.44 5.00 8.35 1.39 0.91 41.24 100.04	: * * D *		
IDE I	ION ABSC RESULTS)rbance								
02 02)3 10 0 10 10 10 10 20 20 7AL=	WT% 43.52 1.43 8.78 0.00 21.73 0.00 0.57 8.33 11.70 1.87 1.10 99.01		· · ·							
# (OF OXY =	23.0	¹⁰ # ⁻	7c (NEAF	2 XTL	RIW)			
SI TI AL CR	క.ర 0.1 1.5 0.0	370 636 774 000	КЪ	= 4.	°14±1					

FE 2.7724 NI 0.0000 MN 0.0730 MG 1.8932 CA 1.9109 IA 0.5521

∼κ 0.2134

13:48:46

N SCHEDULE ZZZ -NOV-87 14:41:40

☞ WDS:SI WDS:TI WDS:AL WDS:CR WDS:FE WDS:NI WDS:MN WDS:MG W ⊣ WDS:NA WDS:K

F CORRECTION 15.00 KV 40.00 Deg

, .	Of Itera	tions	Ø					
	- к	[2]	[A]	[F]	[ZAF]	ATOM	.%. WT.%	
-K	0.142	1.000	1.337	0.998	1.337	15.58	19.10	
-К	0.004	1.137	1.031	0.981	1.152	0.24	0.50	
-К	0.030	1.032	1.480	0.991	1.516	3.97	4.70	÷
-K	0.000	1.142	1.012.	0.954	1.104	0.00	0.00	
К	0.144	1.150	1.003	1.000	1.154	6.82	16.72	
-K	0.000	1.142	1.012	1.000	1.156	0.00	0.00	
-К	0.004	1.167	1.007	1.000	1.175	0.21	0.51	
-К	0.026	0.988.	1.723	0.994	1.694	4.31	4.53	×
-K	0.074	1.032	1.042	0.992	1.068	4.56	7.99	
-К	0.005	1.025	2.262	0.996	2.313	1.30	1.31	÷
-К	0.009	1.054	1.062	0.981	1.099	0.64	1.09	
-K	0.205	0.956	2.226	0.999	2.128	62.36	43.67 [) X
						TOTAL=	100.12	

= 85 (NERF XTL RIM)

+ b- 4.72 ± 1

- High Absorbance

IDE	RESULTS
	WT%
02	40.92
02	0.83
2	8.87
З	0.01
υ	21.50
0	0.00
0	0.66
0	7.55
0	11.19
20	1.76
20 1	1.31
TAI =	94.40

OF OXY = 23.00

SI 6.5763 ΤI 0.1002 ΑL 1.6808 CR 0.0007 FE 2.8906 NI 0.0000 MN 0.0902 MG 1.8083 CA 1.9264 Ά 0.5483 0.2693

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SAMPLE LOCATION (OR STD)...SOUTH CAROLINA, U.S.A. DATE OF ANALYSIS.....21/12/87 LAB IDENTIFICATION CODE....LH2 IRRADIATION CODE.....MM007-16 DISKS (ALIGNMENT/CALCS)..../ DATAFILE.....A:LH2V.A1A MINERAL ANALYZED.....APATITE ANALYSIS BY / NOTES PAGE ≠.VM /

INDIVIDUAL GRAIN DATA

PRISTAL	<u>85</u>	9.I	ALADE	RATIO	ERC S		(T) - FE	1 <u>GE (117)</u>
			4 14 14 14 14 14 14 14 44 14 14 14 14 14				י דוריו ג'רופי מיז פט ימי פין ערבי בערביין ביו דוריפי בערפי יובי או איין איין איין איין איין איין איין	
	246	239			8.01E+05	\$.03E+6E	7	

SUMMARY OF STATISTICS

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-	Variance of Square root of NS = 0.8172
-	Variance of Square root of NI = 0.4396
-	Correlation coefficient (R for NS vs. NI) = 0.9069
_	Chi ² = 8.740728 with 18 degrees of freedom. TEST PASS

AGE CALCULATION

-	Glass dosimeter	=	SRM614		
-	ND counted	=	1200	92.4 ±	8.9 MA
-	RHOD D for glass	=	1.980E+04 ± 0.572E+03		
-	Area of 1 QUAD	=	1.5625E-06		
-	Pooled NS/NI	=	0.823 ± 0.071	MEAN	AGE
-	Mean NS/NI	=	0.795 ± 0.046		
				89.3 ±	6.5 MA

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INDIVIDUAL GRAIN DATA

CERSTAL	53	21 	WIAIS	IATI,	33. 3			1
							 International states of the formation of the states of the	Pro I. S. C.
	543	496			8.47E-15	E.86E+06	: 0	

SUMMARY OF STATISTICS

AGE CALCULATION

-	Variance	of Squa	re root	of.	NS		= 2.17	67	
-	Variance	of Squa	re root	of	NI		= 1.4	849	
-	Correlat	ion coef	ficient	(R	for NS	vs.	NI) = 0.	9036	
-	Chi² =	19.36049	with	27	degrees	οf	freedom.	TEST	PASS

-	Zeta factor used	=	11420.02 380.0	POOLED AGE	
-	Glass dosimeter	=	SRMG14		
-	ND counted	Ξ	1200	123 ± 9.4 MA	
-	RHOD D for glass	=	1.975E+04 = 0.570E+03		
-	Area of 1 QUAD	=	1.21E-06		
-	Pooled NS/NI	=	1.105 ± 0.068	MEAN AGE	=
-	Mean NS/NI	=	1.085 ± 0.062		
				121 - 8.7 MA	

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SAMPLE LOCATION (OR STD)...LIBERTY HILL, SOUTH CAROLINA DATE OF ANALYSIS......01/04/88 LAB IDENTIFICATION CODE....LH5 IRRADIATION CODE.....MM008-3 DISKS (ALIGNMENT/CALCS)..../ DATAFILE.....A:LH5J.A1A MINERAL ANALYZED.....APATITE ANALYSIS BY / NOTES PAGE =.JM /

INDIVIDUAL GRAIN DATA

2127.1	1 53	51	QUALS	EATEC	320 3		. (01 995	40 <u>5 (M</u> 4
	• •	1.6	(†	104		\ 37∓ ⊥ 8F	.	112.3 - 37.2
2	::							11343 - 5440
-	•••	10	- ī	0.000		3.132*.0	÷	12.4 I 13.4
	11	61	25	6.868		2.053+06		- 22.2 ± 12.2
,	\$	15	0:	3 129	1 317.67	1 5 17	-	50 4 4 55 1
1	1	1.2		0,701				
	:	÷ 1	10	3.14:	····		2	5041 I 1641
2	•		25		<u> </u>	1 . C . S		94.T ± 35.9
:			ō.;				3	10 4 - 55 2
2		::	20	4.0.0				02.4 4 60.4
:	.:		21			1		<i>dini</i> ± 30.€
:		0.5	2.5	6.646	5.117.65		11	70.5 ÷ 12.8
		0.5		- 625			::	
	÷1	÷ 1	<u> </u>		0.012-00	2.2.2.2.2	12	(3,4 I 14,1
	29	21 C	25	0.810	2.22 - 01	1.94E-03	13	89.4 ± 10.6
	÷ •		2.5	0 212	2 642-95	7 237.65	12	100 1 + 30 3
:7	73	41	67	6.004		1 1 2 2 1 0 0 1 1 2 7 1 0 0	10	
11	-	÷4			5.201+00	1.1.1.1.1.1.1	11	51.4 I 41.0
14	11	1	25	0.917	0.642-05	0.35E-65	-	100.8 ± 42.1
;	• •	1.5	6.2	1.00	1 167-167	1 101.45	3	101 3 4 10 1
	. *	15			7.301-Ve 0.10 1 .01	1.3927VE		
	:	10	20	5.601	2.301+00	÷.201*00	2	00.1 I 22.9
	-		25	0.000	1.345+05			59.8 ± 15.5
	č.		0.5		- 11 - 14			101 1 4 50 5
::		÷.	60					
.:	<u>1</u> •	1	2.0	50	2.111-01	じんし ユモルマ	11	93.5 I 30.5
	<u><u></u></u>		1.5					24.4 ± 20.4
		5.0	6.5		1 127.08	2 207-04	÷.	27 2 4 20 2
22				2.01		0.000-00		10.0 1 10.0
		1	20		4.035+05	3.542+00	ţ.	139.0 ± 05.2
	15	42	25	1.066	1.191406	1 3726	21	109.9 + 24.6
7)		57	55		0 010 05		1	10 E ± 11 F
73	÷ :	93	-9		0.011100		13	Jui 1 1 1
11	11	19	11		3.64E+05	6.18E+85	11	63.8 ± 24.2
2.2	• •	3 F	9.5	5 221	1 635-05	1 127-05		102 6 + 08 1
		ē.		5 203	1 1001 00			00 1 1 01 2
÷:		<u> </u>	-	C. 100	1.105100		.1	00.1 I L1.0
- 13	12	5.C	25	0.800	5.19 2- 05	6.61E+15		88.0 ± 23.5
6.2	· -	e 1	25	1 996	0.017.05	2 012.03	1.2	110 7 4 11 5
72	7.5	7:		1,200			1-	
		<u>.</u>		9.116	0.002-00		45	<u></u>
	5.30	651			F . F . F . F	5 145175		

SUMMARY OF STATISTICS

-	Variance of	Square root of	NS = (0.9167
-	Variance of	Square root of	NI =	1.3831
-	Correlation	coefficient (R	for NS vs. NI) =	0.8011
-	$Chi^2 = 26.1$	10124 with 29	degrees of freedom	. TEST PASS

-	Zeta factor used	=	11420.0= 380.0	POOLE	ED AGE	=
	Glass dosimeter	=	SRM614			
-	ND counted	Ξ	1225	82.1 ±	6.0 MA	
-	RHOD D for glass	=	$1.940E+04 \pm 0.554E+03$			
-	Area of 1 QUAD	=	1.21E-06			
-	Pooled NS/NI	=	0.746 = 0.044	MEAN	AGE	=
-	Mean NS/NI	=	0.797 ± 0.048			
				87.7 ±	6.5 MA	

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INDIVIDUAL GRAIN DATA

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			E.742-08	foffer f	

SUMMARY OF STATISTICS

-	Variance of	Square root	ωf	NS		= 11.9543
-	Variance of	Square root	01	NI		= 12.7810
-	Correlation	coefficient	(R	for NS v	s. NI;	= 0.3070

- Chi? = 22.01207 with 3 degrees of freedom. TEST FAIL -

AGE CALCULATION

APPENDIX 2

-**-** -

Age spectra of analyzed minerals

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APPENDIX 3

Raw data for fission track age calculation

SAMPLE LOCATION (OR STD)...LIBERTY HILL, SOUTH CAROLINA LAB IDENTIFICATION CODE....LH1 IRRADIATION CODE.....MM008-2 DISKS (ALIGNMENT/CALCS).... / DATAFILE.....A:LH1V.A1A MINERAL ANALYZED.....APATITE ANALYSIS BY / NOTES PAGE #.VM /

INDIVIDUAL GRAIN DATA

CRYSTAL	¥3	ЯЦ	SUADS	<u> RATIQ</u>	SHE S	<u>24011</u>	11 103	AGE (MA)
1515 Charles exercited and the extra contraction of the carter carter of the carter carter of the carter carte	an - 1946 and - 1916 and 1916	Construction of the construction of the second s	e i e i e i e i e i e i e i e i e i e i			1		
	487	5:1			6.298405	7.13E+05	12	

SUMMARY OF STATISTICS

	Variance of	Square root	of	NS	=	0.9111
-	Variance of	Square root	of	NI	=	1.1165
_	Correlation	coefficient	(R	for NS vs	5. NI) =	= 0.8903

- Chi² = 10.8229 with 23 degrees of freedom. TEST PASS

AGE CALCULATION

- Zeta factor used = 11420.01 380.0

- Glass dosimeter = SRM614

POOLED AGE =

- - . . .

	- RHOD D for glass	=	1.940E+04 ± 0.004E+03		
	- Area of 1 QUAD	=	1.21E-06		
	- Pooled NS/NI	Ξ	0.882 ± 0.057	MEAN	AGE
-	- Mean NS/NI	=	0.909 ± 0.044		
an article				99.9 ±	6.6 MA

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